1 ADDITIONAL DETAILS ON METHODS

2 Detailed Methods on the Preparing of the Artificial Sandstone Blocks

We casted mixtures of water, Portland cement and fine sand into boxes of $0.56 \times 0.56 \times 0.3 \text{ m}^3$ to 3 4 create homogeneous artificial sandstone of 0.2 mm grainsize, varying the cement to sand ratio 5 for two different tensile strengths σ_l (Sklar and Dietrich, 2001; Table S1). The block's tensile 6 strengths were measured from samples using a Split-Hopkinson pressure bar (Zhou et al., 2011). 7 For the impact abrasion experiments we placed the blocks on sand bags to reduce shacking and 8 confined them with strapped wood boards to avoid shock wave reflection and block cracking 9 from the free lateral surfaces (Fig. 1B). This procedure served as approximation for lateral 10 infinite massive bedrock (i.e. for constant horizontal lithostatic pressure), at least for the 11 relatively small ratios of impactor size to block size of the experiments.

12

13 Detailed Methods on Effective Impact Energy Calculation of the Drop Abrasion Experiments

Having vertical and horizontal scale bars attached to the target blocks (Fig. 1B), we tracked the impacting grains by means of a laterally positioned time-laps camera with 100 frames per second, and calculated effective kinetic impact energies, ε_{kin} , responsible for block abrasion:

17
$$\varepsilon_{kin} = \varepsilon_{imp} - \varepsilon_{reb} = (M_i g H_d) - (\varepsilon_{kin,reb} + \varepsilon_{rot,reb}), \qquad (1)$$

18 with ε_{imp} being the kinetic impactor energy at the time of impact, ε_{reb} being its rebounding energy 19 afterwards, M_i the impactor mass, g acceleration due to gravity, and H_d the impactor's drop 20 height above the target block (Fig. S1A). Based on time-laps camera observations (Fig. S1B and 21 C), ε_{reb} after the first impact was calculated as the combination of rebound kinetic energy $\varepsilon_{kin,reb}$ 22 and rotational energy $\varepsilon_{rot,reb}$:

23
$$\varepsilon_{kin,reb} = (M_i g H_{reb}) + \left(0.5M_i \left(\frac{dx}{dt}\right)^2\right), \qquad (2)$$

24
$$\varepsilon_{rot,reb} = 0.5 M_i \left(\frac{D_i}{2}\right)^2 v_{rot,reb}^2,$$
 (3)

with maximum impactor rebound height H_{reb} , first rebound hop distance and time dx and dt, the impactor's rebound rotation velocity $v_{rot,reb}$, and the impactor's diameter D_i (Fig. S1A). We could mostly avoid second impacts on the blocks, but the few happening did neither contribute notable impact energies, nor measurable total abrasion volumes V_a or fragment volumes V_{frag} . Further, we assumed dominance of plastic deformation (abrasion) for the brittle target blocks and neglected any energy loss due to elastic energy radiation (Farin et al., 2015).

32 Detailed Methods on Total Slab Abrasion Measurements

- 33 After each drop experiment and after removing abraded sand and fragments from the target block
- 34 (Fig. 2A), we placed sub-millimeter-calibrated CHI photogrammetry scale bars (SfM bars) on the
- 35 block's surface, outside the impact area. Then we took five high-resolution photos
- 36 (42 Megapixels) around and atop the block using a SONY α7RII DSLR camera with a fixed
- 37 55 mm lens (Fig. 1B). Applying structure from motion, *SfM*, and multi view stereo
- 38 photogrammetry in Agisoft Photoscan (Agisoft Photoscan, 2018) resulted in sequential sub-
- 39 millimeter resolved and accurate digital elevation models, *DEMs*. Vertically differencing these
- 40 individual epoch *DEMs* in cloudcompare (CloudCompare, 2019), we calculated total abrasion
- 41 volumes V_a (i.e. the sum of wear and macro-abrasion V_{frag}) for all experiments.

