# Timing of arc construction and metamorphism in the Slate Creek Complex, northern Sierra Nevada, California, USA 

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## DATA REPOSITORY: SAMPLE DESCRIPTIONS AND ${ }^{40} \mathrm{Ar} /{ }^{39} \mathrm{Ar}$ ISOTOPIC RESULTS

Amphiboles were separated from eight rocks collected at seven localities in and adjacent to the Slate Creek Complex (SCC), and were used for ${ }^{40} \mathrm{Ar} /{ }^{39} \mathrm{Ar}$ geochronology to constrain the timing of volcanism and metamorphism of the SCC (Table DR-1). Two samples of relict volcanic hornblende were analyzed to determine the age of volcanism in the Lexington Hill volcanics. Amphiboles from a massive metadiorite, foliated metadiorite, and foliated metagabbro were analyzed in order to evaluate the age of metamorphic crystallization or cooling and to assess the effect of dynamic $v s$. static metamorphic crystallization on argon loss.

Hornblende from the Cascade pluton was analyzed to determine an igneous cooling age for this batholith. Two hornblende separates from unmetamorphosed dikes were analyzed to evaluate ages of rapid cooling of dikes injected into the SCC.

## Petrographic Descriptions

Relict volcanic hornblende (strictly, Mg-hastingsite; Tables DR-2 and DR-3; Fig. DR-1) was separated from two samples (F94-20 and F94-23) of subgreenschist to lowermost greenschist facies tuffs near the base of the Lexington Hill volcanics, the youngest unit in the SCC. Sample F94-20 is a weakly-foliated hornblende-plagioclase crystal tuff with minor lapillisized lithic fragments of volcanic origin. Textural domains include fine-grained matrix, relict plagioclase phenocrysts, relict hornblende phenocrysts, and veins. The matrix is composed of a
fine-grained assemblage of quartz, albite, phengite, chlorite, tremolitic amphibole, and epidote with minor titanite, apatite, and opaque. The matrix is cut by rare veins composed predominantly of chlorite and/or epidote. Volcanic plagioclase is replaced by aggregates of sericite, epidote, and albite. Relict hornblende consists of euhedral prismatic crystals with high relief and tan to olive-green pleochroism (Fig. DR-2A). Prisms up to 1 mm long have length to width ratios of approximately 5:1. Some of the phenocrysts are weakly fractured and exhibit minor alteration to chlorite. Epitaxial overgrowths of acicular actinolite are common, and quartz-rich pressure shadows occur adjacent to some hornblende phenocrysts.

Sample F94-23 also is a massive hornblende-plagioclase crystal tuff. This sample was collected near the base of the Lexington Hill volcanics approximately 200 meters stratigraphically above F94-20. Textural domains in F94-23 consist of matrix, relict plagioclase and hornblende phenocrysts, and thin veinlets of quartz. The matrix consists of fine-grained quartz, albite, phengite, and epidote, with minor chlorite, titanite, and opaques. Plagioclase is replaced completely by aggregates of sausserite, albite, and fine-grained quartz. Hornblende phenocrysts exhibit euhedral crystal margins, prismatic form, and olive-green pleochroism, but are coarser grained and more extensively fractured than the hornblende in F94-20 (Fig. DR-2). Epidote, actinolite, chlorite, and titanite are common in fractures within the hornblende phenocrysts.

Amphiboles were separated from massive dioritic (F94-13), foliated dioritic (F92-117), and foliated gabbroic (F92-36) samples from the Slate Creek metaplutonic unit. Whole-rock Xray fluorescence (XRF) analyses of (XRAL Laboratories, Don Mills, Ont., Canada) indicate $\mathrm{SiO}_{2}$ concentrations of approximately $42.5 \mathrm{wt} . \%$ in F92-36 and $57.1 \mathrm{wt} . \%$ in F94-13 (Fagan, 1997). No XRF analysis was obtained for F94-117, but estimated modal abundances indicate a dioritic composition intermediate between F92-36 and F94-13. Quartz, feldspar, amphibole, epidote,
chlorite, titanite, sericite, and opaques are common to all three samples. The only feldspar identified in F94-13 is albite; both albite and oligoclase are present in F92-36 and F94-117.

Multiple textural and compositional varieties of amphibole are present in the metaplutonic samples. Compositional varieties of amphibole include Mg -hornblende in all three samples, actinolite in F94-13, and pargasite in F92-36 (Tables DR-2 and DR-3; Fig. DR-1). Textural varieties of amphibole include coarse crystals ( $250 \mu \mathrm{~m}$ across or greater) with equant to stubby (length:width no greater than 5:1) prismatic form, and finer crystals (typically $50 \mu \mathrm{~m}$ across) with more elongate (length:width commonly 10:1) tapered to acicular form (Fig. DR-2C, D). These two textural varieties of amphibole form two distinct populations in F94-13, in which fine elongate amphiboles commonly protrude from, and coalesce to form serrated margins on, the coarse equant amphiboles. The elongate amphibole tends to be more Si-rich/Al-poor than the equant amphibole, but both textural varieties have similar hornblendic compositions (Fig. DR-1). Actinolite (Fig. DR-1) occurs in irregular seams and pockets within the equant hornblende, and comprises a minor fraction ( $<5 \%$ ) of the amphibole in F94-13.

Coarse, equant and fine, elongate amphiboles are present in F92-36 and F94-117, but many amphiboles in these samples have textures intermediate between the two extremes. The amphiboles in F94-117 are slightly more tschermakitic than the F94-13 hornblendes, but exhibit the same intra-sample correspondence between texture and morphology: namely, elongate amphibole tends to be Si-rich and Al-poor relative to the equant amphibole (Fig. DR-1). Unlike F94-13, no actinolite was identified in F94-117. In F92-36, most of the coarse equant amphibole is pargasite identical in composition to the fine elongate amphiboles (Fig. DR-1). Some minor patches of Mg-hornblende (Fig. DR-1) occur within the interiors of coarse equant amphiboles in F92-36.

The three metaplutonic samples have undergone different extents of deformation during metamorphism. Relict plutonic textures are preserved best in F94-13, which is massive, has a high ratio of equant:elongate amphibole, and is characterized by tabular feldspar with variable concentrations of sericite and granular epidote. Sample F92-36 exhibits gneissic foliation with layers rich in amphibole and chlorite alternating with layers rich in sausseritized feldspar. The original plagioclase has been altered to aggregates of epidote and sericite with some albite and oligoclase. Most of the amphibole in F92-36 consists of the elongate variety, although some equant amphibole is also present (Fig. DR-2C). The compositional layering in F92-36 may reflect original cumulate layering, but if so, cumulus/intercumulus textures between minerals have been totally destroyed during metamorphism. Sample F94-117 has a moderate foliation defined by parallel orientation of amphibole. The original igneous feldspar was replaced completely by clear oligoclase and cloudy albite with some sericite. Both coarse and fine textural varieties of amphibole are common.

One amphibole separate was collected from a sample (F94-10) of granodiorite from the cross-cutting Cascade pluton. The sample exhibits a moderate foliation defined by parallel orientation of igneous biotite and hornblende. Equant quartz and tabular plagioclase, occasionally with sausseritized cores, are the most abundant felsic minerals. Alkali feldspar occurs as individual crystals and in myrmekitic intergrowths with quartz. Epidote and titanite occur in trace amounts, and the coarse grain size of titanite is distinctive, both in thin section and outcrop. Brown biotite and blue-green amphibole comprise the mafic phases. Amphibole occurs as subhedral prisms up to 1 mm long with typical length:width ratios of 5:2 (Fig. DR-2E) and Mg-hornblende composition (Table DR-3; Fig. DR-1).

Two amphibole separates were collected from dikes that cross-cut units of the Slate Creek Complex. F92-46 is from a dike that cross-cuts metadiorite of the Slate Creek
metaplutonic unit. The interior of the dike consists of hornblende and clinopyroxene phenocrysts in a groundmass consisting of feldspar and hornblende with minor opaque sulfides. Several phenocrysts in the chilled margin have been altered to chlorite. Chlorite and epidote-rich aggregates also occur in the dike interior, suggesting a patchy, incomplete alteration pattern. Amphibole in F92-46 is euhedral pargasite (Table DR-3; Fig. DR-1) with elongate (length:width of approximately 10:1) prismatic to gently tapered prismatic form and brown pleochroism. Tiny epitaxial overgrowths of actinolite are rare (Fig. DR-1).

F94-54 is from a hornblende-phyric dike that cross-cuts clastic metasedimentary rocks in the Gold Run formation. Hornblende (Mg-hastingsite; see Table DR-3 and Fig. DR-1) is the sole phenocryst in this dike. Hornblende also occurs in the matrix, along with plagioclase, quartz, and opaques. Minor concentrations of chlorite, titanite, and carbonate appear to be secondary. Amphibole in this sample exhibits brown pleochroism and euhedral elongate prismatic crystal form (Fig. DR-2F). Length to width ratios range as high as 20:1. Patchy replacement to chlorite occupies up to approximately $3 \%$ of some amphibole prisms.

## Electron Microprobe Analytical Methods

Amphibole compositions were determined by wavelength dispersive analyses on Cameca SX-50 electron microprobes at the University of Calfornia at Davis (UCD) and the University of Hawaii at Manoa (UH). Analyses were normalized to 23 oxygen with site-assignments and ferrous:ferric iron ratios determined by normalization of all cations less $\mathrm{Na}, \mathrm{K}$, and Ca to 13 (13eNKC), all cations less K to $15(15 \mathrm{eK})$, or all cations less Na and K to 15 (15eNK) (Leake and others, 1997). Metamorphic amphiboles were normalized to either 13 eNKC or 15 eK in accordance with the criteria of Laird and Albee (Laird and Albee, 1981); similar normalization of
some of the igneous amphibole analyses resulted in suspiciously high ferric iron concentrations, thus all the igneous amphiboles were normalized to 15 eNK .

Analyses were collected with a 15 keV accelerating voltage, 10 nanoamp current, and 1 micron spot size with well characterized oxide and silicate standards on both electron microprobes. Manufacturer supplied ZAF (UCD) and PAP (UH) programs were used to correct for interferences. Peak and background counting times, along with lower limits of detection and relative counting errors for both microprobes are listed in Table DR-4. The tabulated detection limits are twice the statistical lower limits of detection. Longer counting times for the UH analyses resulted in a small reduction in counting error for most elements; counting errors and detection limits for F and Na were significantly better at UH because the PC0 diffracting crystal resulted in higher count rates than the TAP crystal used at UCD. Interferences of Fe L $\alpha$ peaks with the $\mathrm{F} \mathrm{K} \alpha$ peak in data from the PC 0 diffracting crystal were calibrated and calculated using CAMECA software.

Four compositionally distinct types of amphibole in three thin sections from the SCC were analyzed at both UCD and UH to assess reproducibility of data collected on the two microprobes. Between 5 and 95 analyses were collected from each of the four varieties of amphibole at both laboratories, and the mean results compared. Mean atoms per formula unit (apfu) of all elements, including $\mathrm{Fe}^{2+}$ and $\mathrm{Fe}^{3+}$, from the four replicate analyses at UH are within 2 standard deviations of the means determined at UCD (using the UCD standard deviations). The standard deviations at UH are generally lower than the UCD standard deviations, probably as a consequence of fewer analyses and a more incomplete sampling of compositional heterogeneity within samples. Still, $96 \%$, or all but two (one Na and one $\mathrm{Fe}^{3+}$ value), of the UCD mean apfu from the four replicate analyses fall within two UH standard deviations of the mean UH apfu.

