

Data Repository for Subduction initiation without magmatism: The case of the missing Alpine magmatic arc, Anders McCarthy, Cyril Chelle-Michou, Othmar Müntener, Richard Arculus, Jon Blundy

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Supplementary Methods for Detrital Zircon Treatment

Zircon U-Pb data from 129 Triassic to Quaternary detrital samples were compiled from available sources (See supplementary references below). When data table were not attached to the publications authors were contacted and kindly provided them. The $^{206}\text{Pb}/^{238}\text{U}$, $^{207}\text{Pb}/^{235}\text{U}$, $^{207}\text{Pb}/^{206}\text{Pb}$ and concordia dates for each analysis were computed with IsoplotR (Vermeesch, 2018).

The percentage of concordance was estimated based on the two U-Pb dates as $100 \times (^{206}\text{Pb}/^{238}\text{U} \times ^{207}\text{Pb}/^{235}\text{U})$. This approach is more relevant for Cenozoic to Mesozoic dates (which are the focus of this study) than the commonly used percentage of concordance defined as $100 \times (^{206}\text{Pb}/^{238}\text{U} \times ^{207}\text{Pb}/^{206}\text{Pb})$. This is because for such young dates, very small variations of the $^{207}\text{Pb}/^{206}\text{Pb}$ ratio causes very large variation of the $^{207}\text{Pb}/^{206}\text{Pb}$ date such that perfectly concordant zircon (i.e., overlap with the concordia) would be identified as excessively discordant using a common discordance criteria of 10%. We considered a strict concordance threshold of 95–105%. Out of 12524 analyses, 9268 passed the concordance test. The $^{206}\text{Pb}/^{238}\text{U}$ and $^{207}\text{Pb}/^{206}\text{Pb}$ dates were considered for analysis younger and older than 1200 Ma, respectively. Kernel density estimates of the “concordant” population was plotted with DensityPlotter (Vermeesch, 2012).

