DATA REPOSITORY ITEM: APPENDIX

AGE MODEL

The age models for Holes 865B and 865C are derived from benthic and planktic foraminifers and calcareous nannofossils (Table DR1). The time scale follows the Geological Time Scale (GTS) 2012 (Gradstein et al., 2012). Although previous studies have been used the shape of foraminifer carbon-isotope (δ^{13} C) curves to date and correlate the Paleocene-Eocene Thermal Maximum (PETM) sediments in high resolution (e.g., Bains et al., 1999; Röhl et al., 2007), we do not use the δ^{13} C curves for dating the PETM sediments. At Site 865, the difference in the shape of the δ^{13} C curves between foraminifer taxa was recognized (e.g., Dunkley Jones et al., 2013).

Table DR1. Horizons and geological ages for the datum events. Abbreviation: LO = the last occurrence and FO = the first occurrence.

Datum event	Age (Ma)	Depth (mbsf)		Reference		
		Hole 865B	Hole 865C			
FO: Discoaster sublodensis	49.11	79.6	79.4	Kozdon et al. (2011)		
LO: Tribrachiatus orthostylus	50.5	83.2	79.4	Bralower et al. (1995), Kozdon et al. (2011)		

FO: Discoaster	53.7	89.6	87.4	Bralower et al. (1995),
	0011	00.0	0/11	
lodoensis				Bralower and Mutterlose
				(1995)
LO: Tribrachiatus	54.14	91.61	91.74	Kozdon et al. (2011)
contortus				
contontas				
LO: Morozovella	55.2	97.85	98.3	Bralower et al. (1995), Kelly
	00.2	01.00	00.0	
velascoensis				et al. (2001)
Paleocene/Eocene	55.96	103.6	103.02	Bralower et al. (1995), Kelly
boundary; Benthic				et al. (2001), Dunkley
foraminifer extinction				Jones et al. (2013)
event				
ovont				
FO: Discoaster	57.21	116.21	114.6	Bralower et al. (1995)
multiradiatus				· · · /
muitiradiatus				

METHODS AND MATERIALS, CONTINUED

Sediment samples of 10–20 cc volume from the cores were washed through a sieve with 63 µm opening and then were dried in an oven at 50°C. Ostracodes were picked from the >125 µm fraction of the sediment samples. We obtained ostracode specimens from 117 samples and identified 42 taxa, observing specimens with a binocular microscope and scanning electron microscopy, JSM-6500F (JEOL Ltd.) at Kochi Core Center, Kochi University. Fragments of valves are included in the dataset only if \geq 1/2 of the valve was present. Our identification has not completely harmonized with the taxonomy of Boomer and Whatley (1995) yet. In this study, we

refer to only data about abundance and species richness of Boomer and Whatley (1995). We counted the number of carapace and valve specimens in each species and sample and considered the number of both specimens as ostracode abundance. We store the dataset on the ostracode taxonomic composition and SEM images of selected ostracode taxa in the PANGAEA database

(http://doi.pangaea.de/10.1594/PANGAEA.836086).

To examine temporal changes in ostracode fauna with statistically valid sample sizes, we binned samples in each 500 kyr interval from 55.96 Ma as the starting point. The dataset to make the binned samples does not include data of Boomer and Whatley (1995). We calculated rarefaction at a cutoff of 20 and 30 individuals and their standard errors in nine binned samples, following Hurlbert (1971) and Heck et al. (1975). Calculating rarefied indexes, we used the free software GNU R (http://www.r-project.org/) and its software package, vegan (Oksanen et al., 2013). The count of species richness is highly sensitive to the number of individuals sampled.

To assign statistically extinction ages of taxa, we calculated the 95% confidence intervals (CIs) of the age for the LO of taxa near the PETM, following Marshall (1997)'s method. The fossil recovery potential was represented by the total abundance of ostracodes. We assigned the CIs in three of the disappearing taxa that are found in more than 10 samples. For the calculation we used the R code that was written by Gene Hunt based on Marshall (1997) as implemented by Rivandeira et al. (2009). To assign temporal changes in grain size, we weighed dry sediments in 140 samples before and after the washing with a sieve of 63-µm opening. The dataset of the grain-size is deposited in the above website of the PANGAEA database.

We calculated the percentage of coarse fraction and also estimated the sedimentation rates using the age model and benthic ostracode accumulation rate (specimens cm^{-2} k.y.⁻¹) using the density (g cm⁻³) of 97 sediment samples, the sedimentation rate, and the ostracode abundance.

