# GSA Data Repository Material, Formolo et al. (2007)

### I. Methods

Core extraction and GC-MS analysis. Samples from the NPC core were collected from the Western Michigan University Core Research Laboratory, and samples from the PC core were obtained from Aurora Energy, Ltd. (Traverse City, MI). The exterior of each sample was rinsed in methanol to remove external contaminants introduced during coring and handling. Powdered rock samples were decarbonated with concentrated sulfurous acid until reaction ceased. Total organic carbon concentrations were determined in triplicate using a Costech ECS140 Elemental Analyzer. Approximately 30 grams of crushed Norwood and Lachine Members of the Antrim Shale from each core were solvent extracted in 3:1 dichloromethane (DCM): methanol (MeOH) for 72 hours in a Soxhlet apparatus. The extract was separated into saturated, aromatic and two polar fractions using Si gel column chromatography by the successive elution with hexane, toluene, dichloromethane and methanol, respectively. Fractions were dried under a purified  $N_2$  stream at room temperature. Total extractable organic matter (EOM) was determined by gravimetric measurements following solvent extraction.

The hexane fraction was analyzed on a HP 6890 gas chromatograph interfaced to a HP 5973 mass selective detector, using helium as a carrier gas. The GC was equipped with a split/splitless injector and a 30m Hewlett-Packard HP-5MS fused silica capillary column (0.25mm i.d., 0.25  $\mu$ m film thickness). Saturated extract samples were redissolved in hexane and automatically injected in volumes of 1.0  $\mu$ L using an inlet temperature of 300°C. The oven temperature was initially set at 60°C for 1.5 minutes followed by an increase of 20°C/minute until 130°C and then a 2°C/minute increase to 300°C, at which it was held for the remainder of the run. Total run time = 100.50 minutes. The MS source was operated at 250°C in EI-mode at 70 eV ionization energy.

Biodegradation indices, maturity and source parameters were calculated using extracted ion chromatograms and peak area integration, using m/z = 57 (alkanes and acyclic isoprenoids), m/z = 191 (hopanes), and m/z = 217 (steranes) (Peters and Moldowan, 1993). Biodegradation biomarker indices are listed in Data Repository Table DR1; source and maturity indices are presented in Table DR2.

Water sample collection. Water samples were collected from continuously-producing wells within the Antrim NPT near the location of the PC and NPC0 cores. Temperature and pH of well waters were measured on site, and averaged 14.4°C and 7.2, respectively. Waters were drawn from within the actively pumping well stream to avoid contact with the atmosphere. Samples of cellular material were collected at each well head by drawing well water into a sterile 60 mL disposable syringe and filtering through a 0.22 µm nominal pore size cellulose acetate filter (Millipore, Billerica, MA) housed in a pre-sterilized 25 mm Swinnex filter holder. Filters for DNA extraction were frozen on dry ice, transported to the laboratory and maintained at -80°C until processed. Parallel filters for cell counts were fixed on site with 4% paraformaldehyde in phosphatebuffered saline, and rinsed with 1:1 ethanol/phosphate-buffered saline solution for subsequent direct counts. A total of eight replicate filters from two wells were DAPI stained and examined immediately by epifluoresence microscopy. Cell counts were calculated as the average and standard deviation of 20 field-counts per slide, using a 100  $\mu$ m diameter field area, a 24 mm diameter filter area, and a total volume of 60 mL of filtered water sample. The average cell count among all samples was 6 x 10<sup>3</sup> cells mL<sup>-1</sup>; the standard deviation was 3.4 x 10<sup>3</sup> cells mL<sup>-1</sup>

**Genomic DNA extraction.** Filters were placed in 0.1 M potassium phosphate buffer at pH 7.0, and shaken gently for 5 minutes to remove cells from the surface. Removal of cells from filters was confirmed by microscopic observation of the filters. Cell material was collected by centrifugation, and the protocol for preparation of genomic DNA followed that of (Stout and Nüsslein, 2005), a modification of (Ausubel et al., 1995) with the following modifications: Lysozyme (300  $\mu$ g/ml) was added followed by sodium dodecyl sulfate (0.5%) to facilitate cell lysis, and RNase A (25  $\mu$ g/ml) was added prior to Proteinase K addition to remove any co-extracted RNA.