In terms of absolute concentration, for elements present in concentrations less than 0.5 apfu, means from the UH microprobe duplicate the UCD means to within 0.02 apfu. For elements (including $\mathrm{Fe}^{\mathrm{T}}=\mathrm{Fe}^{2+}+\mathrm{Fe}^{3+}$ ) present in concentrations greater than 0.5 apfu, mean results from the two microprobes are within 0.08 apfu for all paramaters except for one $\mathrm{Fe}^{\mathrm{T}}$ and two Si analyses. The three exceptions include a discrepancy of $0.09 \mathrm{apfu}(6.5 \%$ relative) in one replicate analysis of $\mathrm{Fe}^{\mathrm{T}}$, and discrepancies of 0.13 and 0.10 apfu ( $1.7 \%$ and $1.3 \%$ relative) in two replicate analyses of Si . Estimates of ferrous and ferric iron reflect errors in analyses of all elements, and show a wider range of mismatch between results collected on the two microprobes. The amphibole analyses at UH yield mean ferric and ferrous iron values that are typically within 0.15 apfu of the means determined at UCD.

## ${ }^{40} \mathrm{Ar} /{ }^{39} \mathrm{Ar}$ Isotopic Results

The amphibole separates were analyzed by step heating in a resistance furnace using procedures described in (Hacker, et al., 1996). Results for each heating increment are documented in Table DR-5. Step-heating profiles are presented in the text.

## REFERENCES

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TABLE DR-1. SAMPLES COLLECTED FOR ${ }^{40} \mathrm{Ar}{ }^{39} \mathrm{Ar}$ ANALYSES

| sample | latitude | longitude | 7.5' quadrangle | geologic unit | outcrop description | lithology |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| F92-36 | 39³6'49" | 12103'11" | Strawberry Valley | Slate Creek complex metaplutonic rocks. | Stream polished right bank of Slate Creek 300 meters downstream from tailings dam. | Well-foliated metagabbro with quartz and epidote veins. Foliation defined by amphibole orientation, and amphibole vs. epidote (replacing feldspar)-rich compositional layering. Epidote-amphibolite facies. |
| F92-46 | $39^{\circ} 36^{\prime} 54 "$ | $121^{\circ} 03^{\prime} 04 "$ | Strawberry Valley | Dike intrusive into Slate Creek complex metaplutonic rocks. | Stream polished right bank of Slate Creek at base of tailings dam. | Hornblende-clinopyroxene-phyric dike intrusive into massive metadiorite Well developed chilled margin. Igneous. |
| F94-10 | $39^{\circ} 34^{\prime} 50 "$ | $121^{\circ} 06^{\prime} 46{ }^{\prime \prime}$ | Strawberry Valley | Cascade pluton. Late-stage crosscutting pluton. | Road cut adjacent to spillway at Sly Creek Reservoir Dam. | Moderately to well foliated coarse biotite-hornblende-granodiorite with coarse-grained epidote and titanite. Igneous. |
| F94-13 | $39^{\circ} 36^{\prime} 54 "$ | $121^{\circ} 03^{\prime} 04 "$ | Strawberry Valley | Slate Creek complex metaplutonic rocks. | Same as F92-46. | Massive metadiorite with salt-and-pepper texture. Epidote-amphibolite facies. |
| F94-20 | 3943'32" | $121^{\circ} 01^{\prime} 37^{\prime \prime}$ | American House | Slate Creek complex Lexington Hill volcanics. | Stream polished right bank of South Fork Feather River. | Massive metamorphosed hornblende-plagioclase lapilli tuff. <br> Lowermost greenschist facies. |
| F94-23 | $39^{\circ} 43^{\prime} 30^{\prime \prime}$ | $121^{\circ} 01^{\prime} 26{ }^{\prime \prime}$ | American House | Slate Creek complex Lexington Hill volcanics. | Stream polished right bank of South Fork Feather River. | Massive metamorphosed hornblende-plagioclase crystal tuff. <br> Lowermost greenschist facies. |
| F94-54 | $39^{\circ} 38^{\prime \prime} 16^{\prime \prime}$ | $121^{\circ} 01^{\prime} 48^{\prime \prime}$ | American House | Dike intusive into Slate Creek complex Gold Run formation. | Stream polished outcrop on right bank of Slate Creek. | Hornblende-phyric dike with well-developed chilled margin, intrusive into fine metasandstone and metasiltstone. Igneous. |
| F94-117 | $39^{\circ} 38^{\prime} 46^{\prime \prime}$ | $121^{\circ} 03^{\prime} 55^{\prime \prime}$ | American House | Slate Creek complex metaplutonic rocks. | Stream polished left bank of Lost Creek. | Moderately-foliated metadiorite. Foliation defined by parallel orientation of amphibole. Epidote-amphibolite facies. |

TABLE DR-2. REPRESENTATIVE ELECTRON MICROPROBE ANALYSES OF AMPHIBOLES

| Sample <br> Rock type | F92-36foliated metagabbro |  |  | F92-46 <br> mafic dike |  | F94-10 <br> granodiorite plutonic | F94-13 massive metadiorite |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  |  |  |  |  |  |  |
| Amphibole type | equant | equant <br> Al-poor | elongate | igneous | epitaxial |  | equant | equant <br> Al-poor | elongate |
| Laboratory | UH | UH | UH | UH | UCD | UCD | UH | UH | UH |
| Normalization | 13eNKC | 13eNKC | 13eNKC | 15eNK | 15 eK | 15 eNK | 13eNKC | 13eNKC | 13eNKC |
| $\mathrm{SiO}_{2}$ | 39.68 | 46.10 | 39.71 | 42.29 | 54.49 | 46.24 | 48.44 | 54.11 | 48.72 |
| $\mathrm{TiO}_{2}$ | 0.47 | 0.70 | 0.57 | 2.30 | <0.16 | 0.87 | 1.01 | 0.31 | 0.23 |
| $\mathrm{Al}_{2} \mathrm{O}_{3}$ | 17.45 | 8.75 | 18.02 | 12.56 | 1.07 | 8.64 | 6.71 | 2.24 | 7.76 |
| $\mathrm{Cr}_{2} \mathrm{O}_{3}$ | <0.17 | <0.17 | <0.17 | <0.17 | <0.18 | <0.18 | 0.17 | <0.17 | <0.17 |
| FeO | 16.71 | 13.72 | 17.00 | 8.21 | 12.07 | 16.33 | 14.61 | 11.2 | 14.74 |
| MnO | 0.43 | 0.43 | 0.41 | <0.12 | 0.34 | 0.56 | 0.52 | 0.40 | 0.31 |
| MgO | 8.20 | 12.96 | 7.73 | 16.53 | 16.65 | 12.62 | 13.73 | 17.10 | 12.82 |
| CaO | 11.29 | 11.31 | 11.13 | 12.03 | 11.92 | 11.45 | 11.14 | 12.17 | 11.78 |
| $\mathrm{Na}_{2} \mathrm{O}$ | 2.99 | 2.40 | 2.94 | 2.51 | 0.20 | 1.08 | 1.13 | 0.33 | 0.95 |
| $\mathrm{K}_{2} \mathrm{O}$ | 0.51 | 0.08 | 0.53 | 0.79 | <0.07 | 0.70 | 0.30 | 0.12 | 0.16 |
| F | 0.17 | <0.12 | <0.12 | 0.32 | <0.58 | <0.58 | <0.12 | <0.12 | <0.12 |
| Cl | <0.17 | <0.17 | <0.17 | <0.17 | <0.08 | <0.08 | 0.21 | <0.17 | <0.17 |
| Sum | 97.9 | 96.45 | 98.04 | 97.54 | 96.74 | 98.49 | 97.97 | 97.98 | 97.47 |
| "Excess O" | 0.07 | B.D. | B.D. | 0.13 | B.D. | B.D. | 0.05 | B.D. | B.D. |
| Sum | 97.83 | 96.45 | 98.04 | 97.41 | 96.74 | 98.49 | 97.92 | 97.98 | 97.47 |
| $\mathrm{H}_{2} \mathrm{O}$ | 1.91 | 2.01 | 2.00 | 1.90 | 2.08 | 2.03 | 1.99 | 2.10 | 2.05 |
| Sum | 99.74 | 98.46 | 100.04 | 99.31 | 98.82 | 100.52 | 99.91 | 100.08 | 99.52 |
| Cations normalized to 23 oxygen |  |  |  |  |  |  |  |  |  |
| Si | 5.92 | 6.80 | 5.90 | 6.15 | 7.84 | 6.78 | 6.97 | 7.63 | 7.05 |
| Alv | 2.08 | 1.20 | 2.10 | 1.85 | 0.16 | 1.22 | 1.03 | 0.37 | 0.95 |
| $\mathrm{Fe}^{3+1 \mathrm{~V}}$ | 0.00 | 0.00 | 0.00 | 0.00 | 0.00 | 0.00 | 0.00 | 0.00 | 0.00 |
| Sum | 8.00 | 8.00 | 8.00 | 8.00 | 8.00 | 8.00 | 8.00 | 8.00 | 8.00 |
| $\mathrm{Al}^{\mathrm{VI}}$ | 0.99 | 0.32 | 1.06 | 0.30 | 0.03 | 0.27 | 0.11 | 0.00 | 0.38 |
| $\mathrm{Fe}^{3+\mathrm{VI}}$ | 0.42 | 0.44 | 0.42 | 0.20 | 0.16 | 0.32 | 0.88 | 0.51 | 0.56 |
| Ti | 0.05 | 0.08 | 0.06 | 0.25 | B.D. | 0.10 | 0.11 | 0.03 | 0.02 |
| Cr | B.D. | B.D. | B.D. | B.D. | B.D. | B.D. | 0.02 | B.D. | B.D. |
| Mg | 1.82 | 2.85 | 1.71 | 3.58 | 3.57 | 2.76 | 2.94 | 3.60 | 2.77 |
| $\mathrm{Fe}^{2+\mathrm{VII}}$ | 1.66 | 1.25 | 1.69 | 0.67 | 1.23 | 1.56 | 0.87 | 0.82 | 1.22 |
| $\mathrm{Mn}{ }^{\mathrm{VI}}$ | 0.05 | 0.05 | 0.05 | B.D. | 0.00 | 0.00 | 0.06 | 0.05 | 0.04 |
| Sum | 5.00 | 5.00 | 5.00 | 5.00 | 5.00 | 5.00 | 5.00 | 5.00 | 5.00 |
| $\mathrm{Fe}^{2+\mathrm{M} 4}$ | 0.00 | 0.00 | 0.00 | 0.13 | 0.06 | 0.13 | 0.00 | 0.00 | 0.00 |
| $\mathrm{Mn}^{\text {M4 }}$ | 0.00 | 0.00 | 0.00 | B.D. | 0.04 | 0.07 | 0.00 | 0.00 | 0.00 |
| Ca | 1.80 | 1.79 | 1.77 | 1.87 | 1.84 | 1.80 | 1.72 | 1.84 | 1.83 |
| $\mathrm{Na}^{\text {M4 }}$ | 0.20 | 0.21 | 0.23 | 0.00 | 0.06 | 0.00 | 0.28 | 0.09 | 0.17 |
| Sum | 2.00 | 2.00 | 2.00 | 2.00 | 2.00 | 2.00 | 2.00 | 1.93 | 2.00 |
| $\mathrm{Na}^{\text {A }}$ | 0.67 | 0.47 | 0.62 | 0.71 | 0.00 | 0.31 | 0.03 | 0.00 | 0.09 |
| K | 0.10 | 0.02 | 0.10 | 0.15 | 0.00 | 0.13 | 0.06 | 0.02 | 0.03 |
| Sum | 0.77 | 0.49 | 0.72 | 0.85 | 0.00 | 0.44 | 0.09 | 0.02 | 0.12 |
| F | 0.08 | B.D. | B.D. | 0.15 | B.D. | B.D. | B.D. | B.D. | B.D. |
| Cl | B.D. | B.D. | B.D. | B.D. | B.D. | B.D. | 0.05 | B.D. | B.D. |
| Cations | 15.77 | 15.49 | 15.72 | 15.85 | 15.00 | 15.44 | 15.09 | 14.95 | 15.12 |
| $\mathrm{Mg} /(\mathrm{Mg}+\mathrm{Fe})$ | 0.52 | 0.69 | 0.50 | 0.82 | 0.73 | 0.62 | 0.77 | 0.82 | 0.69 |

Note: Analyses in weight percent; all Fe analyzed as FeO ; $\mathrm{H}_{2} \mathrm{O}$ calculated from stoichiometry; UCD $=$ analysis at
University of California at Davis, UH = analysis at University of Hawaii; normalization and determination of ferrous:ferric iron ratios discussed in Electron Microprobe Methods.

