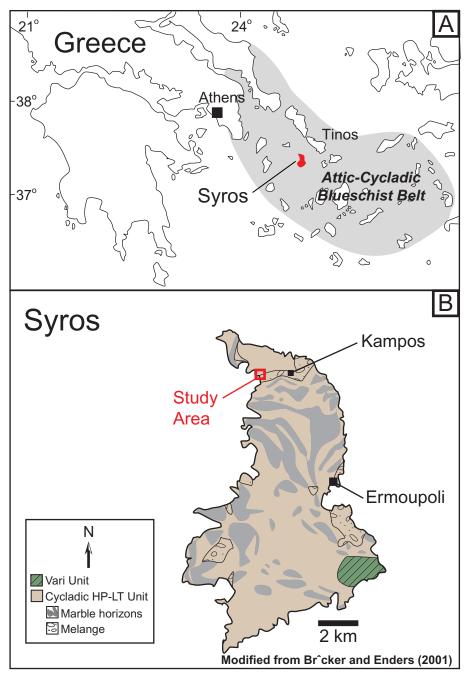
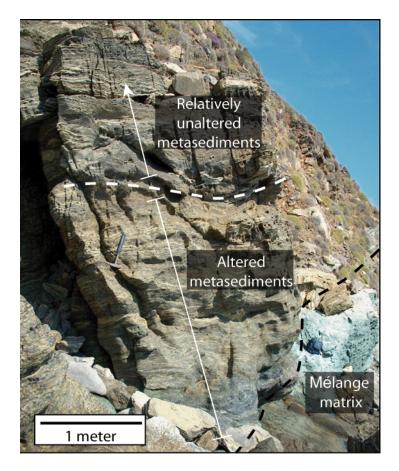
## Breeding et al., Data Repository Material Figure DR1



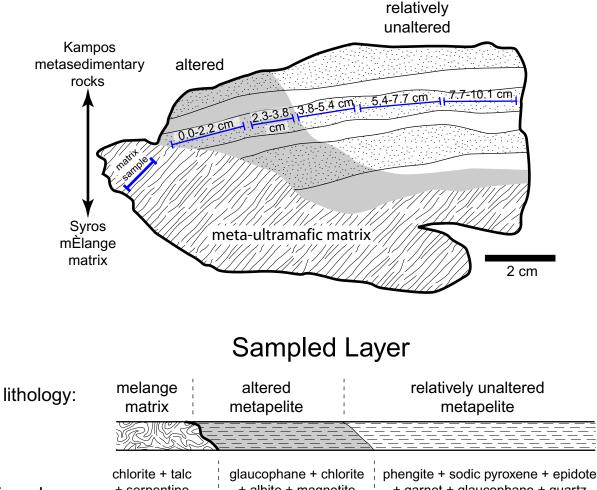
**Figure DR1.** Location and geologic maps of Syros. A. Map of Greece and Aegean showing Syros within HP-LT Attic-Cycladic Blueschist Belt. B. Geology of Syros showing Cycladic HP-LT unit with marbles and mÈlange and structurally overlying Vari Unit (not HP-LT). Study area at edge of mÈlange zone near Kampos marked by red square.

## Breeding et al., Data Repository Material Figure DR2



**Figure DR2.** Photograph of altered metasedimentary rocks at edge of Syros mÈlange zone. Visibly altered regions as thick as 2.5 meters in the metasediments are found adjacent to contacts with mÈlange matrix.

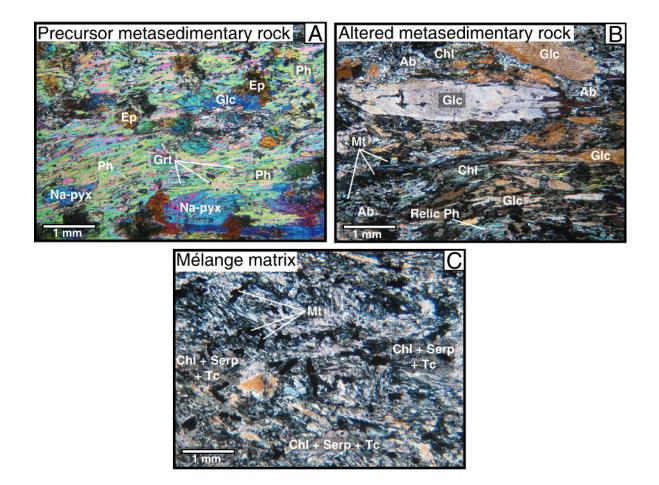
## Breeding et al., Data Repository Material Figure DR3



**Figure DR3.** Schematic diagram of sample showing mineral assemblages and sampling locations. Metasediments preserve primary interlayering of meta-psammite (non-patterned) and metapelite (stippled). Sampled layer consists of metapelite in direct contact with meta-ultramafic mÈlange matrix. Metasomatic alteration (gray shading) is visible near contact. Samples for analysis were taken every ~1.5 to 2 cm along metapelite layer (see sampling subdivisions) to investigate geochemical changes as matrix contact is approached.