PCR amplification of 16S rRNA genes. Bacterial 16S rRNA genes were amplified from bulk community DNA using Bacteria-specific primers 27f (Delong, 1992) in combination with the universal primer 1492r (Lane et al., 1985) in a 50 µL reaction mixture comprising 1X PCR-Buffer, 400 ng/ul bovine serum albumin, 0.25 mM each dNTP, 1 U/10 µl Tag polymerase (all from Promega, Madison, WI), 1.5 mM MgCl<sub>2</sub> (Sigma, St. Louis, MO), 0.5 µM of each primer (Integrated DNA Technologies, Coralville, IA), and 40 ng DNA template in a final volume of 30 µl. The PCR conditions for Bacteria consisted of 30 cycles of 95°C for 30 sec, 56°C for 30 sec, and 72°C for 30 sec with an initial denaturation at 95°C for 5 minutes, and a final extension at 72°C for 7 minutes. PCR amplification of archaeal 16S rRNA genes used the same concentrations of reagents as mentioned above, except an Archaea-specific nested approach was used, with primers 21f (Delong, 1992) and 1492r followed by 25f (listed in Achenbach and Woese, 1995) and 958r (Delong, 1992). In the first PCR employing primers 21F and 1492R, samples were denatured at 95°C for 5 min, followed by 25 cycles of 30 seconds each of 94°C, 55°C, and 72°C, completed by a final extension step at 72°C for 5 minutes. The amplicon was cleaned with the QIAquick PCR purification kit (Qiagen, Valencia, CA), and 1.0 µl of cleaned PCR product from the first PCR was used as template for the second PCR in a fresh reaction cocktail identical to the first reaction, but with the internal Archaea-specific primers 25f and 958r. The second PCR amplified a region of the 16S rRNA gene of around 870 bp in length, and used a program of 30 cycles of 94°C for 30 s, 55°C for 30 s, 72°C for 30 s, initiated by a denaturation step of 95°C for 5 min, and followed by a final extension step at 72°C for 5 min. All PCR reactions were performed in a PTC-200 Peltier thermal cycler (MJ Research, Waltham, MA). In order to minimize PCR bias in subsequent cloning steps, three separate reactions were run for each sample and then pooled together before DNA guantification. PCR products were purified with the QIAguick PCR purification kit (Qiagen, Valencia, CA), quantified by comparison to PCR ladder V (PGC Scientifics, Frederick, MD) using a digital imaging system with LabWorks software (UVP, Upland, CA), and were directly sequenced.

**Phylogenetic analysis.** Amplified 16S rRNA genes were sequenced with a model 3730xl DNA Analyzer (Applied Biosystems, Inc., Foster City, CA). All sequences were manually edited using the software BioEdit v7.0.4 (Hall, 1999), and aligned to the closest relative using ClustalX v1.83 (Thompson et al., 1997). Phylogenetic analyses were restricted to nucleotide positions that could be unambiguously aligned. The sequences were checked for phylogenetic affiliation using the BLAST search program (Altschul et al., 1990) at the National Center for Biotechnology Information (NCBI; www.ncbi.nlm.nih.gov), and sorted based on low BLAST score similarities and discontinuous alignment portions. All sequences were checked for chimeric artifacts using the software Mallard v1.02 (Ashelford et al., 2006), and eleven chimeric 16S rRNA gene sequences were excluded from further phylogenetic analysis. Sequences that showed  $\geq$ 98% nucleotide sequence identity were considered similar strains of the same phylotype and grouped as a separate sample (*i.e.* phylotypes) using the software

DOTUR v1.53 (Schloss et al., 2004). A 16S rRNA gene phylogenetic tree was constructed from the alignments based on the method Minimum Evolution and calculated by the algorithm of the Tamura-Nei model (Tamura and Nei, 1993), using the software MEGA v3.1 (Kumar et al., 2004). Bootstrap confidence values were obtained with 1000 replicates. The tool Classifier and Sequence Match in the Ribosomal Database Project II release 9.37 (Cole et al., 2003) and BLASTn (Altschul et al., 1990) were used to classify the 16S rRNA gene clones and identify the nearest neighbors of the 16S rRNA genes in the GenBank database. The sequences obtained from the clones have been deposited in the GenBank database under accession numbers EF117331 to EF117602.

#### II. Methane generation and isotope mass balance models

Crushed core gas analysis reveals an average  $CH_4$  content in rocks from the Antrim NPT of 25 scf/ton, corresponding to  $6.3 \times 10^{-5}$  moles  $CH_4$  cm<sup>-3</sup><sub>rock</sub>. Extended across the entire Antrim NPT (160 km × 40 km × 40 m), this yields a total existing  $CH_4$  content of 190 Tg C ( $CH_4$ ).

