

## Supplementary Material

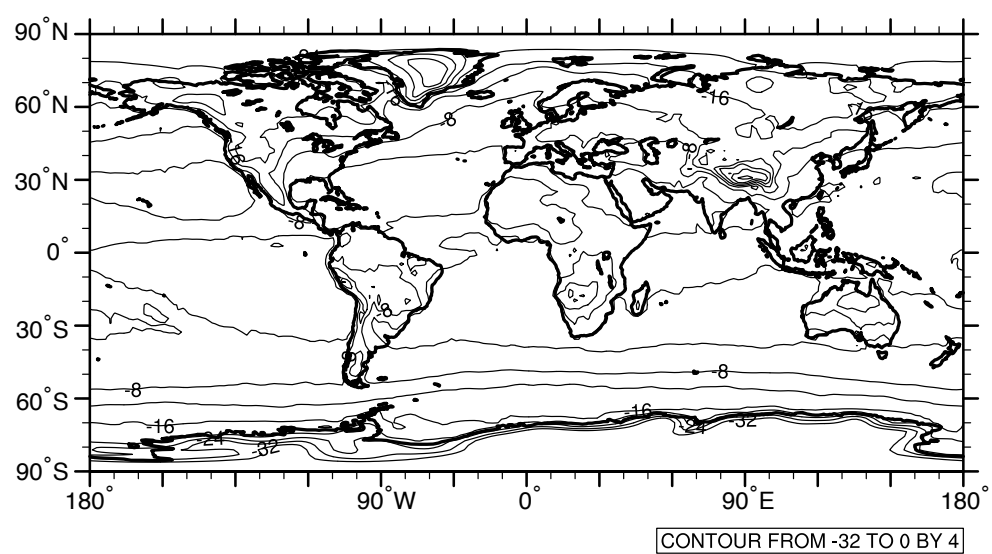
The ability of GENESIS to simulate the modern-day distribution of stable isotope compositions ( $\delta^{18}\text{O}$  and  $\delta\text{D}$ ) in meteoric water has been reported in previous studies (Mathieu et al., 2002; Zhou et al., 2008; Poulsen et al., 2010). These studies conclude that GENESIS successfully captures the large-scale climate-isotope relationships including elevation-isotope relationships (altitude effect). Further documentation of the stable isotope compositions simulated by GENESIS is provided here (Figs. DR1 and DR2).

In general, GENESIS does a reasonable job of simulating high-elevation precipitation  $\delta^{18}\text{O}$  (Figs. DR1 and DR2) over the North American Rockies, South American Andes, and Tibetan Plateau. As summarized in Blisniuk and Stern (2005), typical precipitation  $\delta^{18}\text{O}$  values over these regions are -12 to -16‰ (Sierra Nevada), -9 to -17‰ (central Andes), and -20‰ (southern Tibetan Plateau). Over these same regions, GENESIS simulates values of approximately -13‰ (Sierra Nevada), -14‰ (central Andes), and -22‰ (southern Tibetan Plateau) (Fig. DR2, dashed line). As discussed in the manuscript, the use of a slab-ocean model in GENESIS results in slightly lower precipitation  $\delta^{18}\text{O}$  over high elevations (Fig. DR2, solid line).

