DATA REPOSITORY ITEM 2014244

CARBONATE CARBON AND OXYGEN ISOTOPES

We report $\delta^{13}C_{carb}$ and $\delta^{18}O_{carb}$ measurements of 48 samples from sections J1204 and F930 of the Callison Lake Dolostone (Table DR1). These are parallel measured sections (~1-2 m apart) of the exact same strata from a small drainage south of Mount Gibben at N64°40'53.6" W139°13'55.6" in the Coal Creek inlier, Yukon Territory, Canada. In Fig. 1 of the manuscript, we present $\delta^{13}C_{carb}$ data solely from J1204 because F930 is a lower resolution repetition of the same section. We use the $\delta^{13}C_{org}$ data from F930 based on very straightforward bed-by-bed correlations combined with matching the same $\delta^{13}C_{carb}$ values (Table DR1).

Carbon ($\delta^{13}C_{carb}$) and oxygen ($\delta^{18}O_{carb}$) isotopic results are reported in per mil notation of $^{13}C/^{12}C$ and $^{18}O/^{16}O$, respectively, relative to the standard VPDB (Value of the Pee-Dee Belemnite). Dolostone samples were cut perpendicular to bedding, polished, and carefully microdilled (~2-10 mg of powder) to avoid secondary veins, cements, and siliciclastic components. Carbonate $\delta^{13}C$ and $\delta^{18}O$ isotopic data were acquired simultaneously on a VG Optima dual inlet isotope ratio mass spectrometer coupled with a VG Isocarb preparation device (Micromass, Milford, MA) in the Laboratory for Geochemical Oceanography at Harvard University. Approximately 1 mg of sample powder was reacted in a common, purified phosphoric acid (H₃PO₄) bath at 90°C. The evolved CO₂ was collected cryogenically and analyzed using an in-house reference gas. Measured data were calibrated to VPDB using the Cararra marble standard. Total analytical errors (1 σ) are better than \pm 0.1‰ for both $\delta^{13}C$ and $\delta^{18}O$ based on repeat analysis of standards and samples. Increasing the reaction time to eleven minutes for dolomite samples minimized potential "memory effects" resulting from the common acid-bath system, with the total memory effect estimated at <0.1‰ based on reproducibility of standards run directly after samples.

ORGANIC CARBON ISOTOPES

We report $\delta^{13}C_{org}$ measurements from 13 samples of section F930 of the Callison Lake Dolostone (Table DR1). $\delta^{13}C_{org}$ values were obtained from the total organic carbon (TOC) of insoluble residues. Whole rock samples were initially trimmed to remove weathered material and secondary veins and then crushed into a powder in a SPEX 8500 Shatterbox using a hardened steel grinding container and puck. Analyses were performed on large (\sim 10-20g) samples to accommodate low TOC values. Samples were decalcified with concentrated HCl (6N) for 48 hours, buffered back to a neutral pH (>pH5), filtered, and dried. Care was taken to ensure that acid was added and acidification continued until there was absolutely no visible carbonate dissolution so that the analyses would not be affected by contamination from residual inorganic carbon. Homogenized residues were analyzed in the Harvard University Laboratory for Geochemical Oceanography on a Carlo Erba Elemental Analyzer attached to a ThermoFinnigan Delta V configured in continuous flow mode. Samples and standards were bracketed such that our 13 organic carbon analyses (each run in duplicate) were associated with 6 internal standards. These standards, each with known organic carbon contents and isotope values, were used to calibrate TOC contents and isotopic compositions. The mass of insoluble residue was taken as siliciclastic content, and TOC values for the bulk samples were calculated by combining the carbonate concentration data obtained from the Delta with measurements of the ratio of insoluble residue to original pre-decarbonated powder.