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# Breeding et al., Data Repository Material Figure DR4



**Figure DR4.** Crossed-polarized photomicrographs of mineral assemblages from relatively unaltered metasedimentary rock, altered metasedimentary rock, and the mÈlange matrix. All scale bars are 1 mm long. A: Distal to contact with mÈlange, metasedimentary layer primarily contains phengite (Ph), sodic-pyroxene (Na-pyx), epidote (Ep), garnet (Grt), glaucophane (Glc), and quartz (not labeled). B: Within the alteration zone adjacent to the contact, the mineral assemblage is dominated by glaucophane (Glc), chlorite (ChI), albite (Ab), and magnetite (Mt), and includes relic phengite (Ph). Albite often contains abundant inclusions of magnetite. Note the larger grain size and higher abundance of glaucophane relative to part A. C: The meta-ultramafic matrix consists of fine-grained, intergrown chlorite (ChI), serpentine (Serp), and talc (Tc) with magnetite (Mt).

### Breeding et al., Data Repository Material Appendix DR1

#### Mineral Analysis

Mineral analyses were performed at Yale University on a JEOL JXA-8600 electron microprobe equipped with wavelength dispersive spectrometers using natural and synthetic standards, off-peak background corrections, and  $\varphi(oz)$  matrix corrections. Standard analytical conditions were 15 kV accelerating voltage and 10 um beam spots. Major and minor element acquisitions on all minerals were done with a 15 nA beam current and 10 second on-peak and off-peak count times. Pb and Cu concentrations in most phases were measured using a 50 nA beam current and count times of 300 seconds both on- and off-peak. For trace element analyses of hydrous minerals like phengite and chlorite, a 15 nA beam current and 20 µm beam spots were employed with longer count times of 600 seconds both on- and off-peak. Detection limits for Pb and Cu were 170 ppm and 120 ppm respectively (99 % confidence). Cu and Pb were consistently present in chlorite and epidote, respectively, whereas values were below detection in the other minerals (i.e., phengite, Na-pyroxene, garnet, glaucophane, albite, titanite, rutile, apatite, and magnetite). Detection limits for Sr in epidote and apatite were 210 and 130 ppm, respectively (99 % confidence). Epidotes have measured Sr concentrations as great as ~2600 ppm, whereas apatite Sr ranges from ~300 to ~450 ppm. Sr was below detection in the other minerals.

#### **Thermobarometric calculations**

Thermobarometric calculations made using Thermocalc V. 3.2 (Holland and Powell, 1998) at 500-550  $^{\circ}$ C (Trotet et al., 2001) and the reaction: 4 Albite + 2 Clinochlore = 2

Glaucophane + Amesite + 2 H<sub>2</sub>O. Mineral compositions in alteration assemblage determined from electron microprobe analyses using AX software (Holland and Powell, 1998) and are as follows: albite = NaAlSi<sub>3</sub>O<sub>8</sub>; chlorite = Mg<sub>3.517</sub> Fe<sup>2+</sup><sub>1.156</sub> Mn<sub>0.107</sub> Al<sub>1.175</sub> [Si<sub>2.902</sub>Al<sub>1.098</sub>] O<sub>10</sub> [F<sub>0.032</sub>(OH)<sub>7.968</sub>]; glaucophane = Na<sub>1.992</sub> Ca<sub>0.021</sub> Mg<sub>2.488</sub> Fe<sup>2+</sup><sub>0.435</sub> Fe<sup>3+</sup><sub>0.584</sub> Mn<sub>0.039</sub> Al<sub>1.400</sub> Si<sub>8.020</sub> O<sub>22</sub> [F<sub>0.065</sub>(OH)<sub>1.935</sub>]. Cl contents were below detection in chlorite and glaucophane. H<sub>2</sub>O activity was set at 1.0 based on a lack of significant halogen contents in hydrous minerals and fluid inclusion data from Syros indicating little CO<sub>2</sub> in HP-LT fluids (Barr, H., 1990, Preliminary fluid inclusion studies in a high-grade blueschist terrain, Syros, Greece: Mineralogical Magazine, v. 54, p. 159-168).