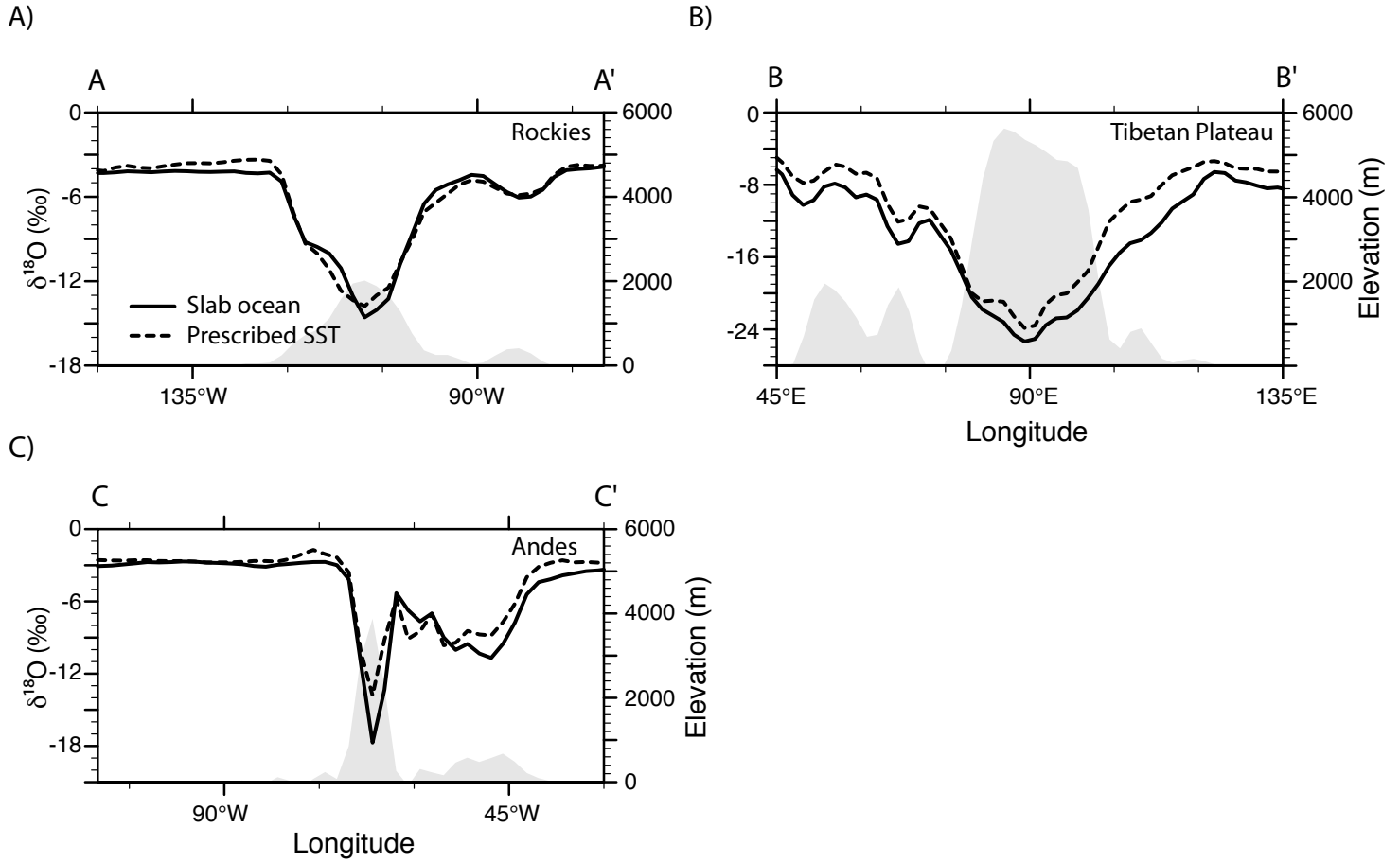
## References

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- Mathieu, R.D., Pollard, D., Cole, J.E., White, J.W.C, Webb, R.S., and Thompson, S.L., 2002, Simulation of stable water isotope variations by the GENESIS GCM for modern conditions: *Journal of Geophysical Research*, v. 107 (D4), p. 2-1 to 2-18.
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**Supplemental Fig. DR1.** Simulated annual amount-weighted precipitation  $\delta^{18}\text{O}_p$  (‰) for the modern climate.



**Supplemental Fig. DR2.** Annual amount-weighted  $\delta^{18}\text{O}_p$  (‰) predicted using prescribed sea-surface temperatures (dashed line) and a slab-ocean model (solid line).  $\delta^{18}\text{O}_p$  is shown for three transects across: (A) Rocky Mountains ( $\sim 35^\circ\text{N}$ ); (B) Tibetan Plateau ( $\sim 31^\circ\text{N}$ ); and (C) Andean Plateau ( $\sim 23^\circ\text{S}$ ) as in Fig. 1. The gray shading indicates the elevation. Note that the use of the slab-ocean model leads to lower  $\delta^{18}\text{O}_p$  over the major orogens due to stronger subsidence.