METHODS

Zircon for U–Pb and U–Th geochronology was separated from ~0.2–2 kg crushed and sized samples using heavy liquids, mounted on glass and cast in 25 mm diameter epoxy plugs along with R33 standard zircon grains, ground to expose grains in cross section, polished, and imaged with reflected light and SEM cathodoluminescence (CL). Inclusions exposed on polished surfaces of zircon crystals were located with SEM backscattered-electron imaging and identified by their x-ray energy spectra. Spots were selected for analysis with the Sensitive High-Resolution Ion Microprobe with Reverse Geometry (SHRIMP-RG) at the Stanford–USGS laboratory at Stanford University on the basis of internal zoning revealed in CL. During the mass scans for U–Pb geochronology concentrations of U, Th, Hf, La, Ce, Nd, Sm, Eu, Gd, Dy, Er, and Yb also were measured.

Prior to analysis with the Sensitive High-Resolution Ion Microprobe with Reverse Geometry (SHRIMP-RG) at the Stanford–USGS laboratory at Stanford University, the epoxy plug was washed with saturated EDTA and dilute HCl solutions, rinsed with distilled water, dried in a vacuum oven, and coated with 100 nm of gold. A primary beam of O_2^- with intensities of ~6 nA for U–Pb analyses and ~10-15 nA for ²³⁸U–²³⁰Th analyses was focused into a 20–30 micrometer spot to generate secondary ions. The primary beam was rastered for 90–120 seconds over the analysis area before data were collected. For U–Pb geochronology, counts for secondary ions were collected sequentially with a single ETP electron multiplier on the following peaks: ⁹⁰Zr₂¹⁶O⁺ (2 s), ²⁰⁴Pb⁺ (8 s), background (8 s at 0.050 mass units above ²⁰⁴Pb⁺), ²⁰⁶Pb⁺ (30 s), ²⁰⁷Pb⁺ (20 s), ²⁰⁸Pb⁺ (8 s), ²³⁸U⁺ (6 s), ²³²Th¹⁶O⁺ (3 s), ²³⁸U¹⁶O⁺ (3 s), ¹³⁹La⁺ (2 s), ¹⁴⁰Ce⁺ (1 s each for rest),

 89 Y⁺, 146 Nd⁺, 147 Sm⁺, 153 Eu⁺, 155 Gd⁺, 163 Dy 16 O⁺, 166 Er 16 O⁺, 172 Yb 16 O⁺, 180 Hf 16 O⁺, and 238 U 16 O₂⁺; measurements were made at mass resolution of ~8000–9500 (10% peak height), which resolves all interfering atomic species. For U–Th analyses, secondary ion intensity was measured over seven scans for ${}^{90}\text{Zr}_{2}{}^{16}\text{O}^{+}$ (2 s), ${}^{180}\text{Hf}{}^{16}\text{O}^{+}$ (2 s), ${}^{238}\text{U}{}^{16}\text{O}^{+}$ (6 s), ${}^{230}\text{Th}{}^{16}\text{O}^{+}$ (60 s), background at 0.05 amu below 230 Th 16 O⁺ (60 s), 232 Th 16 O⁺ (3 s), 235 U 16 O⁺ (6 s), $^{238}U^{16}O^+$ (3 s) and $^{232}Th^{12}C^+$ at ~244 amu (10 s) to monitor for a potential isobaric interference at mass 246 due to partial overlap of the primary beam on surrounding epoxy (see Schmitt, 2006). Relative ionization between U and Th was constrained by repeated intra-session analysis of standards R33 (419 Ma, quartz diorite of Braintree complex, Vermont, Black et al., 2004) and homogenous in-house MAD-green zircon from Madagascar (Barth and Wooden, 2010), both of which are sufficiently old (ca. 419 and 555 Ma, respectively) to have ²³⁸U and ²³⁰Th in secular equilibrium. Concentration data for zircons were standardized against MAD-green zircon relative to intensities for ⁹⁰Zr₂¹⁶O⁺; U–Pb ages were calculated by calibration against zircon standard R33 which was analyzed repeatedly throughout the duration of the analytical sessions. Data reduction followed the methods described by Williams (1997) and Ireland and Williams (2003), and used the Squid 1 and Isoplot programs of Ken Ludwig (Ludwig 2001, 2003).

Whole rock major- and trace-element concentrations were determined by GeoAnalytical Laboratories (Washington State University) on powders of the three largest samples (19B, 24B, 24C); comparable data for sample 5i appear in Nye et al. (1994, table 4). Minerals and glass in the six samples were analyzed with a JEOL 8900 electron microprobe at the USGS in Menlo Park, California.

FIGURES

Figure DR1. Transmitted light images of polished thin sections of xenoliths from Redoubt Volcano. Images are approximately 2.4 x 3.6 cm. A: Crystal-rich andesite 19B. Plagioclase in gabbroic clusters in lower center region is An₈₃₋₉₂ zoned to rims as low as An₄₆; large sieved plagioclase is mainly An₄₈₋₅₉ with rims of similar composition; two plagioclase inclusions in hornblende oikocryst near upper center are An₈₈ and An₆₄ zoned to An₅₇ and An₅₁, respectively. Hornblende oikocryst is identical in composition to those in gabbro 22. Note textural heterogeneity of sample and doublet of olivine xenocrysts (Fo₈₁, 900 ppm Ni) with opaque mantle in right center. B: Gabbro 22. Plagioclase cores are An₇₉₋₈₉, rims are An₅₉₋₆₃, and inclusions in hornblende oikocrysts are An₅₆₋₈₅ zoned to An₅₅₋₇₀. C: Gabbro 24B. Typical plagioclase cores are An₇₅₋₇₉, rims An₅₅₋₅₈. D: Gabbro 21B. Typical plagioclase is An₉₄₋₉₅, with rims An₄₇₋₆₈. Olivine, enclosed in pyroxene, is Fo₇₅₋₇₉, ~300 ppm Ni. E: Gabbro 24C. Typical plagioclase cores are An₇₇₋₈₄, rims An₅₅₋₇₆.

Figure DR2. Concordia diagrams for SHRIMP-RG zircon U–Pb data. Uncertainties for error ellipses are ±2σ. Numbers by blue concordia curves are ages in Ma. Numbers adjacent to error ellipses identify specific spot analyses. A: Crystal-rich andesite 19B.
B: Gabbro 22. C: Gabbro 24B. D: Gabbro 21B. E: Gabbro 24C. F: Gabbro 5i.

