

Zhao, L., Zheng, X., Jiao, J., Han, F., Zhou, B., Fan, P., Jia, L., and Ma, Y., 2023, Geochronology, petrogenesis, and magmatic oxidation state of the Mangling intrusive complex, Northern Qinling Belt, Central China: Implications for magma fertility and tectonic setting: GSA Bulletin, <https://doi.org/10.1130/B36714.1>.

Supplemental Material

Supplemental Text. Analytical Methods.

Figure S1. Cathodoluminescence images of representative zircons for U-Pb dating and Lu-Hf isotope analyses for the Mangling intrusive complex.

Figure S2. Chondrite-normalized rare earth element patterns of analyzed zircons for the Mangling intrusive complex.

Figure S3. Field photographs showing relationships of the granitic rocks of the Mangling intrusive complex.

Figure S4. Harker diagrams of selected major and trace elements against SiO₂ for the Mangling intrusive complex.

Table S1. Zircon U-Pb dating results for the dioritic and granitic rocks of the Mangling intrusive complex.

Table S2. Whole-rock major- (wt%), trace- (ppm), and rare earth (ppm) element compositions for the dioritic and granitic rocks of the Mangling intrusive complex.

Table S3. Zircon trace element compositions (ppm) for the dioritic and granitic rocks of the Mangling intrusive complex.

Table S4. Zircon Lu-Hf isotopic data for the dioritic and granitic rocks of the Mangling intrusive complex.

Table S5. Summary of geochronological data for the Mangling intrusive complex.

Table S6. Zircon trace element compositions (ppm) for the mineralized granitic rocks of porphyry molybdenum deposits in Central and NE China.

1 ANALYTICAL METHODS

2 Zircon U-Pb Dating

3 By using conventional heavy liquid and magnetic techniques, zircons were
4 separated from five representative dioritic (biotite diorite and biotite diorite enclave)
5 and granitic rock samples (medium- to fine-grained and fine-grained monzogranite
6 and K-feldspar granite) collected from the Mangling intrusive complex and then
7 handpicked under a binocular microscope. The selected zircon grains were mounted
8 in epoxy resin and polished to expose the crystal centers to observe their internal
9 structure under cathodoluminescence (CL) imaging that was carried out at the Testing
10 Center, Tuoyan Analytical Technology Co. Ltd. (Guangzhou, China) using a TESCAN
11 MIRA3 field-emission scanning electron microprobe (FE-SEM).

12 In-situ U-Pb dating of zircon grains were performed at the Nanjing FocuMS
13 Technology Co. Ltd. (Nanjing, China), using an Agilent Technologies 7700x
14 quadrupole inductively coupled plasma-mass spectrometry (ICP-MS) coupled with a
15 Resonetics RESolution LR (193 nm ArF excimer) laser ablation system. During
16 analysis, the laser ablation system has a spot diameter of 33 μm with a background
17 acquisition of approximately 20 s followed by 40 s of data acquisition. 91500 zircon
18 was used as external standard to correct instrumental mass discrimination and
19 elemental fractionation during the ablation, whilst GJ-1 and Plešovice zircons were
20 treated as quality control for geochronology. Trace elements of zircons were external
21 calibrated against NIST SRM 610 with Si as internal standard. The detailed analytical

22 technique is described by [Li et al. \(2011\)](#). Isotopic ratios of U-Pb dating were
23 performed using ICPMSDataCal ([Liu et al., 2008](#)), and age calculation and plotting of
24 concordia diagrams were performed using Isoplot/Ex 3.0 ([Ludwig, 2003](#)).

