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Appendix DR1. Sample locations and additional compositional data. The Glacial polish samples studied here were collected from Yosemite National Park (YNP), California. A total of 6 samples from the Lyell canyon and Tuolumne meadows area, were used to collect major and trace element data (Fig. DR1.1). Lyell and MaClure samples were collected from elevation, towards the glacier accumulation area for this catchment, while DAFF is located farther below (Fig. DR1.1).

In figure 2 of the main text the major and trace element data for DAFF dome is presented. For brevity, this figure excluded Al₂0₃ and Th. These additional elements for DAFF dome are presented here as a modified version of the figure 2 included in the main text (Fig. DR1.2). The same compositional analysis was performed on the 5 additional MaClure and Lyell basin samples. These data are presented within this appendix as figures DR1.3-1.7. A k-means cluster analysis from MATLAB was performed on these data to identify the polish composition and underlying minerals for each spot analysis. A k-means cluster analysis partitions the calculated element concentrations at each ablation depth into one of two clusters (defined by user) with the nearest mean. The center of each cluster is a collection of values that are interpreted to best define the resulting group. These data (black circles) are what is presented in the main text's figure 3.

Appendix DR2. LA-ICPMS methods and Data Reduction

All LA-ICPMS analyses were conducted using a Photon Machines Analyte 193H Excimer Laser outfitted with Helex 2-Volume Cell feed into a Thermo Scientific X-series quadrupole ICP-MS. Laser spot sizes were set to 25 µm, a minimum size found to meet instrument detection limits for selected major elements: Na, Mg, Al, Si, Ca, Fe. Elements U and Th were also measured but often below detection limits, in particular at depth within host minerals. Laser parameters include: 80 bursts at a 4Hz shot rate. Sample unknowns were measured along with standard NIST-610. Spot analyses were conducted by ablating from the glacially polished surfaced downwards through the amorphous layer and into the coarse minerals of the underlying host rock. It was not possible to utilize the traditional LA-ICPMS data reduction technique of internal normalization because the proportions of all elements change within this depth profile. We instead adopted a sumnormalization technique, applied to each cycle of masses within the ICPMS data. By dividing the proportions of each cycle by the sum we can roughly determine the relative proportions of each element as a function of depth. This data reduction includes: 1) blank correction; 2) conversion of signal intensities (cps) to elemental concentrations using measured NIST 610 intensities; 3) Conversion of isotopic concentrations to total elemental abundances; 4) Conversion of elemental concentrations to oxides; 5) Normalization of the total oxide for a given mass cycle. We find that the sum normalization technique leaves this final data output from this reduction technique highly insensitive to data reduction steps 2 and 3. Rather the results seemingly deliver a reliable measure of the relative proportion of elements at any depth for a given analyses. It thus remains only to evaluate whether there is inherit bias with a single spot analyses between shallow and deep parts of the laser pit, which we will address below. Our use of this "sum-normalization" techniques makes several assumptions including: 1) That sum normalization is not "missing" any large component of the material (e.g. that Na, Mg, Al, Si, Ca, Fe, U and Th constitute 99% or more of the total material). 2) That element allocation in oxides is valid 3) that elemental fractionation in the NIST standards is identical to the samples, and 4) that down hole changes do occur but they do not cause elemental fractionation (e.g. the relative delivery of different to the mass spectrometer remains constant during spot analyses.

We can evaluate all of these assumptions by through collection and reduction of secondary standard data. Here we present LA-ICPMS analyses from ATHO (Fig. DR1.1) and a analyses collected from multiple minerals on a manually polished section of Cathedral Peak Granodiorite (Fig. DR2.2). The ATHO analyses reveal that there is no inherent bias corresponding with sample depth (e.g. blue and yellow data points cluster in same region). This implies that the compositional trends comparing shallow amorphous material and the underlying host minerals remains a viable conclusion for this paper. This statement is supported by the data collected from a manually polished sample of Cathedral peak granodiorite. The compositional of host minerals are discrete resolved with compositions independent of ablation depth (Fig. DR2.2). These secondary data no not nearly as precise (~10% uncertainties) as other in situ techniques (i.e. electron microprobe), however the technique provides the ability to measure continuous compositional not offered by other these more precise methods.

Appendix DR3: SHRIMP ²³⁰Th, ²³²Th, ²³⁴U, ²³⁸U analytical methods and data reduction

Opal in situ U-Th isotopic measurements were performed using the SHRIMP-RG (sensitive high-resolution ion microprobe with reverse geometry) operated by Stanford University and the U.S. Geological Survey during five analytical sessions between March 2017 and July 2018. Fragments (10 to 24 mm) Daff granite with polish coating was cast in a 25 mm diameter epoxy disc with pre-polished reference materials (BZVV opal and NIST-SRM-611 glass. BZVV is a ca. 2.83 Ma porcelaneous biogenic opal (Virgin Valley, NV, USA), that is relatively high U (748-888 ppm U) used as a secular equilibrium reference material (Paces et al., 2004; Amelin and Back, 2006). The mounted samples were rinsed with a 10% EDTA solution (ethylenediaminetetraacetic acid) and thoroughly rinsed in distilled water and dried in a vacuum oven. The sample surface was coated with ~100 nm Au and inspected to ensure uniformity and conductivity, stored in a pre-evacuation chamber at ~10-7 torr overnight prior to analysis to minimize degassing of the epoxy.