Figure DR3. Whole rock compositions of xenoliths and crystal-rich andesite from Redoubt Volcano. A: Elemental abundances normalized to primitive mantle values (Sun and McDonough, 1989). B: Rare earth element abundances normalized to chondrite values (Sun and McDonough, 1989). Figure DR4. Cathodoluminescence images of zircons from xenoliths from Redoubt Volcano. Images of polished grain mounts were obtained with a JEOL scanning electron microscope. Sample number appears at top of image (e.g., 08RDCRB019B). All CL images are shown, including those for which no zircons were analyzed. Numbers by individual zircons identify analyses, e.g., 72.1 is zircon 72, analysis spot 1 in the sample. Numbers with uncertainties listed, e.g., 1861±8, are 238 U/ 206 Pb (<1000 Ma) or 207 Pb/ 206 Pb (≥1000 Ma) ages in Ma; uncertainties are ±1 σ (D indicates analysis was >10% discordant). For two grains (08RDCRB019B zircon 62 and 90CNR05i zircon 1), 238 U– 230 Th disequilibrium model ages are given; these are identified by units of ka. Uranium concentrations are given in separate CL images of low-U zircons.

Figure DR5. Concentrations of U and Hf (ppm) and Ce/Ce* ratios in zircon obtained during SHRIMP-RG U–Pb geochronology analyses. Symbols indicate age groups for nominal ages of analysis spots for all six samples plotted together. Note largely separate groups for 1800–1900 Ma and 280–350 Ma points. A: U vs. Hf. B: Ce/Ce* vs. Hf.

Figure DR6. Concentration ratios Yb/Gd vs. Th/U in zircon obtained during SHRIMP-RG U–Pb geochronology analyses. Symbols indicate age groups for nominal ages of analysis spots for all six samples plotted together. Note coherent trend in 1800–1900 Ma data and clustering of most 280–350 Ma points.

Figure DR7. Concentration ratio U/Ce vs. Th (ppm) in zircon obtained during SHRIMP-RG U–Pb geochronology analyses. Symbols indicate age groups for nominal ages of analysis spots for all six samples plotted together. Gray lines give reference U/Ce:Th ratios. Non-igneous zircon plots at U/Ce>2xTh. Note separation into two clusters for majority of 1800–1900 Ma and 280–350 Ma points. Spot analyzes with U/Ce >2 likely represent zircon that precipitated from aqueous fluid.

Figure DR8. Concentrations of REE in zircon obtained during SHRIMP-RG U–Pb geochronology analyses normalized to those in chondritic meteorites (Korotev, 1996).
Values for Pr calculated on the basis of those for La and Nd. Symbols indicate age groups for nominal ages of analysis spots. A: Crystal-rich andesite 19B. B: Gabbro 22.
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TABLES

Table DR1. Sample numbers and localities for xenoliths from Redoubt Volcano.

Table DR2. Chemical analyses of xenoliths from Redoubt Volcano.

Table DR3. Electron microprobe analyses of minerals and glass in xenoliths fromRedoubt Volcano.

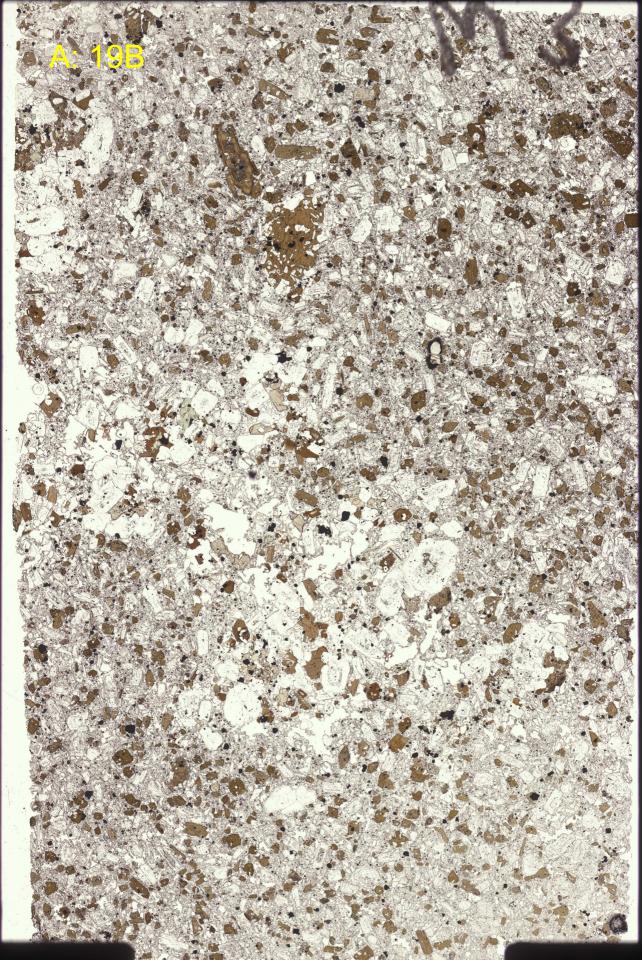
Table DR4. SHRIMP-RG U–Pb geochronology data and calculated ages for zircon from xenoliths from Redoubt Volcano.

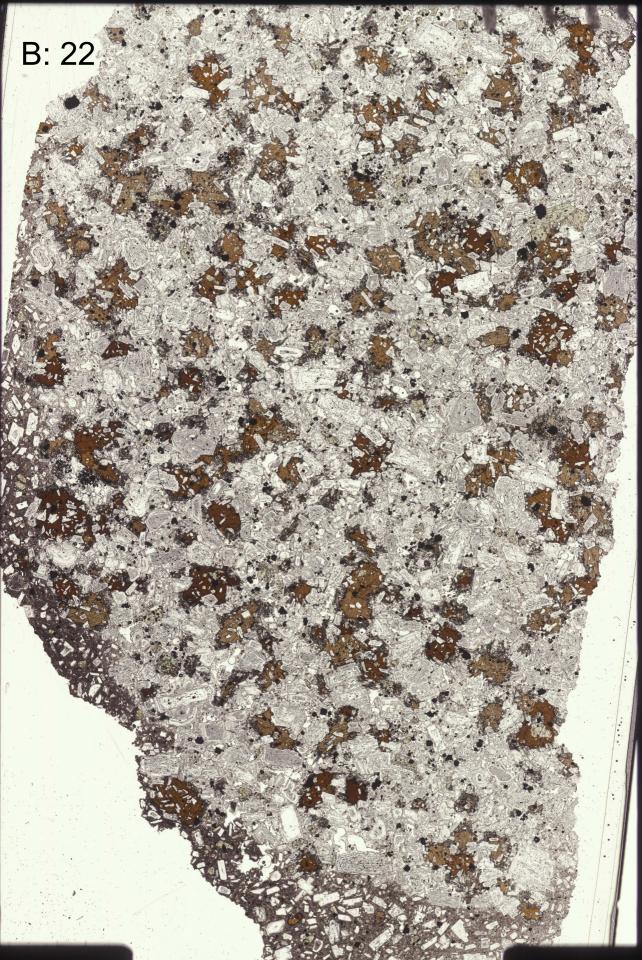
Table DR5. SHRIMP-RG U–Th geochronology data and calculated ages for zircon from xenoliths from Redoubt Volcano.