Table DR2. Abundance and rarefied species richness of ostracodes from the 500 kyr binned samples. NA = Not applicable, SE = standard error.

Interval	Number	Number	Species	<i>E</i> (S ₂₀)	<i>E</i> (S ₃₀)	Even-
(Ma)	of samples	of individuals	richness	±SE (1σ)	±SE (1σ)	ness
49.46-49.96	1	68	11	8.02±1.04	9.00±0.97	0.89
50.46–50.96	1	14	8	NA	NA	0.92
51.46–51.56	1	7	6	NA	NA	0.98
51.96–52.46	1	5	2	NA	NA	0.97
52.96–53.46	1	6	3	NA	NA	0.92
53.46–53.96	1	9	4	NA	NA	0.83
53.96–54.46	1	12	7	NA	NA	0.90
54.46–54.96	13	60	17	9.72±1.47	12.41±1.42	0.78
54.96–55.46	15	28	3	2.43±0.63	NA	0.28
55.46–55.96	20	79	14	6.68±1.43	8.57±1.46	0.58
55.96–56.46	28	254	19	5.77±1.36	7.07±1.51	0.53
56.46–56.96	34	374	22	6.06±1.24	7.10±1.38	0.58

Table DR3. The observed age and upper limits of the 95% confidence interval of the last occurrence in the five disappearing taxa. Abbreviation: LO = the last occurrence.

 Taxon
 Number of
 Age of observed
 Upper limit of the 95% CI

	samples	LO (Ma)	in the age of LO (Ma)
Cytherella sp.1	29	55.98	55.74
<i>Krithe</i> sp.A	62	55.91	55.67
Profundobythere			
multipunctata	18	55.96	52.12

THE SAMPLING EFFECT ON DIVERSITY LOSS

Increasing in ostracode abundance, both the number of species and the diversity loss (%) from the interval of 55.46–55.96 Ma to of 54.96–55.46 Ma increase (Fig. DR1). Not only the number of species but the diversity loss depends on the abundance. The diversity loss between the 55.46–55.96 Ma and 54.96–55.46 Ma intervals shows a plateau with a value of 64% at the abundance as many as 19–28 individuals. At abundance of more than 19, the sampling bias is less effective on the estimated diversity loss.

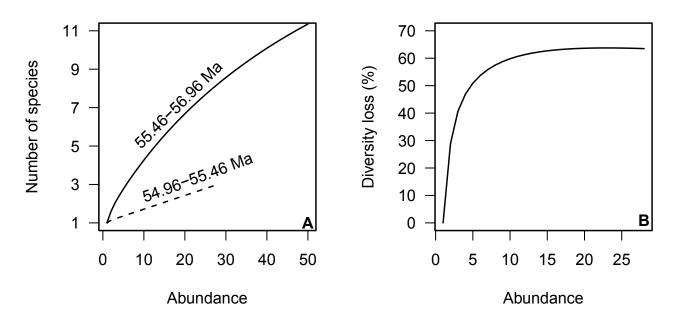


Fig. DR1. A. Species accumulation curve in the intervals of 55.46–55.96 and of 54.96–55.46 Ma. B. Diversity loss (%) from the interval of 55.46–55.96 Ma to of 54.96–55.46 Ma.

PALEOECOLOGY OF OSTRACODE DOMINANT TAXA

In the ostracode faunas, the dominant taxa change from *Krithe* and *Cytherella* to *Nemoceratina (Pariceratina)* near the P/E boundary as Boomer and Whatley (1995) reported.

Nemoceratina (Pariceratina) is an extinct genus. It is found in Paleocene-Eccene sediments at 1000 to 2000 m paleodepth and in the equatorial upwelling and coastal and the subtropical and gyres zones that were associated with high export productivity (Ma et al., 2014; Table DR4). In the ocean zones, the export productivity substantially increased during the PETM (Ma et al., 2014). Throughout the PETM, *Nemoceratina (Pariceratina)* flourished under the high export productivity. After the Eocene, Nemoceratina (Pariceratina) is found in only the South Pacific and Mediterranean Oceans. Oligocene-Miocene taxa are found in sites out of the Pacific equatorial divergence zone (Boomer and Whatley, 1995; Alvarez Zarikian, 2014). The sites may have been below oligotrophic waters. The taxon is found in the early Paleocene sediments in the South Pacific Ocean site (Alvarez Zarikian, 2014). The site would have been under higher biological productivity, because the site is located near the continents and the export productivity in the South Pacific bloomed following the Cretaceous/Paleogene collapse of marine plankton ecosystems (Hollis et al., 2003). An early Pleistocene taxon of the Mediterranean Sea, Nemoceratina (Pariceratina) barrieri Sciuto, 2014 is considered to have dwelled under eutrophic conditions. The surface productivity in the Mediterranean Sea increased from the latest Pliocene to the early Pleistocene (Lourens et al., 1992). We consider Nemoceratina (Pariceratina) as an indicator of high export production. On the other hand, Krithe and Cytherella are both extant. They have been cosmopolitan taxa

