

## GSA DATA REPOSITORY 2013281

### Data repository for Slab flattening trigger for isotopic disturbance and magmatic flare-up in the southernmost Sierra Nevada batholith, California

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### APPENDIX DR1: Sample Locations, Petrography, and Zircon Morphology

Descriptions of analyzed zircon grains and the samples (corresponding to Fig. 1) from which they were extracted are given below.

#### Plutonic Rocks

##### **6 (originally 08SE451; 34° 52' 24.2" N, 119° 6' 40.7" W)**

Antomony Peak tonalite (Chapman and Saleeby, 2012). Light gray, medium- to coarse-grained, porphyritic, massive to penetratively foliated, hornblende-biotite tonalite containing mainly plagioclase, quartz, and hornblende, with minor amounts of garnet, K-feldspar, biotite, epidote, apatite, titanite, and ilmenite. Where hypidiomorphic granular (igneous) textures are not overprinted by metamorphic foliation, the epidote occurs as subhedral replacement products of embayed hornblende and retains distinct core and rim regions, consistent with a magmatic origin. Original petrographic relationships are locally obscured by extensive seritization and chloritization. Hornblende-rich mafic schlieren are found in the same outcrop as 08SE451. See Chapman et al. (2012) for U-Pb zircon geochronologic data.

##### **7 (originally 08SE46; 34° 54' 35.3" N, 118° 58' 30.7" W)**

San Emigdio tonalite (Chapman and Saleeby, 2012). Gray, medium- to coarse-grained, weakly foliated biotite tonalite. Alteration is minor. See Chapman et al. (2012) for U-Pb zircon geochronologic data.

##### **8 (originally 08SE675; 34° 53' 48.8" N, 118° 55' 20.9" W)**

San Emigdio tonalite (Chapman and Saleeby, 2012). Foliated garnet-biotite leucotonalite. Similar in appearance to the 101 Ma “Garnet-biotite tonalite of Grapevine” of Pickett and Saleeby (1993, 1994). Epidote phenocrysts embayed in biotite. Garnet is 0.5–1 mm in diameter with blebbly quartz inclusions. Accessory minerals include zircon, apatite, and opaques. Very little alteration. Radiation-damage halos surrounding zircon inclusions in biotite are common. See Chapman et al. (2012) for U-Pb zircon geochronologic data.

Zircon grains in samples 7 (Fig. DR1a and b), 6 (Fig. DR1c), and 8 (Fig. DR1d) are generally subhedral, ~50 – 350 µm in length, inclusion-poor, and exhibit simple oscillatory zoning patterns in Cathodoluminescence (CL) images that we interpret as magmatic features. Cathodoluminescence-dark cores are commonly overgrown by CL-bright rims up to ~35 µm wide. In situ laser ablation multicollector inductively coupled mass spectrometry analyses of core and rim regions are concordant with weighted mean ages of ca. 134 to 136 Ma and  $98.1 \pm 4.9$  Ma, respectively (Chapman et al., 2012). These ages are interpreted to reflect the timing of crystallization (cores) and deep crustal overgrowth (rims). A single aberrant grain in sample 8, yielding an age of  $485.0 \pm 34.0$  Ma, is interpreted as related to source inheritance and/or entrainment of a Paleozoic or older zircon.

## **San Emigdio Schist**

The San Emigdio Schist is juxtaposed against plutonic rocks along the remains of a locally ductile to brittle low-angle detachment fault system that probably correlates with the type Rand fault exposed in the Rand Mountains (Postlethwaite & Jacobson, 1987). During Pliocene–Quaternary compression, the Rand fault in the San Emigdio Mountains was largely folded with a major south-dipping thrust fault cutting across the folded structural section. Reported sample depths were approximated based on metamorphic grade and the structural distance from the Rand fault in locations where ductile fabrics are well-preserved.

### **1 (originally 08SE467; 34° 54' 29.3" N, 118° 58' 39.2" W)**

San Emigdio Schist metasandstone collected adjacent to the Rand fault. Outcrop contains abundant deformed quartzofeldspathic (milky quartz + plagioclase ± garnet ± muscovite) veins and stromatic leucosomes and melanosomes, interpreted to result from fluid-saturated partial melting (Chapman et al., 2011). Sample is a coarse-grained garnet + plagioclase + biotite + quartz + ilmenite blastomylonitic schist. Garnet porphyroblasts are typically 2 to 4 mm in diameter and subhedral, with inclusions of plagioclase, biotite, and quartz. Diffusive relaxation of major and trace element growth zonation is common in garnet porphyroblasts from this sample. Plagioclase porphyroblasts are twinned, contain little to no graphite and/or carbonaceous material, and exhibit distinct core and rim regions. Quartz grains lack undulatory extinction and display 120° grain boundaries. Minor minerals include apatite, zircon, and rutile. Secondary minerals, including clinzoisite, celadonitic muscovite, and fibrolitic sillimanite, usually occur within or adjacent to biotite. Radiation-damage halos surrounding zircon inclusions in biotite are common. Phase relations suggest that peak temperatures exceeded 700° C at 10 kbar (Chapman et al., 2011).

