

Wang and Liu et al.

2 Compositional transition in natural alkaline lavas through 3 silica-undersaturated melt-lithosphere interaction

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7 This file concludes:

8 1. ANALYTICAL METHODS

9 2. MODELLING DETAILS

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13

14 ANALYTICAL METHODS

15 Zinc isotope ratios were measured at the Isotope Geochemistry Laboratory of
16 the China University of Geosciences, Beijing using the method reported in previous
17 studies ([Liu et al., 2014, 2016](#)), which is modified after Maréchal et al. ([1999](#)).

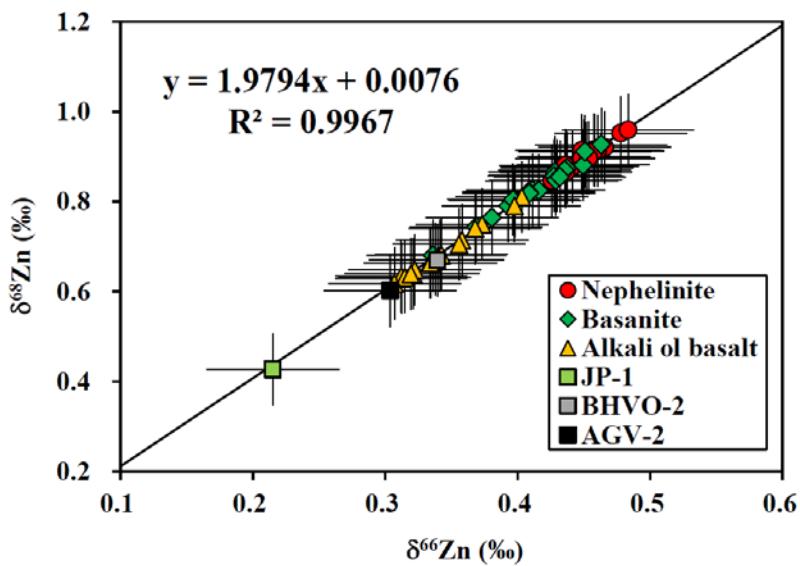
18 Before measurement, chemical purification of Zn was conducted. In brief,
19 approximately 10 mg of samples were totally dissolved in an HF-HNO₃ mixture,
20 followed by re-dissolution in aqua regia and 8 N HCl. Afterwards, 1 ml of sample
21 dissolved in 8 N HCl was loaded onto 2 ml pre-cleaned Bio-Rad® strong anion resin
22 AG-MP-1M. Zinc was eluted in 10 ml of 0.5 N HNO₃ after removal of matrix
23 elements, copper and iron rinsed by 8 N HCl and 2 N HCl, respectively. After
24 evaporation to dryness at 80 °C, the residues were dissolved in 1 ml of 3% HNO₃.

25 The Mg/Zn ratio was checked before isotopic measurement since the presence of
26 argides (Mg²⁴Ar⁺ and Mg²⁶Ar⁺) significantly influences the ratios of ⁶⁶Zn/⁶⁴Zn
27 ([Mason et al., 2004; Wang et al., 2017](#)). The recovery of Zn is over 99% and the
28 whole procedural blank is less than 6 ng and considered to be negligible compared
29 with the collected Zn (>1 µg).

30 Zinc isotopic analysis was performed on a Thermo Scientific *Neptune plus*
31 multiple collector inductively coupled plasma mass spectrometry (MC-ICP-MS)
32 instrument using sample-standard bracketing (SSB) method for mass bias correction.

33 The data are reported in δ-notation in per mil relative to JMC 3-0749L Zn
34 standard: $\delta^{68/66}\text{Zn} = [({}^{68/66}\text{Zn}/{}^{64}\text{Zn})_{\text{sample}}/({}^{68/66}\text{Zn}/{}^{64}\text{Zn})_{\text{JMC 3-0749L}} - 1] \times 1000$. The
35 external reproducibility is ±0.05‰ (2sd) for δ⁶⁶Zn and ±0.10‰ (2sd) for δ⁶⁸Zn

36 based on long-term analyses of synthetic pure Zn solutions and igneous rock
37 standards. All analyzed samples fall on a mass-dependent fractionation line on a
38 three-isotope plot ($\delta^{68}\text{Zn}$ vs. $\delta^{66}\text{Zn}$) with a slope of 1.979 ± 0.007 (1σ ; $R^2 = 0.997$; N
39 = 60), indicating that there are no analytical artifacts from unresolved isobaric
40 interferences. Analyses of three geostandards going through the same chemical
41 procedures as the studied samples yielded $\delta^{66}\text{Zn}$ values of $0.34 \pm 0.05\text{\textperthousand}$ for
42 BHVO-2, $0.30 \pm 0.02\text{\textperthousand}$ for AGV-2 and $0.22 \pm 0.03\text{\textperthousand}$ for JP-1, which are
43 consistent with the values reported in the literature (e.g., Sossi et al., 2015; Chen et
44 al., 2016; Wang et al., 2017).