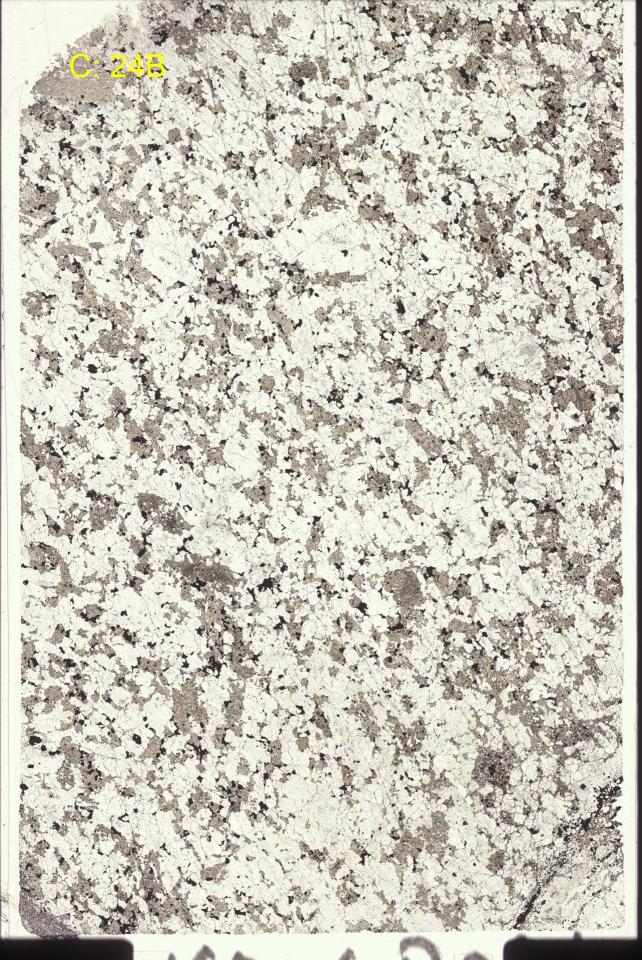
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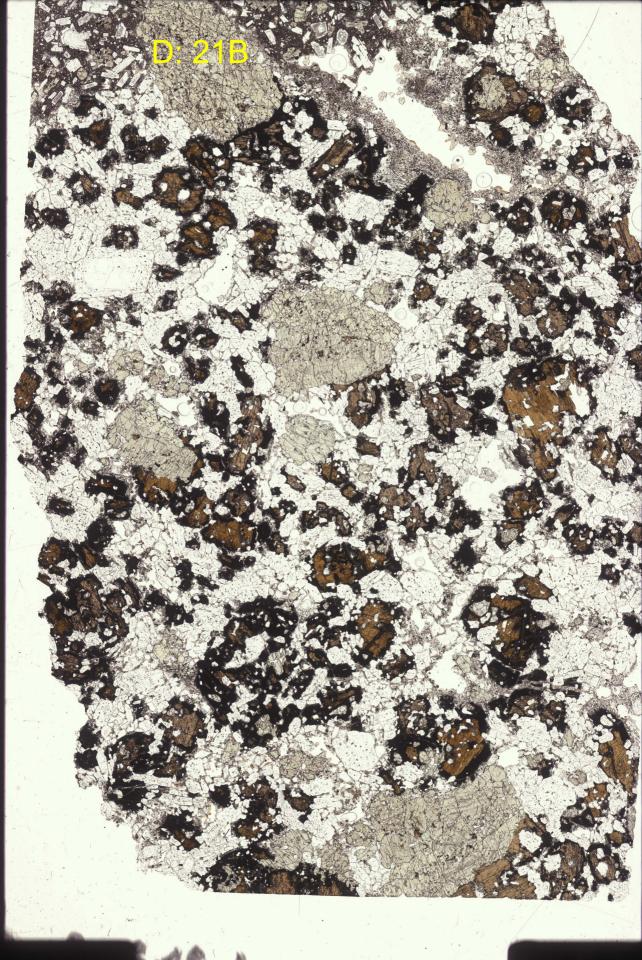
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Wilson, C.J.N., Charlier, B.L.A., Rowland, J.V., and Browne, P.R.L., 2010, U-Pb dating of zircon in subsurface, hydrothermally altered pyroclastic deposits and implications for subsidence in a magmatically active rift: Taupo Volcanic Zone, New Zealand: Journal of Volcanology and Geothermal research, v. 191, p. 69-78. Figure DR1. Transmitted light images of polished thin sections of xenoliths from Redoubt Volcano. Images are approximately 2.4 x 3.6 cm. A: Crystal-rich andesite 19B. Plagioclase in gabbroic clusters in lower center region is An₈₃₋₉₂ zoned to rims as low as An₄₆; large sieved plagioclase is mainly An₄₈₋₅₉ with rims of similar composition; two plagioclase inclusions in hornblende oikocryst near upper center are An₈₈ and An₆₄ zoned to An₅₇ and An₅₁, respectively. Hornblende oikocryst is identical in composition to those in gabbro 22. Note textural heterogeneity of sample and doublet of olivine xenocrysts (Fo₈₁, 900 ppm Ni) with opaque mantle in right center. B: Gabbro 22. Plagioclase cores are An₇₉₋₈₉, rims are An₅₉₋₆₃, and inclusions in hornblende oikocrysts are An₅₆₋₈₅ zoned to An₅₅₋₇₀. C: Gabbro 24B. Typical plagioclase cores are An₇₅₋₇₉, rims An₅₅₋₅₈. D: Gabbro 21B. Typical plagioclase is An₉₄₋₉₅, with rims An₄₇₋₆₈. Olivine, enclosed in pyroxene, is Fo₇₅₋₇₉, ~300 ppm Ni. E: Gabbro 24C. Typical plagioclase cores are An₇₇₋₈₄, rims An₅₅₋₇₆.