All analyses were performed by depth profiling through the surface into the amorphous silica polish. Because the polish is relatively thin and low U, the objective was to obtain a large spot (30-50 microns) that was as flat-bottomed as possible to minimize the sputter rate while maximizing secondary ion intensity. The second analytical session utilized an O- primary ion beam with intensity ranging from 45-55 nA. Although the secondary ion yield was ~35% higher using O-, the sputter rate was higher and thus sputtered through the polish more quickly. Therefore, for all subsequent sessions, we used an O2- primary ion beam with intensity ranging from 12-19 nA. The acquisition routine included the following masses: 28Si416O5+, 238U+,232Th12C+, 232Th16O+, background measured at 0.05 m/z above 232Th16O, 234U16O+ and 238U16O+. All peaks were measured on a single ETP® discrete-dynode electron multiplier operated in pulse counting mode. Analyses were performed with 9-12 scans (peak-hopping cycles from mass 192 through 254) and count times for 232Th16O and 234U16O+ ranging from 90 to 120 seconds and 35 to 50 seconds, respectively. Individual analyses took 45 to 60 minutes and with sputtered pit depths estimated to be 5-8 μ m. Measurements were performed at mass resolutions of M/ Δ M = 7,000-8,500 (10%) peak height) to avoid interfering isobaric interferences.

Measured isotopic ratios were corrected for detector deadtime, background, and calculated cycle-by-cycle (i.e., depth into the polish). As discussed in the text, not all the polish material is amorphous silica or enriched in U; the surface layer (typically <1 μm thick; Fig. 1) can be phyllosilicate alteration and polish can contain micron- to submicron-sized mineral fragments in secular equilibrium. As a result, the initial scan intersecting phyllosilicate material was typically omitted. Additionally, scans lacking U and 230Th or those in secular equilibrium were also omitted, which typically were deeper in the sputter pits (scan number >8), likely due to intersection of mineral fragments or the underlying host minerals that lack U and/or were in secular equilibrium. Based on these criteria, typically 30-80% of the measurement results were used from each spot, and ratios were calculated using methods outlined by Ludwig (2009) with calculated into activity ratios (AR) using 238U, 232Th, 230Th, and 234U decay constants of 1.55125×10-7, 4.9475×10-8, 0.0091577, 0.002834, respectively. BZVV opal reference material were repeatedly measured throughout the analytical sessions.

Measured 238U/232Th AR and 230Th/238U AR were corrected for correct for mass discrimination based on session-averaged measurements of 230Th/238U for BZVV, which should be unity based on U-Th isotopes measurements Amelin and Back (2006) and measurements of a similar opal, M-21277, from the same locality by Paces et al. (2004). Uranium concentrations for glacial polish were calculated relative to BZVV opal, using an average concentration of 840 ppm U from Amelin and Back (2006).

This geochronologic method used in figure 4 assumes that upon formation, this fractionation was complete (e.g. $230 \text{Th}/238 \text{U} \sim 0$), an assumption that is supported by a regressed 230 Th/232 Th activity ratio initial for the Daff Dome polish of 1.6 ± 1.3 (Fig. DR. 3.1).

References

Amelin, Y., and Back, M., 2006, Opal as a U–Pb geochronometer: Search for a standard: Chemical Geology, v. 232, p. 67-86.

Ludwig, K. R., 2009, SQUID 2: A User's Manual, rev. 12. Berkeley Geochron. Ctr. Spec. Pub. 5 110 p.

Paces, J.B., Neymark, L.A., Wooden, J.L. & Persing, H.M., 2004, Improved spatial resolution for U-series dating of opal at Yucca Mountain, Nevada, USA, using ion-microprobe and microdigestion methods, Geochimica et Cosmochimica Acta, v. 68, p. 1591-1606.

Appendix DR4. Ablation rate determination: We can empirically calibrate the depth at which compositional changes occur by ablating samples for a range of laser pulses from 5 to 80 and with inferometer measurements of laser pit depths (Fig. DR 4.1) These data are used to assign approximate depths to compositional data (Fig. 1-2). A total of 80 pulses were used for all analyses presented in this manuscript. The ablation rate varies between 0.08 μ m/pulse at shallow depths decreasing to 0.05-0.06 μ m/pulse at greater depths (Fig. DR4.1). A maximum 0.08 μ m/pulse was assumed to the depth assignments in figures 1 and 2 of the main text. This rate does not influence the data reduction in any way and only permits rough estimates on the depth of chemical transitions to be assigned.