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47 **MODEL OF MELT-LITHOSPHERE INTERACTION**

48 Parameters used in the calculation

Name	$\delta^{66}\text{Zn}$ (‰)	[Zn] ($\mu\text{g/g}$)	ε_{Nd}	[Nd] ($\mu\text{g/g}$)
Basanitic melt ¹	0.45	120	5	40
Orthopyroxene ²	0.15	30		
SCLM-derived melt ³	0.23	75	-10	10

49 ¹Data for basanitic melt are from this study.50 ²Data for orthopyroxene are from Wang et al. (2017).

51 ³Zn isotope data for SCLM-derived melt are based on the >120 Ma mafic rocks
 52 reported in Liu et al. (2016) and Nd isotope data are based on the data presented in
 53 Fig. DR5 (see below).

54 **Melt-opx reaction scenario**

55 The reaction is modelled assuming that peridotite is impregnated by a finite
 56 amount of basanitic melt. The equation of melt-opx reaction is 1 melt⁰ + 0.37 opx →
 57 0.80 melt¹ + 0.20 ol + 0.36 cpx (Lambart et al., 2012). $\delta^{66}\text{Zn}$ and [Zn] of reacted
 58 melt (*melt*¹) can be obtained through the mass balance equation:

59 $[\text{Zn}]_{\text{melt}}^0 + [\text{Zn}]_{\text{opx}} \times 0.37 = [\text{Zn}]_{\text{melt}}^1 \times 0.80 + [\text{Zn}]_{\text{ol}} \times 0.20 + [\text{Zn}]_{\text{cpx}} \times 0.36;$

60 $[\text{Zn}]_{\text{melt}}^0 \times \delta^{66}\text{Zn}_{\text{melt}}^0 + [\text{Zn}]_{\text{opx}} \times \delta^{66}\text{Zn}_{\text{opx}} \times 0.37 = [\text{Zn}]_{\text{melt}}^1 \times \delta^{66}\text{Zn}_{\text{melt}}^1 \times 0.80 + [\text{Zn}]_{\text{ol}} \times \delta^{66}\text{Zn}_{\text{ol}} \times 0.20 + [\text{Zn}]_{\text{cpx}} \times \delta^{66}\text{Zn}_{\text{cpx}} \times 0.36;$

62 $[\text{Zn}]_{\text{ol}} / [\text{Zn}]_{\text{melt}}^1 \approx D_{\text{Zn}}^{\text{ol/melt}} = 0.99, [\text{Zn}]_{\text{cpx}} / [\text{Zn}]_{\text{melt}}^1 \approx D_{\text{Zn}}^{\text{cpx/mel}} = 0.48$ (Le Roux et al.,
 63 2011); $\delta^{66}\text{Zn}_{\text{melt}}^1 = \delta^{66}\text{Zn}_{\text{ol}} = \delta^{66}\text{Zn}_{\text{cpx}}.$

64 Given that *f* mass fraction of melt participates in the reaction, the final $\delta^{66}\text{Zn}$ and
 65 [Zn] of the ascending melt after reactive infiltration into the peridotite should be:

66 $[\text{Zn}]_{\text{final melt}} = [(1-f) \times [\text{Zn}]_{\text{melt}}^0 + f \times [\text{Zn}]_{\text{melt}}^1 \times 0.80] / (1-f + f \times 0.80);$

67 $\delta^{66}\text{Zn}_{final\ melt} = [(1-f) \times [\text{Zn}]_{melt}^0 \times \delta^{66}\text{Zn}_{melt}^0 + f \times [\text{Zn}]_{melt}^I \times \delta^{66}\text{Zn}_{melt}^I \times 0.80] / [(1-f) \times$
68 $[\text{Zn}]_{melt}^0 + f \times [\text{Zn}]_{melt}^I \times 0.80].$

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70 **Melt-melt mixing scenario**

