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Origin and time evolution of subduction polarity reversal from plate kinematics of Southeast Asia

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Materials and Methods:

In this study we propose a regional-scale tectonic reconstruction that fits into the global geodynamic frame, including all known plates. We reconstruct the movements of plates through time, instead of showing selected snapshots. To achieve this goal we use GPlates (www.gplates.org). This software allows both computation and real-time visualization of global plate-tectonic reconstructions (Boyden et al. 2011, Gurnis et al., 2012). GPlates is based on data from spreading activity. It allows to create rotations, and create a mid ocean ridge that moves with a half stage rotation. It is not possible to impose or measure stable mid-ocean ridges. Moreover, currently GPlates does not include intraplate deformation. Consequently absolute plate velocities reconstructed based on opening of oceans may be prone to errors. However, with closed circuit plate reconstructions it is possible to trace subduction polarity reversals through time, which will yield insights on the underlying processes. Our reconstruction southeast Asia is based on the plate motion model by Seton et al. (2012), from which we adopt the absolute reference frame, the plate circuit (linking the relative motions of the tectonic plates), and the geomagnetic polarity timescale. We take into account recent improvements of the model for the (Proto) South China Sea (Zahirovic et a. 2014). In order to track the subduction zones polarity, we modify and adapt locally the geometry of the plate boundaries, which are used to define continuously (temporally and spatially) closed plate polygon (Gurnis et al. 2012). We start our reconstructions in the late Cretaceous, a time when oceanic crust of the now extinct Izanagi Plate (a conjugate to the Pacific Plate) was subducting westward underneath Eurasia, i.e. opposite to present day subduction direction in south of Taiwan. Evidence for westward subduction during that time is evident from rhyolite-basalt assemblages, diabase/lamprophyre dikes and mostly peraluminous granitoids exposed in the South China fold belt.

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27 Several plate reconstructions focused on the Neogene plate kinematics in southeast
28 Asia as well as on the Taiwan area (e.g. Hall, 2002; Malavieille et al., 2002; Sibuet and
29 Hsu, 2004). Figure S1 shows a compilation of existing models. A key player in SE Asian
30 tectonics may be the now extinct Proto South China Sea. Extinct oceans have been
31 shown to play a crucial role in mountain building phases. For instance in the European
32 Alps, existence of the Meliata Ocean, of which only very limited remnants have been

33 suggested is key ingredient for understanding the present day mountain belt. It was only
34 in the continuation of the Alps into the Dinarides, where remnants of this ocean and its
35 associated orogenic events are exposed. Despite extensive research on the Alps since
36 more than 250 years, this ocean and its importance was not discovered until recently
37 (Schmid 2004). In our reconstructions we can show the importance of the Proto South
38 China Sea on evolution of Taiwan (see movie of reconstructions).

39

40 Due to large scale extension the pre-collisional Eurasian passive margin in the Taiwan
41 area is characterized by thick Paleo- to Mesozoic and mostly Cenozoic sediments,
42 overlying a wide zone of transitional lithosphere, which is either thinned, volcanically
43 intruded continental (Nissen et al., 1995), or volcanically thickened oceanic crust
44 (Hsuehshan Range et al., 2004), and potentially includes continental fragments (Shyu et
45 al., 2005), or continental margin promontories that reflect the inherited rift-transform
46 configuration (Lee et al., 2015; Mirakian et al., 2013). Our reconstructions show the
47 north-westward movement of the Luzon Arc starting as late as 6.5 Ma. At this time,
48 rifting ceased and foreland strata witness initiation of collision (Lin et al., 2003). This
49 coincides with fast rollback of the Ryukyu Trench (Yamaji, 2003).

50

51 When collision occurred for the first time remains a matter of debate, and arguments
52 have been put forward for a comparably old (i.e. mid to late Miocene) to close to recent
53 (i.e. mid to latest Pliocene) collision. For instance the transition between pre- and syn-
54 collisional strata in the Taiwan Strait at ~6.5 Ma supports a Miocene collision (Lin et al.,
55 2003). On the other hand, at 4 and 3 Ma the sedimentary suite of the passive margin
56 sediments overlying oceanic crust is incorporated into the wedge, which is witnessed by
57 massive deposition of sediments into the surrounding basins thereafter (Chen et al.,
58 2001; Chi et al., 1981; Nagel et al., 2013), and consequently was interpreted as young
59 onset of collision. Recent thermochronological evidence supports a predominantly
60 young, late Pleistocene exhumation pulse, with first onset of cooling occurring in the
61 late Miocene (Fox et al., 2014; Kirstein et al., 2010; Mesalles et al., 2014).

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63 Figure S1: Compilation of different reconstructions suggested for southeast Asia.
64 Reconstructions focused on evolution of the Philippine Sea Plate, South China Sea,
65 Taiwan, or on the entire region. Regional models often cannot account for details within
66 particular areas, whereas more local studies often are not put in the global plate tectonic
67 framework. Our model refines existing global models with a special focus on evolution

68 of the Taiwan area. Figure based on Seno & Maruyama (1984), Letouzey & Kimura
69 (1985), Kong et al. (2000), Deschamps & Lallemand (2002), Sdrolinas et al. (2004), for
70 Okinawa Trough and Philippine Sea Plate; Mahoney et al. (2011), Barckhausen et al.
71 (2014) and Yeh et al. (2010) for the South China Sea; Sibuet & Hsu (2004), Malavieille et
72 al. 2002, and Teng and Lin (2004) for the Taiwan area; Hall et al. (2008), Richards et al.
73 (2007), Replumaz & Tapponier (2003), Morley (2002), Hall (2002), Yin (2010), Hall
74 (2012), Xu et al. (2014), and Zahirovic et al. (2014) for the Proto South China Sea and
75 regional models.

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Figure DR1

