

Data Repository: Uba, Strecker, and Schmitt**U-Pb radiometric dating**

The age controls for the stable isotopic records presented here are from both published sources and U-Pb geochronology performed in this study. We analyzed six zircon separates by U-Pb geochronology with secondary ionization mass spectrometry (SIMS, ion microprobe) using the Cameca ims 1270 at UCLA. For each of the samples, the volcanic ashes were crushed, sieved and the zircons were separated using heavy liquid procedures and magnetic separation techniques. About 10 - 30 zircons per sample were then picked by hand and placed in epoxy mounts, which were polished and gold-coated for analysis. A ~15 nA O⁻ primary beam with 22.5 keV total impact energy was focused to a ~25 μm diameter spot. Secondary ions were extracted at 10 kV with an energy band pass of 50 eV. O₂ pressure in the sample chamber was adjusted to ~0.002 Pa, which increased Pb sensitivity by ~50 %. In three analytical sessions the ²⁰⁶Pb/²³⁸U age reproducibility of 33, 19, and 24 analyses (July 4th 2006, August 20th 2006, and April 6th 2007, respectively) of reference zircon AS-3 was between 2 and 3 % (1 standard deviation). All U-Pb ages were corrected for initial ²³⁰Th deficit using measured zircon U/Th and whole-rock U/Th = 2.3 (Central Andean ignimbrite average; Schmitt et al., 2002). This correction typically adds ~0.1 Ma to the equilibrium ²⁰⁶Pb/²³⁸U ages. Ion microprobe results for samples are presented in Data Repository Table 1 and Tera-Wasserburg isochron plots are shown in Data Repository Fig. 1.

Due to extremely slow diffusion of U and Pb in zircon, crystallization ages are essentially unaffected by pre-eruptive crystal storage even at elevated temperatures. This is in contrast to K-Ar (Ar-Ar) or zircon fission track dates, in which any memory of pre-

eruptive crystallization is erased. There are also numerous cases where U-Pb zircon ages of young volcanic rocks significantly predate the eruption due to magmatic assimilation of older wall-rocks or mechanical entrainment of zircons from country rock fragments or sediment during eruption and/or deposition, collectively termed “detrital”. Detrital contamination typically results in complex zircon age distributions, whereas zircon age spectra with a single dominant age peak indicate minor assimilation or reworking. Pre-eruptive zircon residence or zircon recycling from just solidified precursor intrusions may cause an overestimation of the eruption age, but crystallization just prior to the eruption is frequently recorded in zircon (e.g., Lanphere et al., 2004), and by targeting zircon rims we aimed to enhance the detection of near-eruption crystallization.