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revised for GSABulletin
Table DR-2
MS \#21828 - data repository arc construction and metamorphism, Sierra Nevada, CA
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| Sample <br> Rock type | F94-20 <br> metatuff |  | F94-23 <br> metatuff | F94-54 <br> mafic | F94-117 <br> foliated metadiorite |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Amphibole type | relict volcanic | epitaxial | relict volcanic | igneous | elongate | equant |
| Laboratory | UCD | UCD | UCD | UCD | UH | UH |
| Normalization | 15 eNK | 13eNKC | 15eNK | 15eNK | 13eNKC | 13eNKC |
| $\mathrm{SiO}_{2}$ | 42.28 | 54.22 | 42.75 | 40.21 | 44.33 | 47.13 |
| $\mathrm{TiO}_{2}$ | 1.62 | <0.16 | 1.84 | 3.14 | 0.37 | 0.37 |
| $\mathrm{Al}_{2} \mathrm{O}_{3}$ | 11.51 | 1.80 | 12.90 | 13.14 | 11.98 | 9.86 |
| $\mathrm{Cr}_{2} \mathrm{O}_{3}$ | <0.17 | <0.18 | <0.17 | <0.17 | <0.17 | <0.17 |
| FeO | 13.11 | 13.77 | 12.08 | 10.21 | 15.45 | 15.65 |
| MnO | 0.25 | 0.44 | 0.30 | B.D. | 0.38 | 0.35 |
| MgO | 14.28 | 15.71 | 14.92 | 15.62 | 11.32 | 12.58 |
| CaO | 11.80 | 12.95 | 11.83 | 11.98 | 11.09 | 11.14 |
| $\mathrm{Na}_{2} \mathrm{O}$ | 2.42 | 0.16 | 2.27 | 2.44 | 1.65 | 1.41 |
| $\mathrm{K}_{2} \mathrm{O}$ | 0.81 | <0.07 | 0.75 | 0.73 | 0.21 | 0.16 |
| F | <0.58 | <0.58 | <0.58 | <0.58 | <0.12 | <0.12 |
| Cl | <0.08 | <0.08 | <0.08 | <0.08 | <0.17 | <0.17 |
| Sum | 98.08 | 99.05 | 99.64 | 97.47 | 96.78 | 98.65 |
| "Excess O" | B.D. | B.D. | B.D. | B.D. | B.D. | B.D. |
| Sum | 98.08 | 99.05 | 99.64 | 97.47 | 96.78 | 98.65 |
| $\mathrm{H}_{2} \mathrm{O}$ | 2.02 | 2.10 | 2.07 | 2.03 | 2.01 | 2.06 |
| Sum | 100.10 | 101.15 | 101.71 | 99.50 | 98.79 | 100.71 |

Cations normalized to 23 oxygen

| Si | 6.21 | 7.69 | 6.13 | 5.87 | 6.49 | 6.72 |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| $\mathrm{Al}^{\text {lV }}$ | 1.79 | 0.30 | 1.87 | 2.13 | 1.51 | 1.28 |
| $\mathrm{Fe}^{3+1 \mathrm{~V}}$ | 0.00 | 0.01 | 0.00 | 0.00 | 0.00 | 0.00 |
| Sum | 8.00 | 8.00 | 8.00 | 8.00 | 8.00 | 8.00 |
| $\mathrm{Al}^{\mathrm{VI}}$ | 0.20 | 0.00 | 0.31 | 0.13 | 0.55 | 0.38 |
| $\mathrm{Fe}^{3+\mathrm{VI}}$ | 0.40 | 0.33 | 0.40 | 0.48 | 0.90 | 1.00 |
| Ti | 0.18 | B.D. | 0.20 | 0.34 | 0.04 | 0.04 |
| Cr | B.D. | B.D. | B.D. | B.D. | B.D. | B.D. |
| Mg | 3.13 | 3.32 | 3.19 | 3.40 | 2.47 | 2.67 |
| $\mathrm{Fe}^{2+\mathrm{VI}}$ | 1.10 | 1.30 | 0.91 | 0.64 | 1.00 | 0.87 |
| $\mathrm{Mn}^{\mathrm{VI}}$ | 0.00 | 0.05 | 0.00 | B.D. | 0.05 | 0.04 |
| Sum | 5.00 | 5.00 | 5.00 | 5.00 | 5.00 | 5.00 |
| $\mathrm{Fe}^{2+\mathrm{M} 4}$ | 0.11 | 0.00 | 0.15 | 0.13 | 0.00 | 0.00 |
| $\mathrm{Mn}^{\text {M4 }}$ | 0.03 | 0.00 | 0.04 | B.D. | 0.00 | 0.00 |
| Ca | 1.86 | 1.97 | 1.82 | 1.87 | 1.74 | 1.70 |
| $\mathrm{Na}^{\text {M4 }}$ | 0.00 | 0.03 | 0.00 | 0.00 | 0.26 | 0.30 |
| Sum | 2.00 | 2.00 | 2.00 | 2.00 | 2.00 | 2.00 |
| $\mathrm{Na}^{\text {A }}$ | 0.69 | 0.01 | 0.63 | 0.69 | 0.21 | 0.09 |
| K | 0.15 | B.D. | 0.14 | 0.14 | 0.04 | 0.03 |
| Sum | 0.84 | 0.01 | 0.77 | 0.83 | 0.25 | 0.12 |
| F | B.D. | B.D. | B.D. | B.D. | B.D. | B.D. |
| Cl | B.D. | B.D. | B.D. | B.D. | B.D. | B.D. |
| Cations | 15.84 | 15.01 | 15.77 | 15.83 | 15.25 | 15.12 |
| $\mathrm{Mg} /(\mathrm{Mg}+\mathrm{Fe})$ | 0.72 | 0.72 | 0.75 | 0.82 | 0.71 | 0.76 |

Note: Analyses in weight percent; all Fe analyzed as $\mathrm{FeO} ; \mathrm{H}_{2} \mathrm{O}$ calculated from stoichiometry; UCD = analysis at University of California at Davis, UH = analysis at University of Hawaii; normalization and determination of ferrous:ferric iron ratios discussed in Electron Microprobe Methods.

TABLE DR-3. MEAN (STANDARD DEVIATION) $\mathrm{Mg} /(\mathrm{Mg}+\mathrm{Fe}), \mathrm{Si}, \mathrm{K} 2 \mathrm{O}$, AND $0.5^{*} \mathrm{~K} / \mathrm{Ca}$ OF AMPHIBOLES

| Sample | Amphibole type | n | Microprobe | Normalization | $\begin{gathered} \mathrm{Mg} / \\ (\mathrm{Mg}+\mathrm{Fe}) \end{gathered}$ | Si | $\mathrm{K}_{2} \mathrm{O}$ | $0.5 * \mathrm{~K} / \mathrm{Ca}$ | Classification* |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Lowermost greenschist facies metatuffs |  |  |  |  |  |  |  |  |  |
| F94-20 | relict volcanic | 27 | UCD | 15eNK | $\begin{gathered} 0.74 \\ (0.04) \end{gathered}$ | $\begin{gathered} 6.17 \\ (0.10) \end{gathered}$ | $\begin{gathered} 0.79 \\ (0.05) \end{gathered}$ | 0.042 (0.009) | Mg-hast |
| F94-20 | epitaxial | 7 | UCD | 13eNKC | $\begin{gathered} 0.73 \\ (0.03) \end{gathered}$ | $\begin{gathered} 7.64 \\ (0.10) \end{gathered}$ | $\begin{gathered} 0.07 \\ (0.06) \end{gathered}$ | 0.003 (0.003) | actinolite |
| F94-23 | relict volcanic | 12 | UCD | 15 eNK | $\begin{gathered} 0.74 \\ (0.01) \end{gathered}$ | 6.25 (0.16) | $\begin{gathered} 0.77 \\ (0.05) \end{gathered}$ | 0.039 (0.002) | Mg-hast |
| Epidote-amphibolite facies metadiorites |  |  |  |  |  |  |  |  |  |
| F92-36 | elongate | 27 | UH | 13eNKC | $\begin{gathered} 0.51 \\ (0.01) \end{gathered}$ | $\begin{gathered} 5.88 \\ (0.06) \end{gathered}$ | $\begin{gathered} 0.47 \\ (0.04) \end{gathered}$ | 0.025 (0.002) | pargasite |
| F92-36 | equant | 19 | UH | 13eNKC | $\begin{gathered} 0.51 \\ (0.02) \end{gathered}$ | 5.90 (0.07) | $\begin{gathered} 0.51 \\ (0.05) \end{gathered}$ | 0.027 (0.002) | pargasite |
| F92-36 | equant, AI- <br> poor | 7 | UH | 13eNKC | $\begin{gathered} 0.84 \\ (0.07) \end{gathered}$ | $\begin{gathered} 6.73 \\ (0.07) \end{gathered}$ | $\begin{gathered} 0.09 \\ (0.01) \end{gathered}$ | 0.005 (0.001) | Mg-hbl |
| F94-13 | elongate | 10 | UH | 13eNKC | $\begin{gathered} 0.70 \\ (0.03) \end{gathered}$ | $\begin{gathered} 7.06 \\ (0.19) \end{gathered}$ | $\begin{gathered} 0.15 \\ (0.04) \end{gathered}$ | 0.008 (0.002) | Mg-hbl |
| F94-13 | equant | 71 | UH | 13eNKC | $\begin{gathered} 0.72 \\ (0.07) \end{gathered}$ | 6.8 (0.2) | $\begin{gathered} 0.29 \\ (0.08) \end{gathered}$ | 0.016 (0.004) | Mg-hbl |
| F94-13 | equant, AI- <br> poor | 14 | UH | 15 eK | $\begin{gathered} 0.76 \\ (0.01) \end{gathered}$ | $\begin{gathered} 7.71 \\ (0.12) \end{gathered}$ | $\begin{gathered} 0.06 \\ (0.03) \end{gathered}$ | 0.003 (0.002) | actinolite |
| F94-117 | elongate | 41 | UH | 13eNKC | $\begin{gathered} 0.71 \\ (0.04) \end{gathered}$ | $\begin{gathered} 6.48 \\ (0.14) \end{gathered}$ | $\begin{gathered} 0.21 \\ (0.05) \end{gathered}$ | 0.012 (0.003) | tscher |
| F94-117 | equant | 56 | UH | 13eNKC | $\begin{gathered} 0.75 \\ (0.07) \end{gathered}$ | 6.7 (0.3) | $\begin{gathered} 0.17 \\ (0.07) \end{gathered}$ | 0.009 (0.004) | Mg-hbl |
| Dikes |  |  |  |  |  |  |  |  |  |
| F92-46 | igneous | 18 | UH | 15eNK | 0.80 (0.02) | $\begin{gathered} 6.15 \\ (0.11) \end{gathered}$ | $\begin{gathered} 0.79 \\ (0.05) \end{gathered}$ | $\begin{gathered} 0.040 \\ (0.003) \end{gathered}$ | pargasite |
| F92-46 | epitaxial | 7 | UCD | 15 eK | $\begin{gathered} 0.75 \\ (0.02) \end{gathered}$ | $\begin{gathered} 7.78 \\ (0.10) \end{gathered}$ | $\begin{gathered} 0.07 \\ (0.05) \end{gathered}$ | $\begin{gathered} 0.004 \\ (0.003) \end{gathered}$ | actinolite |
| F94-54 | igneous | 62 | UCD | 15 eNK | $\begin{gathered} 0.78 \\ (0.05) \end{gathered}$ | $\begin{gathered} 5.89 \\ (0.13) \end{gathered}$ | $\begin{gathered} 0.67 \\ (0.05) \end{gathered}$ | $\begin{gathered} 0.033 \\ (0.003) \end{gathered}$ | Mg-hast |
| Cascade pluton |  |  |  |  |  |  |  |  |  |
| F94-10 | plutonic | 18 | UCD | 15eNK | $\begin{gathered} 0.63 \\ (0.02) \end{gathered}$ | $\begin{gathered} 6.79 \\ (0.14) \end{gathered}$ | $\begin{gathered} 0.73 \\ (0.13) \end{gathered}$ | $\begin{gathered} 0.039 \\ (0.007) \end{gathered}$ | Mg-hbl |

Note: Mg , $\mathrm{Fe}, \mathrm{Si}, \mathrm{K}$, and Ca are atoms per formula unit on a 23-oxygen basis. Fe is ferrous iron and depends on normalization (see Electron Microprobe Methods); $\mathrm{K}_{2} \mathrm{O}$ in weight percent; $\mathrm{n}=$ number of microprobe analyses; UCD = analyzed at University of California at Davis; UH = analyzed at University of Hawaii.
*abbreviations: Mg-hast = Mg-hastingsite; Mg-hbl = Mg-hornblende; tscher = tschermakite. Classification of Leake et al. (1997).