Zircon U-Pb Methodology References

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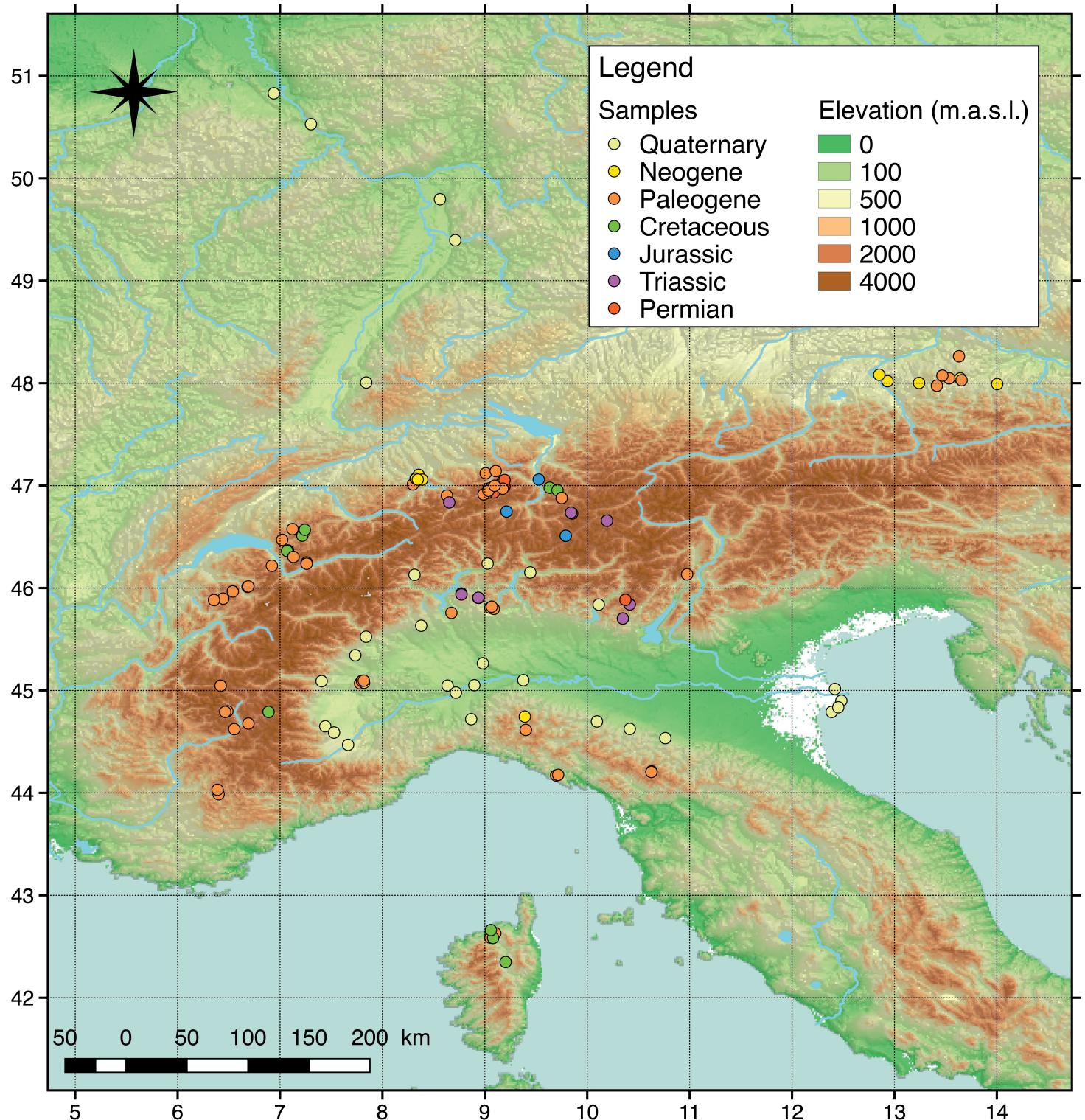
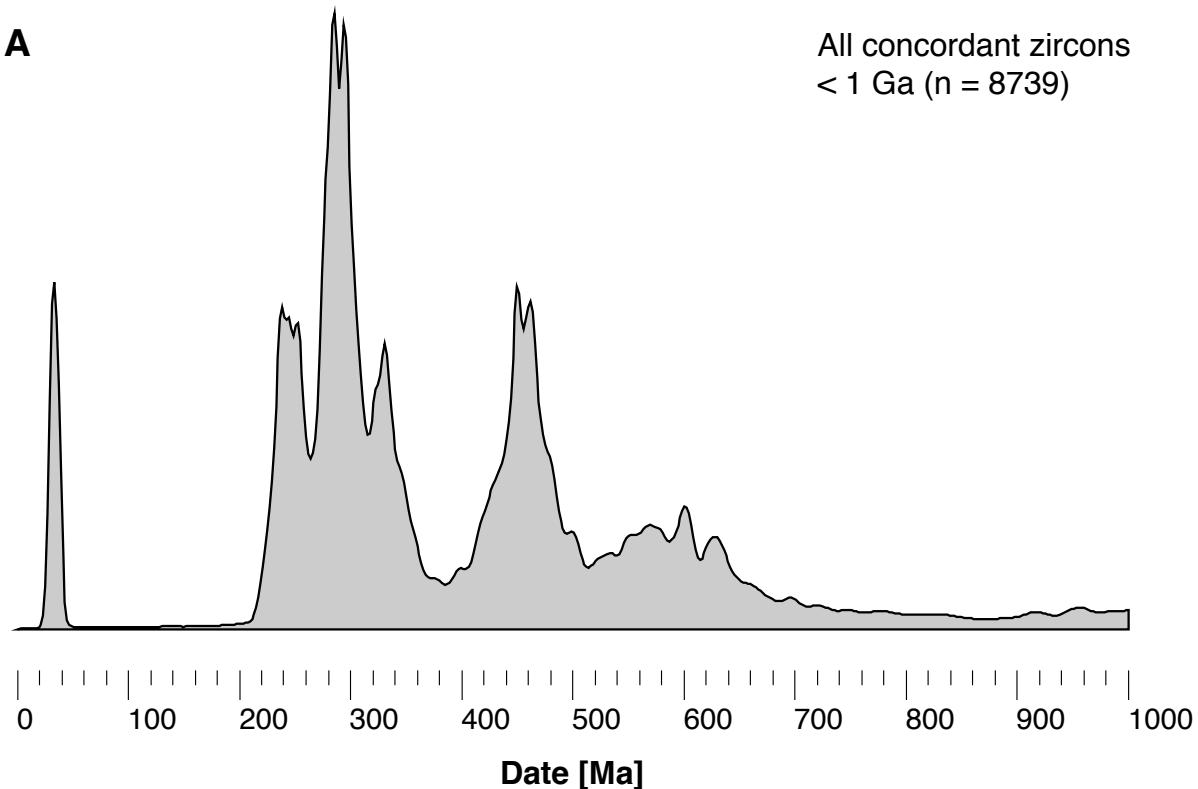


Fig. DR1: Map of the European Alps and Apennines showing the locations of samples used for detritial zircon analysis (see References sections in appendix). Note that Quaternary to Paleogene samples are sampled from all major watersheds: Rhône, Rhine, Po and Danub watersheds.