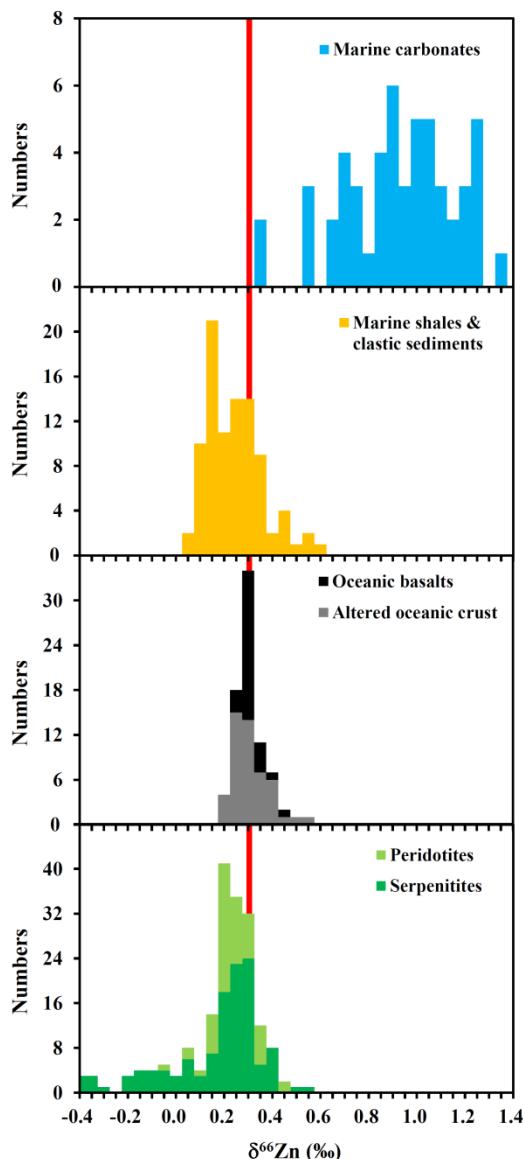
71 If f proportion of SCLM-derived melt is mixing with the basanitic melt, the $\delta^{66}\text{Zn}$
72 and $[\text{Zn}]$ of the mixed melt should be:

73 $[\text{Zn}]_{final\ melt} = ([\text{Zn}]_{melt}^0 \times (1-f) + [\text{Zn}]_{SCLM\ melt} \times f);$

74 $\delta^{66}\text{Zn}_{final\ melt} = (\delta^{66}\text{Zn}_{melt}^0 \times [\text{Zn}]_{melt}^0 \times (1-f) + \delta^{66}\text{Zn}_{SCLM\ melt} \times [\text{Zn}]_{SCLM\ melt} \times f) / [\text{Zn}]_{final\ melt}.$

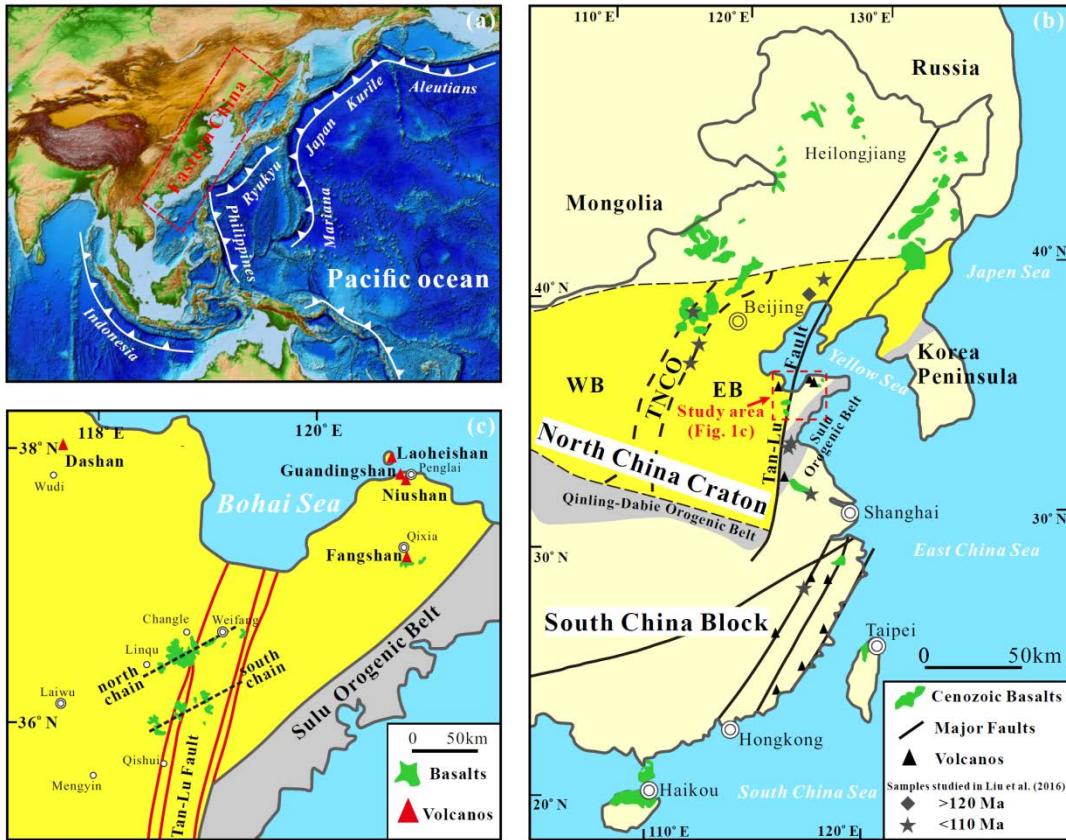
75 Similarly, the ε_{Nd} and $[\text{Nd}]$ of the mixed melt can also be calculated.