The amount of methane cumulatively produced in the Antrim NPT, and thus the amount subsequently lost during deglaciation, can be estimated from an isotope mass balance model that requires  $CH_4$  and all inorganic carbon pools ( $CO_2$ ,  $HCO_3^-$ , and fracture-filling CaCO<sub>3</sub>) to derive from degradation of shale organic matter:

 $m_{\text{CO2}} \cdot \delta_{\text{CO2}} + m_{\text{HCO3}} \cdot \delta_{\text{HCO3}} + m_{\text{CaCO3}} \cdot \delta_{\text{CaCO3}} + m_{\text{CH4}} \cdot \delta_{\text{CH4}} = (m_{\text{CO2}} + m_{\text{HCO3}} + m_{\text{CaCO3}} + m_{\text{CH4}}) \cdot \delta_{\text{kerogen}}$ 

Using  $m_{CO2} = 1.1 \times 10^{-6}$  moles cm<sup>-3</sup><sub>rock</sub> (2% of total gas),  $\delta_{CO2} = 18$  ‰;

 $m_{\text{HCO3}} = 2 \times 10^{-6} \text{ moles cm}^{-3}_{\text{rock}}$  (40 mM HCO<sub>3</sub><sup>-</sup>, 5% porosity),  $\delta_{\text{HCO3}} = 28$  ‰;  $m_{\text{CaCO3}} = 193 \times 10^{-6} \text{ moles cm}^{-3}_{\text{rock}}$  (1.0 % CaCO<sub>3</sub>, rock density = 1.93 g cm<sup>-3</sup>),  $\delta_{\text{CaCO3}} = 28$  ‰;  $\delta_{\text{kerogen}} = -30$  ‰ and  $\delta_{\text{CH4}} = -52$  ‰,

yields an estimated methane content of 517 ×  $10^{-6}$  moles cm<sup>-3</sup><sub>rock</sub> to achieve isotope mass balance. Extended across the entire Antrim NPT, this equals a cumulative CH<sub>4</sub> mass of 1550 Tg C (CH<sub>4</sub>). The 63 ×  $10^{-6}$  moles<sub>CH4</sub> cm<sup>-3</sup><sub>rock</sub> remaining today represents ~12% of this cumulative methane pool.

A similar isotope mass balance can be constructed from a Rayleigh distillation model, using  $R/R_0 = f^{\alpha-1}$  where R is the present  ${}^{13}C/{}^{12}C$  ratio in the Antrim inorganic carbon pool ( $CO_2 + HCO_3^- + CaCO_3$ ),  $R_0$  is the  ${}^{13}C/{}^{12}C$  ratio of the inorganic carbon pool prior to methanogenesis,  $\alpha$  is the fractionation associated with  $CO_2$ -reduction methanogenesis ( $\alpha = 0.93$  between the total inorganic carbon pool and methane), and f is the fraction inorganic carbon remaining. Using  $R/R_0 = (\delta^{13}C + 1000)/(\delta^{13}C_0 + 1000)$ ,  $\delta^{13}C$  at present for inorganic carbon equals 28‰, and  $\delta^{13}C_0 = -30\%$  (assumed to all derive from shale OM), the fraction of inorganic carbon pool. 56% has been converted into methane, yielding a cumulative methane pool of  $260 \times 10^{-6}$  moles cm $^{-3}_{rock}$ . Extended across the entire Antrim NPT, this equals 780 Tg C as CH<sub>4</sub>. The 6.3 ×  $10^{-5}$  moles<sub>CH4</sub> cm $^{-3}_{rock}$  remaining today represents ~25% of this cumulative methane pool. A closed-system Rayleigh isotope distillation approach, which does not accommodate sustained production of CO<sub>2</sub> or methanogenesis via mechanisms other than CO<sub>2</sub> reduction, likely yields a severe underestimate of CH<sub>4</sub> production and CO<sub>2</sub> loss.