43 Detailed Methods on Compilation of Single Grain Impact Abrasion Data from the Literature 44 We compiled published bulk abrasion data from grain drop experiments and from abrasion mill 45 runs (Table S2), all of which didn't report on observed fragment production. For the abrasion 46 mill data, we used an updated version of the total load model (Lamb et al., 2008; neglecting the 47 cover term) to calculate mean effective impact energies, ε_{kin} , and mean eroded volumes, V_a , per 48 single impact in the mills. Thereby, instead of estimating the near-bed volumetric sediment 49 concentration, c_b , from total volumetric sediment flux (Lamb et al., 2008 eq. 18), we based its 50 calculation on the known total sediment mass, M_m , in the mills:

51
$$M_m = \int_0^{H_m} \rho_s \ c(z) \ r_w \ L_c \ dz,$$
 (4)

with H_m being the mill's water depth, ρ_s is the grain sediment density, c(z) is the volumetric sediment concentration at height *z* above the bedrock disk on the mill's bottom, r_w is the total radial width of the bedrock disk, and L_c is the mean circumferential travel distance of a grain in the abrasion mill. By transforming eq. 4 and adopting the thickness of the bedload layer H_b (Sklar and Dietrich, 2004 eq.12), we obtained:

57
$$\frac{M_m}{\rho_s r_w L_c} = c_b \int_0^{H_m} \frac{c(z)}{c_b} dz = c_b \left[H_b + \int_{H_b}^{H_m} \frac{c(z)}{c_b} dz \right],$$
(5)

58 and hence:

59
$$c_b = \frac{M_m}{\rho_s \, r_w \, L_c \left[H_b + \int_{H_b}^{H_m c(z)} dz \right]},\tag{6}$$

60 wherein the second term in the square bracket is the Rouse profile (Lamb et al., 2008 eq.26). For 61 the associated Rouse parameter, we used a proportionally constant of 2.5 and a Stokes number 62 of 90 to account for viscous dampening (generally following Scheingross et al., 2014). Mean 63 kinetic impact energies, ε_{kin} , and mean eroded volumes, V_a , of single grain impacts were then 64 calculated as follows:

$$65 \qquad \varepsilon_{kin} = 0.5 M_i \varepsilon_{imp,eff}^2, \tag{7}$$

$$66 V_a = \frac{V_d}{I_r A_d} = \frac{V_d}{\left(\frac{0.36c_b \,\varepsilon_{imp,eff}}{V_i}\right) * A_d}, (8)$$

67 with V_d being the reported total abraded disk volume, I_r is the particle impact rate per disk area, 68 A_d (Lamb et al., 2008 eq.12), V_i is the mean single impacting grain volume, and $\varepsilon_{imp,eff}$ is the 69 effective, vertical grain impact energy accounting for turbulent fluctuations (Lamb et al., 2008 70 eq.35). Principally, abrasion volume scales with excess impact energy (vertical component of the 71 impact velocity minus a material-specific threshold), but this reduction is neglected here due to 72 generally low thresholds and for the case of mean values for V_a (Neilson, 1968; Engel, 1978; 73 Sklar and Dietrich, 2004).

74

76 Detailed Methods on Compilation of Potential Grain Impact Energy in Bedrock Rivers

The potentially median transportable grain size D_{50} (the grainsize of 50% volumetric abundance in a grainsize distribution) for a given river bed slope *S* and water discharge depths *H* can be calculated as:

80
$$D_{50} = \frac{\tau_b}{(\rho_s - \rho_w)g\tau_b^*},$$
 (9)

81 where ρ_s and ρ_w are the densities of sediment and water, $\tau_b = \rho w g H S$ is bed shear stress and τ_b^* 82 is its non-dimensional form (the Shields number). The latter value for a bankfull discharge is 83 gained from:

84
$$au_h^* = 1.5 au_{hc}^* = 0.225 \, S^{0.25},$$
 (10)

based on two observations: (i) the general relation between τ_b^* and its critical value τ_{bc}^* for the 85 onset of sediment motion (i.e., the Shields stress; Parker 1978), and (ii) the dependence of τ_{bc}^* on 86 channel slope (Lamb et al., 2008b). Following Lamb et al., 2008, we scaled both volumetric 87 88 sediment supply (based on Fernandez Luque and van Beek, 1976) and the flow roughness 89 parameter D_{50} , and calculated potential bed impact energies ε_{kin} of D_{50} using the total load model 90 (Lamb et al., 2008; including the sediment cover term), varying both S and H. This procedure 91 presumes this model's validity also for large grain sizes (i.e. valid calculation of bedload layer 92 thickness, bedload transport and settling velocities), and that the grains follow the fluid.