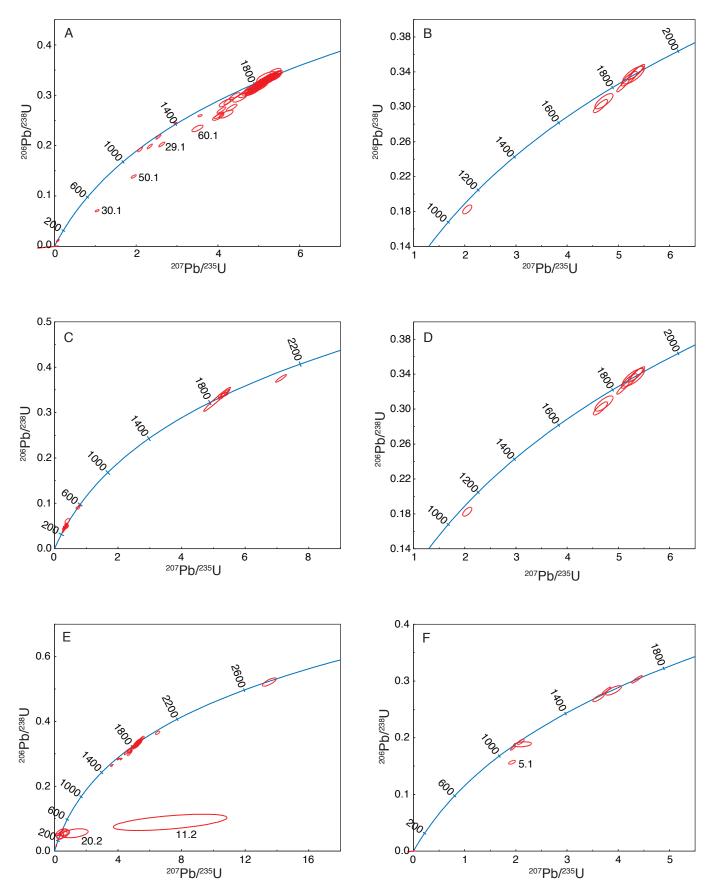


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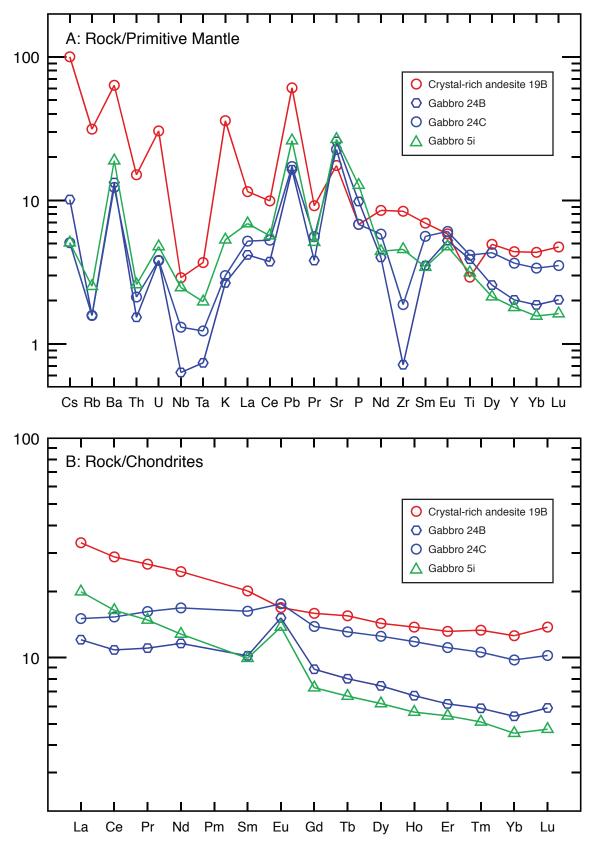
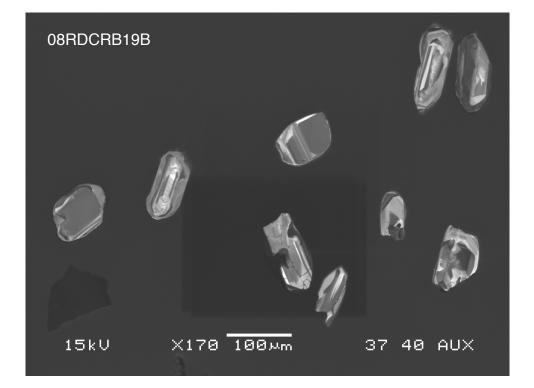
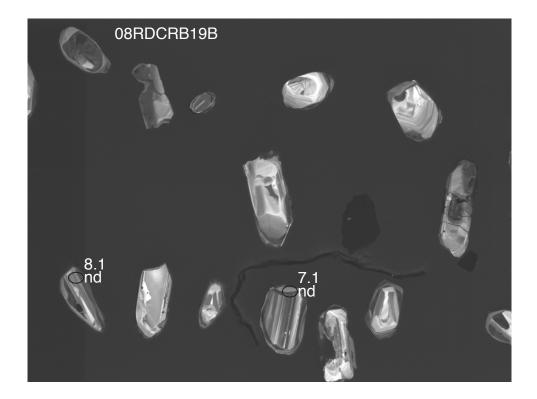
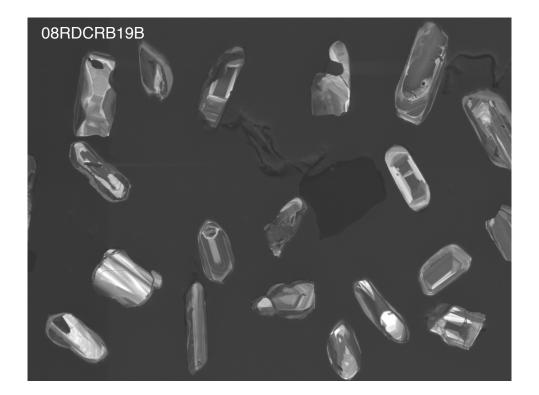