A third means to estimate the cumulative original mass of biogenic  $CH_4$  in the Antrim employs the difference in extract yield measured in the PC and NPC cores. When normalized to TOC content, the biodegraded core reflects a loss of 50

 $mg_{extract}/g_{TOC}$ . If the loss of these hydrocarbons resulted entirely in the conversion to methane and CO<sub>2</sub> as suggested by methanogenic decomposition of hexadecane, this would equal to generation of a cumulative pool of 440 × 10<sup>-6</sup> moles<sub>CH4</sub>/cm<sup>3</sup><sub>rock</sub> and 134× 10<sup>-6</sup> moles<sub>CO2</sub>/cm<sup>3</sup><sub>rock</sub>. Extended across the entire Antrim NPT, this equals 1320 Tg C as CH<sub>4</sub>. In this estimate, the remaining pool of Antrim CH<sub>4</sub> is nearly 20% of the initial pool. The total CH<sub>4</sub> addition to the atmosphere is calculated using 1 ppbv = 2.86 Tg C (CH<sub>4</sub>).

#### III. Minimum maintenance energy calculations

Direct counts of Antrim natural gas well waters collected from the NPT indicate a freefloating population of  $(1\pm0.3)\times10^4$  cells mL<sup>-1</sup>, with an average diameter of 1.5  $\mu$ m. Using cell biomass (in pg C  $\mu$ m<sup>-3</sup>) = 0.09 × V<sup>0.9</sup> (Norland, 1993), we estimate a cellular biomass in the Antrim NPT of 1.5 × 10<sup>-10</sup> g<sub>C-biomass</sub> cm<sup>-3</sup><sub>rock</sub>. Using a minimum maintenance energy requirement equivalent to 0.14 kJ mol<sup>-1</sup><sub>C</sub> hr<sup>-1</sup> (Scholten and Conrad, 2000), we estimate a total community maintenance energy demand of 1.54 × 10<sup>-8</sup> kJ cm<sup>-3</sup><sub>rock</sub> yr<sup>-1</sup>. Lower maintenance energy demand may be indicated for *in situ* subsurface communities (D'Hondt et al., 2002; Biddle et al., 2006); however, freefloating cells may also represent only a small fraction of the subsurface Antrim community.

Methanogenic hydrocarbon decomposition can be written (for hexadecane) as:

$$4 \text{ C}_{16}\text{H}_{34} + 30 \text{ H}_2\text{O} \rightarrow 49 \text{ CH}_4 + 15 \text{ CO}_2$$

Estimated Gibbs free energy yield for this reaction, at concentrations of reactants and products measured in enrichment cultures, is -1596 kJ/reaction (Zengler et al., 1999), which equals -399 kJ mol<sub>hexadecane</sub><sup>-1</sup> or -2.08 kJ g<sub>C-hexadecane</sub><sup>-1</sup>. This energy yield specifically relates to culture conditions described by Zengler et al. (1999) where activities of reactants and products were known. Energy yields for conditions in the Antrim subsurface would need to be modified to reflect in situ conditions where CH<sub>4</sub>, CO<sub>2</sub> and hydrocarbon activities may be different. Community maintenance energy demand divided by free energy yield gives a rate of shale OM degradation of 7.4 × 10<sup>-9</sup> g<sub>C</sub> cm<sub>rock</sub><sup>-3</sup> yr<sup>-1</sup>, if the free energy change for methanogenic decomposition of hexadecane is representative of total shale OM. At this rate, 20.9 million years would be required to remove all OM from a rock containing 8% total organic carbon (0.154 g<sub>ToC</sub> cm<sup>-3</sup>).

The time required to generate the present pool of methane in the Antrim shale  $(6.3 \times 10^{-5} \text{ moles}_{CH4} \text{ cm}_{rock}^{-3})$  can be estimated from the rate of shale OM degradation, using the stoichiometry of the hexadecane decomposition reaction (49 moles CH<sub>4</sub> produced per 64 moles hydrocarbon-C degraded). Hydrocarbon degradation at a rate of 7.4 × 10<sup>-9</sup> g<sub>C</sub> cm<sub>rock</sub><sup>-3</sup> yr<sup>-1</sup> equals a methane production rate of 4.7 × 10<sup>-10</sup> moles<sub>CH4</sub> cm  $_{rock}^{-3}$  yr<sup>-1</sup>. It would require 133 kyr to generate 6.3 × 10<sup>-5</sup> moles<sub>CH4</sub> cm $_{rock}^{-3}$  (the modern inventory of Antrim methane), or 0.55-1.09 Myr to generate the estimated (440-517) × 10<sup>-6</sup> moles<sub>CH4</sub> cm $_{rock}^{-3}$  cumulatively produced in the rock.