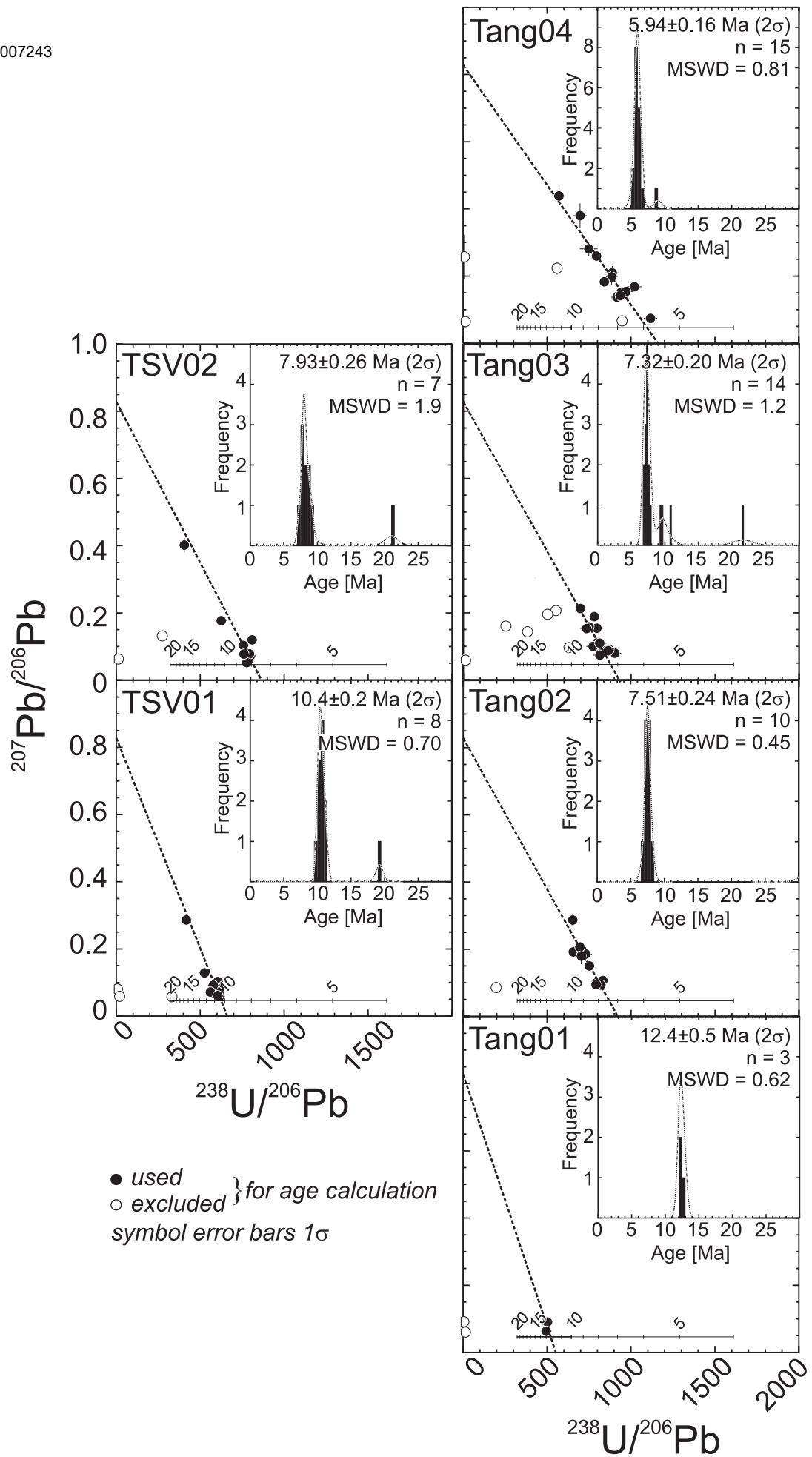
Zircon U-Pb age distributions for four out of six dated tephra samples (TSV 01, TSV 02, Tang 01, Tang 02, Tang 03, Tang 04) show a dominant young age peak with heterogeneous older ages present in varying proportions. After excluding older detrital or xenocrystic zircons, the remaining populations are homogeneous and yield the following average zircon crystallization ages for Puesto Salvación samples: 10.4 ± 0.2 Ma (TSV 01; age uncertainties quoted at 95% confidence) and 7.93 ± 0.26 Ma (TSV 02); and Angosto de Pilcomayo: 7.51 ± 0.24 Ma (Tang 02), 7.31 ± 0.20 (Tang 03), and 5.94 ± 0.16 Ma (Tang 04).

U-Pb ages for tephra samples TSV 02 and Tang 02 overlap within uncertainty, and zircons are also indistinguishable in their U abundance (Data Repository Table 1 and Fig. 1). It is therefore permissible to consider both samples as part of the same tephra, but given the continuous and frequent silicic volcanic activity within the Altiplano - Puna region throughout the Miocene, they may as well have originated from separate

eruptions. The sample Tang 01 (base of the Angosto de Pilcomayo section) has high proportions of detrital zircons. Because we initially screened the samples by rapid collection of the ^{206}Pb signal and excluded clearly identifiable old grains from further analysis, the reported ages in Data Repository Table 1 and Fig. 1 are thus already biased to younger ages. Closely overlapping U-Pb ages are obtained for three zircon grains from Tang 01 (average 12.4 ± 0.5 Ma). We interpret these ages as maximum depositional ages for each tephra. In summary, the U-Pb zircon chronostratigraphy for tephra samples from the Puesto Salvación and Angosto de Pilcomayo sections reveals that deposition of the Yecua formation may have started as early as ~ 12.5 Ma and was well under way by 10.4 Ma. The base of the overlying Tariquia Formation is dated to ~ 7.5 Ma, and the age of the upper Tariquia Formation is constrained by the youngest tephra dated at ~ 6 Ma.

