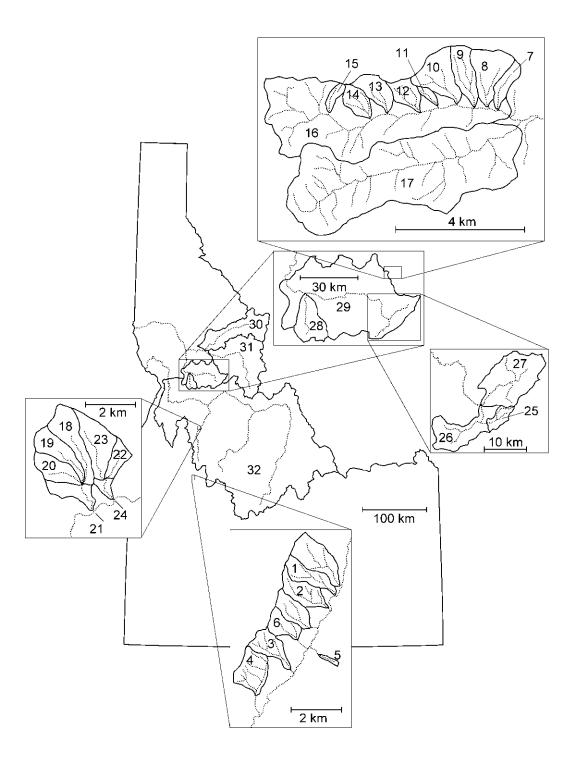
2001064 Kirchner et al., p. 1

GSA Data Repository Items for "Mountain erosion over 10-year, 10 000-year, and 10 000 000-year time scales", by James W. Kirchner, Robert C. Finkel, Clifford S. Riebe, Darryl E. Granger, James L. Clayton, John G. King, and Walter F. Megahan



**Figure A.** Study catchment boundaries (solid lines) and major streams (dotted lines), superimposed on outline map of the state of Idaho. Catchment identification numbers are keyed to Table A.