93

From literature, we compiled a data set on bedrock river sections, comprising riverbed slope *S*, flow depth *H*, bedrock type and tensile strength σ_t , and measured grainsizes, as available (Table S3). For cases of not reported σ_t , but descriptions of the local rock type, we adapted σ_t values of comparable rock samples, as measured by Sklar and Dietrich (2001). Reported values 98 of Schmidt Hammer measurements SH were converted to σ_t using the following conversion

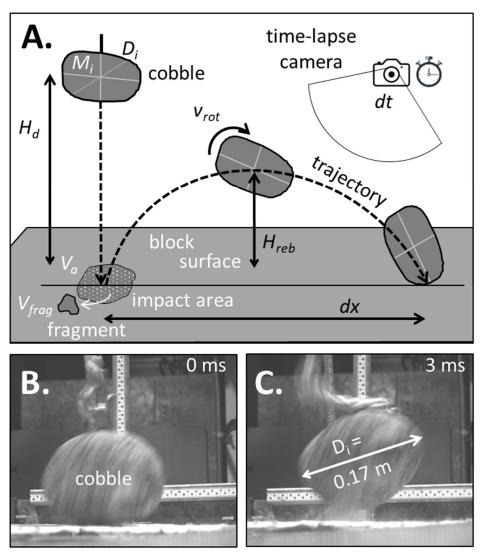
99 (Jamshidi et al., 2018):

100

$$\sigma_t = 0.0347 \ 10^6 \ SH^{1.3783} \tag{11}$$

101 Based on the given slopes S and flow depths H of the bedrock rivers, we then calculated 102 transportable median grainsizes, D_{50} , and related effective impact energies, ε_{kin} , following the total 103 load model procedure described above. The predicted D_{50} mostly exceeded reported D_{50} , but were 104 consistent with maximum observed boulder sizes, D_{max} , which already move under partial 105 submergence (Carling et al., 2002; Fujioka et al., 2015; Alexander and Cooker, 2016). Using the 106 rock tensile strengths, σ_t , we obtained potential bedrock abrasion volumes V_a for the D_{50} grains by 107 means of the extended erosivity regression (Fig. 3). Considering mean kinetic impact energy 108 densities $(\varepsilon_{kin} / D_{50}^2)$ and the assumed threshold for the onset of macro-abrasion (Fig. 2C), we then 109 assigned the river sections to fall into the wear- and macro-abrasion regimes, respectively (Fig. 4).

111 GSA DATA REPOSITORY FIGURES AND TABLES



112 **Figure S1.** Impactor trajectory measurements shown for a dropped cobble (cf. Fig. 1B) with (A)

- 113 definitions of the measured impactor values (mass M_i , diameter D_i , drop height H_d , rebound
- height H_{reb} , rebound rotation velocity $v_{rot,reb}$, hop distance dx and hop time dt), block abrasion
- 115 volume V_a and fragment abrasion volume V_{frag} , and (B-C) time-lapse snapshots in microseconds
- 116 of a cobble before and after its impact on a block's surface.

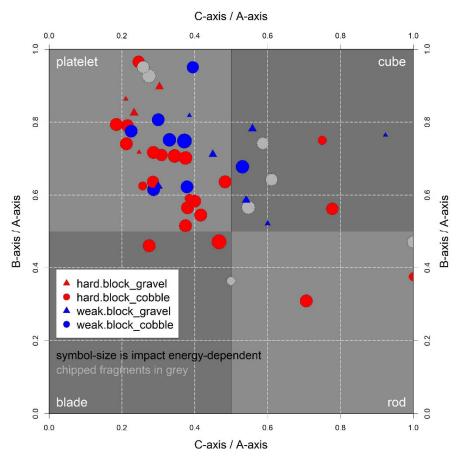


Figure S2. Size relations of the largest fragments abraded per impact experiment, shown in the framework of grain shape types (represented by four sectors with mainly platelet, cube, blade or rod shape). Defining their ABC-axis by decreasing lengths, most abraded fragments were of platelet-shape with size relations of C-axis/A-axis smaller than 0.5 and B-axis/A-axis larger than 0.5.