Table. DR1.1 Sample locations.

Sample name	Coordinate	N	\mathbf{W}
Daff	37°52'39.4"N 119°24'52.3"W	37.877611	-119.414528
Lyell08	37°45'17.5"N 119°15'35.6"W	37.754861	-119.259889
Lyell09	37°45'20.3"N 119°15'35.3"W	37.755639	-119.259806
Lyell10	37°45'24.2"N 119°15'57.5"W	37.756722	-119.265972
Maclure03	37°45'16.1"N 119°16'52.6"W	37.754472	-119.281278
Maclure05	37°45'56.2"N 119°16'25.8"W	37.765611	-119.273833

Fig. DR1.1

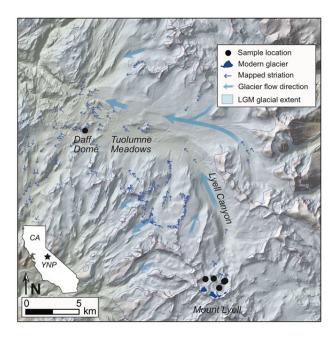


Figure DR1.1 Map of Lyell Canyon and Tuolumne Meadows, Yosemite National Park (YNP), CA. Compositional data were collected on 6 samples near the top of the Lyell canyon catchment, and 1 at Daff dome. U-series analyses were only conducted on the sample collected from Daff dome.

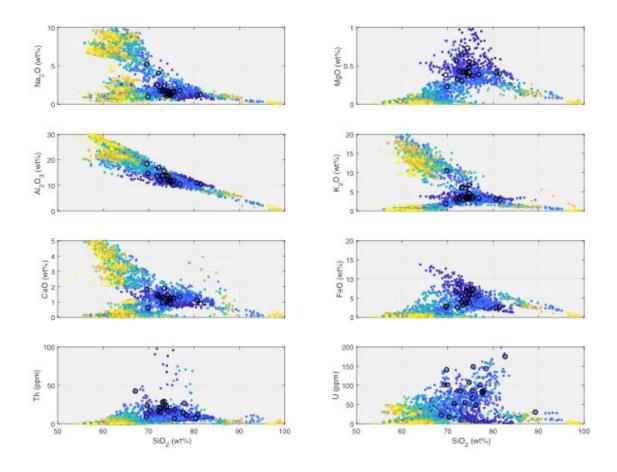


Figure DR1.2 Compositional data from Daff dome. The same data presented in Figure 2 of the main text, presented here with Al_2O_3 and Th data.

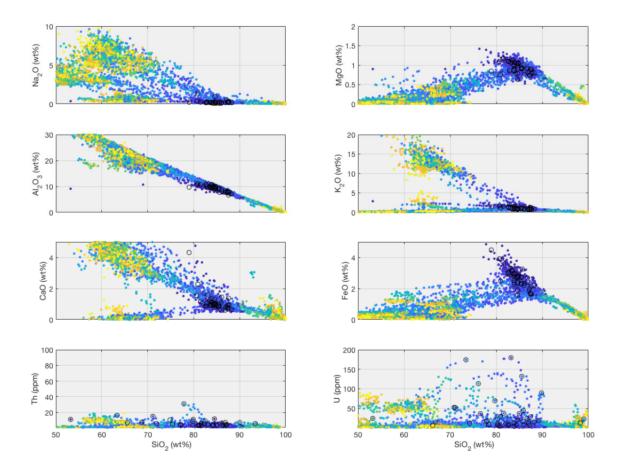


Figure DR1.3 Compositional data from sample MaClure03. Polish compositions for each laser spot analysis are defined as black circles and presented in Figure 3 of the main text. As in figure 2, yellow colors are deep reflecting the bottom of laser pits and the composition of underlying host minerals, while blue colors mark shallow depths and that of the glacial polish. Ablation rate was not calibrated for this sample, thus depths are relative.

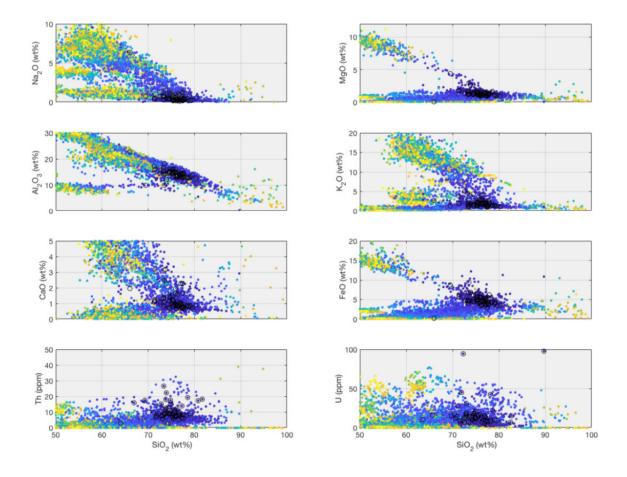


Figure DR1.4 Compositional data from sample MaClure05. Polish compositions for each laser spot analysis are defined as black circles and presented in Figure 3 of the main text. The polish covering high Fe and Mg amphiboles are was measured on this sample, with compositions the same as those above quartz and feldspar. As in figure 2, yellow colors are deep reflecting the bottom of laser pits and the composition of underlying host minerals, while blue colors mark shallow depths and that of the glacial polish. Ablation rate was not calibrated for this sample, thus depths are relative.