6

during the Cenozoic. Both the taxa indicate conditions of low or moderate foodsupply (Didié et al., 2002; Horne et al., 2011; Stepanova and Lyle, 2014).

Table DR4. Deep-sea sites bearing late Paleocene–early Eocene taxa of *Cytherella*, *Krithe*, and *Nemoceratina* (*Pariceratina*). The plus and minus marks indicate the presence and the absence, respectively. The coordinates at 56 Ma were calculated, using an online service of the Ocean Drilling Stratigraphic Network (www.odsn.de).

Site	Modern-		At 56 Ma		Ocean zone (Ma et al.,
	day				2014)
	Longitude	Latitude	Longitude	Latitude	
ODP 865	179.6W	18.4°N	164.1°W	3.0°S	Equatorial upwelling and
					coastal zone
ODP 762	112.3°E	19.9°S	87.5°E	40.6°S	Equatorial upwelling and
					coastal zone
ODP 1260	54.5°W	9.3°N	53.1°W	4.2°N	Equatorial upwelling and
					coastal zone
ODP 1261	54.3°W	9.0°N	52.8°W	4.0°N	Equatorial upwelling and
					coastal zone
ODP 1049	76.1°W	30.1°N	71.4°W	28.0°N	Equatorial upwelling and
					coastal zone
ODP 1050	76.2°W	30.1°N	71.5°W	27.9°N	Equatorial upwelling and
					coastal zone
ODP 1051	76.4°W	30.1°N	71.6°W	27.9°N	Equatorial upwelling and
					coastal zone
ODP 1052	76.6°W	30.0°N	71.9°W	27.9°N	Equatorial upwelling and
					coastal zone
DSDP 549	13.1W	49.1°N	25.6°W	43.0°N	Subtropical and gyres
DSDP 550	13.1W	48.5°N	25.8°W	42.4°N	Subtropical and gyres

DSDP 400	9.2°W	47.4°N	21.9°W	41.6°N	Subtropical and gyres
DSDP 401	8.8°W	47.4°N	21.5°W	41.6°N	Subtropical and gyres
Basque	2.3°W	43.3°N	15.1°W	38.0°N	Subtropical and gyres
Basin					
DSDP 516	35.3°W	30.3°N	47.8°W	27.1°N	North Atlantic and Tethys
Caravaca	1.9°W	38.1°N	14.1°W	32.8°N	North Atlantic and Tethys
ODP 689	3.1°E	64.5°S	9.7°W	65.5°S	Polar
ODP 690	1.2°E	65.2°S	11.5°W	66.2°S	Polar
DSDP 20	26.8°W	28.5°S	23.4°W	32.6°S	South Atlantic and India
					Ocean
DSDP 356	41.1°W	28.3°S	38.3°W	33.0°S	South Atlantic and India
					Ocean
DSDP 357	35.6°W	30.0°S	32.4°W	34.5°S	South Atlantic and India
					Ocean

Site	Ostracode taxon			Reference
	Cytherella	Krithe	Nemoceratina	
ODP 865	+	+	+	Boomer and Whatley (1995);
				This study
ODP 762	+	-	-	Guernet and Gasburn (1992)
ODP 1260	+	-	-	Guernet and Danelian
				(2006)
ODP 1261	-	-	-	Guernet and Danelian
				(2006)
ODP 1049	+	+	+	Guernet and Bellier (2000)
ODP 1050	+	+	-	Guernet and Bellier (2000)
ODP 1051	+	+	+	Guernet and Bellier (2000)
ODP 1052	-	+	-	Guernet and Bellier (2000)

TABLE DR4. (continued)