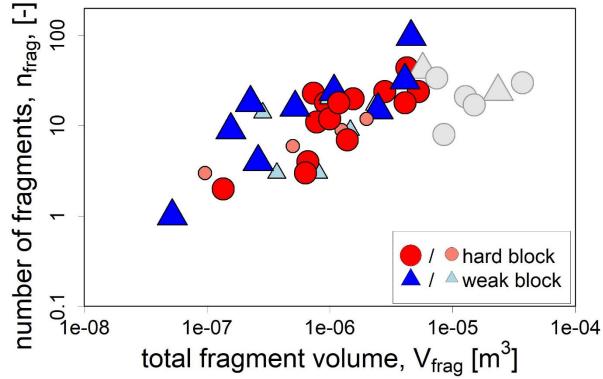


Figure S3. The increase in total fragment volume, V_{frag} , for higher impact energies was due to an increase of the number of fragments, n_{frag} , independent of the impactor size (represented by small and large symbols) or target strength (represented by color). Light grey-shaded symbols show chipped fragments.

Table S1. Overview of the macro-abrasion drop experiments. Individual runs during all seven experimental sets equal the number of structure from motion, *SfM*, models (62). Grey-shaded background denote variables hold constant and bold numbers show varied variables of impactor type (cobble versus pebbles), impactor mass, M_i , and impactor drop height, H_d (Fig. S1A). The tensile strengths, σ_i , of the concrete blocks (conditioned by their cement to sand mixture ratio) was measured using a Split-Hopkinson pressure bar and conform to published values of the same artificial rock material (Sklar and Dietrich, 2001).

	target		impactor(s)			runs		
exp. set	casting ratio	tensile strength	type	single mass	drop height	impacts	repetitions	no. of
	cement : sand	τ [Mpa]		<i>M</i> _i [kg]	<i>H</i> _d [m]	frequency		SfM models
1	1 : 2 (hard)	2.34	cobble	11.8	0.50 - 1.50	1	3	12
2					0.50 - 4.87		1	8
3	(naru)		pebbles	0.3 - 0.5	2.40	5 (1)	2	4
4		1.32	cobble	11.8	0.10 - 4.87	5(1)	1	13
5	1 : 4 (weak)		pebbles	0.1 - 0.5				7
6				0.3	2.40	24	3	3
7				0.3 - 0.5		21	1	4
				constant value	varied value			

136

- **Table S2.** Reported bedrock abrasion experiment data sets with calculated magnitudes of
- 139 effective single grain impact energies, ε_{kin} , and abraded target volumes, V_a . The diameter
- $D_{ero,sphere}$ of equivalent abraded spheres with volume V_a is given for illustration, only.

experiments			orders of magnitudes for single grain impacts					
impactor	target	test principle	impactor mass <i>M</i> ; [kg]	kinetic impact energy ε _{kin} [J]	abraded target volume V_{α} [m ³]	diameter of equiv. sphere with V _a D _{ero,sphere} [m ³]	source	
Crystolon dust	plate glass	sand blasting	1e-9	1e-7 to 1e-8	1e-16 to 1e-15	1e-5	Head and Harr, 1970	
steel balls	concretes with varying aggregates	abrasion mill	1e-3 to 1e-2	1e-3	1e-13 to 1e-12	1e-4 to 1e-5	Liu, 1981	
sand, pebbles	natural rocks, artificial sandstone	abrasion mill	1e-7 to 1e-2	1e-9 to 1e-2	1e-18 to 1e-11	1e-6 to 1e-4	Sklar and Dietrich, 2001	
pebbles	artificial sandstone	grain drop	1e-2	1e-1	1e-9	1e-3	Sklar and Dietrich, 2004	
sand, pebbles	polyurethane foam	abrasion mill	1e-7 to 1e-1	1e-11 to 1e-2	1e-19 to 1e-11	1e-7 to 1e-4	Scheingross et al., 2014	
pebbles, cobble	artificial sandstone	grain drop	1e-1 to 1e1	1e0 to 1e2	1e-7 to 1e-5	1e-3 to 1e-2	this study	

- 142 **Table S3.** Compiled set of published bedrock river data (slope *S*, water depth *H*, measured grain
- 143 sizes *D*, and partially rock tensile strength σ_t) and predicted values of potential streamflow
- 144 erosivity, sorted by year of publication and order in the original publication. Note several
- sections per listed rivers and partly converted or assumed tensile strengths.