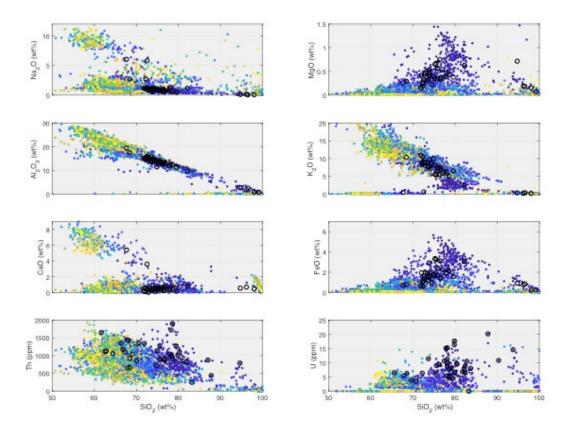


Figure DR1.5 Compositional data from sample Lyell08. Polish compositions for each laser spot analysis are defined as black circles and presented in Figure 3 of the main text. As in figure 2, yellow colors are deep reflecting the bottom of laser pits and the composition of underlying host minerals, while blue colors mark shallow depths and that of the glacial polish. Ablation rate was not calibrated for this sample, thus depths are relative.

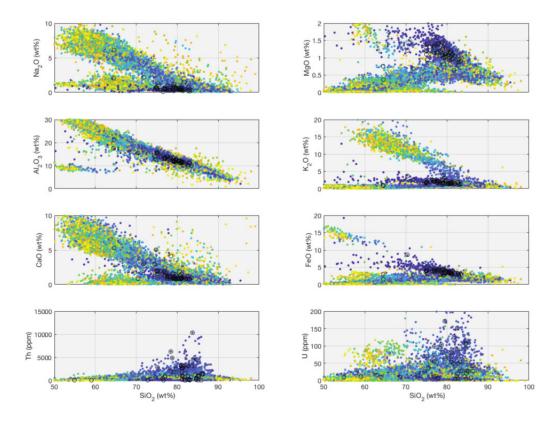


Figure DR1.6 Compositional data from sample Lyell09. Polish compositions for each laser spot analysis are defined as black circles and presented in Figure 3 of the main text. As in figure 2, yellow colors are deep reflecting the bottom of laser pits and the composition of underlying host minerals, while blue colors mark shallow depths and that of the glacial polish. Ablation rate was not calibrated for this sample, thus depths are relative.

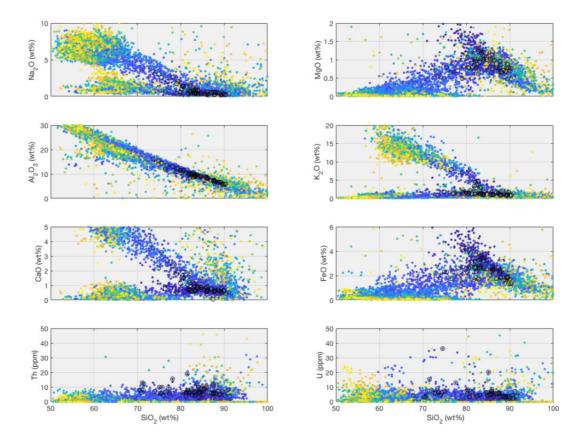


Figure DR1.7 Compositional data from sample Lyell10. Polish compositions for each laser spot analysis are defined as black circles and presented in Figure 3 of the main text. As in figure 2, yellow colors are deep reflecting the bottom of laser pits and the composition of underlying host minerals, while blue colors mark shallow depths and that of the glacial polish. Ablation rate was not calibrated for this sample, thus depths are relative.

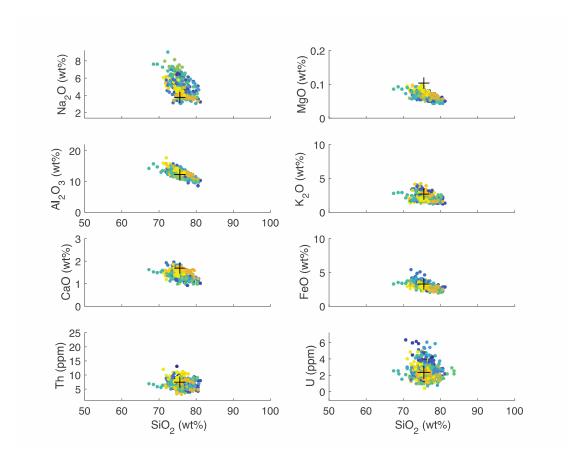


Figure DR2.1 Compositional data for secondary standard data ATHO. Crosses mark accepted values. As in figure 2, yellow colors are deep while blue colors mark shallow depths. Data are reduced using the sum normalization technique discussed in DR2. The apparent lack of depth vs. compositional bias suggests that this technique, while not delivering the most precise data do not deliver data that are inherently biased by laser ablation depth. The apparently inaccurate Mg value likely reflects the use of NIST610 Mg concentrations that are incorrect. Ablation rate was not calibrated for this sample, thus depths are relative.