25 **Whole-Rock Major and Trace Elements**

26 Whole-rock major and trace element analyses were performed at the Key
27 Laboratory of Western China's Mineral Resources and Geological Engineering,
28 Ministry of Education and the State Key Laboratory of Isotope Geochemistry,
29 Guangzhou Institute of Geochemistry, Chinese Academy of Sciences with sample
30 names as "ML21-". Before analysis, following procedures should be performed for all
31 the samples: (1) powder them to less than 200 meshes, (2) add Lithium Borate Flux
32 (50%–50% $\text{Li}_2\text{B}_4\text{O}_7$ – LiBO_2), and (3) mix them well and fuse in an auto fluxer
33 between 1050 °C and 1100 °C. The X-ray fluorescence (XRF) spectrometry was used
34 to analyze major elements with analytical precision better than 1%. For trace elements,
35 HNO_3 + HF acid was used to digest whole-rock powder (50 mg) in steel-bomb coated
36 Teflon beakers for two days in order to assure complete dissolution of the refractory
37 minerals under high pressure. Then ICP-MS was used to analyze the solution, with
38 analytical precision better than 5% for most trace elements.

39 **Magmatic Oxidation State and Crystallization Temperature**

40 Zircon $\text{Ce}^{4+}/\text{Ce}^{3+}$ ratio and Ti-in-zircon thermometer were used to evaluate the
41 magmatic oxidation state and crystallization temperatures of igneous rocks ([Shen et](#)
42 [al., 2015](#); [Zhang et al., 2020](#)), with detailed calculations in [Ballard et al. \(2002\)](#) and

43 [Ferry and Watson \(2007\)](#), respectively. Moreover, the zircon Eu anomalies (Eu/Eu^*)
44 were calculated based on normalized values of Sm and Gd concentrations by a
45 conventional method, similar to the equation of whole-rock Eu anomalies. For the
46 magmatic oxygen fugacity ($f\text{O}_2$), it is commonly estimated as relative
47 fayalite-magnetite-quartz buffer (ΔFMQ) based on the incorporation of cerium into
48 zircon, Ti-in-zircon temperature, and water content using the equation of [Smythe and](#)
49 [Brenan \(2016\)](#). In this study, the Ti-in-zircon thermometer and zircon $\text{Ce}^{4+}/\text{Ce}^{3+}$ and
50 Eu/Eu^* ratios were calculated to evaluate the magma nature of the Mangling intrusive
51 complex using Geo- $f\text{O}_2$ software proposed by [Li et al. \(2019\)](#), with assumed SiO_2 and
52 TiO_2 activity to be 1 and 0.7, respectively, because all samples from the Mangling
53 intrusive complex of this study contain primary magmatic quartz and accessory
54 mineral of sphene ([Hao et al., 2021](#)). For the water content, we use 4.0 wt% for the
55 biotite diorite enclave (BDE), biotite diorite (BD), and medium- to fine-grained
56 monzogranite (MMG) based on occurrence of hornblende in these samples ([Hao et al.,](#)
57 [2021](#)) and 3.0 wt% for the fine-grained monzogranite (FMG) and K-feldspar granite
58 (KG) because the hornblende begins to crystallize when the silicate melts contain
59 more than 3.0 wt% H_2O ([Burnham, 1979](#)).

60 **Zircon Lu-Hf Isotopes**

61 Zircon Lu-Hf isotope analyses were performed at the Nanjing FocuMS
62 Technology Co. Ltd. (Nanjing, China) using a Nu Plasma II multi-collector
63 (MC)-ICP-MS (Wrexham, Wales, UK), coupled to a 193 nm excimer laser ablation