TABLE DR-4. DETECTION LIMITS AND RELATIVE COUNTING ERRORS OF ELECTRON MICROPROBE ANALYSES

| Element | UCD analyses |  |  |  | UH analyses |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  | $\begin{aligned} & \text { Lower Limit } \\ & \text { of } \\ & \text { Detection } \\ & \text { (wt\% oxide) } \end{aligned}$ | Relative Counting Error | @ wt \% oxide | counting time (s) | $\begin{aligned} & \text { Lower Limit } \\ & \text { of } \\ & \text { Detection } \\ & \text { (wt\% oxide) } \end{aligned}$ | Relative Counting Error | @ wt \% oxide | counting time (s) |
| F | 0.58 | 10.5\% | @ 2.88 | 10 | 0.12 | 7.7\% | @ 0.30 | 18 |
| Na | 0.09 | 7.7\% | @ 0.29 | 10 | 0.06 | 2.9\% | @ 0.35 | 18 |
| Na | 0.09 | 3.8\% | @ 1.26 | 10 | 0.06 | 1.4\% | @ 1.26 | 18 |
| Mg | 0.12 | 1.0\% | @ 12.38 | 10 | 0.07 | 1.0\% | @ 11.87 | 12 |
| Mg | 0.12 | 0.9\% | @ 16.58 | 10 | 0.07 | 0.8\% | @ 16.58 | 12 |
| Al | 0.07 | 4.9\% | @ 0.45 | 10 | 0.06 | 3.2\% | @ 1.01 | 12 |
| Al | 0.07 | 1.1\% | @ 8.63 | 10 | 0.06 | 1.1\% | @ 8.63 | 12 |
| Al | 0.07 | 0.9\% | @ 14.32 | 10 | 0.06 | 0.9\% | @ 13.14 | 12 |
| Si | 0.10 | 0.5\% | @ 50.46 | 10 | 0.06 | 0.5\% | @ 50.46 | 12 |
| Cl | 0.08 | 40.8\% | @ 0.02 | 10 | 0.17 | 14.7\% | @ 0.17 | 12 |
| K | 0.07 | 3.1\% | @ 2.14 | 10 | 0.08 | 4.8\% | @ 0.76 | 12 |
| Ca | 0.11 | 1.5\% | @ 9.74 | 10 | 0.06 | 1.0\% | @ 11.69 | 12 |
| Ca | 0.11 | 1.1\% | @ 15.93 | 10 | 0.06 | 0.8\% | @ 15.93 | 18 |
| Ti | 0.16 | 6.1\% | @ 0.79 | 10 | 0.08 | 3.7\% | @ 0.79 | 18 |
| Cr | 0.18 | 11.2\% | @ 0.23 | 10 | 0.17 | 11.6\% | @ 0.23 | 12 |
| Mn | 0.20 | 18.6\% | @ 0.19 | 10 | 0.12 | 18.8\% | @ 0.19 | 18 |
| Fe | 0.20 | 3.0\% | @ 7.02 | 10 | 0.13 | 2.4\% | @ 7.02 | 18 |

Note: UCD and UH refer to University of California at Davis and University of Hawaii; detection limits calculated as twice the statistical limits; relative counting errors vary with concentration and are listed for more than one concentration for several elements; F and Cl " $\mathrm{wt} \%$ oxide" values are simple $\mathrm{wt} \%$ of F and $\mathrm{Cl} ; \mathrm{Fe} \mathrm{wt} \%$ oxide is reported as FeO .

TABLE DR-5. Ar/Ar DATA

| Metamorphic amphibole from foliated metagabbro, Slate Creek metaplutonic unit |  |  |  |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Sample F92-36 |  | $\mathrm{J}=0.0030186$ |  |  | $\mathrm{Wt}=31.4 \mathrm{mg}$ |  |  | 200-250 $\mu \mathrm{m}$ grain size |  |  |
| $\mathrm{T}\left({ }^{\circ} \mathrm{C}\right)$ | ${ }^{40} \mathrm{Ar}(\mathrm{mol})$ | ${ }^{40} \mathrm{Ar}{ }^{\beta 3} \mathrm{Ar}$ | ${ }^{38} \mathrm{Ar}{ }^{39} \mathrm{Ar}$ | ${ }^{37} \mathrm{Ar}{ }^{39} \mathrm{Ar}$ | ${ }^{36} \mathrm{Ar}{ }^{33} \mathrm{Ar}$ | K/Ca | ${ }^{\text {- }}{ }^{39} \mathrm{Ar}$ | ${ }^{40} \mathrm{Ar}^{*}$ | Age (M | a) $\pm 1 \sigma$ |
| 550 | 2.2e-14 | 629.906 | $4.6 \mathrm{e}-2$ | 11.5867 | 2.0224 | 0.042 | 0.00244 | 0.051 | 167.6 | $\pm 149.2$ |
| 650 | $3.0 \mathrm{e}-14$ | 765.368 | 6.1e-2 | 7.3608 | 2.3690 | 0.067 | 0.00521 | 0.085 | 324.6 | $\pm 168.2$ |
| 750 | $6.8 \mathrm{e}-14$ | 884.367 | 2.7e-2 | 9.4296 | 2.8526 | 0.052 | 0.01072 | 0.047 | 212.5 | $\pm 115.4$ |
| 850 | $2.5 \mathrm{e}-14$ | 182.535 | $1.2 \mathrm{e}-2$ | 11.0995 | 0.5181 | 0.044 | 0.02061 | 0.161 | 153.5 | $\pm 14.9$ |
| 900 | $2.4 \mathrm{e}-14$ | 43.421 | $1.8 \mathrm{e}-3$ | 12.1158 | 0.0456 | 0.040 | 0.05948 | 0.690 | 156.2 | $\pm 3.0$ |
| 950 | 7.7e-14 | 32.793 | $0.0 \mathrm{e}+0$ | 10.8076 | 0.0103 | 0.045 | 0.22591 | 0.908 | 155.2 | $\pm 1.2$ |
| 975 | $5.1 \mathrm{e}-14$ | 32.733 | 3.5e-4 | 12.2217 | 0.0098 | 0.040 | 0.33640 | 0.912 | 155.6 | $\pm 1.4$ |
| 1000 | $5.0 \mathrm{e}-14$ | 35.226 | $5.5 \mathrm{e}-3$ | 18.3018 | 0.0175 | 0.027 | 0.43746 | 0.854 | 156.7 | $\pm 2.5$ |
| 1015 | $3.9 \mathrm{e}-14$ | 36.297 | $9.6 \mathrm{e}-3$ | 22.7759 | 0.0195 | 0.022 | 0.51526 | 0.841 | 159.0 | $\pm 2.5$ |
| 1030 | $1.9 \mathrm{e}-14$ | 34.395 | 8.8e-3 | 22.8162 | 0.0163 | 0.021 | 0.55541 | 0.860 | 154.3 | $\pm 3.0$ |
| 1050 | $4.5 \mathrm{e}-15$ | 33.829 | $1.6 \mathrm{e}-3$ | 14.4150 | 0.0189 | 0.034 | 0.56491 | 0.835 | 147.7 | $\pm 7.2$ |
| 1070 | $6.4 \mathrm{e}-15$ | 34.932 | 3.3e-3 | 15.0097 | 0.0209 | 0.033 | 0.57804 | 0.823 | 150.1 | $\pm 5.5$ |
| 1090 | $6.5 \mathrm{e}-15$ | 36.083 | 2.6e-3 | 15.2984 | 0.0169 | 0.032 | 0.59089 | 0.862 | 161.9 | $\pm 6.1$ |
| 1105 | $6.6 \mathrm{e}-15$ | 34.625 | $0.0 \mathrm{e}+0$ | 14.8553 | 0.0131 | 0.033 | 0.60454 | 0.888 | 160.2 | $\pm 5.9$ |
| 1120 | $1.1 \mathrm{e}-14$ | 37.060 | $0.0 \mathrm{e}+0$ | 14.4059 | 0.0231 | 0.034 | 0.62517 | 0.816 | 157.5 | $\pm 4.3$ |
| 1135 | $2.5 \mathrm{e}-14$ | 38.121 | $4.0 \mathrm{e}-3$ | 14.2497 | 0.0279 | 0.034 | 0.67165 | 0.784 | 155.8 | $\pm 2.6$ |
| 1150 | $4.2 \mathrm{e}-14$ | 37.687 | 2.7e-3 | 15.1277 | 0.0228 | 0.032 | 0.75122 | 0.822 | 161.2 | $\pm 1.9$ |
| 1165 | $4.5 \mathrm{e}-14$ | 35.777 | 2.7e-3 | 14.4472 | 0.0190 | 0.034 | 0.84045 | 0.843 | 157.1 | $\pm 1.7$ |
| 1180 | $2.9 \mathrm{e}-14$ | 32.642 | $0.0 \mathrm{e}+0$ | 12.8240 | 0.0107 | 0.038 | 0.90318 | 0.903 | 153.8 | $\pm 2.0$ |
| 1250 | $3.8 \mathrm{e}-14$ | 31.726 | $1.4 \mathrm{e}-3$ | 12.1992 | 0.0072 | 0.040 | 0.98863 | 0.933 | 154.3 | $\pm 1.6$ |
| 1350 | $4.9 \mathrm{e}-15$ | 30.872 | $0.0 \mathrm{e}+0$ | 12.3504 | 0.0031 | 0.040 | 1.00000 | 0.971 | 156.2 | $\pm 6.0$ |