### **2 (originally 99-57 (Grove et al., 2003); 34° 52' 12.6" N, 119° 6' 34.7" W)**

San Emigdio Schist metasandstone collected ~25 m below the Rand fault. This sample shows a protomylonitic texture (Grove et al., 2003).

### **3 (originally 06SE23 and SE1b (Jacobson et al., 2011); 34° 52' 56.0" N, 119° 4' 21.1" W)**

San Emigdio Schist metasandstone collected from ~30 m below the Rand fault. Medium- to coarse-grained garnet + plagioclase + biotite + quartz + rutile blastomylonitic schist. Garnet grains are idioblastic and 0.5 to 2 mm in diameter with well-defined growth zonation and blebby inclusions of exclusively quartz confined to individual zones (Chapman et al., 2011), suggesting epitaxial growth of quartz. Some porphyroblasts have pressure shadows composed of fine quartz, plagioclase, and chlorite. Some garnet grains are torn completely apart or frictionally displaced up to 2 mm along fractured surfaces with no systematic orientation. Plagioclase is poikiloblastic with blebby quartz inclusions and finely disseminated graphite and/or carbonaceous material. Some grains exhibit polysynthetic twinning. Quartz grains lack undulatory extinction and display 120° grain boundaries, indicative of high-temperature annealing. Secondary alteration is mainly restricted to mm-scale shear bands that truncate the peak mineral assemblage and contain retrograde clinzoisite, chlorite, celadonitic muscovite, and tourmaline. Sample contains minor tschermakitic hornblende. Calculated temperatures and pressures for this sample are 665 ± 29 °C and 10.0 ± 0.9 kbar (Chapman et al., 2011).

### **4 (originally SE03 (Jacobson et al., 2011); location approximate: 34° 52' 51.2" N, 119° 3' 21.6" W)**

San Emigdio Schist metasandstone collected from ~100 m below the Rand fault.

### **5 (originally 07SE34; 34° 53' 4.4" N, 119° 1' 38.4" W)**

San Emigdio Schist metasandstone collected from ~400 m below the Rand fault. Fine- to medium-grained plagioclase + biotite + chlorite + muscovite + clinzoisite + quartz granular schist. A faint shape-preferred orientation of biotite defines the schistosity. Plagioclase is poikiloblastic, lacks twinning, and

contains blebby quartz inclusions and finely disseminated graphite and/or carbonaceous material. Garnet porphyroblasts are sparse, with a single ~1 mm skeletal grain in thin section. Biotite is closely associated with ilmenite, tourmaline, and calcite. This sample was not suitable for thermobarometric work since garnet was not in textural equilibrium with plagioclase, biotite, or muscovite. However, phase relations suggest that peak temperatures did not exceed 600 °C (Chapman et al., 2011).

Zircon grains from all structural levels in the schist share similar morphologies. Most grains are subhedral to anhedral, ~10 – 150 µm in length, inclusion-poor, with abundant fractures. Cathodoluminescence imaging indicates a clear relationship between zircon zoning patterns and structural depth (Fig. DR1). At the deepest levels of exposure (> 100 m; e.g., sample 5; Fig. DR1f), zircons typically exhibit simple oscillatory zoning patterns that we interpret as magmatic features. At shallow structural levels (< 100 m), oscillatory zoning is still generally present; however, metamorphic domains truncate older magmatic zoning in several grains and narrow (1 -50 µm) rims of CL-bright zircon commonly crosscut zoning in zircon cores. Adjacent to the Rand fault (e.g., samples 3; Fig. DR1e and 1; Fig. 2), little to no oscillatory zoning is present in zircons; instead, zircons lack cores and display complex metamorphic zoning patterns characterized by multiple crosscutting domains.