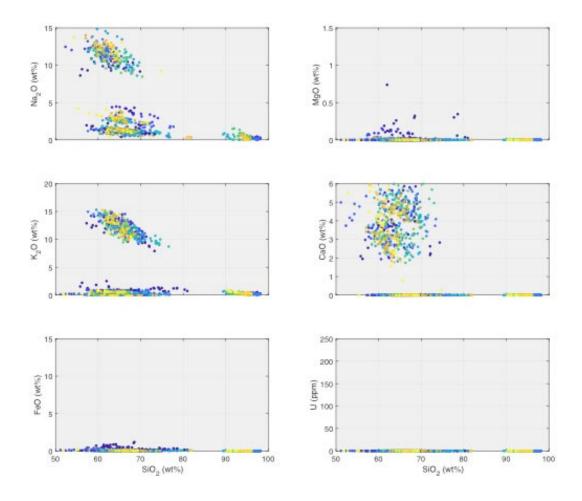


Figure DR2.1 Compositional data from a manually, non-glacially polished sample of cathedral peak shown at the same scale as figure 2 to illustrate the differences between glacially polished samples and those of the host minerals. The apparent lack of depth vs. compositional bias suggests that this technique is not inherently biased by laser ablation depth. Note that the distinct host mineral compositions are separately resolved and that shallow analyses do not artificially yield compositions consistent with the glacial polish as shown in figure 2. For example, the Mg, Fe and U enrichment in the surface of glacial polish (Fig. DR1.2-1.7) is not recorded on the manually polished sample. Ablation rate was not calibrated for this sample, thus depths are relative.

Table DR 3.1: SHRIMP-RG ²³⁰Th, ²³²Th, ²³⁴U, ²³⁸U data, note ²³⁴U was not acquired during sessions 1-2 and thus not included in Figure 4. Spot analyses below 10 ppm U are not included in figure 4 and struck through and red in the tables below. All Daff analyses are from the same Daff01 samples described in the text and figures 1-2.

