

A short, sharp pulse of potassium-rich volcanism during continental collision and subduction

SUPPLEMENTARY FILE

ANALYTICAL METHODS

The analytical method for determining the boron concentration and isotope composition of the samples has been adapted from Wei et al. (2013). All the reagents used were high purity and all the sample processing took place in clean laboratories equipped with boron-free HEPA filters at the University of Southampton.

Approximately 100 mg of rock powder was accurately weighed and transferred into a 7 ml screw top PFA vial, together with 100 μ L of 1% mannitol, 100 μ L H₂O₂ and 1 mL 24 M HF. The vial was sealed and placed on a hot plate for 3 days at 60°C. A white precipitate, consisting largely of fluorite, was present in the sample solution after this time. Dissolution of this precipitate and analysis by inductively coupled plasma mass spectrometry (ICP-MS) revealed that it did not contain any detectable boron, in accordance with previous observations (Nakamura et al., 1992; Wei et al., 2013). After dissolution, 0.5 mL of the sample was added to 1.5 mL of water in a polypropylene centrifuge tube. After centrifuging, the supernatant was transferred to another 7 mL PFA vial and a further 2 mL of water was added to yield a 3 M HF solution.

The boron concentration was determined by taking 0.1 mL of the 3 M HF sample solution and diluting to 5 mL with 3% HNO₃ containing a 10 ng/g Be internal standard. The boron concentration was then determined using a Thermo Element high-resolution (HR) ICP-MS against a range of 0 to 5 ng/g boron standards (also containing 10 ng/g Be).

Boron was separated from the solution using Bio-Rad AG MP-1 anion exchange resin. The columns consisted of a polypropylene funnel fitted with a porous frit. A fresh aliquot of ~100 μL of resin was loaded into the stem of the funnel for each separation. The overall ion exchange chromatography procedure is summarized in Table DR1. The sample load volume was adjusted so that ~100 ng of boron was added to the column. The boron fraction was collected in a 7 mL PFA vial to which 10 μL of 1% mannitol was added and the solution was evaporated to dryness at 55°C.

The sample was taken up in 0.5 mL of 0.5 M HNO_3 to yield a solution with ~200 ng/g boron. Boron isotope analyses were then carried out using a Thermo Neptune multi-collector (MC) ICP-MS following the method described in Foster (2008) that was modified with the use of a 0.3 M HF /0.5 M HNO_3 wash solution to speed up the wash-out time. The NIST SRM 951 boric acid standard (200 ng/g) was used during the sample-standard bracketing and to define the $\delta^{11}\text{B}$ values of the samples. Typical in-run 2σ precisions were <0.2 ‰.

During this study, all samples were analyzed in batches of 10. Each batch consisted of duplicates of 4 powdered rock samples, a powdered rock reference material and a blank. Thus each sample was independently analyzed for both its concentration and $\delta^{11}\text{B}$ value at least twice, and 7 samples were analyzed in triplicate. The blank after sample dissolution was always <0.5 ng and typically <0.1 ng. The blank associated with the ion exchange chromatography was always <1 ng of boron and typically <0.3 ng (compared to a sample analyte of ~100 ng of boron). The largest difference between the highest and lowest $\delta^{11}\text{B}$ replicate analyses of a sample was 1.0 ‰, with the average difference being 0.3 ‰ (Table DR2). The largest difference between the highest and lowest boron concentration [B] replicate analyses of a sample was 10.7 % of the absolute value, with the average difference being 3.7 %. During this study, 10 analyses of standard IAEA-B5 (basalt from Etna volcano; Gonfiantini et al., 2003) yielded an average $\delta^{11}\text{B}$ value of -4.2 ± 0.2 ‰ (2σ) and an average

[B] of $9.2 \pm 0.2 \text{ } \mu\text{g/g}$ (2σ). These values compare well with values previously obtained for this reference material (Table DR3).

