

## ***Data Repository Items***

### **BLUE RIDGE, COLORADO, QUARTZITE SUCCESSION**

#### **Rock descriptions**

The Proterozoic metasedimentary succession and crystalline basement rocks exposed along Blue Ridge were mapped and described by Reuss (1974), and the following unit descriptions are summarized largely from his work with additional data from the UNM Advanced Field Camp. Basement exposures are dominated by weakly foliated to locally mylonitic, coarse-grained granodiorite of the Paleoproterozoic Twin Mountain batholith (ca. 1705 Ma; Bickford et al., 1989) but also locally include feldspathic biotite schist and gneiss. Basement granitoids are overlain by a 15–25 m thick layer of quartzite pebble conglomerate on the northern side of the metasedimentary succession (Figs. 2A and 2D). Quartzite pebbles are 0.5–2.0 cm in size, and the surrounding matrix consists of quartz, muscovite, and well-rounded zircon up to 0.1 mm in size. The lower 5–10 m of conglomerate also contain up to 20% microcline crystals that are 0.5–2.0 cm in size and contain numerous quartz and muscovite inclusions and black tourmaline veinlets. These clasts are interpreted to represent igneous material derived from the underlying granitoid basement. Between the conglomerate and basement granitoid rocks is a thin (5–10 m), poorly-exposed zone of phyllitic material that locally contains large granitic clasts (Fig. 2D). This zone is interpreted to be a regolith that represents part of the underlying granodiorite that was sub-aerially exposed and heavily weathered prior to deposition of the conglomerate.

The quartzite pebble conglomerate is overlain by quartzite and schist that are exposed in eight northeast-trending, subvertical layers ranging in thickness from 15 to 350 m (Fig. 2A). The contact between conglomerate and quartzite ranges from sharp to gradational, and isolated lenses of conglomerate occur within the upper units of the quartzite. Based on the gradational stratigraphic progression from highly-weathered granitoid basement to basal conglomerate to quartzite and schist, the basal conglomerate and, therefore, the entire metasedimentary succession, is interpreted to be in unconformable depositional contact with underlying plutonic basement rocks. Quartzite is relatively pure (80–95% quartz with 1–10% muscovite), and individual quartz grains range in size from 0.5 to 5.0 mm and are strongly recrystallized with interlocking, sutured grain boundaries. Tabular cross-bedding is well preserved (Fig. 2C) with individual bed thicknesses between 0.1 and 0.4 m thick, and paleocurrent analysis corrected for deformation suggests a dominant current direction from the north-northeast. Oscillation ripple marks are also locally preserved and have an average amplitude and wavelength of 0.5–1.5 cm and 2–5 cm, respectively. Schist layers are dominated by quartz, biotite, and muscovite but also presently contain a variety of metamorphic minerals including garnet + sillimanite + andalusite + staurolite ± cordierite ± actinolite. Interlayered beds of quartzite and/or quartz pebble conglomerate are common.

## **Structural relationships**

After deposition, the protoliths of the metasedimentary succession experienced two major phases of deformation and metamorphism. The first phase ( $D_1/M_1$ ) involved development of a bedding-parallel foliation accompanied by growth and/or

recrystallization of biotite, muscovite, staurolite, and locally cordierite (Reuss, 1974).

The second phase of deformation ( $D_2$ ) produced the 4-km-wavelength Gooseberry Gulch syncline (Fig. 2B), the dominant structure at Blue Ridge (Fig. 2A).  $D_2$  folding was accompanied by amphibolite-facies metamorphism ( $M_2$ ) and local development of a subvertical, axial-planar foliation ( $S_2$ ). It is difficult to determine whether or not basement granitoids were folded with the overlying succession. The southern limb of the syncline is cut by a ductile shear zone that juxtaposes steeply-dipping quartzite and schist against strongly-deformed, locally-mylonitic granodiorite, omitting nearly 1 km of stratigraphic thickness from the lower part of the metasedimentary succession (Fig. 2B; Reuss, 1974). Deformation is most intense within 5–10 m of the granodiorite/quartzite contact and grades into weakly-foliated granodiorite across a distance of 20–30 m. The shear zone fabric is subparallel with  $S_2$  in the quartzite succession and, thus, is interpreted to have formed during the late stages of  $D_2$  shortening. The folded succession is also cut by a series of 60–180 m left-lateral, en-echelon folds that postdate development of the main syncline. These folds are interpreted to have formed during the late stages of  $D_2$  or during a later deformation event ( $D_3$ ; Reuss, 1974).

## REFERENCES

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**Table DR1. U-Pb Isotopic Data and Ages for Samples from Blue Ridge, Colorado**

