Supplement to "Volcanic Spreading and Lateral Variations in the Structure of Olympus Mons, Mars", by P. J. McGovern and J. K. Morgan.

## Method.

We use Particle Dynamics models (PD, also known as Distinct Element Method, or DEM) to calculate the deformation of a large volcanic edifice subject to variations in basal slope and friction. The procedure used in our numerical experiments follows that of Morgan and McGovern (2005), and is described below, with the most pertinent boundary conditions described in the main text. The PD code TRUBAL resolves contact forces acting on discrete particles, and solves Newton's equations of motion for each particle to maintain quasi-static equilibrium of the assemblage. Volcanic edifice growth is simulated by "raining" frictional particles on a planar substrate. Once each increment of particles has settled, particle positions, contacts, and forces are recorded for processing. The simulations used 9600 particles, deposited in increments of 300, to yield high-resolution models. Note that in general, the left sides of the models are intended to correspond to the northwest flank of Olympus Mons, and the right sides to the southeast flank.

At the summit of Olympus Mons, a zone of very low slope (also broadest to the southeast and south) encompasses the caldera complex. When the effects of caldera faults are filtered out, the entire summit constitutes a broad low-slope region (Fig. 2). This flatness indicates an inability to sustain broad-scale slopes over mechanically weak magma chambers. To reflect this, in some models we assign the lowest values of basal friction $\mu_{\mathrm{b}}$ to a small basal patch near the center of the edifice, to reflect the inability of a fluid-filled magma chamber to support broad-scale slopes above it. Without it, some
models develop "pointy", high-sloped summits that do not resemble the broad, flat Olympus Mons summit region. This patch is not meant to represent the basal friction in this region per se.

Lateral variations in internal friction are not considered here; effective internal friction is fixed at $\sim 0.6$. While we do not explicitly model the effects of lithospheric flexure on basal slopes, models with outward-decreasing friction may mimic several slipinhibiting effects of flexure, including outward flux of pore fluid (due to gradients in lithostatic pressure (McGovern and Solomon, 1997) and temperature) and inwardincreasing basal slopes.

## Figures.

34 Supplementary Figure DR1. Topography cross-sections and broad-scale slopes derived from MOLA topography for four profiles through the caldera of Olympus Mons. Halfprofiles are defined by azimuth from the caldera center ( $18.35^{\circ} \mathrm{N}, 226.8^{\circ} \mathrm{E}$ ). Left vertical axis: black line denotes topography in km; data interval 500 m ; vertical exaggeration approximately 7:1. Right vertical axis: diamond symbols denote magnitudes of alongtrack slopes derived from a least-squares-fit line to an 80-km-wide flank segment centered on the x-coordinate of the symbol. Slopes with magnitudes greater than 12 degrees are plotted at the 12-degree value. (a) NW-SE profile (azimuths 315 and 135). (b) NW-SE profile (azimuths 300 and 120). (c) NE-SW profile (azimuths 45 and 225). (d) NE-SW profile (azimuths 60 and 240).

44 Supplementary Figure DR2. Filtered broad-scale slopes of a NW-SE-oriented section of Olympus Mons (corresponding to that in Figure 2 of the main text), calculated from MOLA $1 / 128^{\text {th }}$ degree topography. Solid line indicates +- 300 km from center of caldera

52

66Supplementary Figure DR4. Stratigraphy and internal deformation of the same DEM models
complex oriented along $135^{\circ}$ and $315^{\circ}$ azimuths. Broad-scale slopes are calculated by a least-squares planar fit to topography points within an $80-\mathrm{km}$ diameter circle. This figure provides a direct comparison to the slope plots of Figure 3 and Supplementary Figure 3 in terms of the baseline distance ( 80 km ) over which slopes are calculated.

2 Supplementary Figure DR3. Evolution of edifice slopes as functions of load increment for four DEM models of volcanic spreading. Bottom plot: edifice slopes for each load increment, filtered by averaging over 80 km-wide bins; the absolute value of slope is shown. Edifice bases are marked by black curves. Top plot: basal friction as a function of distance. (a) Model M2 (constant friction, no basal slope), to address point raised in first paragraph of "results" section regarding wedge-like deformation and convex morphology being characteristic of constant friction models, regardless of slope. (b) Model M23, outward-decreasing friction with superposed 0.6 degree basal slope, to illustrate point raised in second paragraph of "Results" section that basal slope has only a slight effect on the nature of deformation on the uphill flank. (c) Model M20, leftward-decreasing friction with superposed 0.6 degree basal slope, to illustrate point raised in third paragraph of the "Discussion" section that basal slope has only a slight effect on the convex shape of the uphill flank. of volcanic spreading shown in Figure 3. Top plot: Color-coded stratigraphy denoted by green, gold, and white layers. Each color layer constitutes four 300-element load increments. Bottom plot: Horizontal displacement gradient field: red denotes horizontal
extension, blue denotes horizontal contraction. Strain magnitude is denoted by color intensity. (a-c) correspond to same models shown in Supplementary Figure 3.

## Movies:

Supplementary Movie 1: Surface topography, stratigraphy, and horizontal displacement gradient field for model M8: constant friction with $0.6^{\circ}$ basal slope (the model in Figures 3a and 4a-b). Note: movies use TRUBAL code coordinates, which are in .1 meter increments.

Supplementary Movie 2: Same as above for model M23: friction decreasing outward from center, no basal slope (the model in Figures 3b and 4c-d).

Supplementary Movie 3: Same as above for model M18: friction decreasing to left, no basal slope (the model in Figures 3b and 4e-f).

Supplementary Movie 4: Same as above for model M31: friction decreasing to left, $0.3^{\circ}$ basal slope (the model in Figures 3b and 4g-h).

Figure DR1





Figure DR2


Figure DR3


Figure DR4