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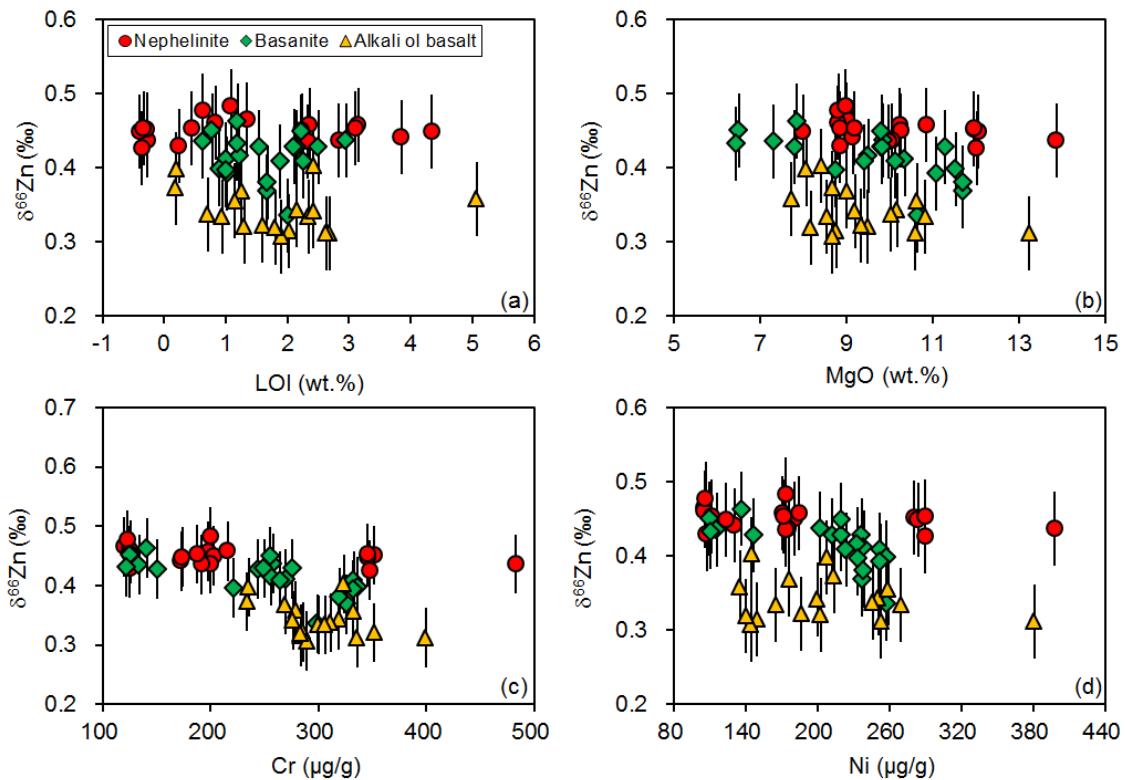
78 **Figure DR1.** A compilation of zinc isotope compositions of marine carbonates
 79 (Pichat et al., 2003), marine shales and clastic sediments (Bentahila et al., 2008;
 80 Pons et al., 2011), oceanic basalts (Herzog et al., 2009; Chen et al., 2013; Wang et
 81 al., 2017), altered oceanic crust (Huang et al., 2016), serpentinites (Pons et al., 2011;
 82 Pons et al., 2016) and peridotites (Doucet et al., 2016; Wang et al., 2017). The red
 83 line refers to the mean value of MORB (Wang et al., 2017).