In addition to the 42 $\delta^{11}\text{B}$ and [B] data listed in Table 2, Tonarini et al. (2005) report data for 16 WA volcanic rocks. They did not analyse the IAEA-B5 reference material, but they report a $\delta^{11}\text{B}$ value of $7.3 \pm 0.4 \text{ } \text{‰}$ for the GSJ-JB2 basalt reference material. Four analyses of this reference material in this study yielded average $\delta^{11}\text{B}$ and [B] values of $7.2 \pm 0.2 \text{ } \text{‰}$ and $29.4 \pm 0.7 \text{ } \mu\text{g/g}$, respectively. Hence, the $\delta^{11}\text{B}$ values within the two data sets are directly comparable.

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FIGURE CAPTIONS

Figure DR1.

Harker diagrams for WA volcanic rocks

(a) [B] $\mu\text{g/g}$ vs. SiO_2 wt%, (b) [B] $\mu\text{g/g}$ vs. MgO wt%, (c) $\delta^{11}\text{B}$ ‰ vs. SiO_2 wt%, (d) $\delta^{11}\text{B}$ vs. MgO wt%. Major element data from Ersoy and Palmer (2013), and references therein.

Figure DR2.

[B] and $\delta^{11}\text{B}$ values versus radiogenic isotope ratios

(a) [B] $\mu\text{g/g}$ vs $^{87}\text{Sr}/^{86}\text{Sr}$, (b) $\delta^{11}\text{B}$ ‰ vs $^{87}\text{Sr}/^{86}\text{Sr}$, (c) $\delta^{11}\text{B}$ ‰ vs $^{143}\text{Nd}/^{144}\text{Nd}$, (d) $\delta^{11}\text{B}$ ‰ vs $^{208}\text{Pb}/^{204}\text{Pb}$. Radiogenic isotope data from Ersoy and Palmer (2013), and references therein.

