

## Data Repository item 2004028

### APPENDIX 1: ANALYTICAL METHODS

U/Pb geochronology was completed at The University of Texas at Austin. Samples were crushed and milled and zircons were separated by standard separation techniques using a Rogers<sup>TM</sup> or Wilfley<sup>TM</sup> table, heavy liquids and a Frantz<sup>TM</sup> Isodynamic Magnetic Separator (Krogh, 1982a). Minerals were hand picked from non-magnetic fractions (and occasionally low-angle magnetic fractions) to segregate distinct morphologies and select the best grains. Representative zircons were imaged by cathodoluminescence (CL) to determine internal morphologies (Connelly, 2000). Similar grains were chosen or partial grains were removed from CL mounts for analysis. All grains were abraded, cleaned and spiked with a <sup>205</sup>Pb/<sup>235</sup>U mixed spike prior to dissolution in HF and HNO<sub>3</sub> in Teflon<sup>TM</sup> dissolution bombs at 210°C (Krogh, 1973; Krogh and Davis, 1975; Krogh, 1982b; Parrish and Krogh, 1987). U and Pb were isolated from dissolved zircon by anion exchange techniques. Both U and Pb were loaded onto zone-refined Re filaments with a mixture of silica gel and phosphoric acid and analyzed with a Finnigan Mat 261<sup>TM</sup> thermal ionization mass spectrometer with 7 Faraday cups and a single ion-counting channel. Due to complex zircon populations, most zircon fractions were analyzed in peak jumping mode using only the ion counter, but a few initial zircon fractions were analyzed in static faraday-ion counter mode with <sup>204</sup>Pb measured by the ion counter. Repeated analyses of the NBS981 Pb standard and U500 U standard in peak jumping mode determined Pb and U fractionation to be 0.1% and 0.07% per amu, respectively. Blanks were typically less than 2 pg Pb and .2 pg U. Ages and errors were generated using an in-house program that incorporates an unpublished error propagation program (L. M. Heaman) and a linear regression calculation program (Davis, 1982). The 2-sigma error is represented by the size and orientation of ellipses on concordia diagrams.

Whole-rock samples were crushed and then pulverized in a Certiprep<sup>TM</sup> Shatterbox Puck Mill using an alumina ceramic grinding assembly. For Pb analysis, 35-100 mg aliquots of the pulverized powder were washed in 1N HNO<sub>3</sub> for 30 minutes on a warm hotplate before dissolution in Teflon<sup>TM</sup> dissolution bombs. Samples were dissolved in two stages using HF/HNO<sub>3</sub> and 6N HCl. Pb was isolated using an HBr anion exchange technique, loaded onto standard Re filaments with a mixture of silica gel and phosphoric acid, and analyzed with a Finnigan Mat 261<sup>TM</sup> thermal ionization mass spectrometer. Twenty-two analyses of NBS981 indicate fractionation of 0.1% per amu with measured isotopic ratios within 0.05%. Replicate analyses of similar sized aliquots of whole-rock powders were typically within 0.1%.

For Sm/Nd analyses, 35-100 mg aliquots of pulverized powder were spiked with <sup>149</sup>Sm/<sup>150</sup>Nd mixed spike and dissolved in the same manner as the Pb samples, except the powders were not washed in HNO<sub>3</sub>. Rare earth elements (REE) were isolated using REE-SPEC<sup>TM</sup> column chemistry. Sm and Nd were subsequently isolated using HDEHP columns. Analyses were normalized to <sup>146</sup>Nd/<sup>144</sup>Nd = .7219. Twelve analyses of Ames Nd yielded an average <sup>143</sup>Nd/<sup>144</sup>Nd of 0.512080+/-7. T<sub>DM</sub> ages are calculated using the theoretical isotopic decay constant (6.54\*10<sup>-12</sup>/yr) of Lugmair and Marti (1978) and the depleted mantle model of DePaolo (1981; E<sub>Nd</sub>=0.25T<sup>2</sup>-3T+8.5). Isochron regressions are determined using Isoplot (Ludwig, 2001).

## Data Repository item 2004028

### REFERENCES

- Connelly, J. N., 2000, Degree of preservation of igneous zonation in zircon as a signpost for concordancy in U/Pb geochronology: *Chemical Geology*, v. 172, p. 25-39.
- Davis, D. W., 1982, Optimum linear regression and error estimation applied to U-Pb data: *Canadian Journal of Earth Sciences*, v. 19, p. 2141-2149.
- DePaolo, D. J., 1981, A neodymium and strontium isotopic study of the Mesozoic calc-alkaline granitic batholiths of the Sierra Nevada and Peninsular Ranges, California: *Journal of Geophysical Research*, v. 86, no. B11, p. 10470-10488.
- Krogh, T. E., 1973, A low-contamination method for hydrothermal decomposition of zircon and extraction of U and Pb for isotopic age determinations: *Geochimica et Cosmochimica Acta*, v. 37, p. 485-494.
- Krogh, T. E., 1982a, Improved accuracy of U-Pb dating by selection of concordant fractions using a high gradient magnetic separation technique: *Geochimica et Cosmochimica Acta*, v. 46, p. 631-636.
- Krogh, T. E., 1982b, Improved accuracy of U-Pb zircon ages by the creation of more concordant systems using an air abrasion technique: *Geochimica et Cosmochimica Acta*, v. 46, p. 637-649.
- Krogh, T. E. and Davis, G. L., 1975, The production and preparation of  $^{205}\text{Pb}$  for use as a tracer for isotopic dilution analyses: *Carnegie Institute Washington, Yearbook*, v. 74, p. 619-623.
- Ludwig, K. R., 2001, Isoplot, Berkeley Geochronology Center.
- Lugmair, G. W. and Marti, K., 1978, Lunar initial  $^{143}\text{Nd}/^{144}\text{Nd}$ ; differential evolution of the lunar crust and mantle: *Earth and Planetary Science Letters*, v. 39, p. 349-357.
- Parrish, R. R. and Krogh, T. E., 1987, Synthesis and purification of  $^{205}\text{Pb}$  for U-Pb geochronology: *Chemical Geology*, v. 66, p. 103-110.