Total fusion age, TFA= $157.1 \pm 1.1 \mathrm{Ma}$ (including J)
Weighted mean age, $\mathrm{WMA}=156.1 \pm 0.6 \mathrm{Ma}$ (including J)
Inverse isochron age $=155.8 \pm 0.6 \mathrm{Ma}$. $($ MSWD $=1.44 ; 40 \mathrm{Ar} / 36 \mathrm{Ar}=299.0 \pm 1.6)$

| Igneous amphibole from dike intrusive into Slate Creek metaplutonic unit |  |  |  |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Sample F92-46 |  | $\mathrm{J}=0.0030001$ |  |  | $\mathrm{Wt}=21.4 \mathrm{mg}$ |  |  | 100-250 $\mu \mathrm{m}$ grain size |  |  |
| $\mathrm{T}\left({ }^{\circ} \mathrm{C}\right)$ | ${ }^{40} \mathrm{Ar}(\mathrm{mol})$ | ${ }^{40} \mathrm{Ar} /{ }^{39} \mathrm{Ar}$ | ${ }^{38} \mathrm{Ar} /{ }^{39} \mathrm{Ar}$ | ${ }^{37} \mathrm{Ar}{ }^{39} \mathrm{Ar}$ | ${ }^{36} \mathrm{Ar}{ }^{39} \mathrm{Ar}$ | K/Ca | ${ }^{\text {• }}{ }^{\text {a }} \mathrm{Ar}$ | ${ }^{40} \mathrm{Ar}^{*}$ | Age ( | a) $\pm 1 \sigma$ |
| 550 | 2.10E-14 | 125.245 | $7.60 \mathrm{E}-03$ | 2.4478 | 0.3499 | 0.200 | 0.00673 | 0.174 | 114.5 | $\pm 8.9$ |
|  |  | 4 |  |  |  |  |  |  |  |  |
| 650 | $2.20 \mathrm{E}-14$ | 46.9968 | $0.00 \mathrm{E}+00$ | 1.0110 | 0.0717 | 0.480 | 0.02558 | 0.549 | 134.6 | $\pm 2.3$ |
| 750 | $4.10 \mathrm{E}-14$ | 62.5336 | $0.00 \mathrm{E}+00$ | 0.8446 | 0.1264 | 0.580 | 0.05224 | 0.403 | 131.4 | $\pm 2.1$ |
| 850 | $4.60 \mathrm{E}-14$ | 55.8838 | $2.40 \mathrm{E}-03$ | 1.1632 | 0.1053 | 0.420 | 0.08549 | 0.443 | 129.3 | $\pm 1.8$ |
| 900 | $2.10 \mathrm{E}-14$ | 38.5479 | $0.00 \mathrm{E}+00$ | 2.9789 | 0.0523 | 0.160 | 0.10801 | 0.599 | 120.8 | $\pm 2.2$ |
| 950 | 2.10E-14 | 39.8008 | $0.00 \mathrm{E}+00$ | 12.8859 | 0.0632 | 0.038 | 0.13006 | 0.531 | 110.9 | $\pm 3.1$ |
| 975 | $1.50 \mathrm{E}-14$ | 42.2994 | $3.00 \mathrm{E}-03$ | 11.5028 | 0.0634 | 0.043 | 0.14444 | 0.557 | 123.3 | $\pm 3.9$ |
| 1000 | $1.30 \mathrm{E}-14$ | 36.9768 | $0.00 \mathrm{E}+00$ | 7.3972 | 0.0376 | 0.066 | 0.15928 | 0.699 | 134.8 | $\pm 3.6$ |
| 1015 | $1.10 \mathrm{E}-14$ | 32.5693 | $0.00 \mathrm{E}+00$ | 6.6649 | 0.0219 | 0.074 | 0.17374 | 0.801 | 136.0 | $\pm 2.7$ |
| 1030 | $1.80 \mathrm{E}-14$ | 30.8248 | $1.20 \mathrm{E}-03$ | 6.5029 | 0.0129 | 0.075 | 0.19774 | 0.876 | 140.6 | $\pm 2.2$ |
| 1050 | $4.10 \mathrm{E}-14$ | 30.3475 | $9.30 \mathrm{E}-04$ | 6.6206 | 0.0099 | 0.074 | 0.25327 | 0.904 | 142.6 | $\pm 1.2$ |
| 1070 | $9.70 \mathrm{E}-14$ | 29.5781 | 1.10E-04 | 6.7499 | 0.0074 | 0.073 | 0.38780 | 0.926 | 142.4 | $\pm 0.7$ |
| 1080 | $5.90 \mathrm{E}-14$ | 29.7212 | $8.20 \mathrm{E}-04$ | 6.7683 | 0.0085 | 0.072 | 0.46864 | 0.916 | 141.6 | $\pm 1.1$ |
| 1090 | $2.20 \mathrm{E}-14$ | 30.5664 | $9.50 \mathrm{E}-04$ | 7.0412 | 0.0119 | 0.070 | 0.49880 | 0.885 | 140.8 | $\pm 2.3$ |
| 1105 | $1.40 \mathrm{E}-14$ | 31.3925 | $4.00 \mathrm{E}-04$ | 6.8474 | 0.0148 | 0.072 | 0.51748 | 0.861 | 140.7 | $\pm 2.8$ |
| 1120 | $1.70 \mathrm{E}-14$ | 32.1679 | $1.20 \mathrm{E}-03$ | 7.2406 | 0.0156 | 0.068 | 0.53935 | 0.857 | 143.3 | $\pm 2.2$ |
| 1135 | $2.30 \mathrm{E}-14$ | 31.3362 | $0.00 \mathrm{E}+00$ | 7.0514 | 0.0132 | 0.069 | 0.56907 | 0.875 | 142.7 | $\pm 1.8$ |
| 1150 | $4.10 \mathrm{E}-14$ | 30.9589 | $0.00 \mathrm{E}+00$ | 6.8672 | 0.0115 | 0.071 | 0.62307 | 0.891 | 143.4 | $\pm 1.3$ |
| 1165 | $8.70 \mathrm{E}-14$ | 30.8576 | 7.10E-05 | 6.7141 | 0.0104 | 0.073 | 0.73902 | 0.900 | 144.4 | $\pm 0.9$ |
| 1180 | $1.40 \mathrm{E}-13$ | 30.4279 | $0.00 \mathrm{E}+00$ | 6.7631 | 0.0090 | 0.072 | 0.93431 | 0.912 | 144.3 | $\pm 0.8$ |
| 1250 | $4.30 \mathrm{E}-14$ | 29.9853 | $0.00 \mathrm{E}+00$ | 6.8128 | 0.0079 | 0.072 | 0.99348 | 0.922 | 143.8 | $\pm 1.2$ |
| 1350 | $4.80 \mathrm{E}-15$ | 30.4788 | $0.00 \mathrm{E}+00$ | 7.2014 | 0.0091 | 0.068 | 1.00000 | 0.911 | 144.4 | $\pm 5.6$ |

Total fusion age, TFA= $140.3 \pm 0.4 \mathrm{Ma}$ (including J)
Weighted mean age, WMA $=143.2 \pm 0.4 \mathrm{Ma}$ (including J)
Inverse isochron age $=141.2 \pm 1.8 \mathrm{Ma} .(M S W D=1.22 ; 40 \mathrm{Ar} / 36 \mathrm{Ar}=340 \pm 40)$

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| Sample F94-10 |  | $J=0.0029814$ |  |  | $\mathrm{Wt}=25.0 \mathrm{mg}$ |  |  | 250-350 $\mu \mathrm{m}$ grain size |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| $\mathrm{T}\left({ }^{\circ} \mathrm{C}\right)$ | ${ }^{40} \mathrm{Ar}(\mathrm{mol})$ | ${ }^{40} \mathrm{Ar} /^{39} \mathrm{Ar}$ | ${ }^{38} \mathrm{Ar} /{ }^{39} \mathrm{Ar}$ | ${ }^{37} \mathrm{Ar} /{ }^{39} \mathrm{Ar}$ | ${ }^{36} \mathrm{Ar}{ }^{39} \mathrm{Ar}$ | K/Ca | ${ }^{\text {- }}{ }^{39} \mathrm{Ar}$ | ${ }^{40} \mathrm{Ar}^{*}$ | Age ( | a) $\pm 1 \sigma$ |
| 550 | 2.70E-14 | $\begin{array}{r} 127.400 \\ 3 \end{array}$ | $0.00 \mathrm{E}+00$ | 1.2880 | 0.3746 | 0.380 | 0.00428 | 0.131 | 87.8 | $\pm 7.7$ |
| 650 | 8.70E-14 | 56.8749 | $0.00 \mathrm{E}+00$ | 0.1582 | 0.1047 | 3.100 | 0.03472 | 0.456 | 134.4 | $\pm 1.3$ |
| 750 | 8.10E-14 | 37.5051 | $0.00 \mathrm{E}+00$ | 0.1187 | 0.0239 | 4.100 | 0.07747 | 0.811 | 156.7 | $\pm 0.6$ |
| 850 | 6.30E-14 | 38.1711 | $0.00 \mathrm{E}+00$ | 0.3900 | 0.0257 | 1.300 | 0.11026 | 0.801 | 157.5 | $\pm 0.8$ |
| 900 | 3.90E-14 | 40.3245 | $0.00 \mathrm{E}+00$ | 0.9338 | 0.0336 | 0.520 | 0.12930 | 0.754 | 156.5 | $\pm 1.2$ |
| 950 | 4.50E-14 | 35.1161 | $0.00 \mathrm{E}+00$ | 2.3865 | 0.0158 | 0.210 | 0.15479 | 0.867 | 156.8 | $\pm 1.1$ |
| 975 | 3.00E-14 | 33.3850 | $2.70 \mathrm{E}-04$ | 2.7658 | 0.0106 | 0.180 | 0.17292 | 0.906 | 155.8 | $\pm 1.2$ |
| 1000 | 4.60E-14 | 33.5415 | $0.00 \mathrm{E}+00$ | 4.6083 | 0.0136 | 0.110 | 0.20055 | 0.880 | 152.2 | $\pm 1.1$ |
| 1015 | 1.70E-13 | 32.2496 | $0.00 \mathrm{E}+00$ | 6.5017 | 0.0107 | 0.075 | 0.30497 | 0.902 | 150.1 | $\pm 0.6$ |
| 1025 | 3.80E-13 | 30.9583 | $1.10 \mathrm{E}-03$ | 6.8581 | 0.0065 | 0.071 | 0.54836 | 0.938 | 149.8 | $\pm 0.2$ |
| 1035 | 1.20E-13 | 29.6512 | $1.10 \mathrm{E}-03$ | 6.5171 | 0.0024 | 0.075 | 0.63219 | 0.976 | 149.3 | $\pm 0.7$ |
| 1045 | 2.60E-14 | 30.1510 | $1.30 \mathrm{E}-03$ | 6.2148 | 0.0037 | 0.079 | 0.64911 | 0.964 | 149.9 | $\pm 1.7$ |
| 1055 | 1.80E-14 | 30.5623 | $0.00 \mathrm{E}+00$ | 6.2382 | 0.0043 | 0.079 | 0.66117 | 0.958 | 151.0 | $\pm 1.8$ |
| 1065 | 1.90E-14 | 30.2387 | $0.00 \mathrm{E}+00$ | 6.4143 | 0.0041 | 0.076 | 0.67385 | 0.960 | 149.7 | $\pm 2.3$ |
| 1080 | 3.30E-14 | 29.9620 | $0.00 \mathrm{E}+00$ | 6.7667 | 0.0039 | 0.072 | 0.69619 | 0.961 | 148.6 | $\pm 1.6$ |
| 1100 | 2.30E-14 | 30.5051 | $0.00 \mathrm{E}+00$ | 7.5731 | 0.0043 | 0.065 | 0.71099 | 0.958 | 150.8 | $\pm 2.0$ |
| 1120 | 4.70E-14 | 30.5848 | $0.00 \mathrm{E}+00$ | 7.3724 | 0.0055 | 0.066 | 0.74157 | 0.947 | 149.4 | $\pm 1.4$ |
| 1135 | 9.00E-14 | 30.7930 | $3.80 \mathrm{E}-04$ | 7.1072 | 0.0053 | 0.069 | 0.80017 | 0.949 | 150.7 | $\pm 0.8$ |
| 1150 | 9.30E-14 | 30.5113 | $5.00 \mathrm{E}-04$ | 7.1163 | 0.0047 | 0.069 | 0.86130 | 0.954 | 150.2 | $\pm 0.8$ |
| 1165 | 7.40E-14 | 30.3492 | $3.70 \mathrm{E}-04$ | 6.8307 | 0.0040 | 0.072 | 0.91020 | 0.961 | 150.4 | $\pm 0.8$ |
| 1180 | 7.20E-14 | 30.3039 | 2.80E-04 | 6.7799 | 0.0039 | 0.072 | 0.95783 | 0.962 | 150.3 | $\pm 0.9$ |
| 1250 | 5.50E-14 | 30.0861 | $1.20 \mathrm{E}-03$ | 6.8157 | 0.0037 | 0.072 | 0.99435 | 0.964 | 149.6 | $\pm 1.0$ |
| 1350 | 8.60E-15 | 30.3315 | $1.90 \mathrm{E}-03$ | 6.7543 | 0.0048 | 0.073 | 1.00000 | 0.954 | 149.2 | $\pm 3.4$ |