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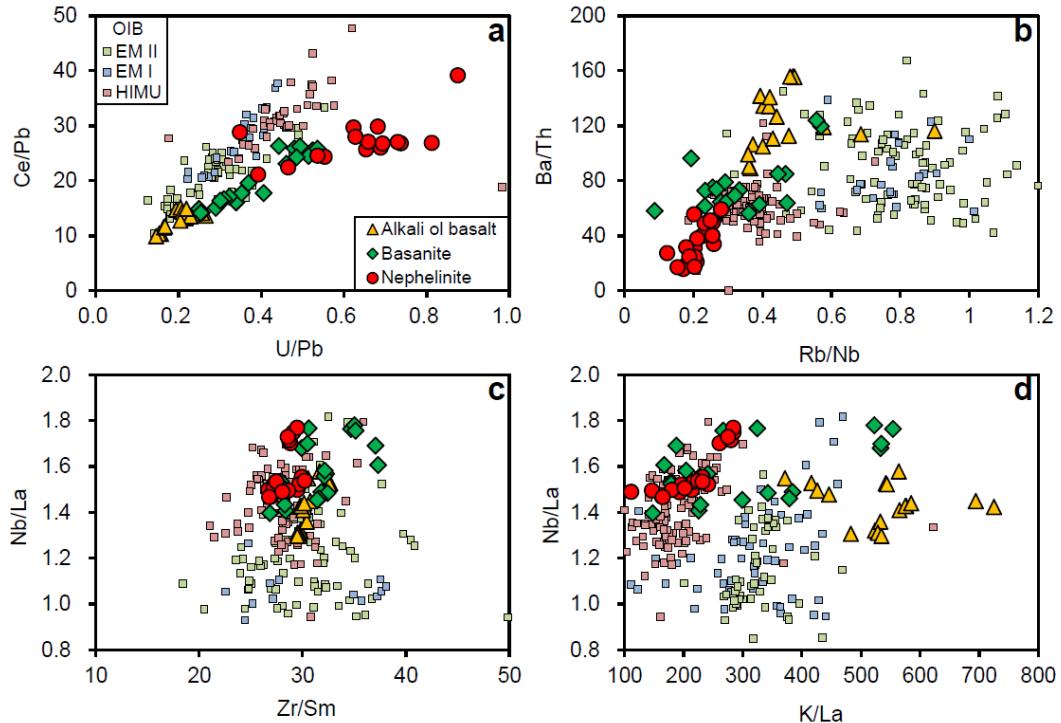
85 **Figure DR2.** Schematic maps showing the distribution of Cenozoic alkali basalts in
 86 eastern China. (a) Digital topographic map of the western Circum-Pacific region
 87 (from <http://www.ngdc.noaa.gov/mgg/global/global.html>). Distribution of Cenozoic
 88 alkali basalts in (b) eastern China and in (c) Shandong Peninsula. Eastern China
 89 comprises two major blocks: North China Craton (NCC) and South China Block
 90 (SCB), which are separated by the Qinling-Dabie-Sulu orogenic belt. The NCC can
 91 be further divided into three parts: Western Block (WB), Eastern Block (EB) and
 92 Trans-North China Orogen (TNCO) ([Zhao et al., 2001](#)). The Shandong Peninsula is
 93 located at the southeastern margin of the EB. Cenozoic alkaline magmatism in
 94 Shandong can be divided into two episodes: 24.0-10.3 Ma and 8.7-0.3 Ma ([Zeng et](#)
 95 [al., 2010](#)). The earlier magmatism is mainly located close to the Tan-Lu Fault, which
 96 is a deep fault that extends into the lithospheric mantle (e.g., [Xu et al., 1987](#)),

97 forming two large parallel volcanic chains: the north chain and the south chain. The
98 later magmatism is widely distributed far from the Tan-Lu Fault in the form of small
99 isolated volcanoes ([Fig. 1c](#)) ([Zeng et al., 2010, 2011](#)). Locations of samples studied
100 in Liu et al. ([2016](#)) are also shown for comparison in [Fig. 1b](#).
101



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103 **Figure DR3.** Plots of $\delta^{66}\text{Zn}$ against loss-on-ignition (LOI), MgO, Cr and Ni
 104 contents, respectively (a-d). The absence of relationships between $\delta^{66}\text{Zn}$ and these
 105 elements suggests that post-magmatic alteration and magmatic differentiation have a
 106 negligible impact on the Zn isotopic compositions of Shandong alkaline lavas.
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109 **Figure DR4.** Comparison of trace elemental ratios between Shandong alkaline lavas

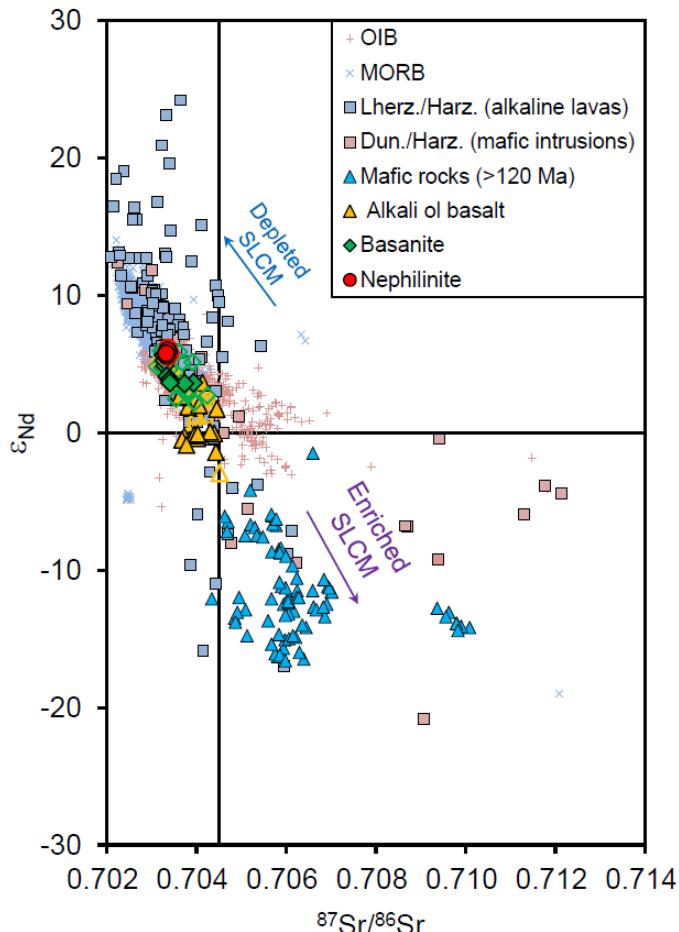
110 and typical HIMU-, EM I- and EM II-type OIBs. (a) Ce/Pb vs. U/Pb; (b) Ba/Th vs.