## APPENDIX DR2: Analytical Methods

### Laser Ablation–Multicollector Inductively Coupled Plasma–Mass Spectrometry

U-Pb geochronology of zircon grains was conducted by laser ablation multicollector inductively coupled plasma mass spectrometry (LA-MC-ICPMS) at the Arizona LaserChron Center following the methods outlined in Gehrels et al. (2006). Zircon grains were extracted from plutonic and metamorphic samples using standard mineral separation techniques of crushing, sieving, magnetic separation, processing through heavy liquids, and hand picking. Separates were then mounted in epoxy, polished, and imaged on the Caltech Zeiss 1550 VP field emission scanning electron microscope before analysis. Representative CL images are shown in Figure DR1. Zircons were ablated using a 193 nm ArF laser with a pit depth of ~12 µm and spot diameters of 25-30 µm depending on grain size. SL2 (564 Ma ± 4 Ma; Gehrels, personal communication, 2011) and R33 (419 ± 1 Ma; Black et al., 2004) zircon standards were used. Data reduction was done using in-house Arizona LaserChron Center Microsoft Excel programs and ISOPLOT/Ex Version 3 (Ludwig, 2003). U-Pb isotopic data is presented in Table SD1. Plots of U-content and U/Th versus U-Pb age are shown in Figure DR2.

### SIMS

Oxygen isotope analyses were conducted with a CAMECA IMS 7f-GEO magnetic sector type secondary ion mass spectrometry (SIMS) instrument at the Caltech Center for Microanalysis. Grain mounts used in U-Pb analysis were polished with 0.25 µm colloidal silica and gold coated prior to SIMS analysis. When possible, zircon grains were analyzed from the same crystal growth domain where U-Pb data were obtained. Zircon grains were sputtered using a 2 nA  $^{133}\text{Cs}^+$  primary ion beam with 20x30 µm diameter, secondary ions were extracted at 10 kV, and  $^{16}\text{O}^-$  and  $^{18}\text{O}^-$  were measured simultaneously using electron multiplier detectors and  $^{18}\text{O}/^{16}\text{O}$  isotope ratios were derived. Each analysis involved 10 to 60 seconds of pre-sputtering, followed by 10 to 20 cycles of data collection, resulting in a total analysis time of 2 to 5 minutes per spot. Instrumental mass fractionation (IMF) was corrected for, using raw isotopic ratios, by bracketing sample and R33 standard ( $5.55 \pm 0.04\text{\textperthousand}$ ; Valley, 2003) analyses every 5 to 10 measurements. Finally,  $\delta^{18}\text{O}_{\text{VSMOW}}$  values were calculated using IMF-corrected isotopic ratios. The internal precision on each standard analysis based on counting statistics was typically between  $\pm 0.25\text{\textperthousand}$  and  $\pm 0.8\text{\textperthousand}$  ( $2\sigma$ ) and

spot-to-spot reproducibility was typically better than 0.6‰. Oxygen isotope data is presented in Table SD2 and Figures 3 and SD3.