TABLE A. CATCHMENT CHARACTERISTICS AND SEDIMENT YIELDS

		Short-term rates				erm rates	Long-term (cosmogenic) rates				
				Drainage	Altitude	Glaciated	Record	Sediment	10Be	Time	Sediment
Cato	chment	Latitude	Longitude	area	range	area	length	yield⁺	concentration§	scale	yield <sup>#</sup>
ID*	Name			(km²)	(m)	(%)	(yr)	(T•km <sup>-2</sup> •yr <sup>-1</sup> )	(10⁵ atom•g⁻¹)	(yr)	(T•km <sup>-2</sup> •yr <sup>-1</sup> )
Silve	er Creek										
1	SC-2	44º22'20"	115°45'59"	1.2	1463-2073	0	27	$13.2 \pm 2.2$	$0.92 \pm 0.04$	5 100	$327 \pm 42$
2	SC-3	44º22'06"	115°46'12"	1.3	1451-2066	0	28	$8.9 \pm 1.4$	$1.61 \pm 0.06$	9 400	$174 \pm 23$
3	SC-5	44°20'46"	115°47'18"	1.1	1390-1772	0	28	$10.9 \pm 1.6$	$1.87 \pm 0.07$	12 000	$136 \pm 18$
4	SC-6	44º20'15"	115°48'20"	1.6	1381-1784	0	27	$9.3 \pm 1.7$	$1.72 \pm 0.11$	11 000	$152 \pm 22$
5	SC-7	44º21'01"	115°46'30"	0.23	1457-1819	0	22	$14.4 \pm 2.5$	$2.88 \pm 0.11$	17 000	$90 \pm 12$
6	SC-8	44º21'21"	115°47'02"	1.1	1415-1720	0	13	$30.0 \pm 10.6$	$2.06 \pm 0.09$	13 000	$121 \pm 16$
Hors	se Creek										
7	HC-2	45°59'37"	115°20'27"	0.57	1268-1744	0	10	$7.3 \pm 1.3$	$2.46\pm0.09$	16 000	$97 \pm 13$
8	HC-4	45°59'44"	115°20'39"	1.4	1280-1747	0	10	$3.5\pm0.6$	$2.71 \pm 0.11$	18 000	$89 \pm 12$
9	HC-6	45°59'37"	115°21'00"	1.0	1293-1726	0	15	$3.3 \pm 0.6$	$2.94 \pm 0.11$	19 000	$80 \pm 11$
10	HC-8	45°59'41"	115º21'26"	1.5	1317-1726	0	12	$11.0 \pm 3.0$	$2.64 \pm 0.13$	17 000	$90 \pm 13$
11	HC-9	45°59'37"	115º21'53"	0.23	1341-1646	0	10	$8.6 \pm 1.3$	$2.86 \pm 0.09$	19 000	$80 \pm 11$
12	HC-10	45°59'32"	115º22'19"	0.65	1354-1729	0	12	$9.9 \pm 2.4$	$2.58 \pm 0.09$	17 000	$92 \pm 13$
13	HC-12	45°59'30"	115°23'07"	0.83	1390-1804	0	14	$8.2 \pm 2.3$	$2.48 \pm 0.08$	16 000	$101 \pm 14$
14	HC-14	45°59'24"	115°23'30"	0.62	1415-1804	0	12	$7.5 \pm 2.3$	$3.09 \pm 0.15$	19 000	$80 \pm 12$
15	HC-16	45°59'27"	115º24'31"	0.21	1524-1768	0	13	$25.1 \pm 6.7$	$4.36 \pm 0.15$	27 000	$55 \pm 8$
16	West Fork	45°59'30"	115º20'10"	17	1250-1804	0	23	$5.0 \pm 0.5$	$2.72 \pm 0.07$	18 000	$87 \pm 12$
17	East Fork	45°59'20"	115°19'58"	14	1241-1835	0	23	$2.5 \pm 0.3$	$3.12 \pm 0.09$	20 000	$76 \pm 11$
Tailholt & Circle End Creeks											
18	Tailholt A	45°03'11"	115°40'54"	2.2	1256-2369	0	21	$11.0 \pm 2.5$	$1.18 \pm 0.07$	6 300	$264 \pm 36$
19	Tailholt B	45°03'10"	115°40'57"	1.6	1256-2141	0	22	$14.6 \pm 3.3$	$1.15 \pm 0.05$	6 400	$262 \pm 34$
20	Tailholt C	45°03'09"	115°41'00"	1.4	1256-2073	0	22	$13.7 \pm 2.4$	$1.41 \pm 0.05$	8 200	$202 \pm 26$
21	Tailholt Main	45°02'34"	115°40'38"	6.6	1091-2369	0	28	$14.0 \pm 2.8$	$1.22 \pm 0.06$	7 000	$239 \pm 32$
22	Circle End A	45°03'17"	115°40'13"	0.8	1296-2015	0	N.D.	N.D.	$1.23 \pm 0.05$	7 300	$226 \pm 29$
23	Circle End B	45°03'16"	115°40'17"	2.3	1296-2369	0	N.D.	N.D.	$1.26 \pm 0.05$	7 300	$229 \pm 30$
24	Circle End Main	45°02'50"	115°40'01"	3.8	1083-2369	0	25	$6.5 \pm 1.1$	$1.30 \pm 0.07$	7 700	$215 \pm 29$
Larg	er Streams and Rivers										
25	Trapper Creek	45°40'13"	115°19'47"	20	1476-2012	0	10	$9.8 \pm 1.6$	$4.67 \pm 0.12$	26 000	$57 \pm 8$
26	South Fk. Red River	45°42'35"	115°20'37"	98	1320-2170	0	14	$8.0 \pm 1.4$	$4.65 \pm 0.13$	25 000	$58 \pm 8$
27	Upper Red River	45°42'38"	115°20'34"	129	1320-2075	0	14	$10.1 \pm 1.6$	$3.03 \pm 0.10$	18 000	$87 \pm 12$
28	Johns Creek	45°49'23"	115°53'18"	293	735-2551	32	10	$7.6 \pm 1.3$	$2.60 \pm 0.09$	15 000	$108 \pm 15$
29	S. Fk. Clearwater Rvr	45°53'15"	116°01'47"	2 149	600-2551	7	25	$7.6 \pm 2.3$	$2.75 \pm 0.09$	17 000	$91 \pm 12$
30	Lochsa River	46°09'04"	115°35'37"	3 055	465-2680	21	72	$26.3 \pm 2.8$	$1.05 \pm 0.04$	6 700	$250 \pm 32$
31	Selway River	46°05'11"	115°30'59"	4 945	490-2709	21	70	$24.5 \pm 3.2$	$1.32 \pm 0.07$	8 100	$205 \pm 28$
32	Salmon River	45°44'38"	116º19'39"	35 079	460-3859	10	84	$13.7 \pm 4.1$	$1.47 \pm 0.08$	6 300	$261 \pm 36$

## **Notes for Table A**

\*Identification numbers are coded to map in Figure A.

