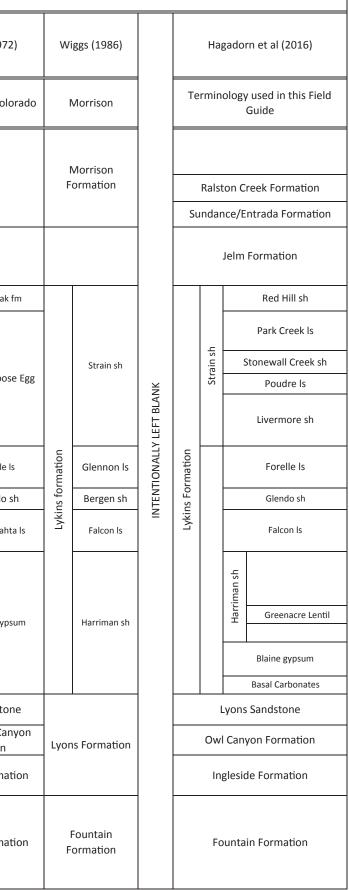
DR2016264 for Hagadorn, J.W., Whiteley, K.R., Lahey, B.L., Henderson, C.M., and Holm-Denoma, C.S., 2016, The Permian–Triassic transition in Colorado, in Keller, S.M., and Morgan, M.L., eds., Unfolding the Geology of the West: Geological Society of America Field Guide 44, doi:10.1130/2016.0044(03)

				1	1		1		1	CORR	ELATI	ION AN	D DEVELOPMENT	OF N	OMENCLATU	RE O	F THE LYKINS FOF	MAT	ION:	FRONT RANGE	1		
		Hayden (1869- 1879)	Marvine (1874)	Cross (1894)	&	nons, Cross Eldridge (1896)		Fenneman (1905)		Butters (1913)			Reed & Scherer ), rev. 1950)	Le	LeRoy (1946)		Burk & Thomas (1956)		Broin (1957)			Pearson (1972	
		Colorado Front Range	South Platte to Little Thompson	Pikes Peak	Den	nver Basin	Boulder		Morrison		Owl Canyon, Northeastern Colorado			Golden		Ce	Central Wyoming		Lykins Gulch			Northeastern Colo	
Jurassic		Jurassic	Jurassic	Morrison Formation		Morrison Formation		Morrison Formation		Morrison Formation		Morrison Formation			Morrison Formation Ralston Creek Formation				Morrison Formation				
								Dr. Bond ss			Sundance Formation								Entrada Sandstone			<u> </u>	
ssic		Trias or "Red	Trias			ion	Lykins Formation					Jelm Formation					Chugwater Formation				Chugwater Formation		
Early Triassic																	red shale, siltstone	e		Red Hill sh	ΟĔ	Red Peak	
Early	p												Spearfish Formation (redefined)				Little Medicine member	tins	Park Creek Is				
										"soft, red shales"					Strain sh		red shale, siltstone	red shale, siltstone	Upper Lykins				
												eout s"	limestone, gray- red, sandy				Ervay member	$\left  \right $	Upp	Stonewall and Livermore shales		Upper Goos	
	idalaupian an Ochoan(?)											"Freezeout heds"	gypsum and shale interbeds				gypsum and red shale and siltstone			Livermore shales			
	Guadalaupian and Ochoan(?)					Upper Division		"Crinkled Sandstone"	Formation	"Crinkled sandstone"	Phosphoria Group	G Forelle Is		Formation	Glennon Is Bergen sh	Forelle ls	Lykins Formation		Forelle Is	rmation	Forelle Is		
									Lykins		sphor		Glendo sh		Bergen sh			Glendo sh		Glendo sh	gg Foi	Glendo s	
Permian				Fountain	ormatio	my one"					Pho		Minnekahta ls	Lyk	Falcon Is Harriman sh	Goose	Minnekahta Is Opeche sh	Lyk		Falcon tongue of the Minnekahta Is	Goose Egg Formation	Minnekaht	
	Leonardian	Beds"		Formation	Wyoming Formation					"soft red Lykin shales"	in	Opeche sh	red shales and sandstone interbeds	_					Lower Lykins	Harriman sh		Blaine gyps	
												0	gypsum and shale interbeds										
	Wolfcampian							•		I		Lyons Formation			I				1	I	Lyons Sandsto		
							Lyo	Lyons Sandstone Fountain Sandstone		Lyons Sandstone Fountain Formation		Owl Canyon Formation		Lyoi	Lyons Formation		Casper Formation		Lyons Sandstone		Satanka/Owl C Formatio		
												Inglesi	side Formation								Ing	leside Format	
Pennsylvanian	Virgilian	-				Lower Division						Founta	ain Formation	Fountain Formation					Fountain Formation		Fountain Format		

**Figure S1.** Correlation and development of the nomenclature of the Lykins Fm. and associated Permian-Triassic strata of Colorado's Front Range (Fig. 1). The age assignments of the units are historical in nature, and do not reflect our current working hypothesis for the age of the various members of the Lykins Fm.. \*The Greenacre Lentil, which overlies the Blaine Gypsum, and the basal carbonates that underlie it, are not formally named units of the succession. But they are used in this contribution, following their common informal use in the oil and gas industry.