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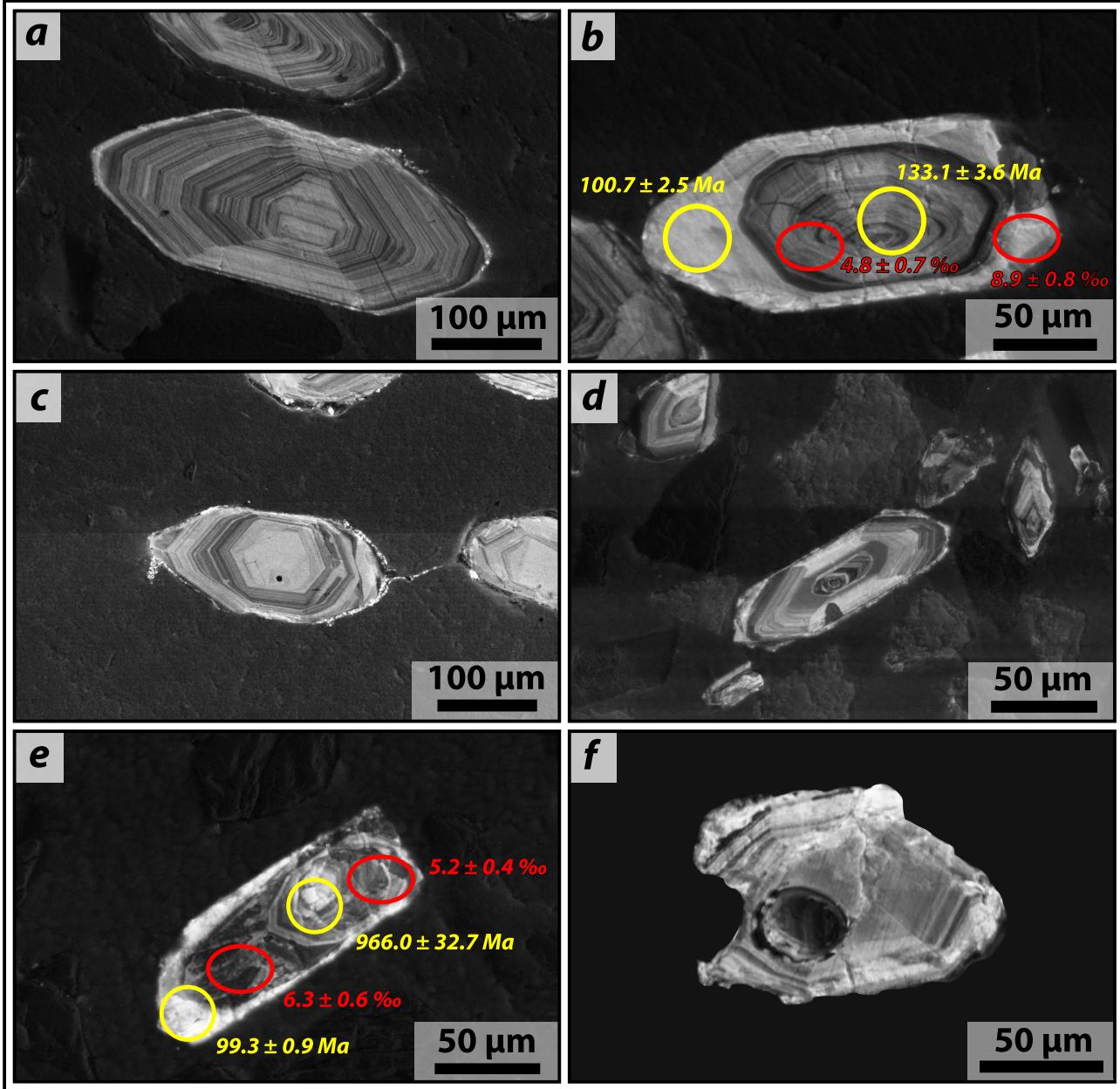
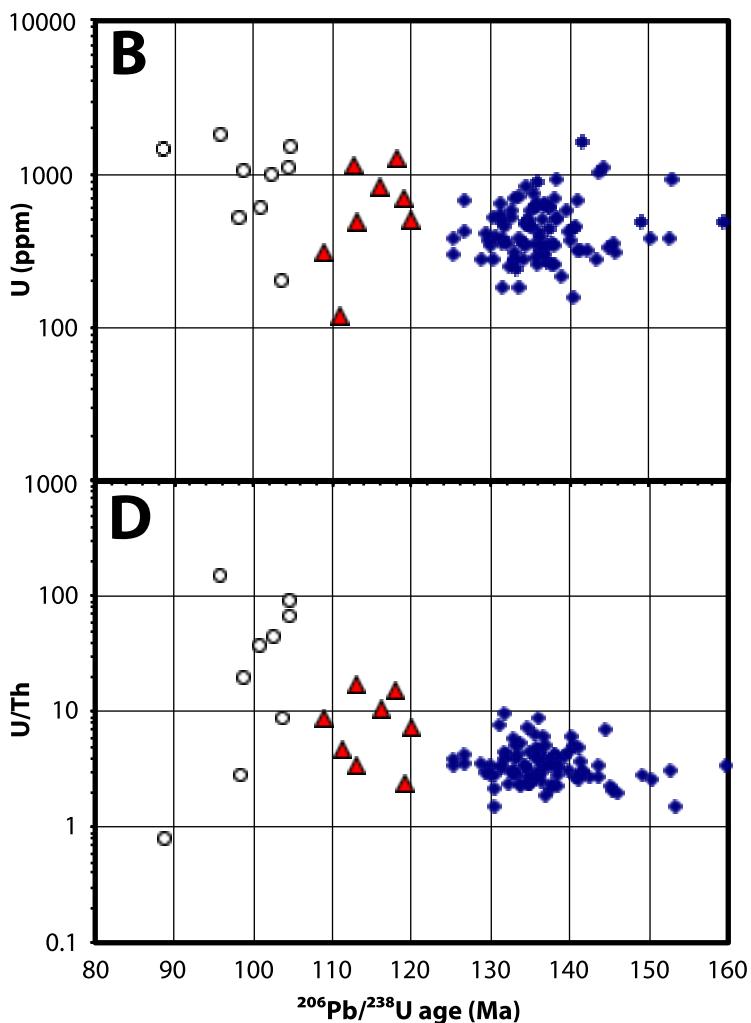
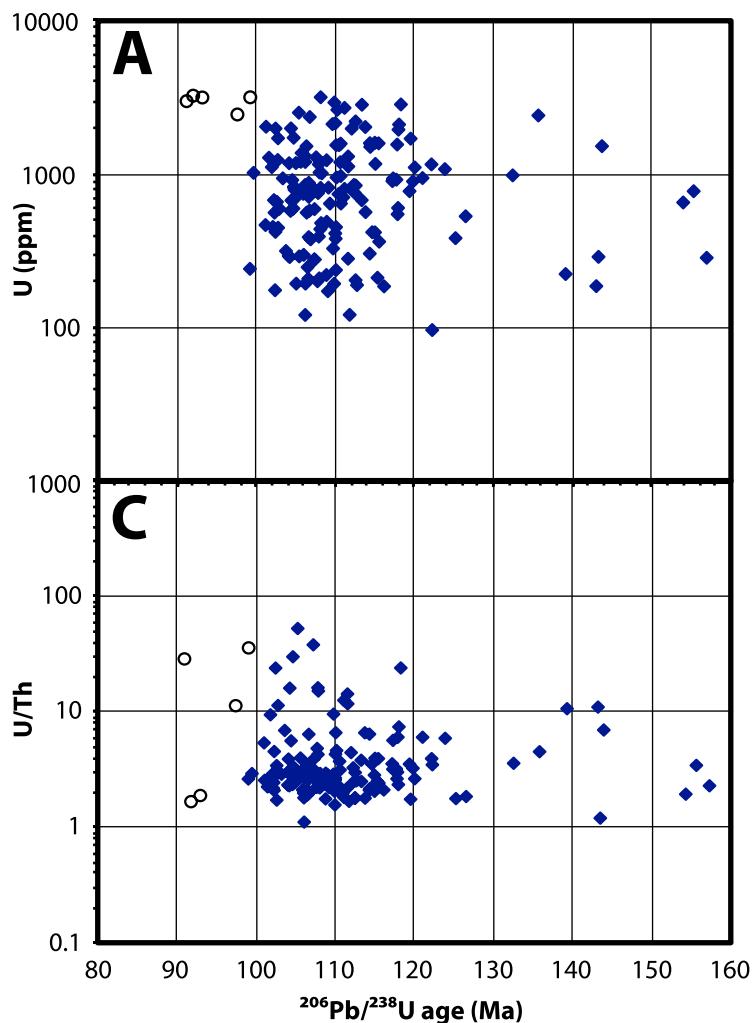