DSDP 549	+	+	+	Whatley and Coles (1991)
DSDP 550	+	+	+	Whatley and Coles (1991)
DSDP 400	-	+	-	Ducasse and Peypouquet
				(1983)
DSDP 401	+	+	-	Ducasse and Peypouquet
				(1979); Yamaguchi and
				Norris (2012)
Basque	+	+	-	Rodriguez-Lazaro and
Basin				Garcia-Zaraga(1996)
DSDP 516	+	-	-	Benson and Peypouquet
				(1983)
Caravaca	+	+	-	Guernet and Molina (1997)
ODP 689	+	+	-	Majoran and Dingle (2002)
ODP 690	+	+	-	Steineck and Thomas (1996)
DSDP 20	+	-	-	Benson (1975)
DSDP 356	-	?	-	Benson (1978)
DSDP 357	-	+	-	Benson (1978)

LIST OF SPECIES IDENTIFIED IN THIS WORK

We grouped ostracode taxa into the Range-through, Lazarus, Disappearing, Late Paleocene, and incoming taxa.

Range-through taxon occurs in most samples.

Nemoceratina (Pariceratina) ubiquita (Boomer, 1994)

Lazarus taxa are sporadically found in less ten samples from the PETM and Crisis

intervals:

Cytherella sp.A

Eucythere sp.B

Hemiparacytheridea sp.A

Argilloecia sp.A

Argilloecia sp.C

Cardobairdia sp.

Cytherella sp.B

Cytheropteron sp.2 of Boomer and Whatley (1995)

Cytheropteron sp.24 of Boomer and Whatley (1995)

Cytheropteron sp.C

Eucythere sp.A

Eucythere sp.D

Eucythere sp.F

Eucytherura sp.B

Neonesidea spp.

Rimacytheropteron sp.24 of Boomer and Whatley (1995)

Disappearing taxa have their last occurrences near the onset of PETM. Their extinction is local. *Profundobythere multipunctata* is reported in the middle Eocene and Oligocene sediments in the Atlantic and Pacific Oceans (Coles and Whatley, 1989; Alvarez Zarikian, 2014). *Cytherella* sp.1 of Boomer and Whatley (1995) *Krithe* sp.A Krithe sp.C

Profundobythere multipunctata Coles and Whatley, 1989

Semicytherura sp.A

below the PETM. Argilloecia sp.D Bythocythere sp. Cytheropteron sp.A Eucythere sp.C Eucytherura sp.A Hemiparacytheridea sp.C Pennyella sp.1 of Boomer and Whatley (1995) Propontocypris sp. Semicytherura sp.B

Late Paleocene taxa found in only Paleocene sediments. Their last occurrences are

Incoming taxa appeared after the end of the Crisis interval.

Argilloecia sp.B

Aversovalva sp.7 of Boomer and Whatley (1995)

Bythocypris sp.

Cytheropteron sp.B

Eucythere sp.E

Eucytherura sp.D

Hemiparacytheridea sp.B

Krithe spp.

Macrocypris sp. Pedicythere sp. Pennyella spp. Pseudocythere sp. Semicytherura sp.C

REFERENCES CITED

- Alvarez Zarikian, C. A., 2014, Cenozoic bathyal and abyssal ostracods beneath the oligotrophic South Pacific Gyre (IODP Expedition 329 Sites U1367, U1368 and U1370): Palaeogeography Palaeoclimatology Palaeoecology, in press
- Bains, S., Corfield, R. M., and Norris, R. D., 1999, Mechanisms of Climate Warming at the End of the Paleocene: Science, v. 285, p. 724–727.
- Benson, R. H., 1975, The origin of the psychrosphere as recorded in changes of deepsea ostracode assemblages: Lethaia, v. 8, p. 69–83.
- Benson, R. H., 1978, The Cenozoic ostracide faunas of the Sao Paulo Plateau and the Rio Grande Rise (DSDP Leg 39, Sites 356 and 357), *in* Supko, P. R., Perch-Nielsen, K., et al., eds., Initial Reports of Deep Sea Drilling Project, Volume 39: Washington, D.C., U.S. Government Printing Office, p. 869–883.
- Benson, R. H., and Peypouquet, J. P., 1983, The upper and mid-bathyal Cenozoic ostracode faunas of the Rio Grande Rise found on Leg 72 Deep Sea Drilling Project, *in* Barker, P. F., Carlson, R. L., Johnson, D. A., et al., eds., Initial Reports of Deep Sea Drilling Project, Volume 72: Washington, D.C., U.S. Government Printing Office, p. 805–818.
- Boomer, I., 1994, On *Pariceratina ubiquita* Boomer sp. nov.: Stereo-Atlas of Ostracod Shells, v. 21, p. 79–86.