64 system (RESOLUTION LR). Each acquisition incorporated 20 s background (gas blank),
65 followed by spot diameter of 50 μm at 9 Hz repetition rate for 40 s. Using exponential
66 correction for mass bias, the measured isotopic ratios of $^{176}\text{Hf}/^{177}\text{Hf}$ were normalized
67 to $^{179}\text{Hf}/^{177}\text{Hf} = 0.7325$. The measured $^{173}\text{Yb}/^{171}\text{Yb}$ and the natural ratio of 1.13268
68 (Chu et al., 2002) were used to calculate the mass bias factor of Yb. In IsotopeMaker,
69 the natural ratios are $^{176}\text{Yb}/^{173}\text{Yb} = 0.79381$ (Segal et al., 2003) and $^{176}\text{Lu}/^{175}\text{Lu} =$
70 0.02656 (Wu et al., 2006). A more detailed analytical technique can be found in Zhang
71 et al. (2015). The initial $^{176}\text{Hf}/^{177}\text{Hf}$ ratios were calculated according to the measured
72 $^{176}\text{Lu}/^{177}\text{Hf}$ ratios and the ^{176}Lu decay constant of $1.867 \times 10^{-11} \text{ year}^{-1}$ (Söderlund et al.,
73 2004), whereas $\varepsilon_{\text{Hf}}(t)$ values were calculated with chondritic values of $^{176}\text{Hf}/^{177}\text{Hf} =$
74 0.0336 and $^{176}\text{Lu}/^{177}\text{Hf} = 0.282785$ (Bouvier et al., 2008). The present-day $^{176}\text{Hf}/^{177}\text{Hf}$
75 $= 0.28325$ and $^{176}\text{Lu}/^{177}\text{Hf} = 0.0384$ (Griffin et al., 2004) were used to define the
76 depleted mantle line. Single-stage Hf model ages (T_{DM}) were calculated with
77 $^{176}\text{Lu}/^{177}\text{Hf} = 0.0384$ (Griffin et al., 2004) relative to the depleted mantle that is
78 assumed to have linear isotopic growth ranging between $^{176}\text{Hf}/^{177}\text{Hf} = 0.279718$ at
79 4.55 Ga and 0.283250 at present, whereas two-stage Hf model ages (T_{DM}^{C}) were
80 calculated assuming a mean value of $^{176}\text{Lu}/^{177}\text{Hf} = 0.015$ for the average continental
81 crust (Griffin et al., 2002).

82 REFERENCES CITED

83 Ballard, J.R., Michael, P., and Campbell, H.I., 2002, Relative oxidation states of
84 magmas inferred from Ce(IV)/Ce(III) in zircon: Application to porphyry copper
85 deposits of northern Chile: Contributions to Mineralogy and Petrology, v. 144, p.

86 347–364, <https://doi.org/10.1007/s00410-002-0402-5>.

87 Bouvier, A., Vervoort, J.D., and Patchett, P.J., 2008, The Lu–Hf and Sm–Nd isotopic
88 composition of CHUR: Constraints from unequilibrated chondrites and
89 implications for the bulk composition of terrestrial planets: *Earth and Planetary
90 Science Letters*, v. 273, p. 48–57, <https://doi.org/10.1016/j.epsl.2008.06.010>.

91 Burnham, C.W., 1979, Magmas and hydrothermal fluids, *in* Barnes, H.L., ed.,
92 *Geochemistry of hydrothermal ore deposits*: New York, JohnWiley, p. 71–136.

93 Chu, N.C., Taylor, R.N., Chavagnae, V., Nesbitt, R.W., Boela, R.M., Milton, J.A.,
94 German, C.R., Bayon, G., and Burton, K., 2002, Hf isotope ratio analysis using
95 multi-collector inductively coupled plasma mass spectrometry: An evaluation of
96 isobaric interference corrections: *Journal of Analytical Atomic Spectrometry*, v.
97 17, p. 1567–1574, <https://doi.org/10.1039/b206707b>.

98 Ferry, J.M., and Watson, E.B., 2007, New thermodynamic models and revised
99 calibrations for the Ti-in-zircon and Zr-in-rutile thermometers: *Contributions to
100 Mineralogy and Petrology*, v. 154, p. 429–437,
101 <https://doi.org/10.1007/s00410-007-0201-0>.

102 Griffin, W.L., Belousova, E.A., Shee, S.R., Pearson, N.J., and O’Reily, S.Y., 2004,
103 Archean crustal evolution in the northern Yilgarn Craton: U–Pb and Hf-isotope
104 evidence from detrital zircons: *Precambrian Research*, v. 131, p. 231–2821,
105 <https://doi.org/10.1016/j.precamres.2003.12.011>.