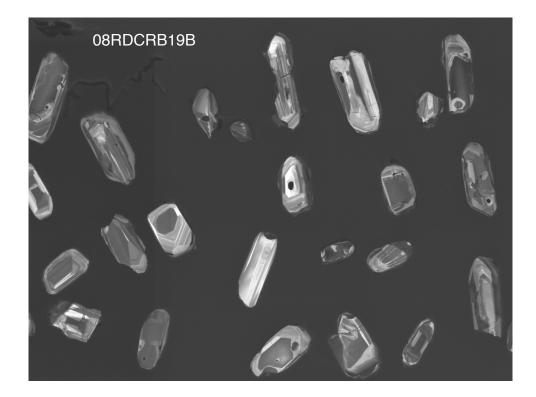
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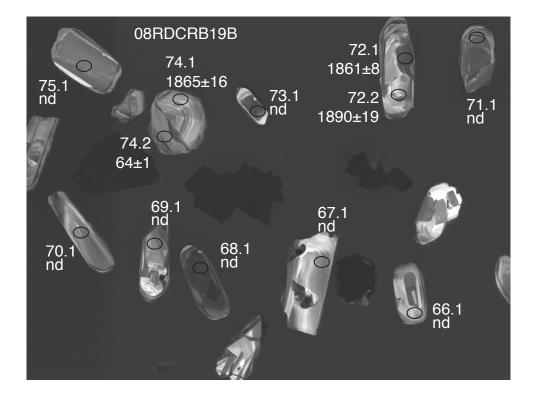


