Brodsky et al.

Documentation of Faults

Table DR1. Scanned faults and roughness processing

Name	Scanner	Location	Lithology	Sense	Slip	Processing
Cascia#	LiDAR	42.719° N 13.002° E	Carbonate	Normal	50m	Small patches ^b + Taper 5% ^d
Gubbio Upper#	LiDAR	43.344° N 12.597° E	Carbonate	Normal	50-100m	Small patches ^b + Taper 5% ^d
Gubbio Lower#	LiDAR	43.344° N 12.597° E	Carbonate	Normal	200m	Small patches ^b + Taper 5% ^d
Monte Coscerno#	LiDAR	42.692° N 12.887° E	Carbonate	Normal	250m	Small patches ^b + Taper 5% ^d
Monte Maggio#	LiDAR	42.762° N 12.941° E	Carbonate	Normal	650m	Small patches ^b + Taper 5% ^d
Val Casana#	LiDAR	42.718° N 12.857° E	Carbonate	Normal	150m	Small patches ^b + Taper 5% ^d
Venere Large#	LiDAR	41.971° N 13.664° E	Carbonate	Normal	>20m	Small patches ^b + Taper 5% ^d
Venere Small#	LiDAR	41.971° N 13.664° E	Carbonate	Normal	4m	Small patches ^b + Taper 5% ^d
West Fucino#	LiDAR	41.940° N 13.362° E	Carbonate	Normal	~80m	Small patches ^b + Taper 5% ^d
Vasquez rocks#	LiDAR	34.483° N 118.316° W	Sandstone	Normal	10±5cm	Small patches ^b + Taper 5% ^d
Yeelim#	LiDAR	31.223° N 35.354° E	Carbonate	Normal	50-80m	Small patches ^b + Taper 5% ^d
Split Mountain#	LiDAR	33.014° N 116.112° W	Sandstone	Strike- slip	30±15cm	Small patches ^b + Taper 5% ^d
Mecca Hills#	LiDAR	33.605° N 115.918° W	Carbonate	Strike- slip	20±10cm	Small patches ^b + Taper 5% ^d
Flowers Pit#	LiDAR	42.077° N 121.856° W	Andesite	Normal	100- 300m	Small patches ^b + Taper 5% ^d
Chimney Rock#	LiDAR	39.227° N 110.514° W	Carbonate	Normal	8m	Small patches ^b + Taper 5% ^d
Lake Mead#	LiDAR	36.062° N 114.831° W	Dacite	Normal	500- 1000m	Small patches ^b + Taper 5% ^d
Corona Heights*	LiDAR + LPa	37.765° N 122.437° E	Chert	Strike- slip	Several m to >1km	Small patches ^b + Large patches ^c + Taper 3% ^d
Vuache- Sillingy*	LiDAR + LPa	45.920° N 6.049° E	Carbonate	Strike- slip	10-30m	Small patches ^b + Taper 3% ^d
Dixie Valley*	LiDAR + LP ^a	39.947° N 117.945° E	Rhyolites	Normal	Several m to >3- 6km	Small patches ^b + Taper 3% ^d
Bolu*	LiDAR	40.685° N 31.568° E	Carbonate	Strike- slip	20m- 85km	Small patches ^b + Taper 3% ^d
Klamath##	LiDAR	42.135° N 121.678° W	Basalt +andesite	Normal	50-300m	Large patches ^c + Taper 3% ^d

Arkitsa§	LiDAR	38.733° N 23.000° E	Carbonate	Normal	>300- 400m	Large patches ^c + Taper 3% ^d
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Sources: # Brodsky et al. (2011), *Candela et al. (2012), ## Sagy et al. (2007), § Resor and Meer (2009) Notes: a Laser Profilometer, b Small clean fault patches free of unwanted objects are selected from the original cloud of points, c Large clean fault patches are obtained by removing locally non-faulting features from the original cloud of points, d During the processing of individual profiles for computing the Fourier spectra we either apply a cosine taper of 3% or 5%.

Scale-dependent roughness in spatial maps

Figure DR1 demonstrates that the scale-dependent roughness is a feature of individual profiles of the raw data in the spatial domain. The parameterization in the Fourier domain usefully captures the same phenomenon shown in Figure DR1.