RHENIUM-OSMIUM GEOCHRONOLOGY

Organic-rich (TOC = 1.15-2.5%), vase-shaped microfossil-bearing black shale was sampled from section J1204. Ten samples large enough for Re-Os geochronology (~50-200 g) were collected over an interval of 2.2 m (J1204-16.8–18.2 m) after an ~25 cm deep trench was dug to remove weathered material from the outcrop. Seven of these ten samples (Table DR2: B–H) were collected horizontally from a very thin (<10 cm) vertical interval (17.7 m), as a large horizontal sampling technique is generally used to maximize the spread of ¹⁸⁷Re/¹⁸⁸Os (Kendall et al., 2009). Given the lack of variability from this small horizon, we incorporated three other vertical samples (Table DR2: A, I, and J) to develop our final isochron age (Fig. 3 of manuscript). This 2.2 m interval is sedimentologically identical and the Osi values (Table DR2) of every sample is practically indistinguishable from one another.

All of the sample weathered surfaces were removed with a diamond-coated rock saw and samples were then hand-polished using a diamond-plated polishing pad to remove cutting marks and eliminate any potential for contamination from the saw blade. The samples were dried overnight at ~60 °C and then crushed to a fine (~30 μ m) powder in a SPEX 8500 Shatterbox using a zirconium grinding container and puck in order to homogenize any Re and Os heterogeneity present in the samples. Re and Os isotopic abundances and compositions were determined at the Department of Earth and Atmospheric Sciences, University of Houston (UH) following methodology developed by van Acken et al. (2012) and Wittig et al. (2013).

0.6 g of sample was digested and equilibrated in 8 ml of $Cr^{VI}O_3$ -H₂SO₄ together with a mixed tracer (spike) solution of ¹⁹⁰Os and ¹⁸⁵Re in carius tubes at 220 °C for 48 hours. Rhenium and osmium was extracted and purified using solvent extraction (NaOH, (CH₃)₂CO, and CHCl₃,), micro-distillation, anion column chromatography methods, and negative mass spectrometry as outlined by Selby and Creaser (2003) and Cumming et al. (2013). The Cr^{VI}O₃-H₂SO₄ digestion method was employed as it has been shown to preferentially liberate hydrogenous Re and Os yielding more accurate and precise age determinations (Selby and Creaser, 2003; Kendall et al., 2004; Rooney et al., 2011). Total procedural blanks during this study were 12.3 ± 0.1 pg and 0.05 ± 0.15 pg for Re and Os respectively, with an average ¹⁸⁷Os/¹⁸⁸Os value of 0.172 ± 0.208 (1 σ , *n* = 3).

Isotopic measurements were performed using the UH ThermoElectron TRITON PLUS mass spectrometer via static Faraday collection for Re and ion-counting using a secondary electron multiplier in peak-hopping mode for Os. In-house Re and Os solutions were continuously analyzed during the course of this study to ensure and monitor long-term mass spectrometer reproducibility. The University of Houston Re standard solution measured on the faraday cups yields an average ¹⁸⁵Re/¹⁸⁷Re value of 0.59827 ± 0.00158 (2σ , n = 10), which is identical to that of (Rooney et al., 2010). The measured difference in ¹⁸⁵Re/¹⁸⁷Re values for the Re solution and the accepted ¹⁸⁵Re/¹⁸⁷Re value (0.5974) (Gramlich et al., 1973) is used to correct the Re sample data. The Os isotope standard solution used at UH is the in-house standard from the University of Maryland. Over the past two years on this Triton, the runs yield a ¹⁸⁷Os/¹⁸⁸Os ratio of

 0.11388 ± 0.00116 (2 σ , n = 41) that is identical, within uncertainty, to the accepted value reported in Brandon et al., (1999).