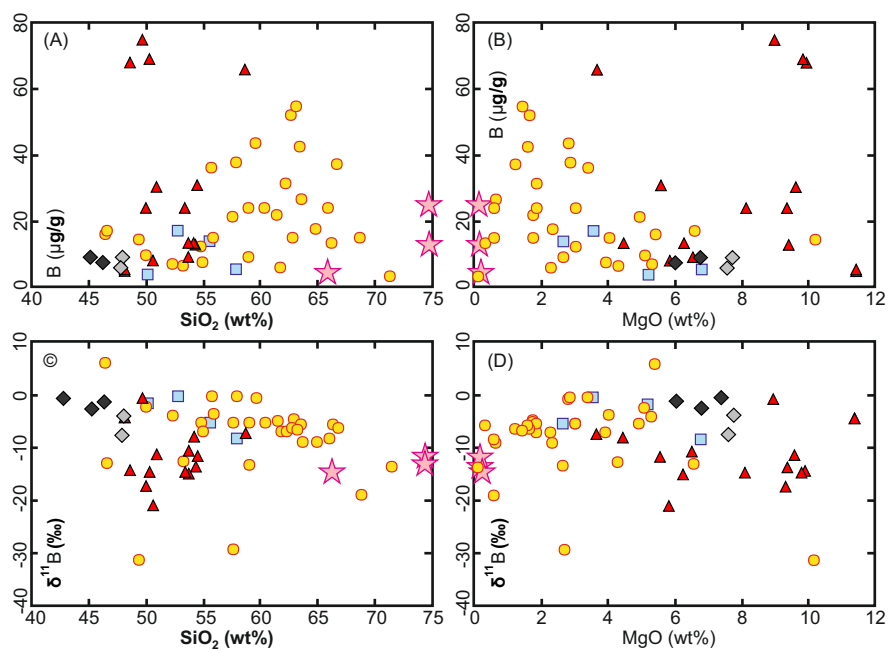


Figure DR1

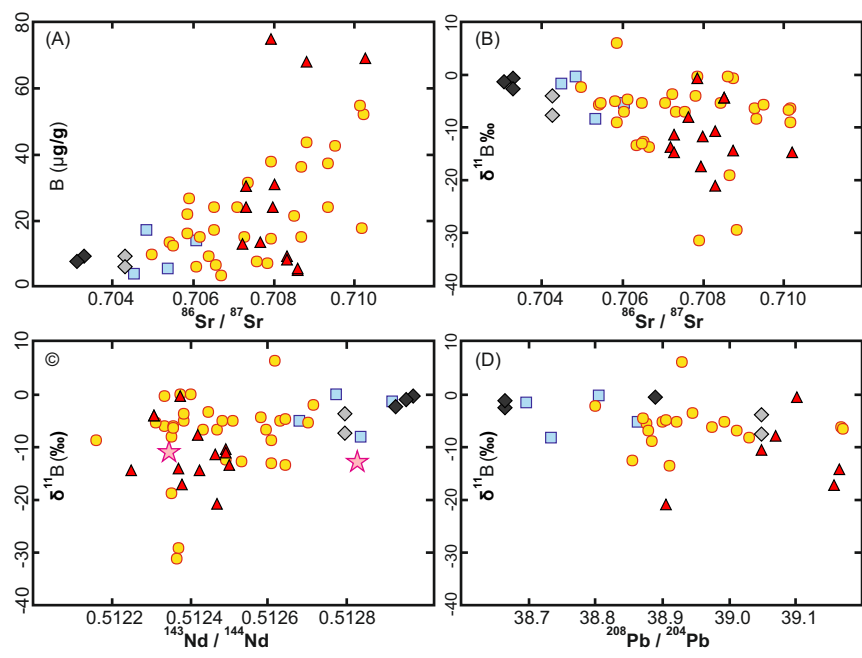


Figure DR2

Table DR1
Ion exchange chromatography procedure

Step	Operation
Step 1	Resin washing and conditioning 2.5 mL H ₂ O 0.5 mL 24 M HF 1.5 mL H ₂ O
Step 2	Sample loading 3 M HF (volume adjusted according to sample)
Step 3	Matrix elution 2 x 0.5 mL H ₂ O 8.5 mL 0.1 M HCl
Step 4	Boron collection 1.2 mL 24 N HF

Boron concentration and isotope data for samples analysed in this study (UTM coordinates give sample location)

Sample (UTM)	B concentration ($\mu\text{g/g}$)				$\delta^{11}\text{B}$ (‰)			
	1 st run	2 nd run	3 rd run	mean	1 st run	2 nd run	3 rd run	mean
<i>Early Eocene orogenic</i> (Kizderbent volcanics)								
G5 (0651592/4488750)	5.2	5.1		5.2	-8.3	-8.3		-8.3
G28 (0651592/4488750)	16.9	16.8	16.6	16.8	-0.6	-0.5	0.2	-0.3
G45 (0749558/4499472)	3.8	3.7		3.8	-1.7	-1.6		-1.6
G48 (0737486/4502470)	14.2	13.6	13.9	13.9	-5.2	-5.3	-5.7	-5.4
<i>Middle Eocene – Miocene orogenic</i> (Late Eocene Balıklıçeşme volcanics)								
B5 (0492827/4448313)	16.1	16.3	16.0	16.1	6.3	6.2	5.8	6.1
B11 (0500469/4459685)	12.9	13.3		13.1	-5.5	-5.8		-5.7
B12 (0486130/4466913)	14.6	14.9		14.7	-4.6	-4.7		-4.7
B17 (0473433/4454350)	23.1	24.0	23.7	23.6	-5.7	-5.3	-4.8	-5.3
B35 (0485295/4455090)	9.9	9.3		9.6	-2.2	-2.2		-2.2
C15 (0481003/4413380)	6.4	6.0		6.2	-12.6	-13.1		-12.8
C42 (0466010/4432014)	6.2	5.6		5.9	-7.0	-7.3		-7.1
(Oligocene Kirazlı volcanics)								
B21 (0477628/4444050)	3.4	3.4		3.4	-13.7	-13.5		-13.6
C7 (0480017/4430858)	11.9	12.0		12.0	-5.2	-5.2		-5.2
C9 (0477916/4420541)	14.6	14.4	14.7	14.6	-3.8	-3.7	-4.1	-3.8
C31 (0497168/4425055)	26.5	26.7		26.6	-8.9	-9.4		-9.1
C36 (0474152/4433482)	21.0	21.9		21.5	-5.1	-4.8		-5.0
(Oligocene Hallaçlar volcanics)								
HG2 (0530735/4407610)	31.0	31.1		31.1	-6.8	-7.0		-6.9
HG9 (0536865/4436733)	24.0	24.3	23.9	24.1	-5.0	-5.4	-5.1	-5.2
(Oligocene Keşan volcanics)								
K20 (0469494/4521168)	8.5	9.4		9.0	-13.2	-13.2		-13.2
(Miocene Yuntdağı volcanics)								