Total fusion age, TFA=150.2 $\pm 0.2 \mathrm{Ma}$ (including J )
Weighted mean age, WMA $=149.8 \pm 0.2 \mathrm{Ma}$ (including J)
Inverse isochron age $=151.0 \pm 0.7 \mathrm{Ma} .(M S W D=34.89 ; 40 \mathrm{Ar} / 36 \mathrm{Ar}=283 \pm 9)$ )

| Metamorphic amphibole massive metadiorite, Slate Creek metaplutonic unit |  |  |  |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Sample F94-13 |  | $\mathrm{J}=0.0029626$ |  |  | $\mathrm{Wt}=28.9 \mathrm{mg}$ |  |  | 125-180 $\mu \mathrm{m}$ grain size |  |  |
| $\mathrm{T}\left({ }^{\circ} \mathrm{C}\right)$ | ${ }^{40} \mathrm{Ar}(\mathrm{mol})$ | ${ }^{40} \mathrm{Ar} /^{39} \mathrm{Ar}$ | ${ }^{38} \mathrm{Ar} /{ }^{39} \mathrm{Ar}$ | ${ }^{37} \mathrm{Ar}{ }^{39} \mathrm{Ar}$ | ${ }^{36} \mathrm{Ar} /{ }^{39} \mathrm{Ar}$ | K/Ca | .$^{39} \mathrm{Ar}$ | ${ }^{40} \mathrm{Ar}^{*}$ | Age ( | a) $\pm 1 \sigma$ |
| 550 | 1.60E-14 | 138.051 | $4.60 \mathrm{E}-02$ | 8.1764 | 0.3937 | 0.060 | 0.00584 | 0.157 | 112.4 | $\pm 17.5$ |
|  |  | 9 |  |  |  |  |  |  |  |  |
| 650 | $1.50 \mathrm{E}-14$ | 67.2288 | $1.00 \mathrm{E}-02$ | 4.8478 | 0.1419 | 0.100 | 0.01697 | 0.376 | 130.3 | $\pm 5.6$ |
| 750 | $2.30 \mathrm{E}-14$ | 75.2487 | $0.00 \mathrm{E}+00$ | 2.7225 | 0.1606 | 0.180 | 0.03229 | 0.369 | 142.7 | $\pm 4.2$ |
| 850 | $1.90 \mathrm{E}-14$ | 47.8583 | $3.10 \mathrm{E}-03$ | 7.5198 | 0.0783 | 0.065 | 0.05248 | 0.517 | 127.6 | $\pm 3.3$ |
| 900 | $2.20 \mathrm{E}-14$ | 40.7089 | $4.20 \mathrm{E}-03$ | 22.2791 | 0.0431 | 0.022 | 0.08041 | 0.687 | 143.6 | $\pm 4.2$ |
| 950 | 7.10E-14 | 41.1909 | $1.00 \mathrm{E}-02$ | 25.5925 | 0.0309 | 0.019 | 0.16886 | 0.778 | 163.7 | $\pm 2.1$ |
| 975 | 1.60E-13 | 40.0328 | $2.10 \mathrm{E}-02$ | 18.2223 | 0.0177 | 0.027 | 0.37825 | 0.870 | 177.0 | $\pm 1.2$ |
| 1000 | $1.70 \mathrm{E}-13$ | 39.8043 | $2.80 \mathrm{E}-02$ | 16.5038 | 0.0126 | 0.030 | 0.59939 | 0.907 | 183.2 | $\pm 0.9$ |
| 1015 | 3.90E-14 | 40.6836 | $2.30 \mathrm{E}-02$ | 16.8879 | 0.0174 | 0.029 | 0.64774 | 0.873 | 180.5 | $\pm 2.5$ |
| 1030 | $1.90 \mathrm{E}-14$ | 42.1204 | $2.20 \mathrm{E}-02$ | 17.6389 | 0.0230 | 0.028 | 0.67067 | 0.838 | 179.4 | $\pm 4.0$ |
| 1050 | 1.60E-14 | 43.2687 | $2.10 \mathrm{E}-02$ | 19.2610 | 0.0247 | 0.025 | 0.68965 | 0.831 | 182.6 | $\pm 4.8$ |
| 1070 | $1.90 \mathrm{E}-14$ | 45.2084 | $1.70 \mathrm{E}-02$ | 20.4928 | 0.0308 | 0.024 | 0.71069 | 0.799 | 183.4 | $\pm 4.4$ |
| 1090 | 1.30E-14 | 48.7940 | $2.30 \mathrm{E}-02$ | 20.6981 | 0.0410 | 0.024 | 0.72467 | 0.751 | 186.0 | $\pm 5.6$ |
| 1105 | $1.50 \mathrm{E}-14$ | 48.1353 | $1.60 \mathrm{E}-02$ | 21.1194 | 0.0376 | 0.023 | 0.74008 | 0.769 | 187.7 | $\pm 5.6$ |
| 1120 | $2.40 \mathrm{E}-14$ | 48.8694 | $2.20 \mathrm{E}-02$ | 20.9294 | 0.0415 | 0.023 | 0.76483 | 0.749 | 185.7 | $\pm 4.3$ |
| 1135 | $3.50 \mathrm{E}-14$ | 48.1671 | $2.40 \mathrm{E}-02$ | 19.0451 | 0.0388 | 0.026 | 0.80197 | 0.762 | 186.2 | $\pm 2.7$ |
| 1150 | $4.30 \mathrm{E}-14$ | 46.4976 | $2.90 \mathrm{E}-02$ | 17.9355 | 0.0311 | 0.027 | 0.84882 | 0.802 | 189.0 | $\pm 2.4$ |
| 1180 | 8.60E-14 | 43.2491 | $2.90 \mathrm{E}-02$ | 17.1733 | 0.0206 | 0.029 | 0.94945 | 0.860 | 188.4 | $\pm 2.6$ |
| 1250 | $4.00 \mathrm{E}-14$ | 42.8797 | $2.90 \mathrm{E}-02$ | 17.0035 | 0.0153 | 0.029 | 0.99674 | 0.895 | 194.2 | $\pm 2.7$ |
| 1350 | $2.70 \mathrm{E}-15$ | 42.4221 | $2.40 \mathrm{E}-02$ | 18.0404 | 0.0284 | 0.027 | 1 | 0.802 | 173.2 | $\pm 14.9$ |

Total fusion age, TFA= $177.7 \pm 0.6 \mathrm{Ma}$ (including J)

Relict volcanic amphibole from the Lexington Hill volcanics, Slate Creek complex

| Sample F94-20 |  | $J=0.0029437$ |  |  | $\mathrm{Wt}=21.7 \mathrm{mg}$ |  |  | 200-250 $\mu \mathrm{m}$ grain size |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| $\mathrm{T}\left({ }^{\circ} \mathrm{C}\right)$ | ${ }^{40} \mathrm{Ar}(\mathrm{mol})$ | ${ }^{40} \mathrm{Ar} /{ }^{39} \mathrm{Ar}$ | ${ }^{38} \mathrm{Ar} /{ }^{39} \mathrm{Ar}$ | ${ }^{37} \mathrm{Ar} /{ }^{39} \mathrm{Ar}$ | ${ }^{36} \mathrm{Ar} /{ }^{\beta 9} \mathrm{Ar}$ | K/Ca | ${ }^{\text {- }}{ }^{39} \mathrm{Ar}$ | ${ }^{40} \mathrm{Ar}^{\text {* }}$ | Age ( | a) $\pm 1 \sigma$ |
| 550 | 2.50E-14 | 75.3638 | $6.60 \mathrm{E}-03$ | 3.1175 | 0.1702 | 0.1600 | 0.00589 | 0.333 | 128.4 | $\pm 4.5$ |
| 650 | 5.90E-14 | 70.1316 | $0.00 \mathrm{E}+00$ | 1.1022 | 0.1179 | 0.4400 | 0.02082 | 0.503 | 178.3 | $\pm 2.1$ |
| 750 | 1.20E-13 | 86.1419 | $0.00 \mathrm{E}+00$ | 1.3097 | 0.1780 | 0.3700 | 0.04504 | 0.390 | 169.9 | $\pm 1.9$ |
| 850 | 1.10E-13 | 41.9453 | $0.00 \mathrm{E}+00$ | 1.4545 | 0.0385 | 0.3400 | 0.09247 | 0.729 | 155.5 | $\pm 1.5$ |
| 900 | 4.80E-14 | 38.1791 | $1.90 \mathrm{E}-04$ | 9.8298 | 0.0363 | 0.0500 | 0.11517 | 0.719 | 140.2 | $\pm 1.6$ |
| 950 | $5.00 \mathrm{E}-14$ | 39.5468 | $2.40 \mathrm{E}-03$ | 27.0934 | 0.0386 | 0.0180 | 0.13825 | 0.712 | 143.6 | $\pm 2.2$ |
| 975 | 4.30E-14 | 39.7724 | $2.30 \mathrm{E}-03$ | 14.7965 | 0.0236 | 0.0330 | 0.15761 | 0.824 | 166.2 | $\pm 2.1$ |
| 1000 | 4.10E-14 | 38.5159 | 4.80E-03 | 11.0145 | 0.0181 | 0.0440 | 0.17678 | 0.861 | 168.0 | $\pm 1.9$ |
| 1015 | 3.60E-14 | 37.6504 | $5.80 \mathrm{E}-03$ | 9.6109 | 0.0130 | 0.0510 | 0.19408 | 0.898 | 171.1 | $\pm 2.1$ |
| 1030 | 5.30E-14 | 36.4641 | $4.10 \mathrm{E}-03$ | 8.2765 | 0.0093 | 0.0590 | 0.22005 | 0.925 | 170.7 | $\pm 1.5$ |
| 1050 | $1.30 \mathrm{E}-13$ | 35.7086 | 3.70E-03 | 7.1409 | 0.0059 | 0.0690 | 0.28433 | 0.951 | 171.9 | $\pm 0.8$ |
| 1070 | 2.20E-13 | 35.3713 | $3.60 \mathrm{E}-03$ | 6.9016 | 0.0047 | 0.0710 | 0.39388 | 0.960 | 171.9 | $\pm 0.7$ |
| 1090 | $1.20 \mathrm{E}-13$ | 35.6998 | $3.10 \mathrm{E}-03$ | 6.9220 | 0.0055 | 0.0710 | 0.45397 | 0.954 | 172.4 | $\pm 0.8$ |
| 1100 | $4.20 \mathrm{E}-14$ | 36.5311 | $3.70 \mathrm{E}-03$ | 7.8467 | 0.0091 | 0.0620 | 0.47437 | 0.927 | 171.3 | $\pm 1.9$ |
| 1110 | $5.00 \mathrm{E}-14$ | 36.4977 | $3.90 \mathrm{E}-03$ | 8.0700 | 0.0089 | 0.0610 | 0.49867 | 0.928 | 171.4 | $\pm 1.4$ |
| 1120 | 6.30E-14 | 36.8221 | 4.40E-03 | 8.0535 | 0.0101 | 0.0610 | 0.52914 | 0.919 | 171.3 | $\pm 1.4$ |
| 1135 | 1.10E-13 | 37.0062 | 3.90E-03 | 7.7263 | 0.0101 | 0.0630 | 0.58295 | 0.920 | 172.2 | $\pm 1.1$ |
| 1150 | $1.90 \mathrm{E}-13$ | 36.7219 | $3.30 \mathrm{E}-03$ | 6.9104 | 0.0089 | 0.0710 | 0.67556 | 0.928 | 172.5 | $\pm 0.7$ |
| 1165 | 4.00E-13 | 36.6370 | $3.70 \mathrm{E}-03$ | 6.3570 | 0.0085 | 0.0770 | 0.87000 | 0.932 | 172.7 | $\pm 0.2$ |
| 1180 | $1.50 \mathrm{E}-13$ | 36.0689 | $2.60 \mathrm{E}-03$ | 6.3146 | 0.0070 | 0.0780 | 0.94328 | 0.943 | 172.1 | $\pm 0.8$ |
| 1250 | $1.00 \mathrm{E}-13$ | 35.9852 | $3.30 \mathrm{E}-03$ | 6.4992 | 0.0067 | 0.0750 | 0.99356 | 0.945 | 172.1 | $\pm 0.9$ |
| 1350 | 1.30E-14 | 36.1747 | $3.00 \mathrm{E}-03$ | 6.3968 | 0.0068 | 0.0770 | 1.00000 | 0.944 | 172.8 | $\pm 2.9$ |