A

All concordant zircons
< 1 Ga (n = 8739)

**B**

All concordant
zircons (n = 9268)

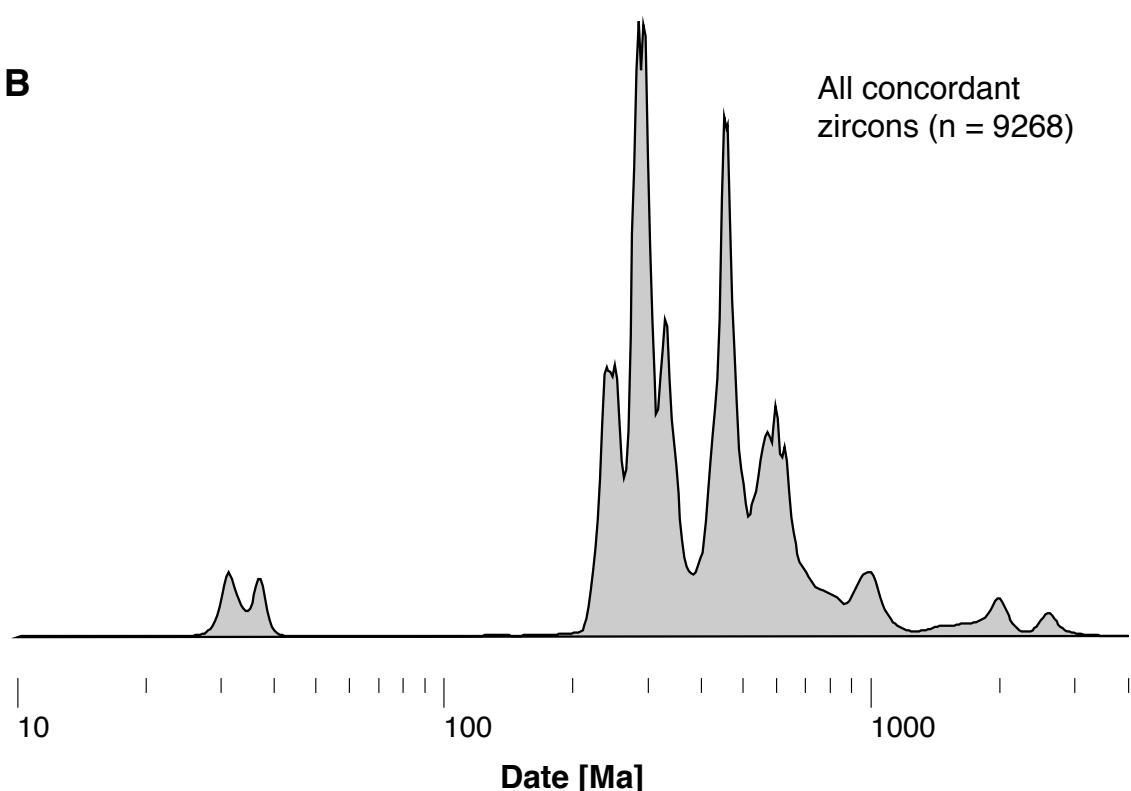


Fig. DR2: Detrital zircon U-Pb date distribution for (A) grains younger than 1 Ga and (B) all grains (note the log scale).