)) is U (ppm) Th/U Scans Notes	0.41 6/9	0.65	1.5 0.34 4/9 low-U	0.45 7/9	0.55 6/8	0.29 5/9	0.67	0.22	0.00	0.20	0.34	0.21	0.14	5 6		0.31	90:0	200	0.0015	0.0002		0.0003		0.0004			ly.*	J) 1s U (ppm) Th/U Scans Notes	1.81 10/10	0.28 10/10	5/10	0.19 4/10	0.46 2/10	0.26 8/10	0.11 4/10	0.73 6/10	0.35 6/10	2.28 8/10	0.12	6/10	0.18			0.10 6/10	60.0	9	62.2 1.09 9/9 840.0 0.0011 8/8	0.0004			1)/* 11/nnm/ Th/II Scane Motos	ora con	2 0.521 2.7 0.23 5/10 low U 2 1.458 4.6 0.23 8/10 low U, secular equilibrium	0.00	200
(234U)/ (238U)	4	æ	36	\$	æ	쉺	24		2 0	±	ቋ	39	33	2 2	2 9	#	22	,	5	g 9	5	Z 8	8 8	67			(234U)/	(238	133	#	st:	\$	#	1 2	9	14	: #	86		92	36	Ξ	75	유 1	-	!	2 2	12			(234U)/		4.432		
Th)/ °	16:0	99	0.686	0.53	698-0	0.3	0.33		5 6	b	0.300	0.0	0 16	5 6	0.0	5	0.0	ć	0.0	0.0	0.0	0.0	0.0	0.029	5		° (hT	18	/ 89	97.5	6.0	4	4.5	4	9-0	d	3	0	0.2	0.0	0.13	0.11	0.0	0.040	0.0	Č	0.017	0.0			Th)/ °	2	0.040		
corr(230Th)/ ° (238U)	1.544	5.427	0.509	1.51	3,155	0.470	0.489	005.0	0.000	0.670	0.852	0.206	0 220	0.443	0.442	919	0.251		L 10. L	1.031	776.0	1.125	0.909	0.99	2.6.0		corr(230Th)/ °	(238U)	-93.356	-5.308	4.779	0.749	2.008	407	0.944	4 523	0.858	0.488	0.596	0.350	0.561	0.535	0.410	0.181	0.413	0	0.990	1.010			corr(230Th)/	(200-)	0.736 0.961		0
18	25.448	41.575	3.595	3.003	8.502	5.465	1 638	1,643	2 6	5.975	3.454	1.488	3 404	100	0.803	1.84	6.003	100	79/	726.8	311./	2337.6	1492.3	1974.1) 101 DZ v v.			1s	461.919	87.120	18.366	606-9	9.522	6.272	25.930	2 977	3 690	0.201	8 035	3.101	2.574	4.516	2.479	1.098	567.7		0.009 818.7	1372.0			5	2	1.466 3.302	1000	
(230Th)/ a	41.700	13.040	1.973	10,716	26.625	6.949	2 444	181		###	8.682	3.758	7 573	777	4.747	11.865	16.079		7977	14562.9	6312.5	12316.7	10224.4	390.3 0934./ average (230Th/2381)	230111/2300		(230Th)/ a	(232Th)	-166.603	-69.993	-6.959	# 1.01	13.576	43.094	33.718	8 038	9.359	0.859	15.892	13.238	10.364	19.816	12.217	5.241	46:914		-0.024 2855.6				(230Th)/ a	(8.219 13.928	-	01.
18	1.380	908-0	1.015	0.218	4.142	986-0	0.483	1 031	56.0	476.0	1.008	1.558	2 007	4 405	. 183	1 999.	3.538	Š	504	1786.1	13/4.4	1885.6	1,40.3	1090.0	average			18	0.290	966-0	869.0	1.703	0.175	0.765	2.323	445	1 543	0.247	4.526	2.965	1.085	3.648	2.804	4.057	1987			1217.9			, t	200	1.422 4.220		. 00
corr(238U)/ ^b (232Th)	2.662	4.845	9.331	6.924	6.714	10,660	5 122	14 108	900	15.822	9:324	15.042	22 597	42.00	12.028	10:049	55.278	0,00	2046	12834.9	5184.8	10008.8	7264.4	1.101.1			corr(238U)/ b	(232Th)	4.728	41.413	44:489	16.306	6.753	44.988	29.748	4.266	8 042	1.378	26.014	33.768	17.513	58.602	24.087	32.588	49999		2744.9	8654.1			corr(238U)/ b	(=)	13.484 13.622		
\$	226.0	3.658	0.686	0.525	698-0	0.374	0 323	188	0.00	919	0.300	0.088	0.162	0.02	0.008	960:0	0.073	0	0.026	0.024	0.024	0.026	0.020	0.024				18	68.433	26.79	0.923	434	4.234	0.475	0.619	0.565	986	0.097	0.274	0.074	0.135	9:40	0.073	0.037	9.048	0	0.003	0.012			-	9	0.040		1000
(230Th)/ ^a (238U)	1.406	4.942	0.464	1.376	2.873	0.428	0.445	0.455	0.5	119	9.776	0.188	0000	0.503	0.402	187.0	0.228	200	0.921	0.939	0.889	1.024	0.90	0.903			(230Th)/ a	(238U)	-85.013	4.834	4.620	0.682	4.828	4.060	0.829	1387	0.782	0.445	0.542	0.319	0.511	0.488	0.374	0.165	9/6-0	0	0.902	0.920			(230Th)/ a	(200-)	0.684 0.892		071
\$	1.380	0.308	1.015	0.217	1.142	0.986	0.483	1021	500	0.921	1.008	1.558	2 997	100.4	1.183	1.868	3.538			1786.1					.0	2		18	0.289	0.995	0.598	1.703	0.174	0.765	2.322	0.341	1.543	0.246	1.526	2.965	1.085	3.648	2.804	4.057	7:82/	0	0.022	1217.9	į	°C	-	2	1.422 4.220		.000
(238U)/ a (232Th)	8.414	5.320	10.247	7.604	6.274	11.706	5 625	15.492	20.407	17.375	10.235	16.518	24814	4000	13.208	11.035	60.702	1	7577	14094.4	5693.6	10990.9	4.00001	0.277.1	7 Drimon	rillary	(238U)/ a	(232Th)	1.897	12.203	15.911	17.906	7.415	13.164	32.667	4 685	9.819	1.513	28.563	37.081	19.231	31.408	26.450	35.786	36.960	9	3.148 3014.2	9503.3	C.	Primary: 02	(238U)/ a	(14.517		000
\$	0.0012	0.0005	0.0039	0.0020	0.008	0.0036	0.0088	00000	0.0000	0.0298	0.0331	0.0560	0.0816	0.00	0.0347	0.0320	0.0316	ļ	//.	1.52	2.88	2.83	40.0	0.00	March 7 0 2017 Dring	, -0, 201,		18	0.000	0.000	0.005	0.017	900.0	0.027	0.027	0.071	0.119	0.046	0.145	0.462	0.201	0.62	0.824	1.398	1.383	0	10.09	23.16		, 2018		9	1.13 2.04		
is/n		1 0.0045	0.0186	0.0198	0.036	0.0364	1 0 0467	0.0571	2000	0.0807	0.0913	3 0.1720	0 1802	0.1002	0.1876	0.1959	0.2616	0	98.9	7.57	79.8	10.18	10.70	0.00	Morch	Marc		n/Si	0.000	0.002														2.198		000	4.292 58.01	153.56	, C 0 min.	June 3-4,	ij	5	- 4 2 2 3		
Spot	Daff-01b-1.12-2	Daff-01b-1.12-4	Daff-01a-1.6-1	Daff-01a-1.12-1	Daff-01b-1.12-1 (Daff-01b-1.8-2	Daff-01a-112-5	Daff-01h-1 12-9	Dall-010-11.12-1	Daff-01b-1.8-1	Daff-01a-1.5-2 0	Daff-01a-1.12-2	Daff_01a_1 7_1	Da# 01a 1 9 1	Dall-01a-1.8-1	Daff-01a-1./-2	Daff-01b-1.8-4	0.70	2.1-VV-3	BZW-3.2	BZVV-4.1	BZVV-3.1	DZVV-1.1	DZ VV-Z. I	Cocion 2	Z HOISSAC		Spot	Daff_01b-2.5	Daff-01A-2.3	Daff_01b-2.2	Daff-01A-2.4	Daff 01b-2.7	Daff 01b-2.4	Daff 01b-2.8	Daff_01b-2.9	Daff 01b-2.6	Daff-01A-17-3	Daff 01b-2.3	Daff-01A-2.6	Daff 01b-2.1	Daff-01b-1.8-5	Daff-01A-2.5	Daff-01A-2.2	Daπ-01A-2.1-1		NIS1611-1.1 BZW-5.4	BZW-5.1	6 4 6 6 6 6 6	session 3	1000		Daff-8.1 Daff-13.5		6