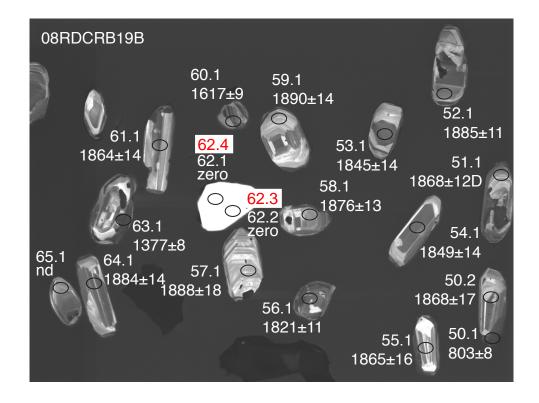


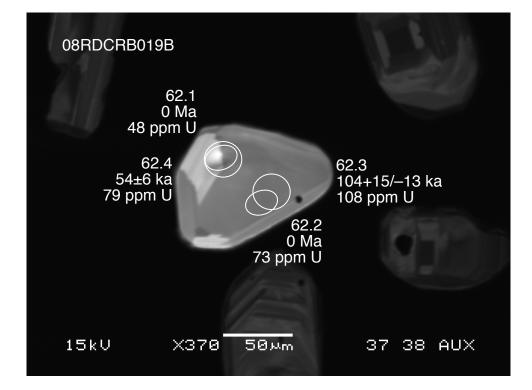


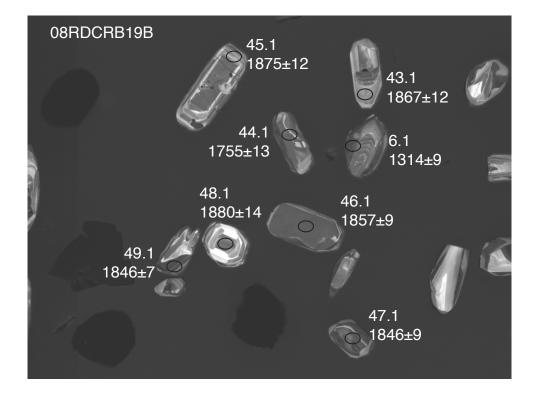




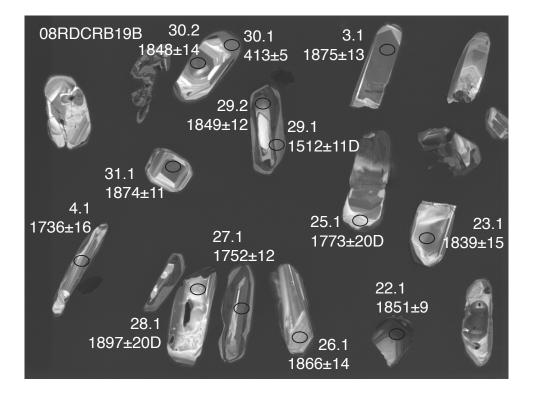


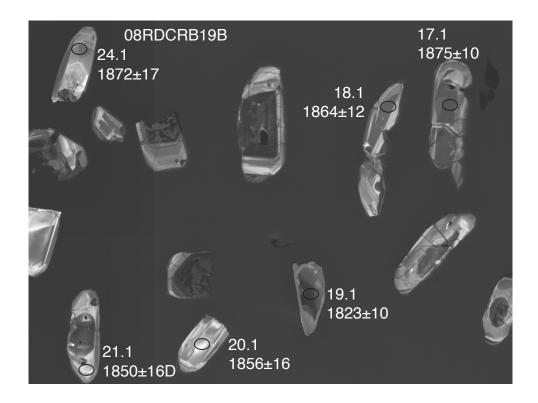


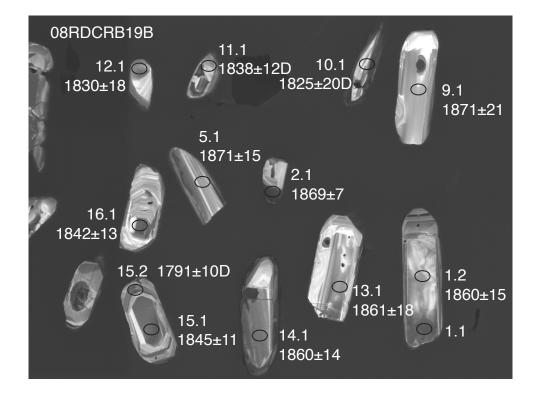










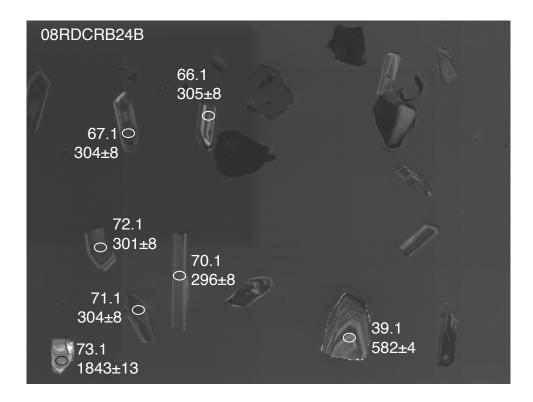


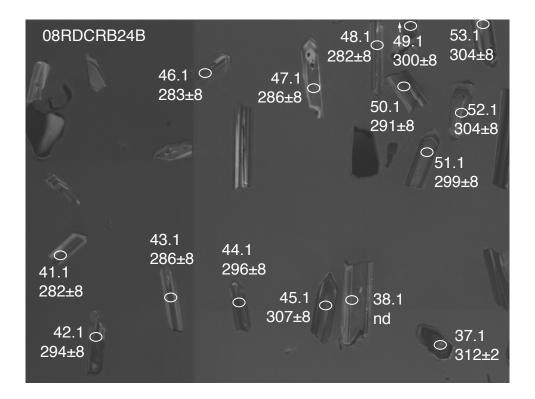


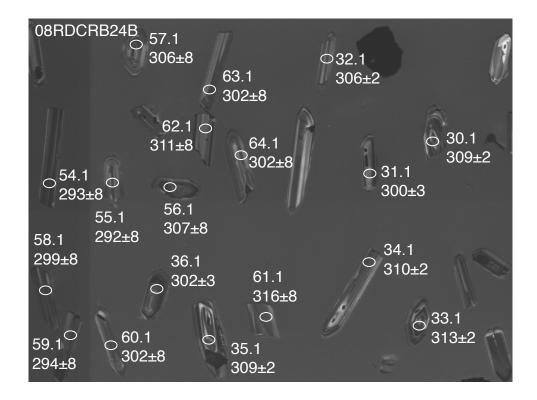


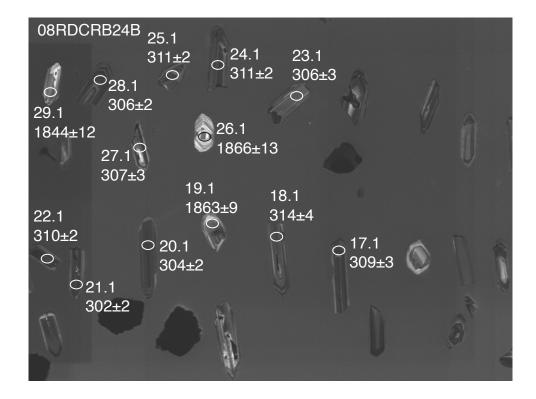


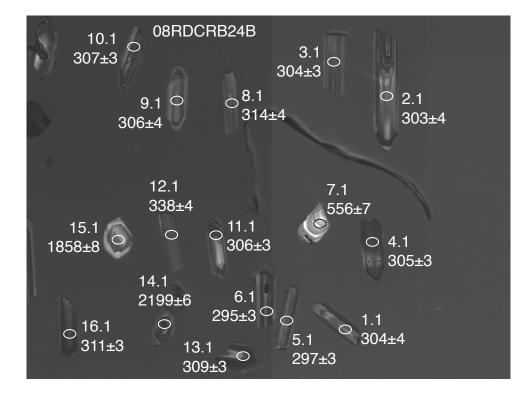






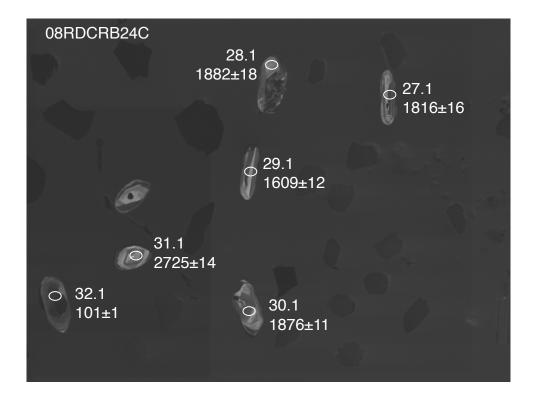