Uncertainties for ¹⁸⁷Re/¹⁸⁸Os and ¹⁸⁷Os/¹⁸⁸Os are determined by error propagation of uncertainties in Re and Os mass spectrometry measurements, blank abundances and isotopic compositions, spike calibrations, and reproducibility of standard Re and Os isotopic values. The Re-Os isotopic data, 2σ calculated uncertainties for ¹⁸⁷Re/¹⁸⁸Os and ¹⁸⁷Os/¹⁸⁸Os, and the associated error correlation function (rho) are regressed to yield a Re-Os date using *Isoplot V*. 4.15 with the λ ¹⁸⁷Re constant of 1.666 x 10 ⁻⁴ a (Ludwig, 1980; Smoliar et al., 1996; Ludwig, 2011). Elemental Re and Os abundances for the J1204 samples range from 0.7 to 9.5 ppb, and 61.8 to 357.6 ppt, respectively, with ¹⁸⁷Re/¹⁸⁸Os and ¹⁸⁷Os/¹⁸⁸Os ratios between 63 and 196, and 1.421 and 3.050 respectively (Table DR2). Regression of the isotopic composition data for these samples yields a Model 1 age of 739.9 ± 6.1 Ma (6.5 if the ¹⁸⁷Re decay constant uncertainty is included, *n* = 10, Mean Square of Weighted Deviates [MSWD] = 0.62, initial ¹⁸⁷Os/¹⁸⁸Os = 0.609 ± 0.01; Fig. 3).

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Table Brizi carbonate carbon and oxygen botopes of the appen cambon zake boloston	Table DR1: Carbonate Carbon	, Organic Carbon and Ox	ygen isotopes of the upp	er Callison Lake Dolostone
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Section	Stratigraphic Height	δ ¹³ C _{carb}	δ ¹⁸ O _{carb}	δ ¹³ C _{org}	Epsilon	Siliclastic Content	Carbonate Content	тос
	(m)	(‰)	(‰)	(‰)	(‰)	(%)	(%)	(%)
J1204	0.3	4.00	-2.84	_	_	_	_	_
J1204	1	3.91	-3.11	_	_	-	-	_
J1204	1.3	3.61	-5.12	_	_	_	-	_
J1204	2	3.77	-4.16	_	_	-	-	_
J1204	2.5	3.61	-4.36	-	_	-	-	-
J1204	3	3.78	-2.45	_	_	-	-	_
J1204	3.4	3.15	-2.96	-	_	-	-	-
J1204	4.1	2.58	0.35	_	_	-	-	_
J1204	4.4	2.26	-0.19	-	_	-	-	-
J1204	5	2.01	-0.93	_	_	-	-	_
J1204	5.5	1.67	0.42	_	_	-	-	_
J1204	6.1	1.63	0.11	-	_	-	-	-
J1204	6.5	1.29	-0.19	_	_	-	-	_
J1204	7	0.73	0.31	-	_	-	-	-
J1204	7.4	0.01	0.92	_	_	_	-	_
J1204	8	1.05	1.33	_	_	-	-	_
J1204	8.3	0.58	1.00	-	_	-	-	-
J1204	19.3	-2.78	-1.19	_	_	-	-	_
J1204	20	-2.67	-1.22	-	_	-	-	-
J1204	20.4	-4.02	-1.56	_	_	-	-	_
J1204	20.9	-3.77	-0.94	-	_	-	-	-
J1204	21.1	-3.60	-0.62	_	_	-	-	_
J1204	21.9	-7.00	-1.30	_	_	-	-	_
J1204	22.3	-2.76	-4.21	-	_	-	-	-
J1204	22.9	-4.21	-1.87	_	_	-	-	_
J1204	23.3	-3.96	-2.01	-	_	-	-	-
J1204	23.9	-3.94	-2.07	_	_	-	-	_
J1204	24.4	-4.07	-2.22	-	_	-	-	-
J1204	25.4	-2.88	-5.85	-	_	-	-	-
J1204	25.9	-6.53	-2.16	-	_	-	-	-
J1204	27.2	-5.64	-2.07	-	-	-	-	-
J1204	27.5	-5.50	-1.89	-	_	-	-	-
J1204	29.6	-3.26	-0.82	-	-	-	-	-
J1204	30.9	-2.97	-2.95	-	_	-	-	-
J1204	32.2	-2.37	-0.28	-	_	-	-	-
F930	2	3.73	-3.65	-25.14	28.87	9.62	90.25	0.13
F930	3	3.38	-1.71	-29.19	32.58	13.82	86.07	0.11
F930	5	2.30	-0.08	-26.93	29.23	13.42	86.44	0.14
F930	6	1.61	0.33	-30.50	32.11	10.20	89.72	0.09
F930	8	1.15	-0.40	-30.08	31.22	13.16	86.74	0.09
F930	9	-0.98	0.98	-31.18	30.20	23.08	76.73	0.19
F930	16	-3.31	-0.16	-31.22	27.91	6.37	93.58	0.04
F930	18	-6.77	-1.91	-31.53	24.76	10.03	89.86	0.11
F930	20	-4.55	-2.72	-27.68	23.13	27.58	72.33	0.09
F930	21	-5.88	-1.88	-30.11	24.23	4.41	95.50	0.08
F930	22	-3.89	0.03	-28.16	24.27	3.64	96.29	0.07
F930	23	-4.83	-0.62	-28.88	24.05	2.01	97.95	0.04
F930	25	-1.39	-0.14	-32.87	31.48	41.57	58.41	0.01
F930	26	-1.62	-0.39	-34.99	33.36	45.42	54.55	0.03