129 (0540981/4353496)	15	14.3		14.7	-19.1	-19.1	5.8	-19.1
(Miocene Yağcıdağ volcanics)								
510 (0657381/4302983)	49.7	54.6		52.1	-6.1	-6.3		-6.2
YF2 (0657381/4302983)	53.5	55.4	54.1	54.3	-6.6	-6.5	-6.3	-6.5
(Miocene Eğreltidağ volcanics)								
521 (0657795/4308366)	25.1	23.1		24.1	-8.6	-8.1		-8.4
(Miocene Sevinçler volcanics)								
718 (0652561/4322172)	36.7	37.4		37.1	-6.4	-6.4		-6.4
(Miocene Miocene Asitepe volcanics)								
754 (0636380/4309275)	42.8	42.1		42.5	-5.7	-5.3		-5.5
(Miocene Yaylaköy volcanics)								
L36 (0453515/4270500)	21.7	20.5		21.1	-5.6	-5.8		-5.4
(Miocene Armağandağ volcanics)								
L47 (0451794/4236840)	43.7	43.2		43.5	-0.8	-0.8		-0.8
(Miocene Kuzayır lamproite)								
518 (0660886/4311660)	9.1	9.3		9.2	-11.2	-10.2		-10.7
528 (0660886/4311660)	8.2	7.8		8.0	-20.8	-21.2		-21.0
(Miocene Doğanca lamproite)								
K17 (0475167/4528879)	4.8	5.1		5.0	-4.0	-4.4		-4.2
K17B (0475167/4528879)	4.9	5.4		5.2	-4.3	-4.1		-4.2
(Miocene Kayacık volcanics)								
756 (0601705/4305141)	16.7	18.0		17.4	-9.1	-8.8		-8.9
(Miocene Güre lamproite)								
822 (0679728/4281490)	66.2	71.2		68.7	-14.7	-14.8		-14.7
(Miocene Karaburun volcanics)								
L30 (0458200/4275780)	14.1	14.7		14.4	-31.4	-31.3		-31.3
(Miocene Orhanlar basalt)								
536 (0680006/4311835)	24.5	23.7		24.1	-17.2	-17.4		-17.3
541 (0679810/4310909)	67.3	68.5		67.9	-14.7	-14.2		-14.4
543 (0679836/4310901)	72.8	76.7		74.8	-0.5	-1.0		-0.8
(Miocene Naşa basalt)								

746 (0671372/4333717)	13.4	13.6	13.5	-8.0	-8.2	-8.1
<i>Transitional – Miocene</i>						
(Late Miocene Kabaklar basalt)						
520-1 (0671542/4298101)	5.9	5.7	5.8	-7.4	-7.6	-7.5
<i>Basement Rocks</i>						
(Menderes Massif - gneiss)						
G01 (0673572/4300943)	13.4	13.1	13.3	-13.3	-12.9	-13.1
G02 (0673572/4300943)	24.9	26.7	25.8	-11.3	-10.8	-11.0
(Menderes Massif - Rahmanlar granite)						
Y23-2 (0674406/4302063)	4.1	4.4	4.3	-14.1	-15.1	-14.6

Table DR3.

Boron concentration and isotope data for IAEA-B5 reference material

Source	[B] ($\mu\text{g/g}$) $\pm 2\sigma$	$\delta^{11}\text{B}$ (‰) $\pm 2\sigma$
This study	9.2 ± 0.2	-4.2 ± 0.2
Gonfiantini et al., 2003	10.0 ± 1.2	-4.1 ± 2.7
Pi et al., 2014	9.3 ± 0.1	-3.9 ± 0.2
Romer et al., 2014	10.0 ± 1.0	-3.7 ± 0.7
Hansen et al., 2017		-4.3 ± 0.2
Berryman et al., 2017	9.3 ± 0.4	-4.3 ± 0.2