147 **REFERENCES CITED**

- AgiSoft PhotoScan Professional (Version 1.4.5; Software), 2018: retrieved from
 http://www.agisoft.com/downloads/installer/
- Alexander, J., and Cooker, M.J., 2016, Moving boulders in flash floods and estimating flow
 conditions using boulders in ancient deposits: Sedimentology, v. 63, p. 1582-1595,
 <u>https://doi.org/10.1111/sed.12274</u>
- Alho, P., Russel, A.J., Carrivick, J.L., and Kaeyhkoe, J., 2005, Reconstruction of the largest
 Holocene joekulhlaup within Joekulsa a Fjoellum, NE Iceland: Quaternary Science
 Reviews, v. 24, p. 2319-2334, <u>https://doi.org/10.1016/j.quascirev.2004.11.021</u>
- Baker, V.R. and Pickup, G., 1987, Flood geomorphology of the Katherine Gorge, Northern
 Territory, Australia, GSA Bulletin, v. 98, p. 635-646, <u>https://doi.org/10.1130/0016-</u>
 <u>7606(1987)98<635:FGOTKG>2.0.CO;2</u>
- Baker, V.R. and Kale, V.S., 1998, The Role of Extreme Floods in Shaping Bedrock Channels, *in* Tinkler, K.J., and Wohl, E.E., eds., Rivers Over Rock: Fluvial Processes in Bedrock
 Channels: American Geophysical Union Geophysical Monograph 107, p. 153-165,
 <u>https://doi.org/10.1029/GM107p0153</u>
- 163 Carling, P.A., Hoffmann, M., and Blatter, A.S., 2002, Initial Motion of Boulders in Bedrock
 164 Channels, *in* House, P.K., Webb, R.H., Baker, V.R., and Levish, D.R., eds, Ancient Floods,
 165 Modern Hazards, American Geophysical Union Water Science and Application, v. 5, p.
 166 147-160, <u>https://doi.org/10.1029/WS005p0147</u>
- 167 CloudCompare (version 2.11; GPL software), 2019: retrieved from <u>http://www.cloudcompare.org</u>
- Farin, M., Mangeney, A., de Rosny, J., Toussaint, R., Sainte-Marie, J., and Shapiro, N.M., 2016,
 Experimental validation of theoretical methods to estimate the energy radiated by elastic
 waves during an impact: Journal of Sound and Vibration, v. 362, p. 176-202,
 <u>https://doi.org/10.1016/j.jsv.2015.10.003</u>
- Fernandez Luque, R., and van Beek, R., 1976, Erosion and transport of bed-load sediment: Journal of
 Hydraulic Research, v. 14 (2), p. 127–144, <u>https://doi.org/10.1080/00221687609499677</u>
- Fujioka, T., Fink, D., Nanson, G., Mifsud, C., and Wende, R., 2015, Flood-flipped boulders: In-situ
 cosmogenic nuclide modeling of flood deposits in the monsoon tropics of Australia:
 Geology, v.43 (1), p. 43-46, <u>https://doi.org/10.1130/G35856.1</u>
- Jamshidi, A., Yazarloo, R., and Gheiji, S., 2018, Comparative evaluation of Schmidt hammer test
 procedures for prediction of rock strength: International Journal of Mining and Geo Engineering, v. 52 (2), p. 199-206, <u>https://doi.org/10.22059/ijmge.2018.244154.594702</u>
- Lamb, M.P., Dietrich, W.E., and Venditti, J.G., 2008b, Is the critical Shields stress for incipient
 sediment motion dependent on channel-bed slope?: Journal of Geophysical Research, v. 113,
 F02008, <u>https://doi.org/10.1029/2007JF000831</u>
- 183 Neilson, J.H. and Gilchrist, A., 1968, Erosion by a stream of solid particles: Wear, v. 11 (2), p. 111 184 122, <u>https://doi.org/10.1016/0043-1648(68)90591-7</u>

Parker, G., 1978, Self-formed straight rivers with equilibrium banks and mobile bed. Part 2. The gravel river: Journal of Fluid Mechanics, v. 89 (1), p. 127–146, <u>https://doi.org/10.1017/S0022112078002505</u>

188	Zhou, Y.X., Xia, K., Li, X.B., Li, H.B., Ma, G.W., Zhao, J., Zhou, Z.L., and Dai, F., 2011, Suggested
189	Methods for Determining the Dynamic Strength Parameters and Mode-I Fracture Toughness
190	of Rock Materials, in Ulusay, R., eds, The ISRM Suggested Methods for Rock
191	Characterization, Testing and Monitoring: 2007–2014. Springer, Cham, p. 35-44,
192	https://doi.org/10.1007/978-3-319-07713-0_3
193	