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	Depth	TOC	Extractable OM	Extract yield	<u>n-all</u>	kanes	acyclic isoprenoids						
Sample name	(m)	(wt. %)	(mg <sub>extract</sub> /g <sub>rock</sub> )	(mg <sub>extract</sub> /g <sub>TOC</sub> )	Σ <b>nC</b> <sub>14</sub> -nC <sub>18</sub> *	Σ <b>nC</b> <sub>19</sub> -nC <sub>36</sub> *	norpristane*	pristane*	phytane*	Pr/ <i>n-</i> C <sub>17</sub> <sup>†</sup>	Ph/ <i>n-</i> C <sub>18</sub> §		
non-producin	ig core												
NPC L-1 <sup>†</sup>	658	4.6	285	190	621	808	375	138	95	0.79	0.58		
NPC L-2	667	5.3	185	117	529	883	296	114	87	0.61	0.53		
NPC N-1	695	5.3	219	110	638	1324	329	125	97	0.86	0.51		
NPC N-2	698	6.2	162	87	1476	1556	603	218	142	0.52	0.43		
average NPC		5.3±0.6	213±54	126±44	816 ± 443	1143 ± 357	401 ± 140	149 ± 47	105 ± 25	0.70 ± 0.16	0.51 ± 0.06		
producing core													
PC L-1	242	6.9	218	89	21	136	122	53	36	7.36	4.27		
PC L-2	251	5.8	143	81	48	235	175	82	44	5.04	2.48		
PC N-1	263	9.9	195	60	214	739	227	93	69	1.2	0.77		
PC N-2	265	11.1	291	75	52	586	649	243	132	3.46	2.5		
average PC		8.4±2.5	212±61	76±12	83 ± 88	424 ± 285	293 ± 242	118 ± 85	70 ± 44	4.27 ± 2.60	2.51 ± 1.43		

### TABLE DR1: SAMPLE DEPTH, TOC, EXTRACTION YIELDS, AND BIODEGRADATION INDICES

<sup>T</sup>The Antrim Shale Formation includes, from the base, the Norwood Member (black shale), the Paxton Member (lime mudstone and gray shale), the Lachine Member (black shale) and Upper Antrim (interbedded black and gray shale). "L-" samples are from Lachine Member, "N-" samples are from the Norwood Member.

\*abundance normalized to C29 20S diasterane.

<sup>†</sup>Pr/n-C17 = pristane/C17 n-alkane.

<sup>§</sup>Ph/n-C18 = phytane/C18 n-alkane.

Sample							Hopan	es*						
	C <sub>34</sub> 22R	C <sub>34</sub> 22S	C <sub>33</sub> 22R	C <sub>33</sub> 22S	C <sub>32</sub> 22R	C <sub>32</sub> 22S	C <sub>31</sub> 22R	C <sub>31</sub> 22S	C <sub>30</sub> βα	C <sub>30</sub> αβ	C <sub>29</sub> βα	C <sub>29</sub> αβ	Tm	Ts
NPC L-1	0.26	0.06	0.18	0.38	0.23	0.37	0.36	0.55	0.34	1.61	0.25	0.94	0.54	0.70
NPC L-2	0.14	0.29	0.25	0.45	0.31	0.38	0.44	0.57	0.29	1.96	0.20	1.08	0.59	0.86
NPC N-1	0.21	0.12	0.22	0.40	0.26	0.42	0.35	0.48	0.18	1.76	0.16	0.87	0.29	0.46
NPC N-2	0.28	0.36	0.52	0.72	0.41	0.80	0.62	1.03	0.53	2.34	0.20	1.26	0.57	0.92
Avg. NPC	0.22	0.21	0.29	0.49	0.30	0.49	0.44	0.66	0.34	1.92	0.20	1.04	0.50	0.73
PC L-1	n.d.	n.d.	0.26	0.07	0.26	0.55	0.36	0.71	1.83	2.28	1.32	3.52	0.58	2.89
PC L-2	0.16	0.67	0.19	0.57	0.33	0.40	0.46	0.79	0.49	2.15	0.34	1.34	0.59	0.69
PC N-1	0.08	0.56	0.35	0.37	0.27	0.47	0.42	0.55	0.61	1.72	0.42	1.56	0.61	1.29
PC N-2	0.24	0.44	0.34	0.55	0.44	0.63	0.67	0.78	0.32	2.28	0.48	1.53	0.76	0.93
Avg. PC	0.16	0.56	0.28	0.39	0.33	0.51	0.48	0.71	0.81	2.11	0.64	1.99	0.63	1.45