Figure DR1. Supplementary representative cathodoluminescence images of selected zircon grains from upper plate samples 7 (originally 08SE46; a and b), 6 (originally 08SE451; c), and 8 (originally 08SE675; d), and schist samples 3 (originally 06SE23; e) and 5 (originally 07SE34; f). Analysis 06SE23-3b ( $966.0 \pm 32.7 \text{ Ma}$ ) was included (e) despite being  $> 30\%$  discordant to illustrate a core-rim age difference. Yellow circles indicate laser ablation multicollector inductively coupled plasma mass spectrometry (LA-MC-ICP-MS) analyzed areas labeled with  $^{206}\text{Pb}/^{238}\text{U}$  ages. Red ellipses indicate Secondary Ion Mass Spectrometry (SIMS) analyzed areas labeled with  $\delta^{18}\text{O}$ . All uncertainties are  $2\sigma$ .



● igneous domain      ○ metamorphic domain      ▲ overlap

Figure DR2. Plots of zircon trace element chemistry versus U-Pb age for all available data from the San Ermidio Schist (A and C) and upper plate batholithic rocks (B and D). Data from Grove et al. (2003) (original sample 9957), Jacobson et al. (2011) (original samples SE1b and SE03), Chapman et al. (2012) (original samples 08SE46, 08SE451, and 08SE675), and this study (original samples 06SE23, 07SE34, and 08SE467).