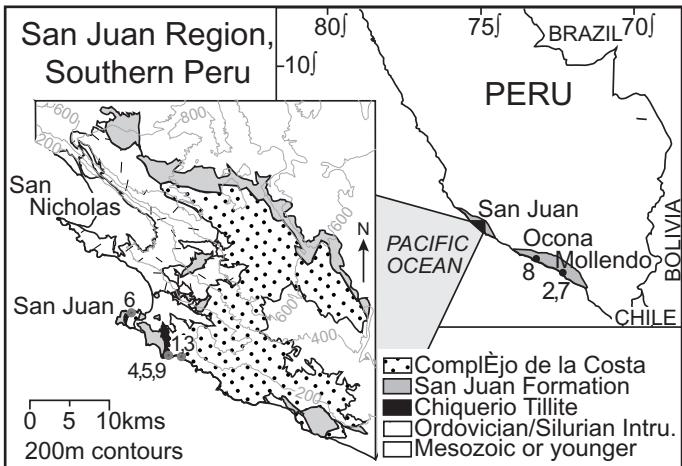


Fig. DR1. Map of southern Peru with coastal exposure of the AAB inlier (gray region) (INGEMMET, 2000). Inset is the geologic map of the San Juan region based on the mapping of Caldas (1978). Numbers indicate U/Pb sample locations.

[INGEMMET Republic del Peru Sector Energía and Minas, 2000, 501 Cuadrángulos Geológicos Digitales de la Carta Nacional 1960-1999.]

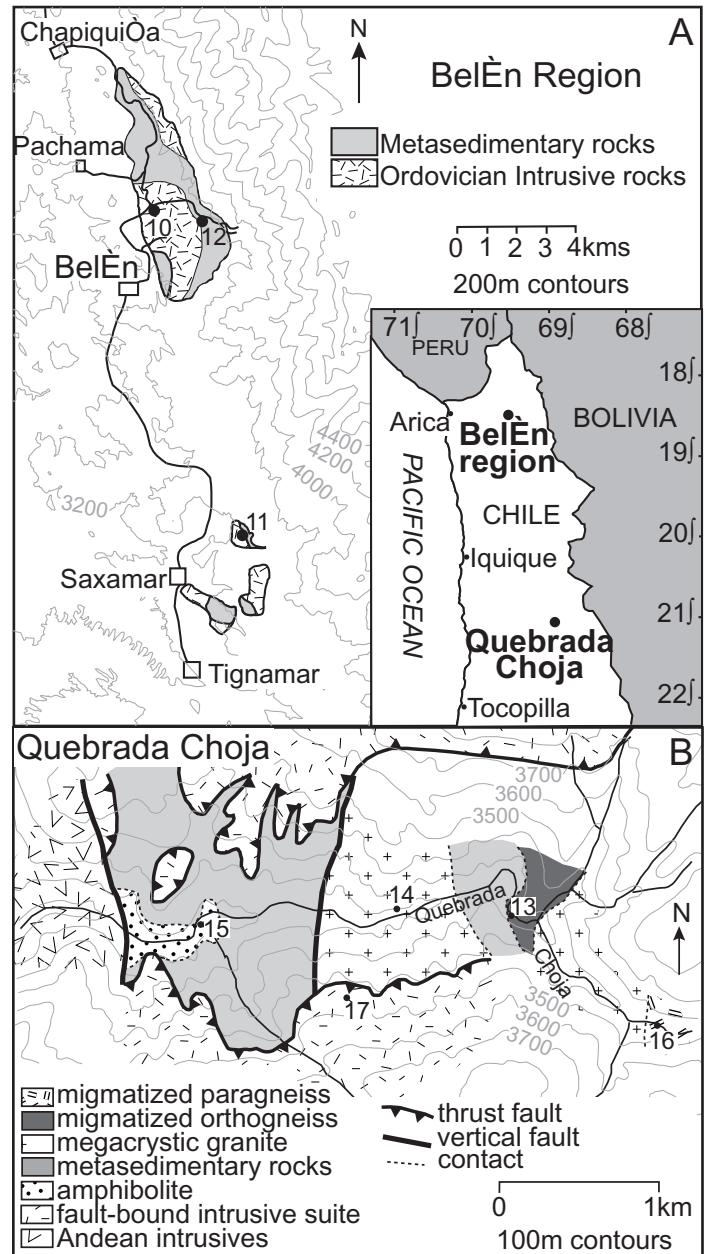


Fig. DR2. a) Map of BelEn area inliers. Numbers indicate U/Pb sample locations. Topographic basemaps are BelEn, Cerro Tejene, Pampa de Oxaya and Tignamar (1:50,000; Instituto Geográfico Militar de Chile). b) Geology of Quebrada Choja. Numbers indicate U/Pb sample locations. Topographic basemap is Quehuita (1:50,000; Instituto Geográfico Militar de Chile).