		Hill (1899)	Lee (1917)	Maher & Collins (1952)	Shaw (1956)	M	aher & Collins (1952)	Shaw (1956)		Mudge (1967)	Maughan (1980)		Kauffman (1986)	Merewether (1987)
		Las Animas County Colorado Raton Basin Colorado		Huerfano County Colorado (section #20)	Sangre de Cristo Mtns (section #1)	Baca County Colorado (section #12)		Southeastern Colorado (section #8)	Eastern Colorado Western Kansas		Eastern Colorado Western Kansas	Pinon Canyon Maneuver Site, Las Animas County, Colorado		Southeastern Colorado
		(Dakota: Cretaceous)	(Morrison: Cretaceous)		(Santa Rosa Fm: basal unit of Dockum, Upper Triassic)		um Group: Upper Triassic)	(Dockum Group: Upper Triassic)	(Doc	kum Group: Upper Triassic)		Entrada Formation: Jurassic However, based on research published by Hecker et al, 2012, this is currently thought to be the Red Draw member of the Jelm Fm: Middle Triassic		(Dockum Group: Upper Triassic)
	ian	Sangre de Cristo Fm	Undifferentiated "red beds" ranging in age					Taloga Fm		Taloga Fm	Taloga Fm	5	Strain sh	Taloga Fm
	Guadalupian		from Pennsylvanian or older, to possibly Triassic; sitting on	Lykins Fm	unnamed unit	-	Day Creek dolomite	Day Creek dolomite	-	Day Creek dolomite	Day Creek dolomite	Forelle ls shales and sandstones lower contact not	Day Creek dolomite	
			PreCambrian basement							Whitehorse Ss			Whitehorse Ss	
				Lyons Fm		Group	Blaine Fm	Blaine Fm	Group	Dog Creek Sh Blaine Fm	Dog Creek Sh Blaine Fm		exposed	
	Leonardian				Glorieta Fm	Nippewalla		Lyons Fm	Nippewlla (	undifferentiated units (Harper St, Salt Plain Fm, Cedar Hills Ss & Flowerpot Sh)	Nippewalla Group		Nippewalla Group	
				Fountain Fm	Yeso Fm	Sumner Group	Stone Corral	Sumner Group		Stone Corral Fm Ninnescah Sh Wellington Fm	Stone Corral Fm Ninnescah Sh Wellington Fm	-		Sumner Group
PERMIAN					Sangre de Cristo Fm	Chase Group	Wreford Ls		Council Grove Group Chase Group	Nolans Ls Odell Sh Winfield Ls Doyle Sh Barneston Ls Matfield Sh Wreford Ls	Chase Group			Chase Group
	Wolfcampian					Council Grove Group	Grove Group			Speiser Sh Funston Ls Blue Rapids Sh Crouse Ls Easly Creek Sh Bader Ls Sterne Sh Beattie Ls Eksridge Sh Grenola Ls Roca Sh Red Eagle Ls Johnson Sh Foraker Ls	Council Grove Group	_	?	Council Grove Group
						Admire Group			Admire Group	Janesville Sh Falls City Ls Onaga Sh	Admire Group			Admire Group
PENNSYLVANIAN														

Figure S2. Correlation and development of the nomenclature of Permian-Triassic strata exposed in southeast Colorado, with emphasis on the Whitehorse Sandstone, Day Creek Dolomite, and Taloga Formation (Fig. 1).

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# **GSA DATA REPOSITORY ITEM S3: METHODS**

## **Stratigraphy & Sedimentology**

To assess the Permian-Triassic transition in Colorado in the Lykins Fm. and its equivalents, we conducted fieldwork at 32 localities and examined 21 partial or complete cores of the units. Only data from the Front Range and southeast Colorado surface exposures are presented here, because this contribution is an early progress report. Except where noted, all images as well as collected fossils, rock samples and thin-section billets are keyed to Figs. 1-2. Many of our study sites are on private land or sensitive public lands where permits are required for fieldwork; thus throughout the text we refer to locality numbers rather than to sites (e.g., referring to "section F4" rather than "Miriam Smith's Ranch"). Detailed locality coordinates are available upon request. For the 16 Front Range and southeast Colorado reference sections, 420 standard thin sections were cut, stained and examined to petrographically characterize sedimentology, paleontology, depositional environments, and diagenesis of the Lykins Fm. and its equivalents.