# TABLE DR1 (continued): BIODEGRADATION INDICES<sup>\*</sup>

		Regular Sterane	S*	
C27ααα (20R)	C <sub>29</sub> ααα (20R)	C <sub>27</sub> ααα (20S)	C <sub>29</sub> ααα (20S)	C <sub>29</sub> αββ
0.22	0.60	0.70	0.78	0.53
0.19	0.64	0.65	0.76	0.37
0.10	0.36	0.46	0.50	0.47
1.02	0.36	0.57	0.35	0.41
0.38	0.49	0.60	0.60	0.44
1.91	5.00	0.99	13.40	5.29
0.39	1.36	0.53	0.86	0.48
0.79	1.55	0.40	0.40	1.47
0.57	1.66	0.53	2.95	0.85
0.91	2.39	0.61	4.40	2.02
-	C <sub>27</sub> ααα (20R) 0.22 0.19 0.10 1.02 <b>0.38</b> 1.91 0.39 0.79 0.57 <b>0.91</b>	C <sub>27</sub> ααα (20R)         C <sub>29</sub> ααα (20R)           0.22         0.60           0.19         0.64           0.10         0.36           1.02         0.36           0.38         0.49           1.91         5.00           0.39         1.36           0.79         1.55           0.57         1.66           0.91         2.39	Regular SteraneC27ααα (20R)C29 ααα (20R)C27ααα (20S)0.220.600.700.190.640.650.100.360.461.020.360.570.380.490.601.915.000.990.391.360.530.791.550.400.571.660.530.912.390.61	Regular Steranes*C27ααα (20R)C29 ααα (20R)C27ααα (20S)C29 ααα (20S)0.220.600.700.780.190.640.650.760.100.360.460.501.020.360.570.350.380.490.600.601.915.000.9913.400.391.360.530.860.791.550.400.400.571.660.532.950.912.390.614.40

\*abundance normalized to C29 20S diasterane.

Alkanes										
SAMPLE	CPI <sup>a</sup>	OEP(17) <sup>b</sup>	OEP <sub>(23)</sub> <sup>b</sup>	OEP(77) <sup>b</sup>	OEP <sub>(29</sub> <sup>b</sup>		OEP <sub>(31)</sub> <sup>b</sup>	OEP(33) <sup>b</sup>	Pr/Ph <sup>c</sup>	
NPC L-1	1.00	1.09	1.00	0.96	1.05		1.11	1.03	1.70	
NPC L-2	0.98	1.17	0.99	0.91	0.97		1.03	0.99	1.36	
NPC N-1	1.04	1.01	1.01	0.99	1.09		1.10 1.08		1.55	
NPC N-2	1.00	1.05	1.03	0.98	1.01		1.00	1.07	1.45	
Avg. NPC	1.00±0.03	1.08±0.07	1.01±0.02	0.96±0.03	1.03±0	.05 1	1.06±0.05 1.04±0.		1.51±0.15	
PC L-1	1.06	1.13	1.01	1.14	0.87		0.94	0.98	1.53	
PC L-2	0.64	1.11	0.51	0.97	1.05		1.11	1.14	1.87	
PC N-1	1.00	1.18	1.02	0.93	0.98		1.03	0.98	1.44	
PC N-2	1.06	1.29	1.04	0.95	1.03		1.07	1.07	1.80	
Avg. PC	0.94±0.20	1.18±0.08	0.90±0.26	1.00±0.10	0.98±0	.08 1	.04±0.07	1.04±0.08	1.66±0.21	
Hopanes										
			22S/(	22S+22R) <sup>d</sup>		βα/(	$(\beta \alpha + \alpha \beta)^{e}$	βα/α	$\beta$ (moretane/h	opane)
SAMPLE	Ts/(Ts+Tm)	C <sub>31</sub>	C <sub>32</sub>	C <sub>33</sub>	C <sub>34</sub>	C <sub>29</sub>	C <sub>30</sub>	C <sub>29</sub>	C <sub>30</sub>	Ts/C <sub>30</sub> 17α
NPC L-1	0.57	0.61	0.62	0.68	0.18	0.21	0.17	0.26	0.21	0.44
NPC L-2	0.59	0.56	0.55	0.64	0.68	0.16	0.13	0.19	0.15	0.44
NPC N-1	0.61	0.58	0.61	0.65	0.63	0.16	0.09	0.19	0.10	0.26
NPC N-2	0.39	0.62	0.66	0.58	0.57	0.14	0.18	0.16	0.23	0.39
Avg. NPC	0.54±0.09	0.59±0.02	0.61±0.04	0.64±0.04	0.52±0.20	0.17±0.03	0.14±0.04	0.20±0.04	0.17±0.05	0.38±0.07
PC L-1	0.83	0.66	0.68	0.21	n.d.	0.27	0.45	0.38	0.8	1.26
PC L-2	0.32	0.63	0.55	0.75	0.81	0.2	0.19	0.26	0.23	0.32
PC N-1	0.68	0.57	0.63	0.52	0.87	0.21	0.26	0.27	0.35	0.75
PC N-2	0.55	0.54	0.59	0.62	0.65	0.24	0.12	0.32	0.14	0.41
Avg. PC	0.60±0.22	0.60±0.05	0.61±0.06	0.53±0.23	0.78±0.11	0.23±0.03	0.26±0.14	0.31±0.06	0.38±0.29	0.69±0.43