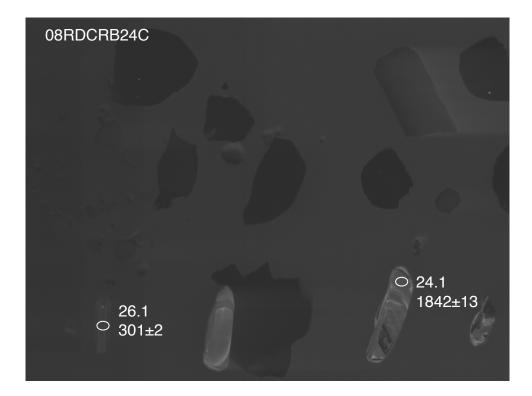




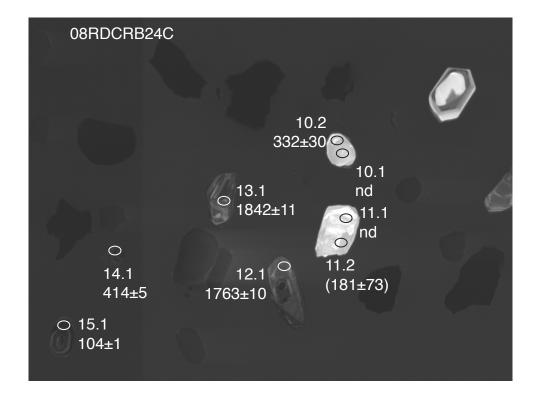






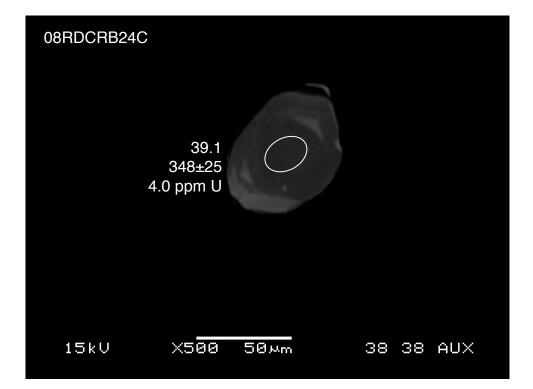


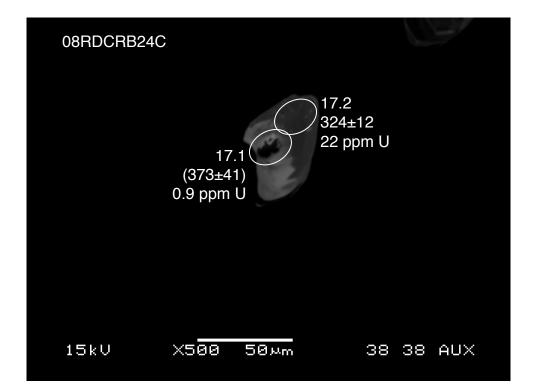


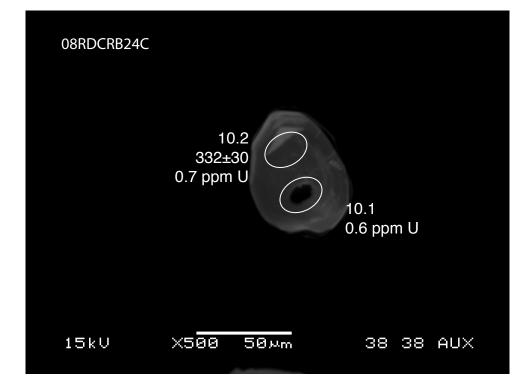


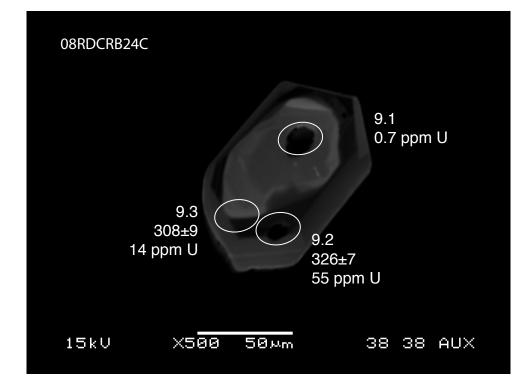


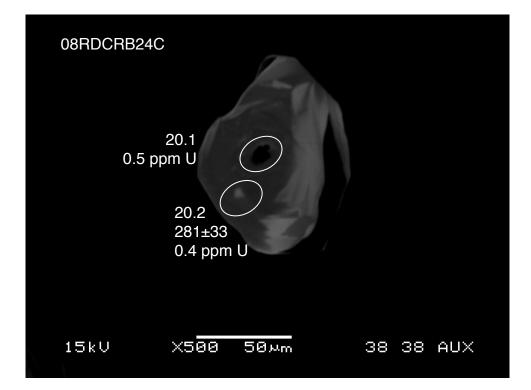


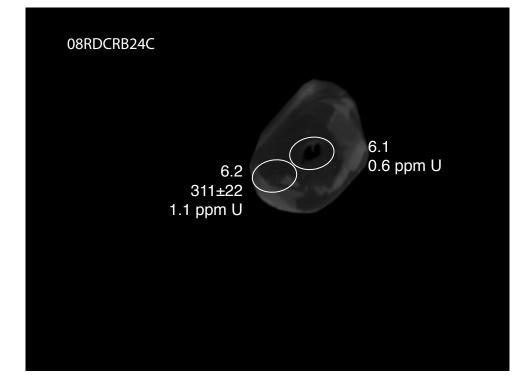


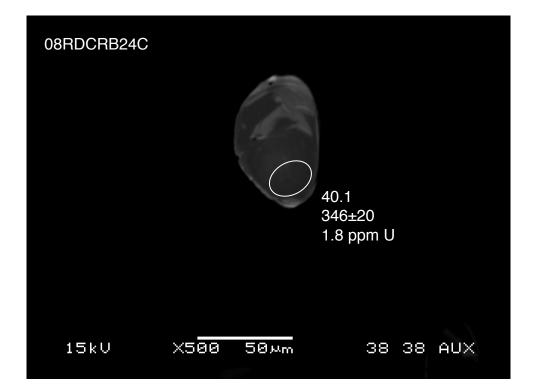


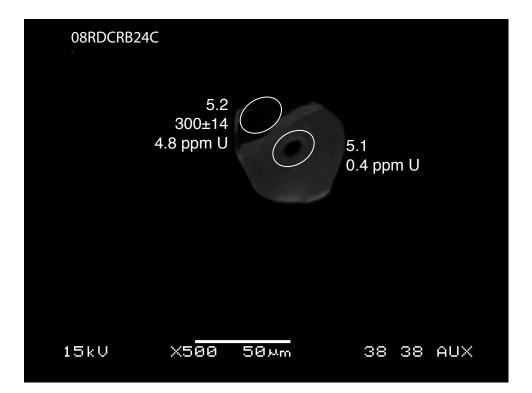




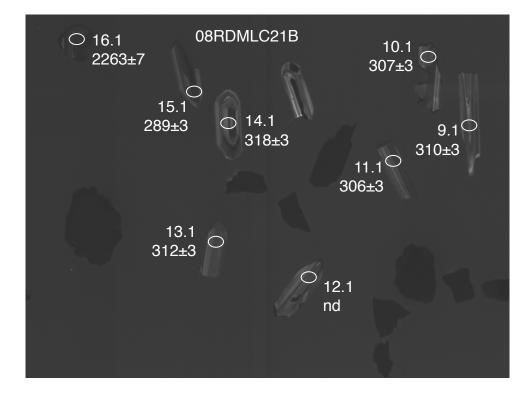


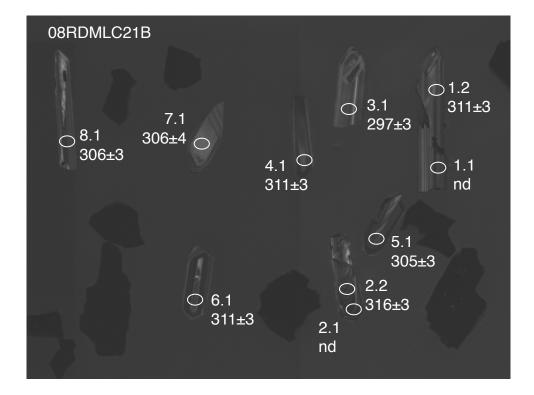


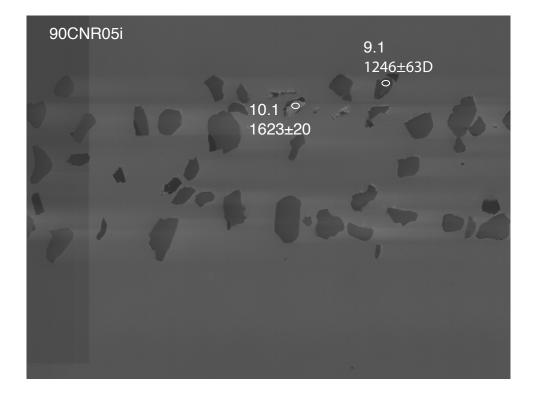