#### Geochronology

A 5-10 kg sandstone sample from each target horizon (solid circles in section F13, Fig. 1) was processed using conventional mineral separation techniques including: crushing, grinding, sieving, Wilfley table density separation, magnetic separation, and heavy liquids separation. Concentrated zircon splits were poured onto double-sided tape and mounted in epoxy, ground and polished to half thickness to expose their internal structure, and imaged in transmitted and reflected light as well as SEM-based cathodoluminescence. Detrital zircons were ablated using a Photon Machines 193 nm Analyte Excite laser ablation system with spot sizes between 20 and

25 um at 5 Hz and  $\sim$ 5 mJ laser energy with an energy density of 6 J/cm<sup>2</sup>. Ablated material was transported to a Nu Instruments AttoM high-resolution, double focusing, single-collector ICPMS. With the magnet set at a constant mass, the flat tops of the isotope peaks were measured at the following masses by rapidly deflecting the ion beam: <sup>202</sup>Hg, <sup>204</sup>(Hg+Pb), <sup>206</sup>Pb, <sup>207</sup>Pb, <sup>208</sup>Pb, <sup>232</sup>Th, <sup>235</sup>U, and <sup>238</sup>U with a 30 s on-peak background measured prior to each 30 s analysis. Raw data were reduced using Iolite<sup>TM</sup> (Paton et al., 2011) to subtract on-peak background signals, correct for U-Pb downhole fractionation, and normalize the instrumental mass bias using external, well-characterized, mineral standards. Ages are corrected by standard sample bracketing with the primary zircon standard Temora2 (417 Ma; Black et al., 2004) and secondary standards Plešovice (337 Ma; Slama et al., 2008), and an in-house standard WRP-63-08 (1707 Ma; Wayne Premo, personal commun., 2015) used at the USGS Central Mineral and Environmental Resources Science Center LA-ICPMS Isotope Lab. Reduced data was compiled into concordia and probability density plots using Isoplot 4.15 (Ludwig, 2012). Analyses with discordance greater than 15% were excluded from probability density plots. Analyses younger than 500 Ma were also visually inspected for degree of concordance. In general the  $^{207}$ Pb/ $^{206}$ Pb ages become more precise for older zircons due to the relative abundance of <sup>235</sup>U (parent to <sup>207</sup>Pb daughter product) in the early Earth. <sup>207</sup>Pb/<sup>206</sup>Pb ages are equally imprecise as <sup>206</sup>Pb/<sup>238</sup>U and <sup>207</sup>Pb/<sup>235</sup>U ages at about 1400 Ma, therefore, the preferred age for zircon greater than 1400 Ma is the  $^{207}$ Pb/ $^{206}$ Pb age.