111 Rb/Nb; (c) Nd/La vs. Zr/Sm; (d) Nb/La vs. K/La. Trace element data of OIBs are

112 from Willbold and Stracke (2006) and those of Shandong alkaline lavas are from

113 Zeng et al. (2010).

114



115

116 **Figure DR5.** ϵ_{Nd} versus $^{87}\text{Sr}/^{86}\text{Sr}$; Solid colorful symbols represent samples of this
 117 study, and open symbols represent samples from other literature (Xu et al., 2012;
 118 Sakuyama et al., 2013; Li et al., 2014; Li et al., 2016; Li et al., 2017). Data for OIB
 119 and MORB are compiled from GeoRoc (<http://georoc.mpch-mainz.gwdg.de/georoc>)
 120 and PetDB (www.earthchem.org/petdb) respectively. Peridotite xenoliths and > 120
 121 Ma mafic rocks are from literature (Song and Frey, 1989; Fan et al., 2000; Zhang et
 122 al., 2002; Zhang et al., 2003; Rudnick et al., 2004; Zhang et al., 2004; Tang et al.,
 123 2008; Xu et al., 2008a; Xu et al., 2008b; Yang and Li, 2008; Chu et al., 2009; Xiao
 124 et al., 2010; Xu et al., 2010; Tang et al., 2011; Zhang et al., 2011; Liu et al., 2012;
 125 Tang et al., 2013).

126 **Table DR1.** Zinc isotopic compositions (‰), selected element concentrations (major elements in wt. ‰; trace elements in µg/g), element ratios

127 (w/w) and initial Sr-Nd isotopic ratios for Shandong alkaline lavas.

Sample no.	SiO ₂	MgO	Na ₂ O	K ₂ O	LOI	Cr	Ni	Zn	Nb	Th	Ce/Pb	Ba/Th	Zn/Fe	La/Yb	$\delta^{66}\text{Zn}$	2SD	$\delta^{68}\text{Zn}$	2SD	(⁸⁷ Sr/ ⁸⁶ Sr) _i	$\varepsilon_{\text{Nd(t)}}$
<i>Nephelineite</i>																				
07JSS01	40.9	12.0	4.7	2.5	1.3	346	290	154	148	10.2	28.0	51.1	14.3	42.3	0.45	0.03	0.90	0.05	0.70336	5.99
07JSS02	40.7	13.9	4.5	2.2	0.8	483	398	149	153	9.7	21.1	59.3	13.9	43.1	0.44	0.03	0.87	0.04	0.70339	5.93
07JSS03	41.1	12.0	4.7	2.4	0.6	348	290	153	147	10	24.6	40.2	14.1	42.8	0.43	0.02	0.85	0.05	0.70334	5.75
07JSS04	40.9	12.0	4.8	2.4	0.2	352	281	154	148	9.8	22.5	34.0	14.1	41.6	0.45	0.04	0.90	0.07	0.70334	5.6
07JSS05	40.9	12.1	4.7	2.4	2.4	346	284	153	147	10.1	24.4	49.8	14.1	41.7	0.45	0.02	0.90	0.03	0.70334	5.75
08LHS01	39.4	8.8	5.5	2.7	0.5	123	107	174	151	13.6	166.3	31.5	15.6	62.8	0.48	0.06	0.95	0.13		
08LHS02	39.5	8.9	5.5	2.7	2.2	126	112	184	146	13.8	120.3	37.9	16.4	63.8	0.45	0.01	0.90	0.04		
08LHS03	39.4	8.9	5.7	2.9	2.8	125	108	176	149	13.6	131.3	49.1	15.7	62.8	0.43	0.04	0.86	0.06		
08LHS04	39.2	9.0	4.7	2.6	3.2	120	106	178	144	13.7	39.2	39.7	15.9	65.7	0.47	0.01	0.92	0.01		
08LHS05	39.4	8.8	4.9	2.7	3.9	126	106	181	132	13.9	29.9	21.0	16.1	64.2	0.46	0.04	0.91	0.07		
08LHS06	39.5	9.2	3.4	1.6	2.3	173	131	169	146	11.9	28.8	24.6	15.8	56.4	0.44	0.03	0.88	0.05		
08LHS07	40.4	8.0	2.9	1.3	3.1	174	124	183	144	13.1	29.7	27.4	16.3	59.5	0.45	0.02	0.91	0.04		
08QJFS01	40.2	9.0	5.2	2.8	1.1	200	174	174	150	15.7	27.1	17.5	16.5	59.5	0.48	0.04	0.96	0.07		
08QJFS02	39.7	9.2	3.4	2.0	4.3	188	172	168	144	15.7	26.8	55.6	16.4	60.2	0.45	0.04	0.90	0.05		
08QJFS03	39.6	10.0	4.1	2.4	-0.3	192	174	172	116	15.6	27.1	24.8	16.6	58.6	0.44	0.03	0.88	0.06		
08QJFS04	39.5	10.3	4.2	2.3	-0.3	198	171	173	119	16.0	26.9	18.2	16.7	57.0	0.5	0.05	0.92	0.09		
08QJFS05	39.5	10.3	4.2	2.3	-0.4	203	181	170	121	16.3	25.7	31.6	16.5	59.0	0.5	0.02	0.90	0.06		
08QJFS06	39.5	10.9	3.7	2.0	-0.3	216	185	167	124	16.5	26.8	17.2	16.5	57.2	0.5	0.02	0.91	0.08		
08QJFS07	39.5	10.1	4.0	2.2	-0.4	200	176	171	120	16.7	26.0	15.8	16.5	59.4	0.4	0.06	0.88	0.04		