# TABLE DR2. SOURCE AND MATURITY INDICES

TABLE DR2. SOURCE AND MATURITY INDICES (continued)												
						Steranes						
	20S/(20S+20R) <sup>f</sup> 20S/(20S+20R) <sup>g</sup>			C <sub>27</sub> /(C <sub>27</sub> +C <sub>29</sub> ) C <sub>27</sub> (dia)/ Dia <sub>total</sub> /						C <sub>27</sub> diasterane		
SAMPLE	C <sub>27</sub> ααα	C <sub>29</sub> ααα	C <sub>27</sub> (dia) αβ	C <sub>27</sub> (dia) βα	C <sub>29</sub> (dia)	C <sub>27</sub> (dia)+ C <sub>27</sub> (reg)	C <sub>29</sub> (dia)+ C <sub>29</sub> (reg)	(Dia <sub>total</sub> + Reg <sub>total</sub> )	diasteranes	C <sub>27</sub> (dia) / C <sub>27</sub> (reg)	C <sub>29</sub> (dia) / C <sub>29</sub> (reg)	βα/αβ
NPC L-1	0.76	0.56	0.47	0.54	0.56	0.67	0.49	0.56	0.5	1.99	0.95	1.69
NPC L-2	0.78	0.54	0.59	0.66	0.59	0.7	0.49	0.58	0.54	2.35	0.95	2.45
NPC N-1	0.83	0.58	0.47	0.61	0.53	0.73	0.59	0.64	0.44	2.69	1.42	1.82
NPC N-2	0.36	0.49	0.50	0.57	0.45	0.67	0.54	0.6	0.46	1.2	2.00	1.67
Avg. NPC	0.68±0.22	0.54±0.04	0.51±0.06	0.60±0.05	0.53±0.06	0.69±0.03	0.53±0.05	0.60±0.03	0.49±0.04	2.06±0.64	1.33±.050	1.91±0.37
PC L-1	0.34	0.73	0.85	0.40	0.53	0.6	0.07	0.19	0.69	1.49	0.08	0.83
PC L-2	0.58	0.39	0.47	0.64	0.53	0.63	0.41	0.49	0.45	1.69	0.7	2.08
PC N-1	0.34	0.71	0.39	0.48	0.55	0.51	0.21	0.28	0.41	1.05	0.27	1.81
PC N-2	0.48	0.64	0.49	0.63	0.56	0.71	0.25	0.4	0.6	2.42	0.33	1.53
Avg. PC	0.44±0.12	0.62±.016	0.55±.020	0.54±.012	0.54±0.02	0.61±0.08	0.24±0.14	0.34±0.13	0.540.13	1.66±0.57	0.35±0.26	1.56±0.54

<sup>a</sup> CPI =  $2[(C_{23}+C_{25}+C_{27}+C_{29}]/[C_{22}+(C_{24}+C_{26}+C_{28})+C_{30}]^{21}$ 

<sup>b</sup> OEP<sub>x</sub> =  $[C_{(x-2)}+6C_x+C_{(x+2)}]/[4C_{(X-1)}+4C_{(x+1)}]^{21}$ 

<sup>c</sup> Pr/Ph = pristane/phytane<sup>21</sup>

 $^{d}22S/(22S+22R) = 17\alpha(H), 22\beta(H)-22S \text{ homohopane} / [17\alpha(H), 22\beta(H)-22S \text{ homohopane} + 17\alpha(H), 22\beta(H)-22R \text{ homohopane}]^{21}$ 