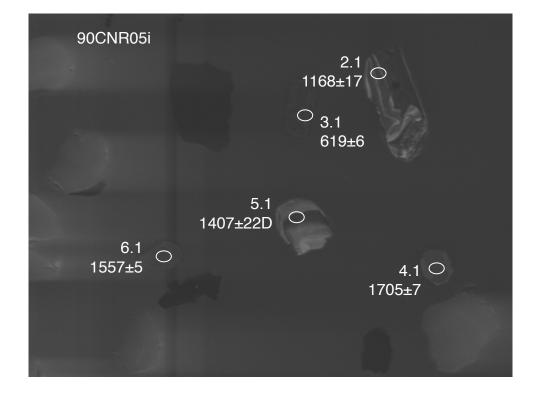


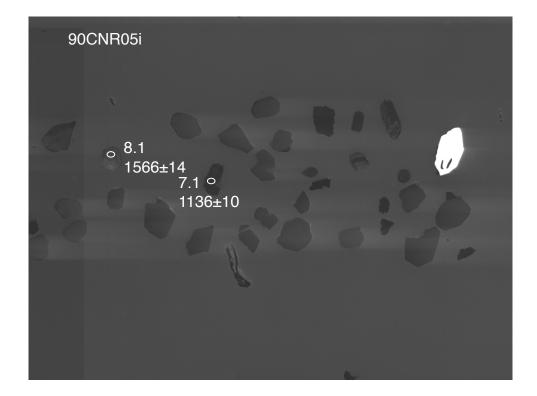


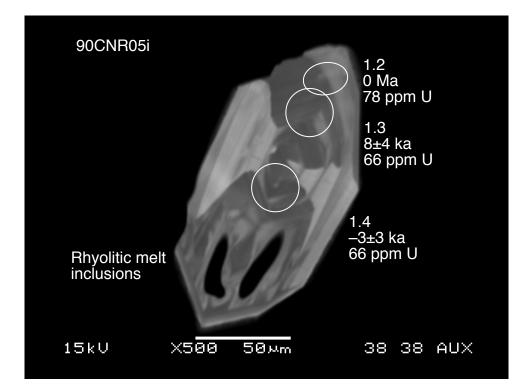












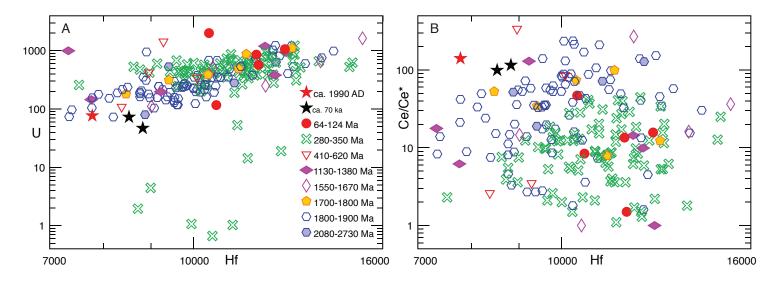


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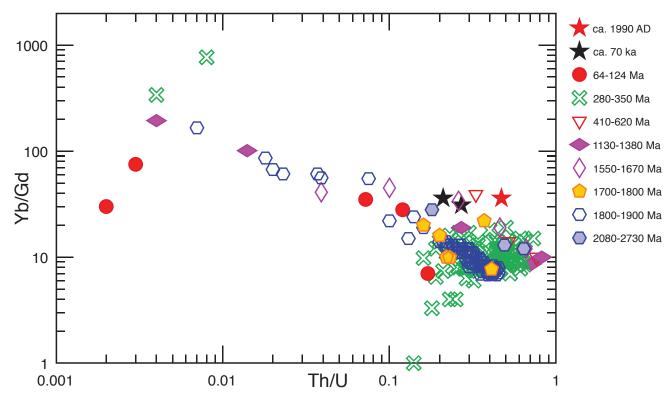


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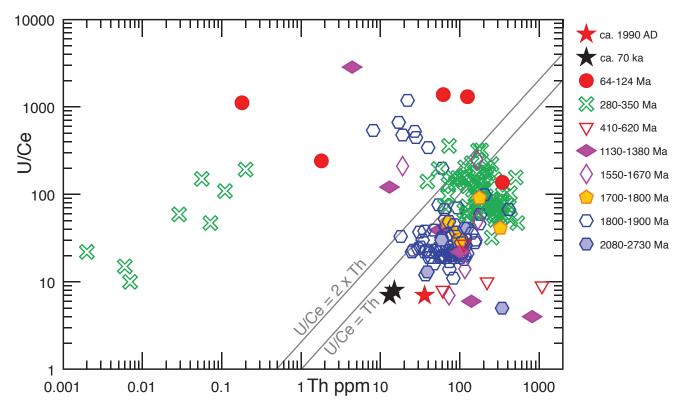


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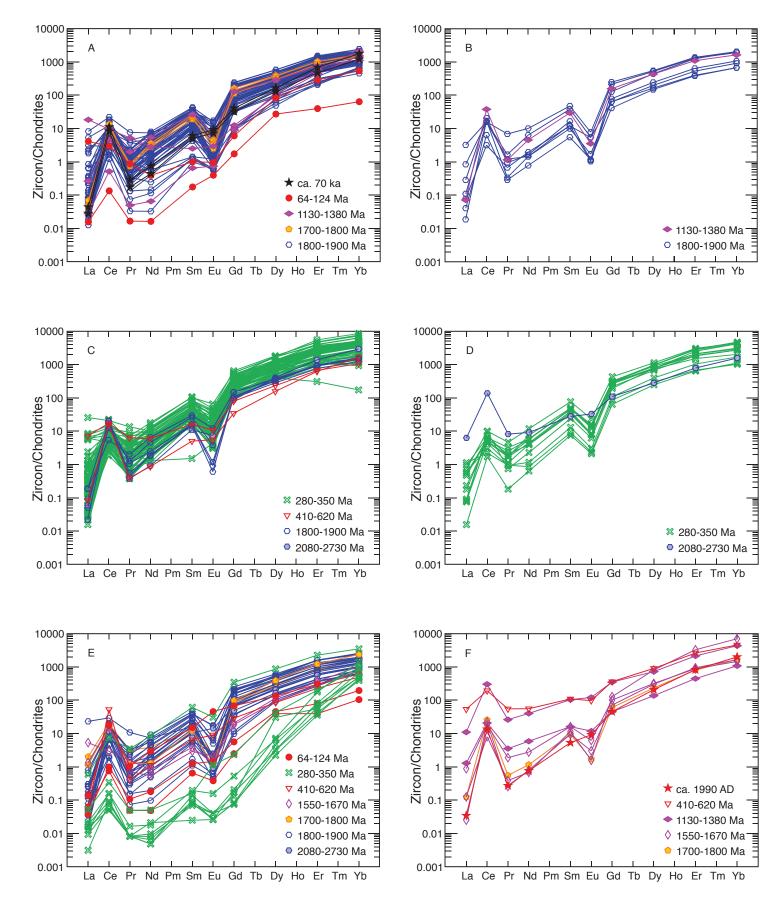


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