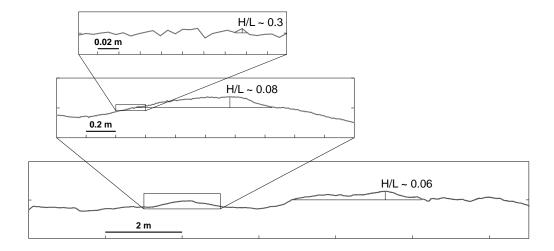


Figure DR1. Example cross sectional profiles at different magnifications showing the scale dependence of the aspect ratio, H/L. Profiles are taken from a ground-based LiDAR scan of the Corona Heights fault surface. Panels show sections through the fault at different magnifications (Note the scale bar is different in each panel).

Fourier Transform Computation

From each fault patch (LiDAR or profilometer), hundreds to thousands of profiles are extracted in the slip direction. Most fault patches have more than 500 profiles in the slip direction and all patches have at least 100. The four steps in the procedure to compute the spectrum of each profile are as follows: (1) Each profile is detrended by subtracting the best-fit trend. (2) Either a 3% or 5% cosine taper is applied to each rough profile to ensure that there are no step functions at the end of the finite window. (3) The discrete Fourier transform is calculated, and the power spectrum is equal to the square of the amplitudes of the coefficients. (4) The power spectrum is normalized by the profile length to obtain the power spectral density. The mean Fourier spectrum of each fault patch is then computed by averaging the spectra of the profiles and restricting the results to well-resolved wavelengths that are more than a factor of 2 less than the profile length. Finally we smooth the spectra in frequency space by binning each averaged spectrum into 20 logarithmicallyspaced intervals in the well-resolved frequency space and averaging the power spectral density within the bin. Logarithmic binning provides a constant density of data points in the logarithmic representation and therefore avoids giving more weight to smaller scales in subsequent fitting procedures.

Maximum Shear Strain in Hertzian Contacts

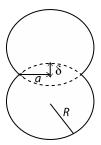


Figure DR2. Cartoon of elastic interaction between two spheres. Dashed lines show undeformed outlines.

The elastic deformation between two identical spheres was first modeled by Hertz (Hertz, 1881). We use the solution to illustrate the general form of the elastic stress field between contacting asperities starting from the general results summarized by Johnson, 1985, Appendix 3. If two elastic spheres of radius R are in contact as shown in Figure DR2 and the contact between them has no shear stress, the solution for the radius a of the contact area is

$$a = (3WR/4E)^{1/3}$$
 (A1)

where W is the loading force on the sphere and E' is the modified Young's modulus (2(1- v^2)/E)-1 and E is the Young's modulus and v is the Poisson ratio. The approach distance between the two spheres is related to the contact radius by

$$\delta = a^2/R \tag{A2}$$

Slightly manipulating eq. A1 yields

$$a/R = \frac{3}{4} \pi W/(E'\pi a^2)$$
 (A3)

Since $W/\pi a^2$ is by definition average normal stress σ on the contact, then combining eq. (A2) and (A3) yields.

$$\delta/a = \frac{3}{4} \pi \sigma/E'. \tag{A4}$$

Complete solutions show that the maximum shear stress τ within the asperity is 0.31 the maximum normal stress, and the maximum normal stress is 3/2 σ (Johnson, 1985, Appendix 3). For complete flattening of a contact of height H and length L, $\delta = H$ and a = L/2. The shear stress is related to the shear strain by $\tau = 2G\varepsilon$ where G is the shear modulus. Therefore,

$$H/L = \pi/0.62 \quad \varepsilon G/E' \tag{A5}$$

and the shear strain is proportional to the aspect ratio as expected. The maximum shear strain occurs in the interior of the asperity and therefore the asperity fails rather than the contact surface. Including friction at the interface does significantly affect the internal stress field of the contacting spheres (Johnson, 1985; Sect. 5.4).

Similar solutions are recovered for more complex geometries. For a sinusoidal, single-wavelength surface, the normal stress required for complete flattening of the asperity is $2^{1/2} \pi E' H/L$, which implies a similar scaling with strain as in (A5) (Johnson et al., 1985). Multi-scale models also preserve the proportionality between calculated stresses, strains and aspect ratio (Krithivasan and Jackson, 2007; Jackson et al., 2012).

The role of surface slope in determining surface failure during shearing has long been recognized in tribology. For instance, the plastic yield criteria for an indentor depends on the cotangent of the apex angle, which is the aspect ratio H/L (Johnson et al., 1985; Sect. 6.1). As a result, wear mechanisms are predicted by using the plasticity index $\psi=(H/L)E'/\mathcal{H}$ where \mathcal{H} is the hardness (Mikic, 1974; Johnson, 1985). The related approach presented here focuses on the inverse problem of determining strength of a material given the observed surface roughness.

References

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