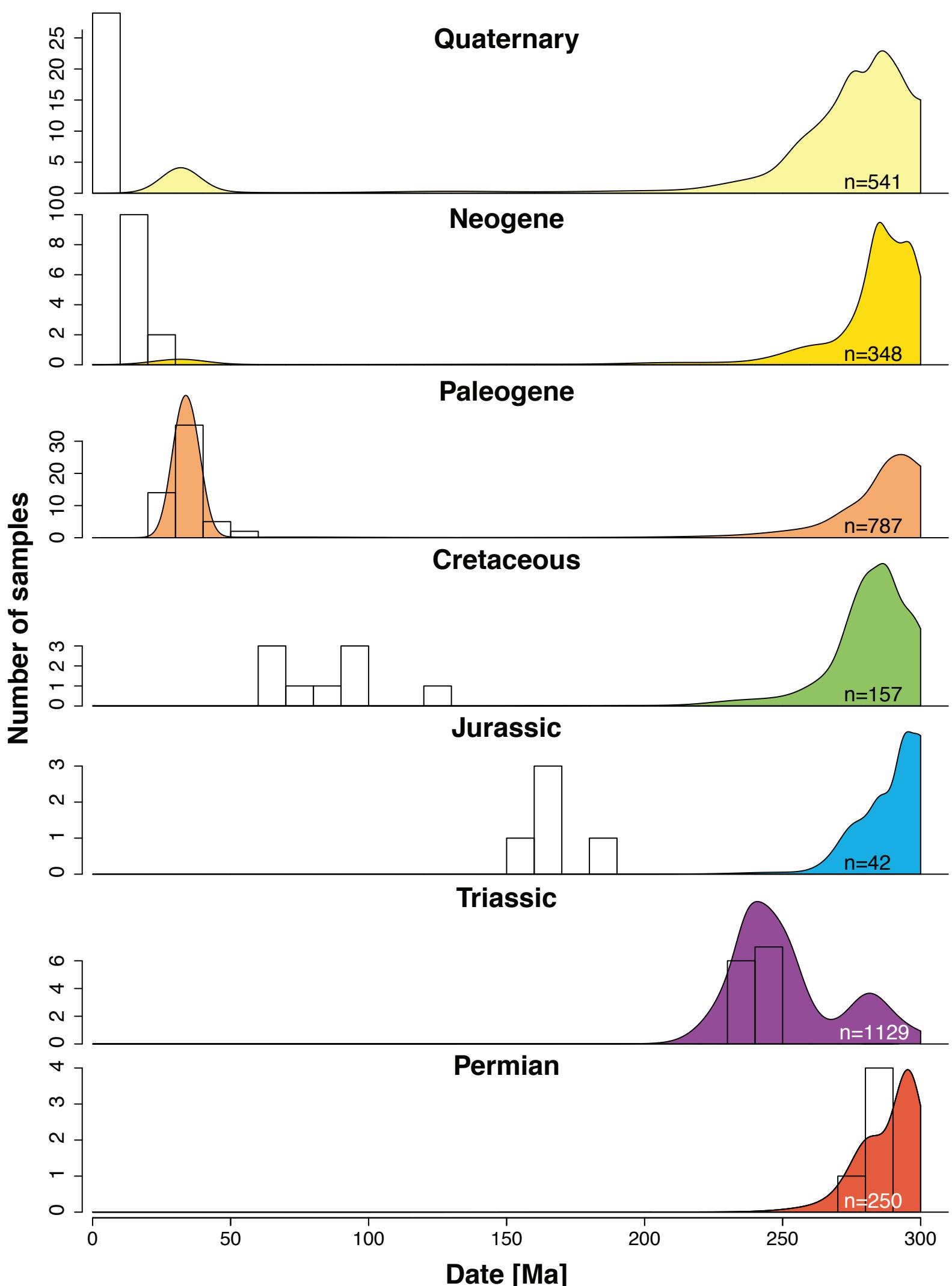


Fig.DR3: Detrital zircon U-Pb date distribution (colored curves) compared to depositional age (white histograms) from Permian (bottom) to Present (Top).