F930-2–9 m = J1204-0.3–8.3 m and F930-16–26m = J1204-19.3–32.2 m

Table DR2: Re and Os abundance and isotope composition data for the Callison Lake Dolostone

Sample	Isochron point	Re (ppb)	±	Os (ppt)	±	¹⁹² Os (ppt)	±	¹⁸⁷ Re / ¹⁸⁷ Os	±	¹⁸⁷ Os/ ¹⁸⁸ Os	±	rhoª	Os initial ^b
16.8	А	0.71	0.01	61.8	0.2	21.8	0.1	64.6	1.1	1.421	0.016	0.489	0.620
17.7#1	В	3.29	0.02	138.8	0.4	44.0	0.1	148.8	1.0	2.453	0.013	0.659	0.608
17.7#2	С	1.93	0.01	80.3	0.3	25.4	0.1	151.5	1.5	2.479	0.022	0.732	0.600
17.7#3	D	3.55	0.02	174.1	0.8	56.9	0.1	124.1	1.0	2.152	0.032	0.395	0.613
17.7#4	E	5.67	0.02	203.3	2.1	63.9	0.4	182.8	0.5	2.880	0.018	0.668	0.613
17.7#5	F	9.51	0.04	357.6	4.9	77.9	0.3	171.8	0.5	2.738	0.029	0.384	0.608
17.7#6	G	4.71	0.03	159.7	1.8	53.6	0.3	196.2	0.7	3.050	0.022	0.617	0.617
17.7#7	Н	4.87	0.02	178.3	1.8	42.2	0.1	177.5	0.5	2.814	0.017	0.680	0.612
18.1	I	1.87	0.01	92.6	0.3	30.3	0.2	123.1	1.1	2.131	0.015	0.704	0.605
18.2	J	1.68	0.01	148.3	0.5	52.6	0.1	63.4	0.7	1.393	0.015	0.462	0.606

Uncertainties are given as 2σ for ¹⁸⁷Re/¹⁸⁸Os, ¹⁸⁷Os/¹⁸⁸Os and ¹⁹²Os

For the latter the uncertainty includes the 2 SE uncertainty for mass spectrometer analysis plus uncertainties for Os blank abundance and isotopic composition.

Ages are calculated using the λ^{187} Re = 1.666 x 10⁻¹¹y^{r-1} (Smoliar et al., 1996).

a Rho is the associated error correlation (Ludwig, 1980)

b Os initial = ¹⁸⁷Os/¹⁸⁸Os isotope composition calculated at 739 Ma

Table DR3: Locations where VSMs have been described in the literature (after Porter and Knoll, 2000). Age constraints are rough and noted from italisized references when neccessary.