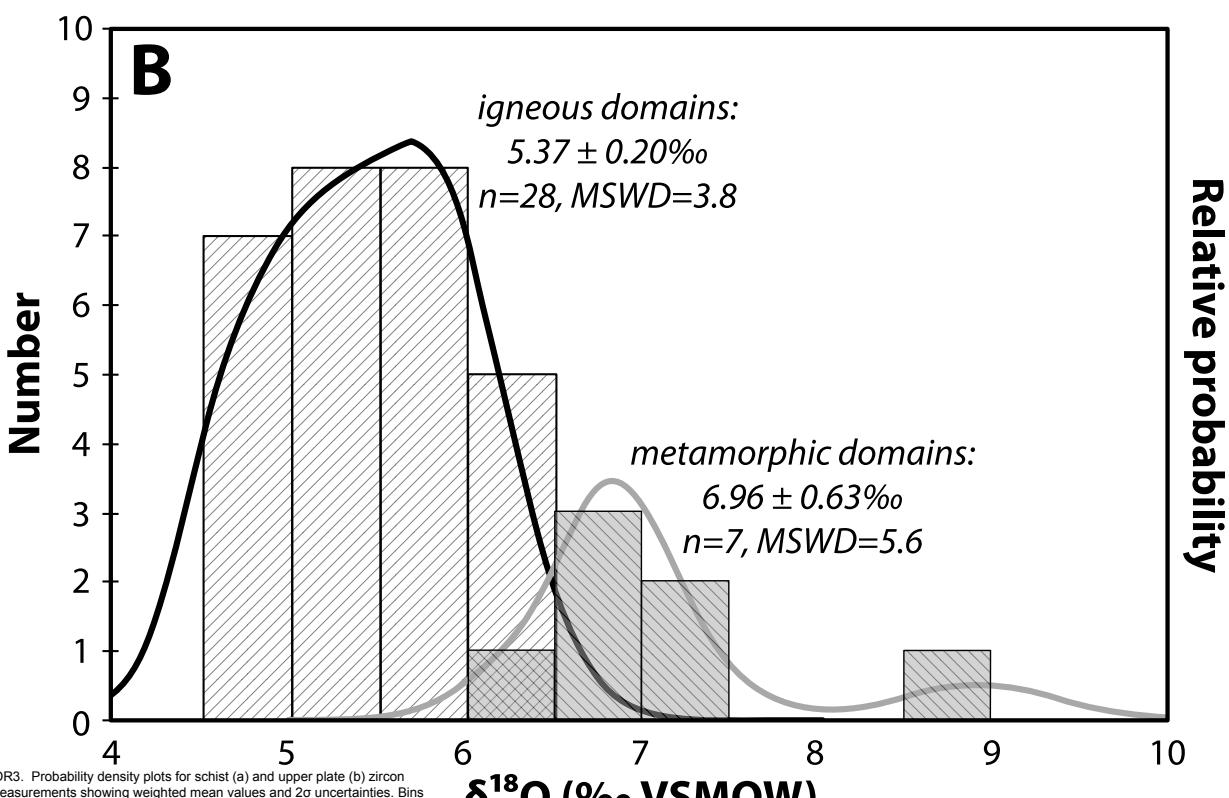
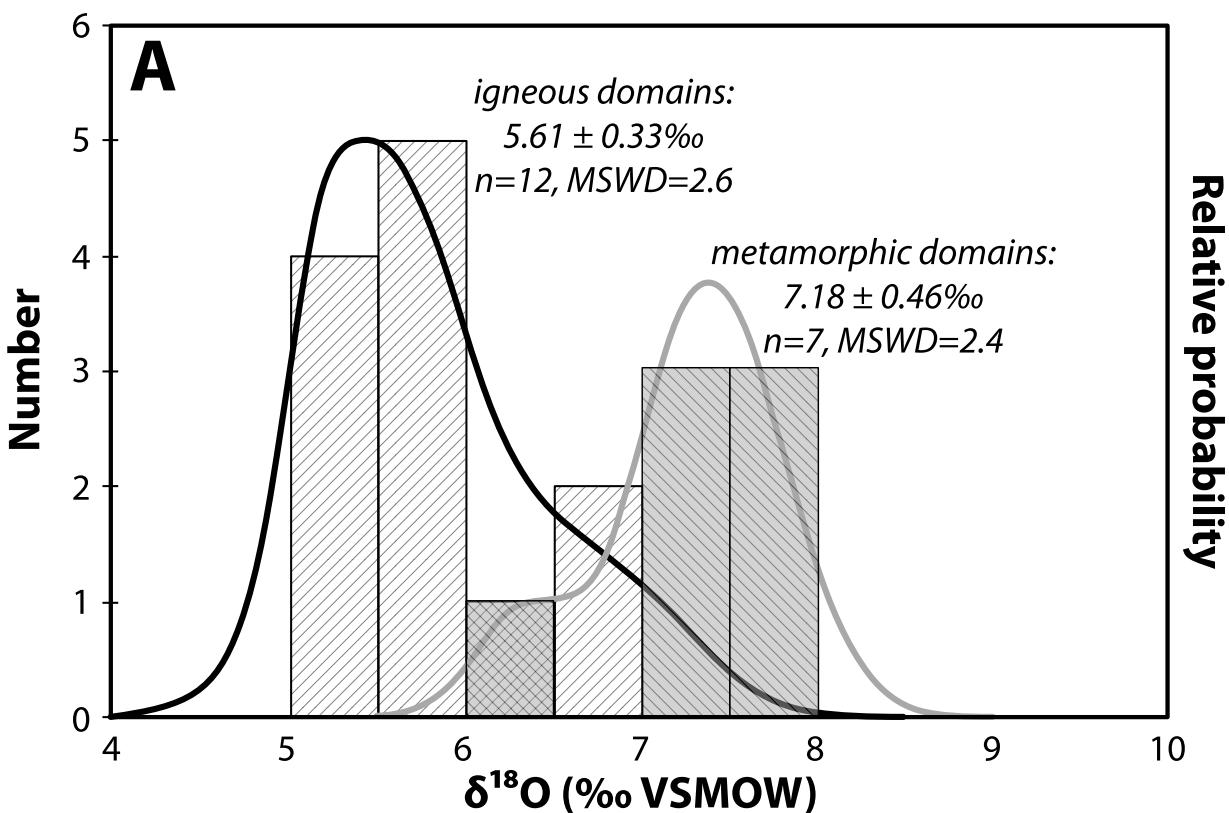


Figure DR3. Probability density plots for schist (a) and upper plate (b) zircon  $\delta^{18}\text{O}$  measurements showing weighted mean values and  $2\sigma$  uncertainties. Bins of analyses from igneous and metamorphic domains are white and gray, respectively.