Total fusion age, TFA= $169.6 \pm 0.3 \mathrm{Ma}$ (including J)
Weighted mean age, WMA $=172.5 \pm 0.2 \mathrm{Ma}$ (including J)
Inverse isochron age $=171.4 \pm 0.9 \mathrm{Ma} .(M S W D=0.72 ; 40 \mathrm{Ar} / 36 \mathrm{Ar}=324 \pm 20)$

| Relict volcanic amphibole from the Lexington Hill volcanics, Slate Creek complex |  |  |  |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Sample F94-23 |  | $J=0.0029247$ |  |  | $\mathrm{Wt}=23.5 \mathrm{mg}$ |  |  | 200-250 $\mu \mathrm{m}$ grain size |  |  |
| $\mathrm{T}\left({ }^{\circ} \mathrm{C}\right)$ | ${ }^{40} \mathrm{Ar}(\mathrm{mol})$ | ${ }^{40} \mathrm{Ar} /{ }^{39} \mathrm{Ar}$ | ${ }^{38} \mathrm{Ar} /{ }^{39} \mathrm{Ar}$ | ${ }^{37} \mathrm{Ar}{ }^{39} \mathrm{Ar}$ | ${ }^{36} \mathrm{Ar} /{ }^{39} \mathrm{Ar}$ | K/Ca | .$^{39} \mathrm{Ar}$ | ${ }^{40} \mathrm{Ar}^{*}$ | Age (M | a) $\pm 1 \sigma$ |
| 550 | 2.30E-14 | 216.894 | $1.70 \mathrm{E}-02$ | 4.7985 | 0.6666 | 0.1000 | 0.00367 | 0.092 | 102.1 | $\pm 20.1$ |
| 9 |  |  |  |  |  |  |  |  |  |  |
| 650 | 4.20E-14 | 127.771 | $3.30 \mathrm{E}-03$ | 1.5950 | 0.3416 | 0.3100 | 0.01535 | 0.210 | 136.2 | $\pm 6.0$ |
| 9 |  |  |  |  |  |  |  |  |  |  |
| 750 | $1.10 \mathrm{E}-13$ | 116.797 | $2.10 \mathrm{E}-03$ | 1.4329 | 0.3185 | 0.3400 | 0.04984 | 0.194 | 115.8 | $\pm 3.1$ |
| 3 |  |  |  |  |  |  |  |  |  |  |
| 850 | 8.50E-14 | 57.2649 | $0.00 \mathrm{E}+00$ | 3.0907 | 0.1074 | 0.1600 | 0.10207 | 0.446 | 129.8 | $\pm 1.4$ |
| 900 | 4.10E-14 | 51.4690 | $1.40 \mathrm{E}-03$ | 20.9493 | 0.0852 | 0.0230 | 0.13072 | 0.511 | 133.6 | $\pm 3.0$ |
| 950 | 4.90E-14 | 55.4390 | $2.40 \mathrm{E}-03$ | 54.5304 | 0.0994 | 0.0090 | 0.16356 | 0.470 | 132.6 | $\pm 4.1$ |
| 975 | 2.50E-14 | 41.0993 | $3.60 \mathrm{E}-03$ | 18.4708 | 0.0451 | 0.0270 | 0.18554 | 0.676 | 140.8 | $\pm 3.0$ |
| 1000 | 2.30E-14 | 35.4363 | $2.10 \mathrm{E}-03$ | 11.0001 | 0.0251 | 0.0450 | 0.20852 | 0.791 | 142.1 | $\pm 2.2$ |
| 1015 | 1.60E-14 | 33.1198 | $1.30 \mathrm{E}-03$ | 9.2614 | 0.0221 | 0.0530 | 0.22573 | 0.803 | 135.1 | $\pm 3.0$ |
| 1030 | 1.70E-14 | 33.3157 | $1.90 \mathrm{E}-03$ | 8.2808 | 0.0170 | 0.0590 | 0.24423 | 0.849 | 143.4 | $\pm 3.0$ |
| 1050 | 4.80E-14 | 35.0276 | $3.00 \mathrm{E}-03$ | 7.3334 | 0.0120 | 0.0670 | 0.29231 | 0.898 | 158.8 | $\pm 1.3$ |
| 1070 | $1.60 \mathrm{E}-13$ | 35.4891 | $3.10 \mathrm{E}-03$ | 7.3516 | 0.0084 | 0.0670 | 0.44850 | 0.930 | 166.2 | $\pm 0.7$ |
| 1090 | 7.80E-14 | 36.1749 | $2.80 \mathrm{E}-03$ | 9.2732 | 0.0116 | 0.0530 | 0.52485 | 0.905 | 165.0 | $\pm 1.1$ |
| 1105 | 3.10E-14 | 49.8117 | $2.20 \mathrm{E}-03$ | 13.0897 | 0.0566 | 0.0370 | 0.54699 | 0.664 | 166.6 | $\pm 2.9$ |
| 1120 | 2.90E-14 | 39.6802 | $4.60 \mathrm{E}-03$ | 15.0001 | 0.0222 | 0.0330 | 0.57327 | 0.835 | 166.8 | $\pm 2.8$ |
| 1135 | 4.10E-14 | 41.2061 | $3.10 \mathrm{E}-03$ | 13.3861 | 0.0242 | 0.0370 | 0.60910 | 0.826 | 171.2 | $\pm 2.5$ |
| 1150 | 7.00E-14 | 40.2847 | $3.50 \mathrm{E}-03$ | 10.3414 | 0.0229 | 0.0470 | 0.67061 | 0.832 | 168.7 | $\pm 1.5$ |
| 1165 | $1.50 \mathrm{E}-13$ | 39.0749 | $3.50 \mathrm{E}-03$ | 8.8265 | 0.0182 | 0.0560 | 0.80393 | 0.863 | 169.6 | $\pm 0.8$ |
| 1180 | $1.50 \mathrm{E}-13$ | 38.9287 | $3.10 \mathrm{E}-03$ | 8.9458 | 0.0171 | 0.0550 | 0.94004 | 0.870 | 170.4 | $\pm 0.9$ |
| 1250 | 6.20E-14 | 38.8652 | $2.60 \mathrm{E}-03$ | 10.2970 | 0.0164 | 0.0480 | 0.99626 | 0.875 | 171.1 | $\pm 1.6$ |
| 1350 | $4.00 \mathrm{E}-15$ | 38.5116 | $8.50 \mathrm{E}-04$ | 10.2764 | 0.0154 | 0.0480 | 1.00000 | 0.882 | 170.9 | $\pm 8.8$ |

Total fusion age, TFA $=159.2 \pm 0.4 \mathrm{Ma}$ (including J)
Weighted mean plateau age, WMPA $=170.0 \pm 0.5 \mathrm{Ma}$ (including J)
Inverse isochron age $=171.9 \pm 3.3 \mathrm{Ma} .(\mathrm{MSWD}=0.65 ; 40 \mathrm{Ar} / 36 \mathrm{Ar}=270 \pm 30)$

Igneous amphibole from dike intruding Gold Run formation

| Sample F94-54 |  | $J=0.0029056$ |  |  | $\mathrm{Wt}=23.9 \mathrm{mg}$ |  |  | 200-250 $\mu \mathrm{m}$ grain size |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| $\mathrm{T}\left({ }^{\circ} \mathrm{C}\right)$ | ${ }^{40} \mathrm{Ar}(\mathrm{mol})$ | ${ }^{40} \mathrm{Ar}{ }^{39} \mathrm{Ar}$ | ${ }^{38} \mathrm{Ar} /{ }^{39} \mathrm{Ar}$ | ${ }^{37} \mathrm{Ar} /{ }^{39} \mathrm{Ar}$ | ${ }^{36} \mathrm{Ar} /{ }^{39} \mathrm{Ar}$ | K/Ca | .$^{39} \mathrm{Ar}$ | ${ }^{40} \mathrm{Ar}^{*}$ | Age (M | a) $\pm 1 \sigma$ |
| 550 | 6.00E-14 | 154.7403 | $1.30 \mathrm{E}-02$ | 3.0487 | 0.4365 | 0.1600 | 0.00846 | 0.166 | 130.1 | $\pm 6.3$ |
| 650 | 1.20E-13 | 117.1801 | $1.60 \mathrm{E}-03$ | 2.0789 | 0.3014 | 0.2400 | 0.02990 | 0.240 | 141.6 | $\pm 3.1$ |
| 700 | 9.50E-14 | 128.9696 | $0.00 \mathrm{E}+00$ | 1.1959 | 0.3415 | 0.4100 | 0.04598 | 0.218 | 141.4 | $\pm 3.7$ |
| 750 | 1.20E-13 | 181.2289 | $0.00 \mathrm{E}+00$ | 1.0648 | 0.5163 | 0.4600 | 0.06083 | 0.158 | 144.3 | $\pm 5.0$ |
| 850 | 1.40E-13 | 102.6742 | $1.50 \mathrm{E}-03$ | 2.4675 | 0.2542 | 0.2000 | 0.09151 | 0.268 | 138.9 | $\pm 2.3$ |
| 900 | 1.10E-13 | 96.7385 | 1.90E-03 | 11.3612 | 0.2285 | 0.0430 | 0.11754 | 0.302 | 147.0 | $\pm 3.0$ |
| 950 | 1.00E-13 | 91.0668 | $1.50 \mathrm{E}-03$ | 13.4511 | 0.2066 | 0.0360 | 0.14195 | 0.330 | 150.9 | $\pm 2.7$ |
| 975 | 4.80E-14 | 70.4266 | 2.80E-03 | 8.0397 | 0.1384 | 0.0610 | 0.15699 | 0.419 | 148.5 | $\pm 2.9$ |
| 1000 | 4.30E-14 | 50.6892 | 2.50E-03 | 7.2645 | 0.0764 | 0.0670 | 0.17559 | 0.554 | 141.6 | $\pm 2.4$ |
| 1015 | 3.00E-14 | 42.5100 | $1.70 \mathrm{E}-03$ | 6.8700 | 0.0507 | 0.0710 | 0.19091 | 0.648 | 138.8 | $\pm 2.4$ |
| 1030 | 2.90E-14 | 35.7665 | $1.00 \mathrm{E}-03$ | 6.9651 | 0.0279 | 0.0700 | 0.20843 | 0.770 | 138.8 | $\pm 1.9$ |
| 1050 | 8.20E-14 | 32.1556 | $1.60 \mathrm{E}-03$ | 7.8327 | 0.0146 | 0.0630 | 0.26424 | 0.866 | 140.3 | $\pm 1.1$ |
| 1070 | 1.80E-13 | 31.5602 | 8.20E-04 | 8.1008 | 0.0119 | 0.0600 | 0.38917 | 0.888 | 141.3 | $\pm 0.8$ |
| 1090 | 1.40E-13 | 32.5574 | $1.50 \mathrm{E}-03$ | 7.9713 | 0.0152 | 0.0610 | 0.48384 | 0.862 | 141.5 | $\pm 0.8$ |
| 1105 | 8.80E-14 | 32.3828 | 5.70E-04 | 7.8593 | 0.0158 | 0.0620 | 0.54317 | 0.856 | 139.7 | $\pm 1.1$ |
| 1120 | 6.90E-14 | 35.5359 | 1.30E-03 | 8.3487 | 0.0246 | 0.0590 | 0.58565 | 0.795 | 142.3 | $\pm 1.2$ |
| 1135 | 9.70E-14 | 38.4091 | $1.30 \mathrm{E}-03$ | 8.9180 | 0.0316 | 0.0550 | 0.64073 | 0.757 | 146.3 | $\pm 1.4$ |
| 1150 | 1.50E-13 | 39.5951 | 1.10E-03 | 8.8827 | 0.0345 | 0.0550 | 0.72161 | 0.742 | 147.8 | $\pm 1.1$ |
| 1165 | 2.70E-13 | 37.7062 | $1.20 \mathrm{E}-03$ | 8.5059 | 0.0286 | 0.0580 | 0.87740 | 0.776 | 147.2 | $\pm 0.6$ |
| 1180 | 1.40E-13 | 35.9327 | 7.40E-04 | 8.2048 | 0.0239 | 0.0600 | 0.96297 | 0.803 | 145.3 | $\pm 0.9$ |
| 1250 | $4.60 \mathrm{E}-14$ | 35.8636 | 3.80E-04 | 8.3778 | 0.0233 | 0.0580 | 0.99088 | 0.808 | 145.8 | $\pm 1.9$ |
| 1350 | $1.50 \mathrm{E}-14$ | 36.4702 | 2.30E-03 | 8.2647 | 0.0251 | 0.0590 | 1.00000 | 0.797 | 146.2 | $\pm 3.3$ |