Formation/Rock Unit	Location	Paleocontinent	Relationship to Islay anomaly?	Approximate Age	Reference
A. Assemblages with multiple VSM species in	common:				
Backlundtoppen Formation	Spitsbergen, Svalbard	Laurentia	Older	780–716 Ma	Knoll et al. (1989); Halverson et al. (2004)
Callison Lake Dolostone	Yukon, Canada	Laurentia	Yes	ca. 740 Ma	This paper
Draken Formation	Spitsbergen, Svalbard	Laurentia	Older	780–716 Ma	Knoll et al. (1991); Halverson et al. (2004)
Chuar Group	Grand Canyon, Arizona	Laurentia	Older?	785–742 Ma	Bloeser (1985); Porter et al. (2003); Karlstrom et al. (2000); Dehler et al. (2012)
Beck Springs Dolomite	Death Valley, California	Laurentia	Yes	770–716 Ma	Licari (1978); Horodyski (1993); Macdonald et al. (2013); Mahon et al. (2014)
Russøya Member, Elbobreen Formation	Nordhaustlandet, Svalbard	Laurentia	Yes	780–716 Ma	Knoll and Calder (1983); Halverson et al. (2004)
Uinta Mountain Group	Uinta Mountains, Utah	Laurentia	Older?	785–716 Ma	Link et al. (1993); Dehler et al. (2010)
B. Assemblages with broad morphological sir	milarity to the Callison Lake Dolos	tone, but with dist	inct taxa or insufficient systematic res	earch	
Bed 18, Eleonore Bay Group	East Greenland	Laurentia	Unknown	780–716 Ma	Vidal (1979); Green et al. (1988); Hoffman et al. (2012)
Chatkaragai Suite	Tien Shan, Russia/Kyrgyzstan	Kazakstan	Unknown	800–766 Ma	Kraskov (1985); Yankaouskas (1989); Sergeev and Schopf (2010); Meert et al. (2011)
Chichkan Formation	Kazakstan	Kazakstan	Unknown	800–766 Ma	Sergeev and Schopf (2010); Meert et al. (2011)
Jabal Rockham	Saudi Arabia	Arabia	Unknown	>650 Ma	Binda and Bokhari (1980); Johnson (2003)
Togari Group	Tasmania	Australia	Yes	780–716 Ma	Saito et al. (1988); Turner et al. (1998)
Visingsö Beds	Sweden	Baltica	Unknown	805–663 Ma	Ewetz (1933); Knoll and Vidal (1980); Vidal and Siedlecka (1983); Martí Mus and Moczydlowska (2000)
C. Reported vasiform microfossils that are dis	stinct from the Callison Lake Dolo	stone, insufficientl	y illustrated, or poorly preserved		
Bonahaven Formation, Dalradian Supergroup	p Scotland	Laurentia?	Younger	665–635 Ma	Anderson et al. (2013)
Dengying Formation	China	South China	Younger	551–542 Ma	Zhang and Li (1991); Ding et al. (1992); Duan et al. (1993); Zhang (1994); <i>Condon et al. (2005)</i>
Doushantuo Formation	China	South China	Younger	635–551 Ma	Duan (1986); Duan et al. (1993); Li et al. (2008); Condon et al. (2005)
Jacadigo Group	Brazil	Amazonia	Younger	716–635 Ma	Fairchild et al. (1978); Freitos et al. (2011)
Rasthof Formation	Namibia	Congo	Younger	665–635 Ma	Bosak et al. (2011)
Simla Slates	India	India	Unknown	823–716 Ma	Nautiyal (1978); Jiang et al. (2003)
Tanafjorden Group	Norway	Baltica	Unknown	807–716 Ma	Vidal and Siedlecka (1983); Vidal and Moczydlowska (1995)
Tindir Group	Alaska	Laurentia	Unknown	811–716 Ma	Allison and Awramik (1994); Macdonald et al. (2010)
Tsagaan Oloom Formation	Mongolia	Mongolia	Younger	665–635 Ma	Bosak et al. (2011)
Upper Min'yar Formation	Urals, Russia	Baltica	Unknown	820687 Ma	Maslov et al. (1994); Maslov (2004)
Vaishnodevi Limestone	Himalaya, India	India	Unknown	<950 Ma	Venkatachala and Kumar (1998); McKenzie et al. (2011)
Vindhyan Supergroup	India	India	Unknown	>650 Ma?	Maithy and Babu (1988); Ray et al. (2002)
Virgin Springs Limestone	Death Valley, California	Laurentia	Younger	740–716 Ma	Macdonald et al. (2013); This paper

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