Probability density curves for igneous and metamorphic domains are black and gray, respectively.





TABLE DR2. SIMS zircon oxygen isotope data.

**3 (originally 06SE23) San Emigdio Schist - San Emigdio Mountains**

Analysis	CL notes <sup>§</sup>	Measured		$\delta^{18}\text{O}_{\text{VSMOW}} (\text{\textperthousand})$	$\pm (\text{\textperthousand})^*$	U-Pb age (Ma) <sup>†</sup>	$\pm (\text{Ma})^*$
		$^{18}\text{O}/^{16}\text{O} (\times 10^{-3})$	$\pm (\times 10^{-3})^*$				
06SE23_det1IG.asc	IG	2.0099	0.000428	5.2	0.4	966.0	32.7
06SE23_det1b.asc	IG	2.0112	0.000817	5.5	0.5	105.0	1.8
06SE23_det1MM2.asc	RXL	2.0136	0.000900	6.3	0.5	99.3	0.9
06SE23_2.asc	RXL	2.0162	0.001026	7.3	0.6	nd	nd

**1 (originally 08SE467) San Emigdio Schist - San Emigdio Mountains**

Analysis	CL notes <sup>§</sup>	Measured		$\delta^{18}\text{O}_{\text{VSMOW}} (\text{\textperthousand})$	$\pm (\text{\textperthousand})^*$	U-Pb age (Ma) <sup>†</sup>	$\pm (\text{Ma})^*$
		$^{18}\text{O}/^{16}\text{O} (\times 10^{-3})$	$\pm (\times 10^{-3})^*$				
08SE467_1.asc	IG	2.0230	0.000785	6.3	0.8	nd	nd
08SE467_2.asc	IG	2.0214	0.000750	5.6	0.8	157.1	1.9
08SE467_8.asc	IG	2.0237	0.001382	6.8	1.0	nd	nd
08SE467_9.asc	IG	2.0236	0.000921	6.9	0.8	nd	nd
08SE467_9.asc	IG	2.0212	0.000461	5.8	0.7	119.4	2.7
08SE467_9b.asc	IG	2.0213	0.000576	5.9	0.7	119.4	2.7
08SE467_12.asc	IG	2.0209	0.000590	5.8	0.7	118.0	2.6
08SE467_12b.asc	IG	2.0195	0.000583	5.2	0.7	118.0	2.6
08SE467_14.asc	IG	2.0194	0.000920	5.3	0.8	108.2	6.7
08SE467_14.asc	IG	2.0194	0.000836	5.4	0.7	123.9	1.7
08SE467_14MM.asc	RXL	2.0238	0.000738	7.7	0.7	91.2	0.7
08SE467_14MM2.asc	RXL	2.0233	0.001028	7.5	0.8	93.2	5.8
08SE467_15.asc	RXL	2.0222	0.000635	7.1	0.6	91.2	2.7
08SE467_23.asc	OL	2.0212	0.000900	6.7	0.7	105.4	1.2
08SE467_25.asc	OL	2.0204	0.001128	6.4	0.8	102.6	1.2
08SE467_31.asc	RXL	2.0226	0.001040	7.6	0.7	97.6	1.2
08SE467_32.asc	RXL	2.0218	0.000910	7.3	0.7	92.0	4.9