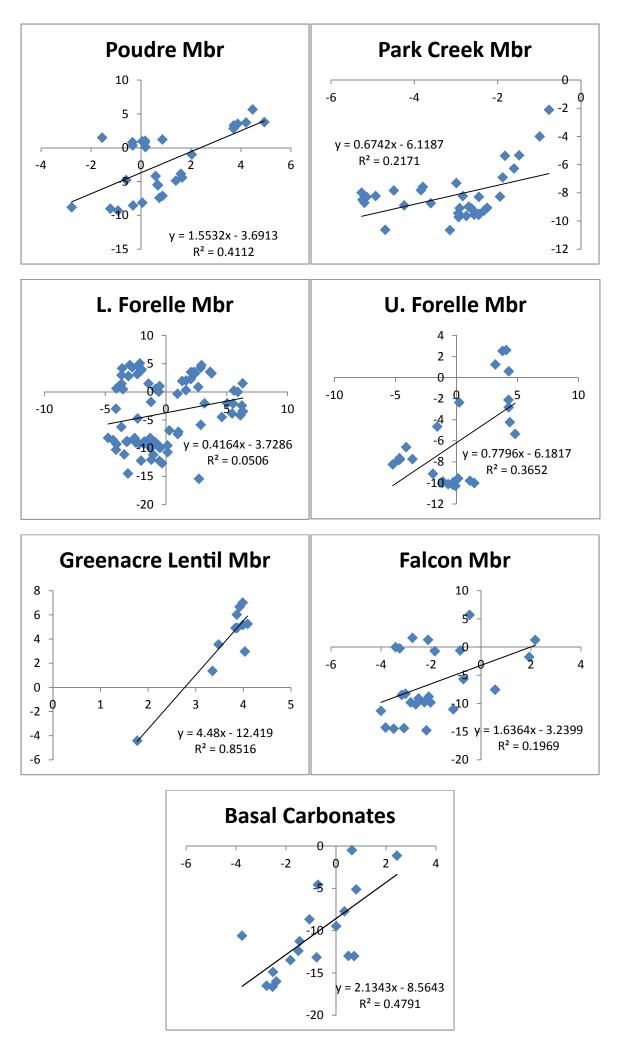
# Paleontology

Only samples that were collected in-situ and logged into reference sections were considered for interpretation, with the exception of a vertebrate trackway, which was collected from float ~60 m above the Park Creek Member near section F13 (Figs. 1, 2). Thin-sections spanning all the representative lithologies in sections F1-F13 and S1-S2 were systematically examined for microfossils at 40x and 100x on a petrographic microscope. We did not search for nannofossils or palynomorphs, except in two rare grey-black calcareous shales from the Lykins Fm., where they mantle the Falcon Member at a locality near section F13. An ~15 kg bulk limestone or dolostone sample was processed for microfossils from each of the six Lykins Fm. carbonates at locality F12 as well as from each of the Falcon coquinas in sections F3 and F13. These carbonates were cleaned then dissolved in 10% buffered acetic acid for 3-4 days, with all particles greater than 75  $\mu$ m retained by sieving. Residues were rinsed, dried and then separated in tetrabromoethane. The resulting heavy fraction was hand-picked under a binocular microscope. Identified microfossils mounted on slides and imaged via SEM.

## Chemostratigraphy

Petrographic screening of all limestones, dolostones, and calcareous siliciclastics was conducted under polarized and cross-polarized light on a petrographic microscope, and where diagenetically least-altered phases (dominantly micrite, dolomicrite, or ooids) were identified, counterpart billets of those samples were microdrilled for analysis. Samples from sections F13-F12 were analyzed at the University of Arizona Environmental Isotope Laboratory where the  $\delta^{18}$ O and  $\delta^{13}$ C of least-altered carbonate phases was measured using an automated carbonate preparation device (KIEL-III) coupled to a gas-ratio mass spectrometer (Finnigan MAT 252). Powdered samples were reacted with dehydrated phosphoric acid under vacuum at 70°C. The isotope ratio measurement was calibrated based on repeated measurements of NBS-19 and NBS-18 and precision was  $\pm 0.10$  ‰ for  $\delta^{18}$ O and  $\pm 0.08$ ‰ for  $\delta^{13}$ C (1 sigma).

The remaining samples were analyzed at the University of Utah SIRFER lab, where powders were weighed using a Sartorious microbalance, loaded into 4.5 ml flat-bottomed borosilicate vials (Labco) and capped with Labco butyl rubber septa. These were transferred to a Thermo Fisher Scientific system comprised of a GasBench II, PAL autosampler, ConFlow IV interface and a MAT 253 mass spectrometer. Vials were flushed for 6 mins on a PAL autosampler, with Ultra High Purity grade helium (99.999% He) at a flow rate of 50 mL/min. During flushing the vials were kept in a 50° C aluminum block. Samples were reacted with 10 droplets of 104% phosphoric acid (H<sub>3</sub>PO<sub>4</sub>) (kept at 50° C) to CO<sub>2</sub> gas. Samples were allowed to equilibrate for 12 hours before analysis. During analysis using the PAL autosampler the produced CO<sub>2</sub> gas in the vials was collected using a sampling loop of 100  $\mu$ L and transported to the mass spectrometer. Nine individual gas injections per analysis were made for each sample and for the reference materials; their average was taken as the final number. Three sets of reference materials were used to calibrate the system and unknown samples, including Carrara marble, LSVEC and Marble-Std. Carrara and LSVEC was used as primary reference materials and Marble -Std used as secondary reference material to cross check the final number. The oxygen fraction factor was calculated using the alpha value proposed by Swart et al. (1991). The average analytical uncertainty for analyzed samples was <0.08 for  $\delta^{13}$ C and <0.10 for  $\delta^{18}$ O; samples yielding low CO<sub>2</sub> output were excluded from reported results.



**Figure S4.** Cross-plots of <sup>13</sup>C (horizontal axis) vs. <sup>18</sup>O (vertical axis) for each of the carbonatedominated intervals of the Lykins Formation. The three thin basal carbonates are plotted together whereas the Lower and Upper Forelle Member are separated. Results are preliminary.