References

- Lanphere, M.A., Champion, D.E., Clyne, M.A., Lowenstern, J.B., Sarna-Wojcicki, A.M., and Wooden, J.L., 2004, Age of the Rockland tephra, western USA: Quaternary Research, v. 62, p. 94– 104.
- Schmitt, A.K., Lindsay, J.M., de Silva, S.L., and Trumbull, R.B., 2002, In-situ U-Pb zircon ages of compositionally contrasting ignimbrites from La Pacana, North Chile: implications for the formation of stratified magma chambers: Journal of Volcanology and Geothermal Research, v. 120, p. 43-53.



sample	grain	spot	analysis date	$^{238}U / ^{206}Pb$	$^{238}U / ^{206}Pb$	$^{207}Pb / ^{206}Pb$	$^{207}Pb / ^{206}Pb$	Correlation of Concordia Ellipses	$^{206}Pb / ^{238}U$ Age [Ma]	$\pm 1s.e.$	U ppm	UO/U	% $^{206}Pb^*$	remarks
				1 s.e.		1 s.e.								
								$\pm 1s$	0.10					
								MSWD	1.16					
								n	14					
Tang04	1	1	7/4/2006	1115.9	40.6	0.0742	0.0045	-0.15	5.65	0.21	2329	9.2	96.4	
Tang04	2	1	7/4/2006	747.9	52.7	0.2811	0.0212	-0.16	6.10	0.70	265	9.0	70.0	
Tang04	3	1	7/4/2006	559.9	28.8	0.2234	0.0192	-0.10	8.98	0.68	219	8.9	77.3	†
Tang04	5	1	7/4/2006	888.1	43.9	0.2092	0.0204	0.06	5.83	0.41	267	9.3	79.1	
Tang04	7	1	7/4/2006	1019.9	43.2	0.1684	0.0109	0.01	5.41	0.28	277	9.2	84.4	
Tang04	8	1	7/4/2006	885.0	34.1	0.1977	0.0246	-0.31	5.95	0.37	324	8.8	80.6	
Tang04	21	1	7/4/2006	841.0	31.0	0.1827	0.0101	0.09	6.41	0.30	1508	9.0	82.6	
Tang04	30	1	7/4/2006	571.4	26.7	0.4387	0.0237	-0.19	5.69	0.72	820	9.3	49.8	
Tang04	31	1	7/4/2006	794.9	34.6	0.2598	0.0135	-0.04	5.97	0.39	393	9.4	72.7	
Tang04	32	1	7/4/2006	939.0	35.4	0.1507	0.0112	0.22	6.03	0.27	636	9.2	86.6	
Tang04	33	1	7/4/2006	916.6	34.8	0.1372	0.0111	-0.14	6.29	0.29	1212	9.4	88.4	
Tang04	34	1	7/4/2006	13.1	0.5	0.0643	0.0012	0.00	471	17	268	9.3	99.0	†
Tang04	22	1	7/4/2006	957.9	38.7	0.1456	0.0096	-0.24	5.96	0.29	858	9.2	87.3	
Tang04	23	1	7/4/2006	954.2	34.5	0.1440	0.0082	-0.08	5.98	0.26	552	9.2	87.5	
Tang04	24	1	7/4/2006	8.4	0.4	0.2574	0.0653	0.02	547	66	46	9.1	74.2	†
Tang04	25	1	7/4/2006	947.0	26.5	0.0664	0.0018	0.02	6.71	0.19	5001	9.5	97.4	†
Tang04	26	1	7/4/2006	969.0	42.4	0.1542	0.0079	0.17	5.81	0.30	603	9.5	86.2	
Tang04	27	1	7/4/2006	935.5	27.3	0.1419	0.0065	0.13	6.12	0.21	1380	9.1	87.8	
Tang04	28	1	7/4/2006	696.9	38.6	0.3799	0.0365	0.22	5.39	0.66	197	9.3	57.3	
								w.m.	5.94					
								$\pm 1s$	0.08					
								MSWD	0.81					
								n	15					

* radiogenic

† not used for weighted average calculations

UO⁺/U⁺ calibration range

7/4/2006 8.4-9.7

8/20/2006 7.9-8.6

4/6/2007 6.5-7.3