106 Griffin, W.L., Wang, X., Jackson, S.E., Pearson, N.J., O’Reilly, S.Y., Xu, X., and
107 Zhou, X., 2002, Zircon chemistry and magma mixing SE China: In-situ analysis

108 of Hf isotopes. Tonglu and Pingtan igneous complexes, *Lithos*, v. 61, p. 237–269,
109 [https://doi.org/10.1016/S0024-4937\(02\)00082-8](https://doi.org/10.1016/S0024-4937(02)00082-8).

110 Hao, H.D., Campbell, I.H., Cooke, D.R., Nakamura, E., and Sakaguchi, C., 2021,
111 Geochronology, petrogenesis and oxidation state of the Northparkes igneous suite,
112 New South Wales, Australia: Implications for magma fertility: *Economic Geology*,
113 v. 116, p. 1161–1187, <https://doi.org/10.5382/econgeo.4825>.

114 Li, H., Zhang, H., Ling, M.X., Wang, F.Y., Ding, X., Zhou, J.B., Yang, X.Y., Tu,
115 X.L., and Sun, W.D., 2011, Geochemical and zircon U–Pb study of the
116 Huangmeijian A-type granite: Implications for geological evolution of the Lower
117 Yangtze River belt: *International Geology Review*, v. 53, p. 499–525,
118 <https://doi.org/10.1080/00206814.2010.496202>.

119 Li, W.K., Cheng, Y.Q., and Yang, Z.M., 2019. Geo- fO_2 : Integrated software for
120 analysis of magmatic oxygen fugacity: *Geochemistry, Geophysics, Geosystems*, v.
121 20, p. 2542–2555, <https://doi.org/10.1029/2019GC008273>.

122 Liu, Y.S., Hu, Z.C., Gao, S., Gunther, D., Xu, J., Gao, C.G., and Chen, H.H., 2008, In
123 situ analysis of major and trace elements of anhydrous minerals by LA–ICP–MS
124 without applying an internal standard: *Chemical Geology*, v. 257, p. 34–43,
125 <https://doi.org/10.1016/j.chemgeo.2008.08.004>.

126 Ludwig, K.R., 2003, User’s Manual for Isoplot 3.00. A Geochronological Toolkit for
127 Microsoft Excel: Berkeley Geochronology Center Special Publication 4, 71 p.

128 Segal, I., Halicz, L., and Platzner, I.T., 2003, Accurate isotope ratio measurements of
129 ytterbium by multiple collection inductively coupled plasma mass spectrometry

130 applying erbium and hafnium in an improved double external normalization
131 procedure: *Journal of Analytical Atomic Spectrometry*, v. 18, p. 1217–1223,
132 <https://doi.org/10.1039/b307016f>.

133 Shen, P., Hattori, K., Pan, H.D., Jackson, S., and Seitmuratova, E., 2015, Oxidation
134 condition and metal fertility of granitic magmas: Zircon trace-element data from
135 porphyry Cu deposits in the Central Asian Orogenic Belt: *Economic Geology*, v.
136 110, p. 1861–1878, <https://doi.org/10.2113/econgeo.110.7.1861>.

137 Smythe, D.J., and Brenan, J.M., 2016, Magmatic oxygen fugacity estimated using
138 zircon-melt partitioning of cerium: *Earth and Planetary Science Letters*, v. 453, p.
139 260–266, <https://doi.org/10.1016/j.epsl.2016.08.013>.

140 Söderlund, U., Patchett, P.J., Vervoort, J.D., and Isachsen, C.E., 2004, The ^{176}Lu
141 decay constant determined by Lu–Hf and U–Pb isotope systematics of
142 Precambrian mafic intrusions: *Earth and Planetary Science Letters*, v. 219, p.
143 311–324, [https://doi.org/10.1016/S0012-821X\(04\)00012-3](https://doi.org/10.1016/S0012-821X(04)00012-3).

144 Wu, F.Y., Yang, Y.H., Xie, L.W., Yang, J.H., and Xu, P., 2006, Hf isotopic
145 compositions of the standard zircons and baddeleyites used in U–Pb
146 geochronology: *Chemical Geology*, v. 234, p. 105–126,
147 <https://doi.org/10.1016/j.chemgeo.2006.05.003>.