**7 (originally 08SE46) San Emigdio tonalite - San Emigdio Mountains**

Analysis	CL notes <sup>§</sup>	Measured		$\delta^{18}\text{O}_{\text{VSMOW}} (\text{\textperthousand})$	$\pm (\text{\textperthousand})^*$	U-Pb age (Ma) <sup>†</sup>	$\pm (\text{Ma})^*$
		$^{18}\text{O}/^{16}\text{O} (\times 10^{-3})$	$\pm (\times 10^{-3})^*$				
08SE46_1.asc	IG	2.0203	0.001275	4.8	0.7	133.1	3.6
08SE46_1b.asc	RXL	2.0287	0.001408	9.0	0.8	100.7	2.5
08SE46_1c.asc	RXL	2.0252	0.001473	7.2	0.8	102.3	3.7
08SE46_1d.asc	IG	2.0232	0.000995	6.2	0.6	144.8	7.3
08SE46_2.asc	IG	2.0230	0.001921	6.1	1.0	153.0	4.1
08SE46_3.asc	IG	2.0209	0.000775	5.1	0.5	130.3	5.3
08SE46_3b.asc	RXL>OL	2.0251	0.000887	7.1	0.6	104.5	1.8
08SE46_4.asc	IG	2.0204	0.000750	4.8	0.5	143.3	3.4
08SE46_4b.asc	RXL>OL	2.0246	0.000731	6.9	0.5	104.5	2.8
08SE46_5.asc	IG	2.0198	0.000978	4.5	0.6	145.8	3.9
08SE46_5b.asc	RXL	2.0233	0.000815	6.3	0.6	103.5	4.6
08SE46_19.asc	IG	2.0155	0.000705	6.0	0.5	159.4	7.2
08SE46_20.asc	IG	2.0160	0.000546	6.3	0.5	145.3	3.8
08SE46_21.asc	IG	2.0146	0.000676	5.6	0.5	150.2	2.8
08SE46_21b.asc	IG	2.0151	0.000778	5.9	0.6	132.8	2.6
08SE46_22.asc	IG	2.0138	0.000765	5.2	0.6	133.6	4.9
08SE46_23.asc	IG	2.0143	0.000636	5.4	0.5	134.5	2.8

Notes:

§CL: Cathodoluminescence; IG: Oscillatory zoned zircon, interpreted

to be detrital igneous grain; RXL: Convolute zoning domain, interpreted to be recrystallized zircon; OL: beam overlapped igneous and recrystallized domains.

\*Uncertainties given at  $2\sigma$ .†Corresponding  $206\text{Pb}/238\text{U}$  ages determined within the same crystal growth domain; nd, not determined.

TABLE DR2. SIMS zircon oxygen isotope data.

08SE46_24.asc	IG	2.0150	0.000770	5.8	0.6	140.8	3.7
08SE46_25.asc	IG	2.0132	0.001239	4.9	0.7	136.1	5.6
08SE46_26.asc	IG	2.0148	0.000851	5.7	0.6	135.9	3.2
08SE46_27.asc	IG	2.0151	0.000616	5.8	0.5	131.1	1.4
08SE46_28.asc	IG	2.0135	0.000469	5.0	0.5	134.1	4.7
08SE46_29.asc	IG	2.0148	0.000758	5.7	0.6	131.5	3.2
08SE46_30.asc	IG	2.0145	0.000877	5.5	0.6	136.8	2.3
08SE46_31.asc	IG	2.0144	0.000690	5.5	0.5	137.7	3.8

**8 (originally 08SE675) San Emigdio tonalite - San Emigdio Mountains**

Analysis	CL notes <sup>§</sup>	measured		$\delta^{18}\text{O}_{\text{VSMOW}} (\text{\textperthousand})$	$\pm (\text{\textperthousand})^*$	U-Pb age (Ma) <sup>†</sup>	$\pm (\text{Ma})^*$
		$^{18}\text{O}/^{16}\text{O} (\times 10^{-3})$	$\pm (\times 10^{-3})^*$				
08SE675_det1a.asc	IG	2.0159	0.000424	5.8	0.5	133.6	3.4
08SE675_det1b.asc	IG	2.0161	0.000750	6.1	0.6	133.6	3.4
08SE675_det2a.asc	IG	2.0142	0.000636	5.2	0.5	135.5	4.3
08SE675_det2b.asc	RXL	2.0171	0.000713	6.8	0.5	98.2	1.7
08SE675_det3a.asc	IG	2.0128	0.000641	4.7	0.5	132.6	4.6
08SE675_det3b@1.asc	OL	2.0146	0.000742	5.6	0.5	116.1	9.1
08SE675_det4a.asc	IG	2.0123	0.000715	5.1	0.5	134.4	4.2
08SE675_det4c.asc	RXL	2.0135	0.000500	6.7	0.5	95.8	6.3
08SE675_5.asc	IG	2.0168	0.000494	4.7	0.5	140.4	4.2
08SE675_6.asc	IG	2.0151	0.000667	4.5	0.4	137.3	7.4
08SE675_7.asc	IG	2.0148	0.000507	5.3	0.5	144.3	3.8

Notes:

§CL: Cathodoluminescence; IG: Oscillatory zoned zircon, interpreted

to be detrital igneous grain; RXL: Convolute zoning domain, interpreted to be recrystallized zircon; OL: beam overlapped igneous and recrystallized domains.

\*Uncertainties given at  $2\sigma$ .

†Corresponding 206Pb/238U ages determined within the same crystal growth domain; nd, not determined.