Total fusion age, TFA $=143.7 \pm 0.3 \mathrm{Ma}$ (including J)
Weighted mean plateau age, WMPA $=144.0 \pm 0.3 \mathrm{Ma}$ (including J)
Inverse isochron age (based on 10 highest-T steps) $=135.8 \pm 1.2 \mathrm{Ma}$. (MSWD $=1.71 ; 40 \mathrm{Ar} / 36 \mathrm{Ar}=372 \pm 10$ )

| Metamorphic amphibole from foliated metadiorite, Slate Creek metaplutonic unit |  |  |  |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Sample F94-117 |  | $J=0.0028863$ |  |  | $\mathrm{Wt}=28.4 \mathrm{mg}$ |  |  | 200-250 $\mu \mathrm{m}$ grain size |  |  |
|  |  | ${ }^{40} \mathrm{Ar} /{ }^{39} \mathrm{Ar}$ | ${ }^{38} \mathrm{Ar} /{ }^{39} \mathrm{Ar}$ | ${ }^{37} \mathrm{Ar} /{ }^{39} \mathrm{Ar}$ | ${ }^{36} \mathrm{Ar}{ }^{39} \mathrm{Ar}$ | K/Ca | .$^{39} \mathrm{Ar}$ | ${ }^{40} \mathrm{Ar}^{*}$ | Age (M | a) $\pm 1 \sigma$ |
| 550 | 4.60E-14 | 774.1777 | $2.00 \mathrm{E}-02$ | 10.0458 | 2.5620 | 0.0490 | 0.00304 | 0.022 | 86.8 | $\pm 92.6$ |
| 650 | 6.30E-14 | 461.6852 | $1.50 \mathrm{E}-02$ | 7.2925 | 1.4415 | 0.0670 | 0.01002 | 0.077 | 177.0 | $\pm 28.0$ |
| 750 | $1.00 \mathrm{E}-13$ | 313.5518 | $7.00 \mathrm{E}-03$ | 7.4568 | 0.9571 | 0.0660 | 0.02646 | 0.098 | 153.3 | $\pm 12.8$ |
| 850 | 9.80E-14 | 112.3196 | $1.30 \mathrm{E}-03$ | 5.4555 | 0.2819 | 0.0900 | 0.07113 | 0.258 | 145.1 | $\pm 3.4$ |
| 900 | $2.40 \mathrm{E}-14$ | 53.1565 | $1.50 \mathrm{E}-03$ | 7.6835 | 0.0844 | 0.0640 | 0.09376 | 0.531 | 141.3 | $\pm 3.4$ |
| 950 | $1.90 \mathrm{E}-14$ | 52.5041 | $1.90 \mathrm{E}-03$ | 19.5341 | 0.0855 | 0.0250 | 0.11285 | 0.519 | 136.6 | $\pm 5.1$ |
| 975 | $2.00 \mathrm{E}-14$ | 53.5116 | $5.10 \mathrm{E}-03$ | 25.9895 | 0.0828 | 0.0190 | 0.13248 | 0.543 | 145.1 | $\pm 5.6$ |
| 1000 | $1.80 \mathrm{E}-13$ | 48.6577 | $6.70 \mathrm{E}-03$ | 30.5206 | 0.0603 | 0.0160 | 0.32241 | 0.634 | 153.8 | $\pm 1.6$ |
| 1010 | 7.30E-14 | 40.8908 | $3.80 \mathrm{E}-03$ | 26.5568 | 0.0355 | 0.0180 | 0.41468 | 0.743 | 151.7 | $\pm 2.1$ |
| 1020 | $2.30 \mathrm{E}-14$ | 41.2372 | $3.70 \mathrm{E}-03$ | 25.0774 | 0.0392 | 0.0200 | 0.44374 | 0.719 | 148.1 | $\pm 3.7$ |
| 1030 | $1.00 \mathrm{E}-14$ | 41.9034 | $3.10 \mathrm{E}-03$ | 24.0399 | 0.0425 | 0.0200 | 0.45607 | 0.700 | 146.6 | $\pm 7.5$ |
| 1050 | 7.60E-15 | 43.4456 | $1.50 \mathrm{E}-03$ | 23.6547 | 0.0474 | 0.0210 | 0.46514 | 0.678 | 147.1 | $\pm 8.6$ |
| 1070 | 1.50E-14 | 42.3225 | $2.80 \mathrm{E}-03$ | 23.1927 | 0.0429 | 0.0210 | 0.48398 | 0.700 | 148.1 | $\pm 5.7$ |
| 1090 | $4.60 \mathrm{E}-14$ | 38.1770 | 1.20E-03 | 22.7327 | 0.0286 | 0.0220 | 0.54630 | 0.779 | 148.5 | $\pm 2.8$ |
| 1105 | $2.10 \mathrm{E}-14$ | 39.0005 | $0.00 \mathrm{E}+00$ | 23.5352 | 0.0313 | 0.0210 | 0.57431 | 0.763 | 148.7 | $\pm 4.3$ |
| 1120 | $2.10 \mathrm{E}-14$ | 39.5413 | $1.50 \mathrm{E}-03$ | 24.4220 | 0.0333 | 0.0200 | 0.60245 | 0.751 | 148.4 | $\pm 4.3$ |
| 1135 | 5.10E-14 | 40.8900 | $1.50 \mathrm{E}-03$ | 24.0392 | 0.0357 | 0.0200 | 0.66688 | 0.742 | 151.4 | $\pm 2.7$ |
| 1150 | $1.00 \mathrm{E}-13$ | 42.1061 | $2.50 \mathrm{E}-03$ | 24.0688 | 0.0388 | 0.0200 | 0.79277 | 0.728 | 152.8 | $\pm 1.5$ |
| 1165 | 1.20E-13 | 43.2344 | 2.70E-03 | 24.2277 | 0.0422 | 0.0200 | 0.93710 | 0.711 | 153.4 | $\pm 1.8$ |
| 1180 | $2.60 \mathrm{E}-14$ | 43.4733 | $2.70 \mathrm{E}-03$ | 25.2561 | 0.0441 | 0.0190 | 0.96878 | 0.701 | 152.0 | $\pm 3.9$ |
| 1250 | $2.00 \mathrm{E}-14$ | 43.0448 | $1.70 \mathrm{E}-03$ | 25.5411 | 0.0401 | 0.0190 | 0.99252 | 0.724 | 155.5 | $\pm 4.3$ |
| 1350 | 6.00E-15 | 41.5832 | $4.80 \mathrm{E}-03$ | 24.7638 | 0.0380 | 0.0200 | 1.00000 | 0.730 | 151.5 | $\pm 9.9$ |

Total fusion age, TFA $=151.0 \pm 0.8 \mathrm{Ma}$ (including J)
Weighted mean plateau age, WMPA $=152.0 \pm 0.7$ (including J)
Inverse isochron age $=151.4 \pm 1.1 \mathrm{Ma} .(\mathrm{MSWD}=2.93 ; 40 \mathrm{Ar} / 36 \mathrm{Ar}=294.6 \pm 1.5)$

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## DATA REPOSITORY FIGURES CAPTIONS

Figure DR-1. Mean Si and $\mathrm{Mg} /(\mathrm{Mg}+\mathrm{Fe})$ of amphiboles in samples from which ${ }^{40} \mathrm{Ar} /{ }^{39} \mathrm{Ar}$ data were collected, after Leake et al. (1997). Labelled lines connect the amphibole teypes identifies in the metaplutonic samples (see Tables DR-2; DR-3). Error bars are $\pm 1 \sigma$ (see Table DR-3). Field labels: tr, tremolite; ac, actinolite; mhb, magnesiohornblende; ts, tschermakite; pa, pargasite; mhs, magnesiohastingsite. Tschermakite, pargasite, and magnesiohastingsite distinguished by concentration of alkalies and ratio of ferric iron to aluminum.

Figure DR-2. Photomicrographs of amphiboles separated for $\left.{ }^{40} \mathrm{Ar}\right)^{39} \mathrm{Ar}$ analyses. All in transmitted, plane-polarized light except for A, which is in cross-polarized light. Fields of view are 1 mm for A and B , and 2.6 mm for C, D, E, and F. (A) Relict volcanic phenocryst, F94-20, Lexington Hill volcanics. Light-colored, low-relief fringe at lower right end of prism is actinolitic overgrowth. (B) Relict volcanic phenocryst, F94-23, Lexington Hill volcanics. Note abundance of inclusions compared to F94-20. (C) Elongate to acicular (a) and equant ( q ) varieties of amphibole from foliated metadiorite, F92-36, Slate Creek complex metaplutonic unit. Foliation is parallel to length of photograph. (D) Equant metaplutonic amphibole overgrown by prominent elongate amphibole at upper left, F94-13, Slate Creek complex metaplutonic section. Light zone in middle of equant amphibole is caused by thinning of amphibole crystal and zoning to more Mg-rich compositions. (E) Igneous plutonic amphibole, F94-10, Cascade pluton. (F) Amphibole phenocrysts in dike, F94-54, intrusive into Gold Run formation.


Fig. DR-1
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Fig. DR-2*
March 10, 2001
*Final versions of A-F all to be submitted as hard-copy photographs.


[^0]:    Note: T: temperature; ${ }^{40} \mathrm{Ar}(\mathrm{mol})=$ moles corrected for blank and reactor-produced ${ }^{40} \mathrm{Ar}$. Ratios are corrected for blanks, decay, and interference. ${ }^{39} \mathrm{Ar}$ is cumulative, ${ }^{40} \mathrm{Ar}{ }^{*}=$ radiogenic fraction of ${ }^{40} \mathrm{Ar}$; MSWD: mean sum of weighted deviates; expresses the goodness of fit of the isochron and has an expected value of $\sim 1.94$ (Wendt and Carl, 1991).