148 Zhang, L., Ren, Z.Y., Xia, X.P., Li, J., and Zhang, Z.F., 2015, IsotopeMaker: A
149 Matlab program for isotopic data reduction: *International Journal of Mass
150 Spectrometry*, v. 392, p. 118–124, <https://doi.org/10.1016/j.ijms.2015.09.019>.

151 Zhang, Y., Tian, J., Hollings, P., Gong, L., Albuero, I., Berador, A.E., Francisco, D.G.,

152 Li, J., and Chen, H.Y., 2020, Mesozoic porphyry Cu–Au mineralization and
153 associated adakite-like magmatism in the Philippines: Insights from the giant
154 Atlas deposit: *Mineralium Deposita*, v. 55, p. 881–900,
155 <https://doi.org/10.1007/s00126-019-00907-2>.



Figure S1. Cathodoluminescence images of representative zircons for U-Pb dating and Lu-Hf isotope analyses for the Mangling intrusive complex.

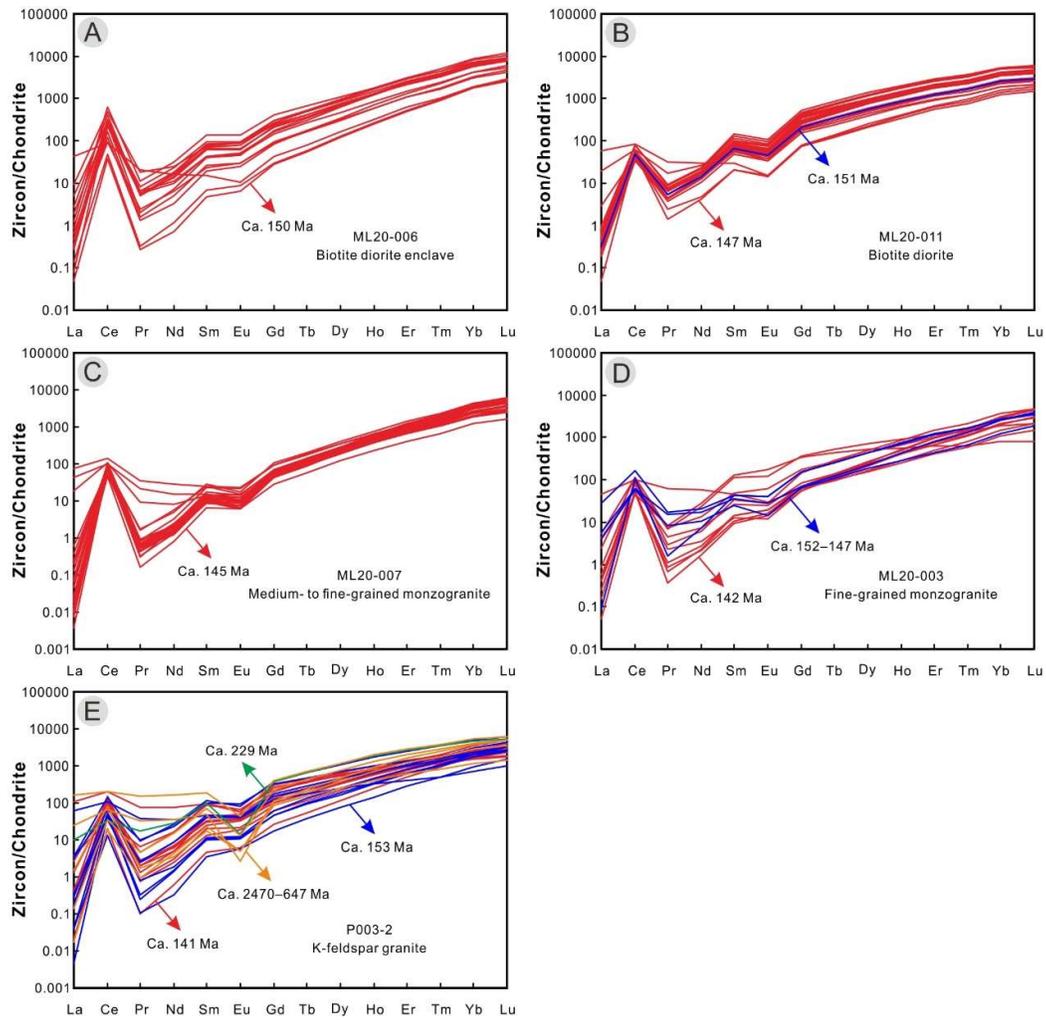


Figure S2. Chondrite-normalized rare earth element patterns of analyzed zircons for the Mangling intrusive complex. Normalizing values are from [Sun and McDonough \(1989\)](#).

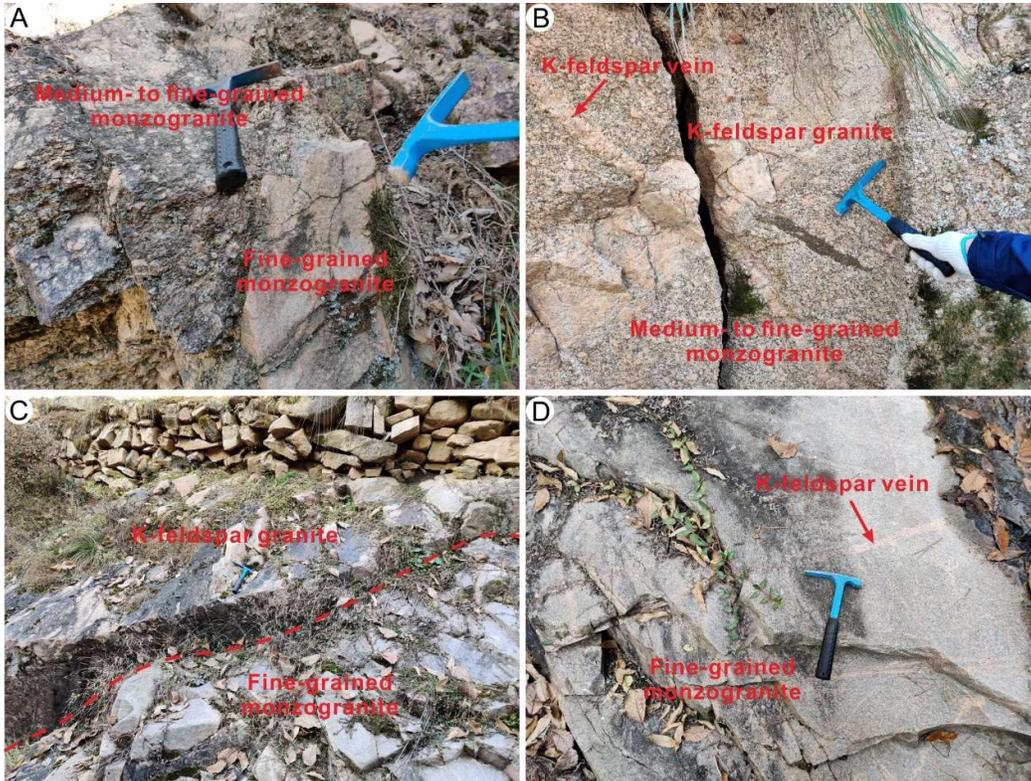


Figure S3. Field photographs showing relationships of the granitic rocks of the Mangling intrusive complex. (A) Medium- to fine-grained monzogranite was intruded by the fine-grained monzogranite. (B) Medium- to fine-grained monzogranite was crosscut by the K-feldspar granite and a K-feldspar vein. (C) K-feldspar granite intrudes into the fine-grained monzogranite. (D) Fine-grained monzogranite was crosscut by K-feldspar veins.