<sup>†</sup>Sediment yields for catchments smaller than 20 km<sup>2</sup> were measured by trapping sediment behind small dams; the trapped sediment was measured and removed once or twice per year (Clayton and Megahan, 1986). Measurements of suspended sediment passing over the dams were used to correct for trap efficiency. Sediment yields for catchments larger than 20 km<sup>2</sup> were calculated from daily measurements of streamflow, scaled by bias-corrected (Ferguson, 1987) sediment rating curves derived from measurements of suspended sediment concentrations and bedload fluxes over a wide range of flows. Periods of record for sediment concentrations (and thus for construction of rating curves) were 10, 14, 14, 10, 5, 4, 4, and 16 years for catchments no. 25 through 32, respectively. Uncertainties (expressed as standard errors of means) were calculated from variability of annual fluxes. For catchments no. 32 and no. 29, bedload was estimated as 5% and 10% of total sediment flux, respectively, and standard errors were estimated at 30 percent of the mean. N.D. = not determined.

\$40-gram samples of quartz were purified from sand-sized sediment by magnetic separation and by acid leaching (Kohl and Nishiizumi, 1992) (which also eliminates meteoric "garden variety" 10Be), spiked with ~0.5 mg Be, and dissolved in HF. Be was separated by ion chromatography and analyzed as 10Be/Be by accelerator mass spectrometry at Lawrence Livermore National Laboratory.

\*Long-term sediment yields were calculated from equation 1. Area-averaged  $P_{\rm n}$  and  $P_{\rm m}$  values were calculated from sea-level high-latitude <sup>10</sup>Be production rates 4.72±0.38 atoms•g<sup>-1</sup>•yr<sup>-1</sup> and 0.11±0.01 atoms•g<sup>-1</sup>•yr<sup>-1</sup> respectively (Riebe et al., 2000; Riebe et al., 2001), scaled for latitude and for the altitude distribution within each catchment (Lal, 1991). Using Nishiizumi et al.'s (Nishiizumi et al., 1996) sea-level high-latitude production rate of 5.8 atoms•g<sup>-1</sup>•yr<sup>-1</sup> would increase calculated sediment yields by ~20 percent. Production rates were corrected for quartz enrichment by chemical weathering (Small et al., 1999), and for topographic shielding (Dunne et al., 1999). We corrected for shielding by snow cover using site-specific relationships between altitude and average snow water equivalent measured at nearby snow survey sites. Weathering fluxes were estimated at W=10.7±0.4 T•km<sup>-2</sup>•yr<sup>-1</sup>, based on 11-yr chemical mass balances at four Silver Creek catchments (Clayton and Megahan, 1986).

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## **APPENDIX A**

To test whether shallow debris flows have a measurable effect on cosmogenic nuclide concentrations, we sampled and analyzed sand-sized material at two points along the debris flow scar at Circle End Creek (samples no. 22 and no. 24 in Table A). These samples yielded cosmogenic nuclide concentrations and long-term erosion rates that were indistinguishable from those of an undisturbed Circle End tributary (catchment no. 23) and adjacent undisturbed Tailholt Creek (catchments 18-21). This indicates that cosmogenic nuclide measurements of long-term erosion rates are not substantially affected by recent debris-flow activity, because the time required for a typical sediment grain to be exhumed from a hillslope and transported to a small channel or hollow is much longer than the timescale of its subsequent storage there. Thus the cosmogenic nuclide signature is primarily acquired during hillslope exhumation and transport, with channel storage having little effect on either nuclide production or decay (the latter requiring storage at depths >1 m for timescales >1 Myr, which is implausible at our sites).