Basanite

07FS01	43.7	8.8	4.2	1.3	2.5	221	234	110	68	4.6	17.7	73.2	11.4	30.1	0.4	0.05	0.80	0.06	0.70329	5.23
07FS03	43.2	9.5	4.1	0.9	1	256	233	115	66	4.6	19.6	64.5	11.6	31.3	0.42	0.04	0.83	0.07	0.7033	4.86
07FYS07	44.3	10.7	2.8	0.9	3	299	258	100	38	2.3	14.8	96.2	10.8	16.3	0.34	0.02	0.68	0.05	0.70343	5.81
07LIS01	42.8	9.9	3.9	1.0	1.5	258	203	97	75	5.5	17.7	57.9	11.1	29.1	0.44	0.05	0.87	0.04	0.70332	5.15
07LS04	44	10.1	2.6	1.8	1.2	265	224	109	57	4.1	16.6	69.2	12.0	28.7	0.41	0.04	0.82	0.08	0.70341	3.84
07LS05	43.8	9.8	3.1	1.8	1.2	250	220	113	64	4.8	15.0	61.4	12.2	33.3	0.4	0.00	0.85	0.01	0.7034	3.86
07LS06	44.5	9.8	2.9	1.4	1.9	255	220	107	54	3.9	16.4	73.7	11.9	26.6	0.5	0.03	0.88	0.02	0.70341	3.63
07LS07	43.7	9.8	2.8	2.0	0.9	244	213	111	61	4.4	17.1	65.4	12.1	31.2	0.4	0.01	0.86	0.04	0.70335	4.08
07NIS02	44.3	11.7	2.3	1.8	1.0	320	239	97	46	3.2	14.2	123.8	11.1	19.2	0.4	0.04	0.76	0.10	0.70374	3.57
07NIS04	44.2	11.7	2.4	1.8	1.0	326	238	96	46	3.3	14.2	119.7	10.9	19.0	0.4	0.02	0.74	0.04	0.70394	3.63
07THS02	43.6	11.5	3.0	2.3	2.2	336	258	95	59	4.1	16.1	84.8	10.7	21.4	0.4	0.03	0.80	0.04	0.70327	5.64
07THS03	43.7	11.1	2.9	2.2	2.1	332	253	96	61	4.2	16.1	84.8	10.8	21.5	0.4	0.02	0.79	0.02	0.70333	5.77
08GDS01	43	7.8	4.8	2.2	2.3	151	147	160	110	10.6	26.2	75.1	15.8	49.6	0.4	0.03	0.87	0.04		
08GDS02	43.1	7.9	5.0	2.2	1.7	141	137	162	112	10.7	26.3	72.7	16.1	50.1	0.5	0.04	0.93	0.01		
08GDS03	43.4	9.4	4.8	1.3	2.0	333	252	129	96	9.2	24.5	78.8	13.8	41.1	0.4	0.01	0.82	0.05		
08GDS04	41.7	11.3	4.6	1.5	1.7	276	237	127	106	9.5	25.5	51.1	14.2	41.8	0.4	0.03	0.86	0.06		
08GDS05	42.7	10.4	5.3	1.6	0.6	270	239	127	108	9.6	25.8	58.7	13.8	42.1	0.4	0.03	0.83	0.09		
08NIUS01	44.4	7.3	5.6	1.9	2.2	134	116	163	102	9.3	24.2	56.3	16.2	44.2	0.44	0.03	0.87	0.02		
08NIUS02	45	6.5	5.5	1.8	0.8	125	110	163	101	9.1	25.8	62.7	16.1	43.6	0.45	0.01	0.91	0.01		
08NIUS03	45.1	6.5	5.5	1.6	1.2	122	111	160	98	8.7	23.1	63.7	16.3	42.3	0.43	0.03	0.86	0.04		
<i>Alkali olivine basalt</i>																				
07CDS01	45.9	10.6	2.3	1.6	2.6	336	253	94	32	2.1	14.7	134.4	11.3	17.5	0.31	0.02	0.63	0.07	0.70373	-0.29
07CDS02	46.3	10.8	2.2	1.7	2.3	300	270	99	35	2.4	13.9	126.5	11.6	18.5	0.33	0.03	0.66	0.06	0.70397	-0.18
07CDS03	45.4	13.3	1.7	1.3	2.7	399	381	92	30	1.9	13.2	89.1	10.8	15.6	0.31	0.06	0.63	0.12		
07CDS04	46	7.7	1.9	1.5	5.1	279	135	89	33	2.2	15.1	134.1	11.5	17.3	0.36	0.03	0.71	0.06	0.70402	-0.41
07CDS05	47.1	9.5	2.7	1.5	1.3	352	203	93	31	2.0	14.4	141.7	11.3	16.9	0.32	0.01	0.64	0.01	0.70369	-0.51