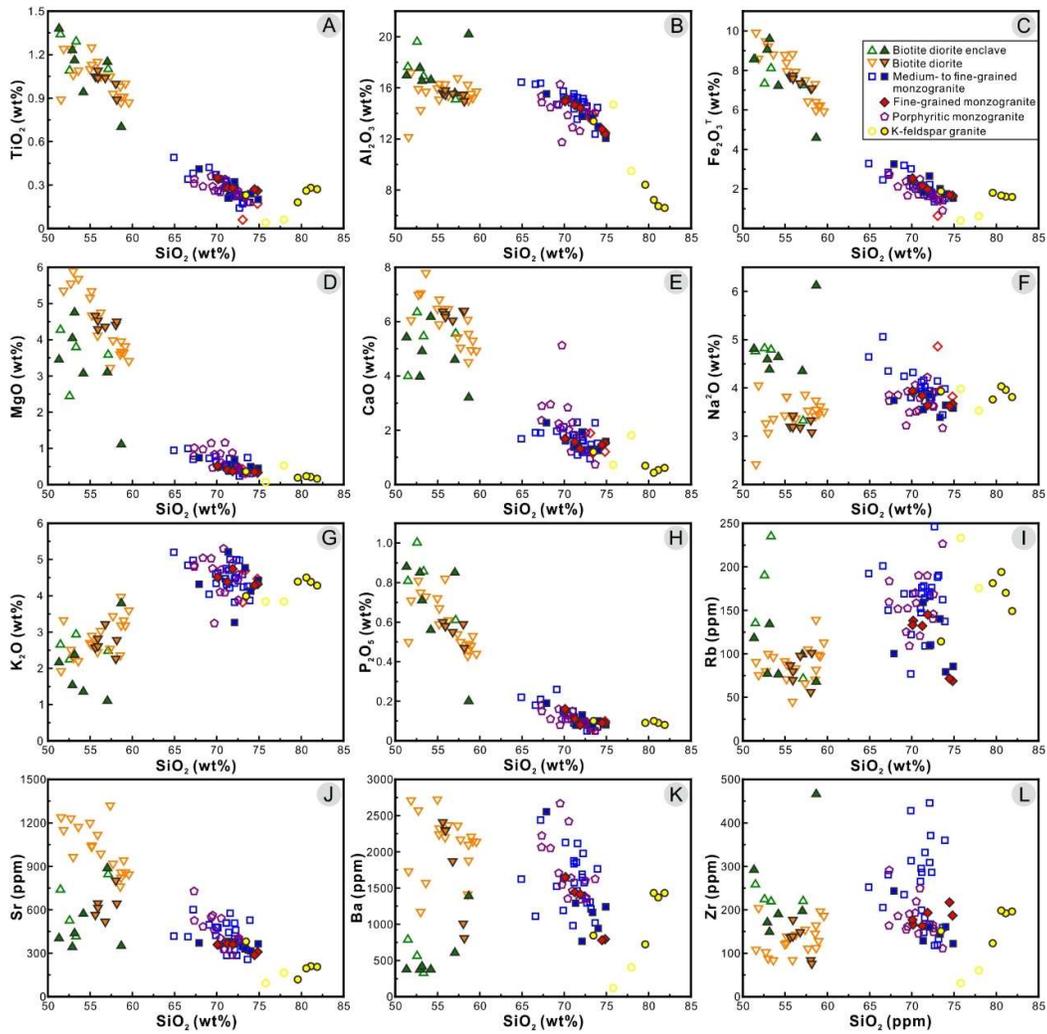


Figure S4. Harker diagrams of selected major and trace elements against SiO_2 for the Mangling intrusive complex. Published data of Mangling dioritic and granitic rocks are same as Fig. 5.

References

Sun, S.S., and McDonough, W.F., 1989, Chemical and isotopic systematics of oceanic basalts: implications for mantle composition and processes, In Saunders, A.D., Norry, M.J., eds., *Magmatism in Ocean Basins*: Geological Society, London,

Special Publications 42, p. 313–345,

<https://doi.org/10.1144/GSL.SP.1989.042.01.19>.