Fraction	Weight [mg]	Concentration			Measured			*Corrected Atomic Ratios				Ages (Ma)				
		U [ppm]	Pb <sup>R</sup> Pb <sup>T</sup> [pg]	Common Pb <sup>T</sup> [pg]	206Pb 204Pb	208Pb 206Pb	206Pb 238U	207Pb 235U	207Pb 206Pb	206Pb 238U	207Pb 235U	207Pb 206Pb				
<b>Foliated granodiorite (J01-BR3; Lat N 38°35.041', Long W 105° 18.258')</b>																
Z1	single clr euh prsm abr	0.002	249	73.7	1	8048	0.0614	0.29561	78	4.2446	114	0.10414	14	1669	1683	1699
Z2	md-sm clr euh prsm abr	0.002	197	59.6	7	1113	0.0654	0.30106	78	4.3315	124	0.10435	18	1697	1699	1703
Z3	md clr euh prsm abr	0.003	181	54.8	3	3667	0.071	0.30026	92	4.3256	132	0.10448	16	1693	1698	1705
<b>Sheared granodiorite (K00-BR-25; Lat N 38°33.697', Long W 105° 18.258')</b>																
Z1	sm-md pnk sbhd prsm abr	0.001	125	38	2	1559	0.0807	0.29885	80	4.3113	132	0.10463	18	1686	1696	1708
Z2	sm-md clr sbhd abr	0.002	90	28.7	2	2640	0.1275	0.30225	118	4.3263	172	0.10381	28	1702	1698	1693
Z3	sm-md euh prsm clr abr	0.003	158	49.8	13	786	0.1213	0.29929	74	4.2958	122	0.10410	18	1688	1693	1698
Z4	md euh clr abr	0.004	185	58.4	10	1419	0.1307	0.2976	70	4.2775	110	0.10425	16	1679	1689	1701
<b>Cross-cutting pegmatite dike (K00-BR-26; Lat N 38°33.697', Long W 105° 18.258')</b>																
Z1	v sm clr-pnk sbhd prsm abr	0.001	210	59.1	4	467	0.064	0.28056	94	4.0002	268	0.10341	56	1594	1634	1686
Z2	v sm pnk-brn sbhd prsm abr	0.001	116	33	6	272	0.1089	0.27299	98	3.9369	496	0.10459	114	1556	1621	1707
Z3	sm pnk euh prsm abr	0.002	335	82.6	6	2130	0.0993	0.24061	56	3.0009	76	0.09045	12	1390	1408	1435
<b>Cross-cutting pegmatite dike (J03-BR4; Lat N 38°33.697', Long W 105° 18.258')</b>																
Z1	md clr-bge sbhd abr	0.002	480	145	20	1084	0.0776	0.29664	70	4.333	106	0.10594	14	1675	1700	1731
Z2	sm euh clr abr	0.002	92	29.2	2	1469	0.1184	0.30166	76	4.3307	136	0.10412	22	1700	1699	1699
Z3	md clr sbhe nabr	0.001	119	30.3	21	98	0.0937	0.25038	76	3.0991	316	0.08977	80	1440	1432	1421

Abbreviations: abr=abraded; brn=brown; bge=beige; clr=clear; euh=euhedral; lg=large; md=medium; nabr=not abraded; pnk=pink; prsm=prisms; sm=small; sbhd=subhedral; v=very

\*Ratios corrected for fractionation, 1 pg and 0.25 pg laboratory Pb and U blanks respectively and initial common Pb calculated using Pb isotopic compositions of Stacey and Kramers (1975). All fractions of zircon are extensively abraded (Krogh, 1982) unless otherwise noted. Two-sigma uncertainties on isotopic ratios are reported after the ratios and refer to the final digits. Pb<sup>R</sup> refers to radiogenic Pb; Common Pb<sup>T</sup> refers to total common Pb.

**Table DR2. U-Pb Isotopic Data and Ages for Detrital Zircon Sample from Phantom Canyon, Colorado**

Fraction	Weight [mg]	Concentration [ppm]	Measured Common $^{206}\text{Pb}$ $^{204}\text{Pb}$	*Corrected Atomic Ratios				Ages (Ma)											
				$^{208}\text{Pb}$	$^{206}\text{Pb}$ $^{238}\text{U}$	$^{207}\text{Pb}$ $^{235}\text{U}$	$^{207}\text{Pb}$ $^{206}\text{Pb}$	$^{206}\text{Pb}$	$^{207}\text{Pb}$	$^{207}\text{Pb}$									
				$^{204}\text{Pb}$	$^{206}\text{Pb}$	$^{235}\text{U}$	$^{238}\text{U}$	$^{235}\text{U}$	$^{238}\text{U}$	$^{235}\text{U}$									
<b>Phantom Canyon quartzite (J03-PC1; Lat N 38°35.041', Long W 105° 18.258')</b>																			
Z9	sm-md	clr	pcs	abr	0.004	239	52	88	171	0.1421	0.2028	50	2.9159	134	0.10428	36	1190	1386	1702
Z18	single	md-lg	clr	abr	0.006	115	35.8	7	1578	0.1919	2.7767	66	3.9914	102	0.10425	16	1580	1632	1701
Z43	single	md	clr	abr	0.003	219	38.9	3	2585	0.0799	0.17465	42	2.5093	64	0.1042	14	1038	1275	1700

Abbreviations: abr=abraded; clr=clear; g=large; md=medium; pcs=pieces; sm=small

\*Ratios corrected for fractionation, 1 pg and 0.25 pg laboratory Pb and U blanks respectively and initial common Pb calculated using Pb isotopic compositions of Stacey and Kramers (1975). All fractions of zircon are extensively abraded (Krogh, 1982) unless otherwise noted. Two-sigma uncertainties on isotopic ratios are reported after the ratios and refer to the final digits.  $\text{Pb}^{\text{R}}$  refers to radiogenic Pb; Common  $\text{Pb}^{\text{T}}$  refers to total common Pb. Zircon fraction numbers correspond to LA-ICP-MS analysis numbers reported in Jones (2005).













