Session 1 March 6-7, 2017 Primary: O₂

Table DR 3.1 continued:

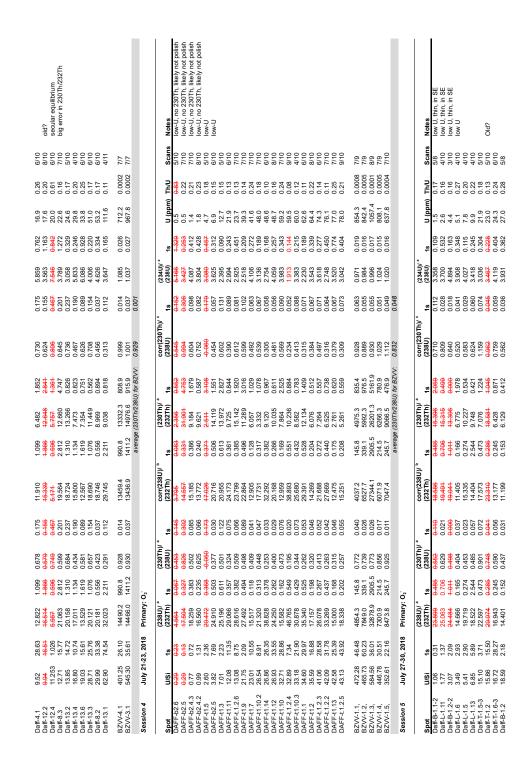


Table DR 3.1 continued:

last five scans	first 5 scans		
5/10 5/10 6/10 3/10 5/10	6/10 5/10 6/10 7/10 6/10	677 677 677 677 677 677	
0.16 0.11 0.18 0.27 0.04	0.21 0.13 0.17 0.12 0.14	1.0108 0.0038 0.0006 0.0003 0.0004 0.0004 0.0003 0.0005 0.0005	
37.3 40.3 55.2 58.8 60.6 65.0	103.8 129.1 131.5 151.8 194.8	4.9 569.0 712.1 792.8 887.6 1008.6	
0.245 0.082 0.303 0.059 0.048	0.223 0.369 0.457 0.100 0.121 0.333	0.52 0.028 0.014 0.004 0.032 0.042 0.013 0.013	
3.967 2.832 3.982 3.977 1.624 3.659	4.689 2.974 3.740 3.073 2.504 3.400	19.98 1.1063 1.063 1.052 0.987 0.874 0.979 0.993	
0.084 0.043 0.041 0.168 0.084 0.027	0.101 0.038 0.123 0.032 0.123	20.743 0.087 0.030 0.019 0.025 0.036 0.027 0.027	0.018
0.564 0.509 0.552 0.672 0.288	0.477 0.519 0.375 0.454 0.568 0.353	-0.036 0.993 0.988 0.989 0.969 1.031 1.033 1.033	
1.384 0.909 1.831 1.694 0.996 0.416	0.729 1.217 0.658 0.635 1.988 0.599	32.99 137.3 1460.2 1428.4 636.7 1954.6 8350.4 4276.5 890.2	J) for BZVV
6.488 14.267 9.966 7.420 36.275 7.436	5.944 13.764 5.630 4.461 8.794 7.813	-0.08 918.9 5877.9 10321.2 5523.7 9521.6 59778.9 14362.3 6007.3	(230Th/238U)
0.250 0.430 0.238 1.645 3.222 0.199	0.282 0.259 0.152 0.352 0.393 0.429	18.9 246.3 619.6 234.2 596.4 7218.8 1492.7 326.3	average
19.647 28.938 17.518 11.778 88.249 18.511	15.241 23.941 14.349 18.989 25.426 22.851	820.4 5353.5 9982.2 5659.8 8732.9 60857.8 12485.4 5786.4 95199.7	
0.082 0.039 0.036 0.167 0.082	0.099 0.033 0.122 0.027 0.122 0.051	20.74 0.085 0.024 0.005 0.018 0.032 0.011 0.021 0.018	
0.438 0.396 0.429 0.523 0.224	0.371 0.404 0.291 0.353 0.442 0.275	-0.03 0.772 0.769 0.769 0.769 0.802 0.777 0.794 0.803	
0.249 0.430 0.238 1.644 3.222 0.199	0.282 0.258 0.151 0.351 0.393	0.072 18.9 246.3 619.6 596.4 7218.8 1492.7 326.336	
25.264 37.210 22.526 15.145 113.476 23.803	19.597 30.785 18.450 24.417 32.695 29.383	3.102 1054.9 6883.9 12835.8 7277.748 11229.3 78255.2 16054.6 7440.516	
21.89 50.37 85.79 25.60 4.07 32.06	61.81 37.40 41.63 54.23 30.92 73.25	1.03 56.1 76.19 48.03 88.44 50.78 81.5	
25.71 27.78 38.04 40.56 41.77	71.54 88.99 90.64 104.62 128.40	3.37 392.1 490.76 546.34 611.74 695.10 737.4	
Daff-T-1.6 Daff-T-1.6-2 Daff-T-1.1-4 Daff-T-1.8 Daff-T-1.1	Daff-T-1.8-3 Daff-T-1.1 Daff-T-1.8-2 Daff-T-1.8-6 Daff-L-1.2 Daff-T-1.8-5	NIST611-1.1 BZWs-1.1 BZW-2.1 BZW-2.1 BZW-1.9 BZWs-1.2 BZWs-1.2 BZW-1.10 BZW-1.3 BZW-1.3 BZW-1.3	

ratios highlighted red and crossed-out were omitted from Fig. 4 (see Notes and text or criteria)
234J was not measured in sessions 1 and 2, and therefore the 234U/239U values are blank
parentheses denote adviviratios
... measured ratios
b.: measured (238U)/222Th*; (session average (230Th)/(238U) for BZVVJ to correct for mass discrimination
c.: measured (230Th)/(238U) / [session average (230Th)/(238U) for BZVVJ to correct for mass discrimination

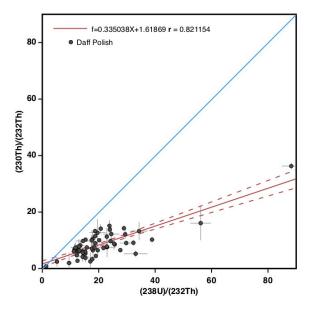


Figure DR 3.1 ²³⁰Th/²³²Th vs ²³⁸U/²³²U for Daff dome sample set (Table 3). ²³⁰Th/²³²Th intercept near zero justifies age determinations presented in figure 4 and that ²³⁰Th as excluded from polish formation

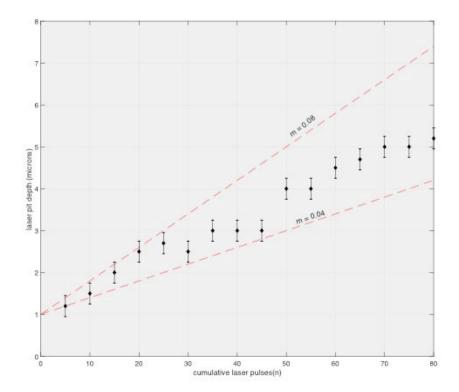


Figure DR4.1 Laser pit depth vs. number of laser bursts for a glacial polish sample from Daff dome. Measurement of the laser pit depth at 5 burst intervals provides a rough estimate of the ablation rate and is used to assign approximate depths to compositional data (Fig. 1-2). A total of 80 pulses were used for all analyses presented in this manuscript. The ablation rate varies between 0.08 μ m/pulse at shallow depths decreasing to 0.05-0.06 μ m/pulse at greater depths. A maximum 0.08 μ m/pulse was assumed to the depth assignments in figures 1 and 2 of the main text. This rate does not influence the data reduction in any way and only permits rough estimates on the depth of chemical transitions to be assigned.