- Boomer, I., and Whatley, R., 1995, Cenozoic Ostracoda from Guyots in the western
 Pacific: Holes 865B and 866B (Leg 143), *in* Winterer, E. L., Sager, W. W.,
 Firth, J. V., and Sinton, J. M., eds., Proceedings of the Ocean Drilling
 Program, Scientific Results, p. 75–86.
- Bralower, T. J., and Mutterlose, J., 1995, Calcareous nannofossil biostratigraphy of Site 865, Allison Guyot, central Pacific Ocean: a tropical Paleogene reference section, *in* Winterer, E. L., Sager, W. W., Firth, J. V., and Sinton, J. M., eds., Proceedings of the Ocean Drilling Program, Scientific Results Volume 134: College Station, Texas, p. 31–74.
- Bralower, T. J., Zachos, J. C., Thomas, E., Parrow, M., Paul, C. K., Kelly, D. C.,
 Premoli Silva, I., Sliter, W. V., and Lohmann, K. C., 1995, Late Paleocene to
 Eocene paleoceanography of the equatorial Pacific Ocean: Stable isotopes
 recorded at Ocean Drilling Program Site 865, Allison Guyot:
 Paleoceanography, v. 10, no. 4, p. 841–866.
- Coles, G., and Whatley, R., 1989, New Paleocene to Miocene genera and species of Ostracoda from DSDP sites in the North Atlantic: Revista Española de Micropaleontología, v. 21, p. 81–124.
- Didié, C., Bauch, H. A., and Helmke, J. P., 2002, Late Quaternary deep-sea ostracodes in the polar and subpolar North Atlantic: paleoecological and paleoenvironmental implications: Palaeogeography Palaeoclimatology Palaeoecology, v. 184, p. 195–212.
- Ducasse, O., and Peypouquet, J. P., 1979, Cenozoic Ostracodes: Their importance for bathymetry, hydrology, and biogeography, *in* Montadert, L., Roberts, D. G., et al., eds., Initial Reports of the Deep Sea Drilling Project, Volume 48:
 Washington, D.C., U.S. Government Printing Office, p. 343–363.

- Dunkley Jones, T., Lunt, D. J., Schmidt, D. N., Ridgwell, A., Sluijs, A., Valdes, P. J., and Maslin, M., 2013, Climate model and proxy data constraints on ocean warming across the Paleocene–Eocene Thermal Maximum: Earth Science Reviews, v. 125, p. 123–145.
- Gradstein, F., Ogg, J. G., Schmitz, M., and Ogg, G.,2012, The Geologic Time Scale 2012, Elsevier, 1176 p.
- Guernet, C., and Bellier, J.-P., 2000, Ostracodes Paléocènes et éocènes du Blake Nose (Leg ODP 171B) et Évolution des envrionmentsbathyaux au large de la Floride Revue de Micropaleontologie, v. 43, no. 4, p. 31.
- Guernet, C., and Danelian, T., 2006, Ostracodes bathyaux du Crétacé terminal Éocène moyen en Atlantique tropical (Plateau de Demerara, Leg 207): Revue de micropaléontologie, v. 49, p. 215–225.
- Guernet, C., and Galbrun, B., 1992, Data report: Preliminary report on the ostracodes of Leg 122 (Exmouth Plateau, Indian Ocean), *in* von Rad, U., and Haq, B. U. et al., eds., Proceedings of the Ocean Drilling Program, Scientific Results, Volume 122: Colleage Station Ocean Drilling Program, p. 835–838.
- Guernet, C., and Molina, E., 1997, Les ostracodes et le passage Paléocène-Eocène dans les Cordillères bétiques (Coupe de Caravaca, Espagne): Geobios, v. 30, no. 1, p. 31–43.
- Heck, K. L., van Belle, G., and Simberloff, D., 1975, Explicit calculation of the rarefaction diversity measurement and the determination of sufficient sample size: Ecology, v. 56, p. 1459–1461.
- Hollis, C. J., Strong, C. P., Rodgers, K.A., and Rogers, K. M., 2003, Paleoenvironmental changes across the Cretaceous/Tertiary boundary at Flaxbourne River and Woodside Creek, eastern Marlborough, New Zealand: New Zealand Journal of Geology and Geophysics, v. 46, no. 2, p. 177–197.