<sup>e</sup> βα/ (βα +αβ) = 17β(H),21α(H) hopane/ [17β(H),21α(H) hopane + 17α(H),21β(H) hopane] <sup>21</sup>

 ${}^{f}20S/(20S+20R) = 5\alpha(H), 14\alpha(H), 17\alpha(H) - 20S \text{ stigmastane} / [5\alpha(H), 14\alpha(H), 17\alpha(H) - 20S \text{ stigmastane} + 5\alpha(H), 14\alpha(H), 17\alpha(H) - 20R \text{ stigmastane} ]^{21}$ 

 $^{g}$  20S/(20S+20R) = C<sub>29</sub> 13 $\beta$ (H),17 $\alpha$ (H)-20S diasterane/ [C<sub>29</sub> 13 $\beta$ (H),17 $\alpha$ (H)-20S diasterane + C<sub>29</sub> 13 $\beta$ (H),17 $\alpha$ (H)-20R diasterane]<sup>21</sup>

# TABLE DR3. BIOGENIC NATURAL-GAS PRODUCTION UNDER FORMERLY GLACIATED OR PERIGLACIAL TERRAINS

	Basin	location	reference
1	Alberta	Alberta (Canada)	Adams et al., (1984); McIntosh et al. (2004)
2	Anadarko	OK (USA)	McIntosh et al. (2004)
3	Appalachian	NY-PA-OH-WV-KY (USA)	McIntosh et al. (2004)
4	Bellingham	WA (USA)	Savoie et al., (2002)
5	Campine	Belgium	Hildenbrand et al. (2006)
6	Columbia	WA (USA)	Johnson et al. (1993)
7	Cook Inlet	AK (USA)	Claypool et al. (1980)
8	Denver	CO (USA)	Shurr and Ridley (2002)
9	Dnieper-Donets	Ukraine	McIntosh et al. (2004)
10	Eastern Siberia	Russia	Nakagawa et al., (2002)
11	Forest City	MO-KS (USA)	McIntosh et al. (2004)
12	Illinois	IL-IN-KY (USA)	McIntosh et al. (2004)
13	Kattegat	Denmark	lversen and Jorgensen (1985)
14	Michigan	MI (USA)	McIntosh et al. (2004); Martini et al., (1998)
15	Po	Italy	Elliot et al. (1993)
16	Powder River	WY-MT (USA)	Ayers (1986)
17	Pripyat	Belarus	McIntosh et al. (2004)
18	Ria de Vigo	Spain	Garcia-Garcia et al. (2003)
19	Ruhr	Germany	Thielemann et al., (2004)
20	San Juan	NM-CO (USA)	Rice et al. (1989); Scott et al. (1994)
21	Skagerrak	Denmark	Dando et al. (1994); Iversen and Jorgensen (1985)
22	Silesian	Poland	Kotarba and Rice (2001)
23	Southern North Sea	Germany, Netherlands	Schroot et al., (2005)
24	Timan-Pechora	Russia	McIntosh et al. (2004)
25	Valenje	Slovenia	Kenduc and Pezdic (2005)
26	Wasatch Plateau	UT (USA)	Lamarre (2003)
27	Williston	MT (USA), AB (Canada)	McIntosh et al. (2004)
28	Yukon	Canada	Martens et al. (1992)



**Figure DR1.** Phylogenetic relationship and distribution of bacterial clones as identified by family/class(\*)/phylum (\*\*). The comparative analysis was inferred by Minimum Evolution analysis of aligned 16S rRNA gene sequences from 91 bacterial clones compared with public nucleotide databases. Clones are designated by the prefixes • and  $\circ$  (reflecting two different Antrim wells adjacent to PC core), followed by the number of representatives [in brackets] in their respective affiliation. The scale-bars represent 5% estimated sequence divergence. Bootstrap values shown are the percentage of occurrences for nodes with  $\geq$ 50% support based on an analysis of 1000 iterations.



**Figure DR2.** Phylogenetic relationship and distribution of archaeal clones as identified by family/class(\*)/phylum (\*\*). The comparative analysis was inferred by Minimum Evolution analysis of aligned 16S rRNA gene sequences from 181 archaeal clones compared with public nucleotide databases. Clones are designated by the prefixes • and  $\circ$  (reflecting two different Antrim wells adjacent to PC core), followed by the number of representatives [in brackets] in their respective affiliation. The scale-bars represent 5% estimated sequence divergence. Bootstrap values shown are the percentage of occurrences for nodes with  $\geq$ 50% support based on an analysis of 1000 iterations.