07CDS07	47.6	8.6	2.7	1.6	0.9	306	166	92	32	2.2	14.9	140.7	11.0	17.0	0.33	0.05	0.66	0.03	0.70378	-0.9
07FYS01	47.4	8.8	2.5	1.4	2	284	150	92	28	1.8	11.5	155.8	11.4	14.5	0.31	0.03	0.63	0.07	0.70396	-0.2
07FYS02	47.6	8.2	2.7	1.5	1.8	284	141	93	29	1.9	11.7	155.7	11.4	14.4	0.32	0.03	0.64	0.04	0.70404	-0.1
07FYS03	47.6	8.7	2.5	1.4	1.9	289	145	91	28	1.8	11.5	155.5	11.3	14.7	0.31	0.05	0.62	0.03	0.70396	-0.16
07FYS04	46.4	9.2	2.6	1.5	2.4	276	200	106	44	3.2	13.6	105.1	12.2	22.4	0.34	0.07	0.68	0.12	0.70382	1.89
07FYS06	46	8.4	2.6	1.5	2.4	324	145	91	50	3.4	13.7	106.4	11.7	25.0	0.4	0.01	0.81	0.03	0.70414	3.63
07FYS09	47.4	9.0	2.5	1.5	1.3	269	177	95	41	2.9	13.5	112.6	11.6	18.7	0.37	0.00	0.74	0.00	0.70446	1.72
07FYS10	47.2	9.4	2.5	1.4	1.6	286	187	94	41	2.9	13.5	110.5	11.4	18.3	0.32	0.04	0.65	0.01	0.70406	2.01
07LIS02	46.9	8.7	3.5	2.0	0.2	234	214	103	46	3.2	14.2	90.4	11.8	22.3	0.37	0.01	0.75	0.07	0.70361	2.57
07LIS04	47.5	8.1	3.5	2.0	0.2	236	208	105	46	3.2	12.7	99.0	11.8	22.3	0.4	0.03	0.79	0.06	0.70363	2.65
07SW01	47.4	10.0	2.8	1.6	0.7	311	246	101	26	2.2	10.4	113.7	11.6	12.9	0.34	0.01	0.67	0.00	0.70431	0.06
07SW02	47	10.2	2.5	1.6	2.2	319	251	100	27	2.2	10.3	116.2	11.6	13.2	0.34	0.02	0.68	0.03	0.7044	-0.06
07SW03	47.4	10.6	2.6	1.2	1.2	332	259	101	24	2.1	9.9	118.9	11.6	12.6	0.36	0.04	0.71	0.08	0.70444	-1.42

Rock standards

JP-1	0.22	0.03	0.43	0.01
BHVO-2	0.34	0.05	0.67	0.08
AGV-2	0.3	0.02	0.6	0.06

128 The data of major-trace elemental concentrations and Sr-Nd isotopic ratios are from Zeng et al. (2010, 2011). 2SD refers to two standard

129 deviations of the population of three repeat runs of the same solution.

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