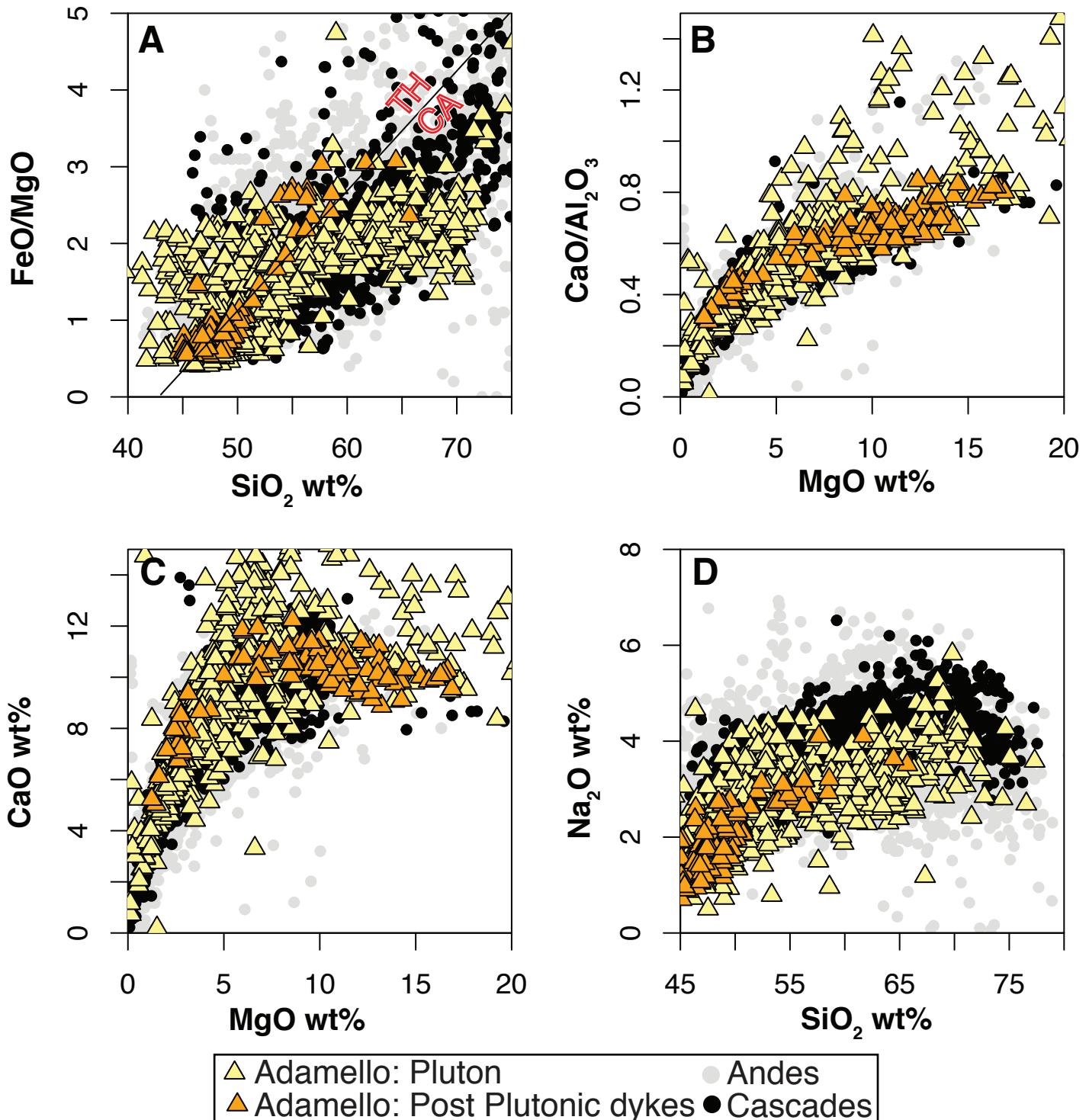


Fig. DR4: Geochemical plots showing the compositional evolution of the Adamello Pluton and post-plutonic dykes compared to the Andes (GEOROC - Geochemical database: n = 13912 volcanic rocks) and Cascades (n = 1679 published volcanic rocks). A: FeO*/MgO vs SiO₂ wt% of Myashiro (1974) showing the arc tholeiitic (TH) to low-Fe «calc-alkaline» (CA) compositional trends of the Adamello Pluton; B-D: CaO/Al₂O₃ and CaO wt% vs MgO wt% and Na₂O wt% vs SiO₂ wt%. Adamello post-plutonic dykes from Hürlimann *et al.*, (2016). Adamello pluton compilation (n = 675) includes the composition of granitoids-tonalites-diorites-gabbros (authors' unpublished data) and data from Macera *et al.*, (1985) and Blundy and Sparks (1992)

Supplementary References for Figures 2 to 4

Figure 2 Subduction initiation:

Data for the Alps include papers cited in the manuscript and :

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Figure 3: Alpine Compilation:

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Figure 4

Present-day Alpine-cross-section after:

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