#### Calculation of amount of metasediment alteration needed for Ba extraction

The amount of fluid alteration of metasedimentary rocks necessary to remove enough Ba to enrich a volcanic eruption equivalent in size to the 1980 Mount St. Helens eruption, USA, by 100 ppm was done using the following parameters: Ba removed from metasediments during alteration = ~400 ppm (see Table DR1), Volume of Mount St. Helens eruption = ~0.4 km<sup>3</sup> (combined ash and pyroclastic volume taken from Tilling, R. I., Topinka, L. & Swanson, D. A., 1990, Eruptions of Mount St. Helens: Past, Present, and Future: USGS Special Interest Publication, 56 p.) Average densities of metasedimentary rock and erupted volcanic ash assumed to be 3.0 g/cm<sup>3</sup> and 2.3 g/cm<sup>3</sup>, respectively.

#### Table DR1: Chemical analyses were performed by X-Ray Assay Laboratories (XRAL) in Don Mills, Ontario. Major elements were determined by XRF and are reported as wt% oxide. Trace elements were determined by ICP-MS and are reported as ppm. Total Fe is reported as Fe<sub>2</sub>O<sub>3</sub>. LOI is loss on ignition. Sample SY 51E-1 SY 51E-2 SY 51E-3 SY 51E-4 SY 51E-5 SY 51E-7 3.8 - 5.1 cm 2.3 - 3.8 cm Dist. from ultramafic contact 0 - 2.2 cm 5.1 - 7.7 cm 7.7 - 10.1 cm Description Altered Altered Some alt. Least alt. Least alt. Matrix $SiO_2$ 45.51 50.54 51.87 50.22 50.71 32.14 TiO<sub>2</sub>\* 0.90 0.80 0.57 0.65 0.58 1.09 $Al_2O_3$ 13.99 14.16 15.77 15.68 15.23 14.64 Fe<sub>2</sub>O<sub>3</sub> 17.40 15.34 12.24 11.78 12.16 15.04 MnO 0.56 0.57 0.62 0.79 0.79 0.51 7.02 MgO 9.50 4.95 4.34 4.45 25.65 CaO 1.97 1.45 2.42 3.36 3.40 1.02 Na<sub>2</sub>O 5.70 6.79 6.32 5.58 5.81 0.49 $K_2O$ 0.13 0.19 2.12 3.64 3.27 0.04 $P_2O_5$ 0.70 0.25 0.18 0.25 0.15 0.19 LOI 3.40 2.25 1.95 1.90 1.70 9.45 Total (wt %) 99.83 99.43 99.17 98.24 98.35 100.40 Rb 13 6.8 96.8 157 153 0.7 Sr 50.1 71.6 119 152 169 10.3 Y 47.1 45.2 40.8 39.5 37.4 47.8 Zr\* 152 168 120 115 108 175 Nb\* 15 13 11 10 15 251 Ba 4.7 18.7 418 392 2.5 La\* 63.3 59.7 45.8 51.5 50.3 64.5 Ce\* 125 114 80.5 80 77.6 124 Pr\* 15.4 14.7 12.3 12.2 17 11 Nd\* 58.3 55.5 42.7 46.6 46.5 62.4 Sm\* 11.6 11.3 8.9 9.7 9.4 11.9 2.82 2.83 2.25 2.35 2.23 1.78 Eu\* 11.8 9.68 Gd\* 11.4 9.35 9.4 11.6 Tb\* 1.7 1.75 1.51 1.57 1.47 1.81 9.7 9.32 8.46 8.27 7.77 9.47 Dv\* 2.05 1.92 1.74 2.01 Ho\* 1.74 1.64 Er\* 5.59 5.57 4.77 4.71 4.36 5.39 Tm\* 0.87 0.79 0.71 0.67 0.62 0.82 Yb\* 5.3 5.2 4.4 4.2 4.2 5.3 0.78 0.81 0.69 0.64 0.61 0.73 Lu\* Th\* 19.6 17.5 13.2 12.6 12.5 14 U 0.9 1.48 2.01 2.15 2.06 0.96 Ph 59 2.5 13 40 66 297 281 238 247 184 V 247 Co 88.1 86.2 72.9 79.2 79.4 103 Cu 56 45 18 61 66 18 18 18 Ga 18 16 17 24 286 246 190 Ni 214 184 581 Cs 0.1 0.6 7.8 11.2 0.05 12

#### Breeding et al., Data Repository Table DR1

167 Note: Totals include major elements and Rb, Sr, Y, Zr, Nb, and Ba summed as oxides.

Hf\*

Zn

\*Mass balance calculations indicate that metapelite lost ~36% rock mass at the contact with mélange matrix during fluid alteration. Raw concentrations of the limited mobility elements Zr, Th, Nb, Ti, Hf, and the REE increase slightly toward the contact, indicating residual enrichment due to the change in total rock mass rather than mass transfer of those elements (cf. Ague and van Haren, 1996; Ague, 2003).

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