- Horne, D. J., Brandão, S. N., and Slipper, I. J., 2011, The Platycopid Signal deciphered: Responses of ostracod taxa to environmental change during the Cenomanian–Turonian Boundary Event (Late Cretaceous) in SE England: Palaeogeography Palaeoclimatology Palaeoecology, v. 308, no. 3–4, p. 304–312.
- Hurlbert, S. H., 1971, The nonconcept of species diversity: a critique and alternative parameters: Ecology, v. 52, p. 577–586.
- Kelly, D. C., Bralower, T. J., and Zachos, J. C., 2001, On the Demise of the early Paleogene *Morozovella velascoensis* lineage: Therminal Progenesis in the planktonic foraminifera: Palaios, v. 16, p. 507–523.
- Kozdon, R., Kelly, D. C., Kita, N. T., Fournelle, J. H., and Valley, J. W., 2011,
 Planktonic foraminiferal oxygen isotope analysis by ion microprobe technique suggests warm tropical sea surface temperatures during the Early Paleogene:
 Paleoceanography, v. 26, p. PA3206.
- Oksanen, J. Blanchet, J.F., Kindt, R., Legendre, P., Minchin, P.R., O'Hara, R.B., Simpson, G.L., Solymos, P., Stevens M.H.H, and Wagner H., 2013, vegan: Community Ecology Package: R package version 2.0-7. http://CRAN.Rproject.org/package=vegan
- Lourens, L.J., Hilgen, F.J., Gudjonsson, L., Zachariasse, W.J., 1992, Late Pliocene to early Pleistocene astronomically forced sea surface productivity and temperature variations in the Mediterranean: Marine Micropaleontology, v. 19, 49–78.
- Ma, Z., Gray, E., Thomas, E., Murphy, B., Zachos, J., and Paytan, A., 2014, Carbon sequestration during the Palaeocene– Eocene Thermal Maximum by an efficient biological pump: nature geoscience, v. 7, p. 382–388.

- Majoran, S., and Dingle, R. V., 2002, Cenozoic deep-sea ostracods from Maud Rise, Weddell Sea, Antarctica (ODP Site 689): a palaeoceanographical perspective: Geobios, v. 35, p. 137–152.
- Marshall, C. R., 1997, Confidence intervals on stratigraphic ranges with nonrandom distributions of fossil horizons: Paleobiology, v. 23, p. 165–173.
- Rivadeneira, M. M., Hunt, G., and Roy, K., 2009, The use of sighting records to infer species extinctions: an evaluation of different methods: Ecology, v. 90, no. 5, p. 1291–1300.
- Rodriguez-Lazaro, J., and Garcia-Zaraga, E., 1996, Paleogene deep-marine ostracodes from the Basque Basin, *in* Keen, M. C., ed., Proceedings of the 2nd European Ostracodologists Meeting: London, British Micropalaeontological Society, p. 79–85.
- Röhl, U., Westerhold, T., Bralower, T. J., and Zachos, J. C., 2007, On the duration of the Paleocene-Eocene thermal maximum (PETM): Geochemisty, Geophysics, Geosystems, v. 8, p. Q12002.
- Sciuto, F., 2014, First report on the palaeopsychrospheric ostracod genus *Nemoceratina (Pariceratina)* Gründel and Kozur, 1972 (Ostracoda,
 Bythocytheridae) from the Quaternary of the Mediterranean basin and
 description of a new species: Palaeontologia Electronica v. 17, p. 32A.
- Steineck, P. L., and Thomas, E., 1996, The latest Paleocene crisis in the deep sea: Ostracode succession at Maud Rise, Southern Ocean: Geology, v. 24, p. 583– 587.
- Stepanova, A., and Lyle, M., 2014, Deep-sea Ostracoda from the Eastern Equatorial Pacific (ODP Site 1238) over the last 460 ka: Marine Micropaleontology, v. 111, p. 100–117.

Whatley, R. C., and Coles, G. P., 1991, Global change and the biostratigraphy of North Atlantic Cainozoic deep water Ostracoda: Journal of micropalaeontology, v. 9, no. 2, p. 119–132.

Yamaguchi, T., and Norris, R. D., 2012, Deep-sea ostracode turnovers through the Paleocene–Eocene thermal maximum in DSDP Site 401, Bay of Biscay, North Atlantic: Marine Micropaleontology, v. 86–87, p. 32–44.