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Supplemental Material

Supplemental Text. Analytical Methods.

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Figure S2. Chondrite-normalized rare earth element patterns of analyzed zircons for the Mangling intrusive complex.

Figure S3. Field photographs showing relationships of the granitic rocks of the Mangling intrusive complex.

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Table S1. Zircon U-Pb dating results for the dioritic and granitic rocks of the Mangling intrusive complex.

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ANALYTICAL METHODS

Zircon U-Pb Dating

By using conventional heavy liquid and magnetic techniques, zircons were separated from five representative dioritic (biotite diorite and biotite diorite enclave) and granitic rock samples (medium- to fine-grained and fine-grained monzogranite and K-feldspar granite) collected from the Mangling intrusive complex and then handpicked under a binocular microscope. The selected zircon grains were mounted in epoxy resin and polished to expose the crystal centers to observe their internal structure under cathodoluminescence (CL) imaging that was carried out at the Testing Center, Tuoyan Analytical Technology Co. Ltd. (Guangzhou, China) using a TESCAN MIRA3 field-emission scanning electron microprobe (FE-SEM).

In-situ U-Pb dating of zircon grains were performed at the Nanjing FocuMS Technology Co. Ltd. (Nanjing, China), using an Agilent Technologies 7700x quadrupole inductively coupled plasma-mass spectrometry (ICP-MS) coupled with a Resonetics RESolution LR (193 nm ArF excimer) laser ablation system. During analysis, the laser ablation system has a spot diameter of 33 μm with a background acquisition of approximately 20 s followed by 40 s of data acquisition. 91500 zircon was used as external standard to correct instrumental mass discrimination and elemental fractionation during the ablation, whilst GJ-1 and Plešovice zircons were treated as quality control for geochronology. Trace elements of zircons were external calibrated against NIST SRM 610 with Si as internal standard. The detailed analytical

technique is described by [Li et al. \(2011\)](#). Isotopic ratios of U-Pb dating were performed using ICPMSDataCal ([Liu et al., 2008](#)), and age calculation and plotting of concordia diagrams were performed using Isoplot/Ex 3.0 ([Ludwig, 2003](#)).

Whole-Rock Major and Trace Elements

Whole-rock major and trace element analyses were performed at the Key Laboratory of Western China's Mineral Resources and Geological Engineering, Ministry of Education and the State Key Laboratory of Isotope Geochemistry, Guangzhou Institute of Geochemistry, Chinese Academy of Sciences with sample names as "ML21-". Before analysis, following procedures should be performed for all the samples: (1) powder them to less than 200 meshes, (2) add Lithium Borate Flux (50%–50% $\text{Li}_2\text{B}_4\text{O}_7$ – LiBO_2), and (3) mix them well and fuse in an auto fluxer between 1050 °C and 1100 °C. The X-ray fluorescence (XRF) spectrometry was used to analyze major elements with analytical precision better than 1%. For trace elements, HNO_3 + HF acid was used to digest whole-rock powder (50 mg) in steel-bomb coated Teflon beakers for two days in order to assure complete dissolution of the refractory minerals under high pressure. Then ICP-MS was used to analyze the solution, with analytical precision better than 5% for most trace elements.

Magmatic Oxidation State and Crystallization Temperature

Zircon $\text{Ce}^{4+}/\text{Ce}^{3+}$ ratio and Ti-in-zircon thermometer were used to evaluate the magmatic oxidation state and crystallization temperatures of igneous rocks ([Shen et al., 2015](#); [Zhang et al., 2020](#)), with detailed calculations in [Ballard et al. \(2002\)](#) and

43 [Ferry and Watson \(2007\)](#), respectively. Moreover, the zircon Eu anomalies (Eu/Eu^*)
44 were calculated based on normalized values of Sm and Gd concentrations by a
45 conventional method, similar to the equation of whole-rock Eu anomalies. For the
46 magmatic oxygen fugacity ($f\text{O}_2$), it is commonly estimated as relative
47 fayalite-magnetite-quartz buffer (ΔFMQ) based on the incorporation of cerium into
48 zircon, Ti-in-zircon temperature, and water content using the equation of [Smythe and](#)
49 [Brenan \(2016\)](#). In this study, the Ti-in-zircon thermometer and zircon $\text{Ce}^{4+}/\text{Ce}^{3+}$ and
50 Eu/Eu^* ratios were calculated to evaluate the magma nature of the Mangling intrusive
51 complex using Geo- $f\text{O}_2$ software proposed by [Li et al. \(2019\)](#), with assumed SiO_2 and
52 TiO_2 activity to be 1 and 0.7, respectively, because all samples from the Mangling
53 intrusive complex of this study contain primary magmatic quartz and accessory
54 mineral of sphene ([Hao et al., 2021](#)). For the water content, we use 4.0 wt% for the
55 biotite diorite enclave (BDE), biotite diorite (BD), and medium- to fine-grained
56 monzogranite (MMG) based on occurrence of hornblende in these samples ([Hao et al.,](#)
57 [2021](#)) and 3.0 wt% for the fine-grained monzogranite (FMG) and K-feldspar granite
58 (KG) because the hornblende begins to crystallize when the silicate melts contain
59 more than 3.0 wt% H_2O ([Burnham, 1979](#)).

60 **Zircon Lu-Hf Isotopes**

61 Zircon Lu-Hf isotope analyses were performed at the Nanjing FocuMS
62 Technology Co. Ltd. (Nanjing, China) using a Nu Plasma II multi-collector
63 (MC)-ICP-MS (Wrexham, Wales, UK), coupled to a 193 nm excimer laser ablation

64 system (RESOLUTION LR). Each acquisition incorporated 20 s background (gas blank),
65 followed by spot diameter of 50 μm at 9 Hz repetition rate for 40 s. Using exponential
66 correction for mass bias, the measured isotopic ratios of $^{176}\text{Hf}/^{177}\text{Hf}$ were normalized
67 to $^{179}\text{Hf}/^{177}\text{Hf} = 0.7325$. The measured $^{173}\text{Yb}/^{171}\text{Yb}$ and the natural ratio of 1.13268
68 (Chu et al., 2002) were used to calculate the mass bias factor of Yb. In IsotopeMaker,
69 the natural ratios are $^{176}\text{Yb}/^{173}\text{Yb} = 0.79381$ (Segal et al., 2003) and $^{176}\text{Lu}/^{175}\text{Lu} =$
70 0.02656 (Wu et al., 2006). A more detailed analytical technique can be found in Zhang
71 et al. (2015). The initial $^{176}\text{Hf}/^{177}\text{Hf}$ ratios were calculated according to the measured
72 $^{176}\text{Lu}/^{177}\text{Hf}$ ratios and the ^{176}Lu decay constant of $1.867 \times 10^{-11} \text{ year}^{-1}$ (Söderlund et al.,
73 2004), whereas $\varepsilon_{\text{Hf}}(t)$ values were calculated with chondritic values of $^{176}\text{Hf}/^{177}\text{Hf} =$
74 0.0336 and $^{176}\text{Lu}/^{177}\text{Hf} = 0.282785$ (Bouvier et al., 2008). The present-day $^{176}\text{Hf}/^{177}\text{Hf}$
75 $= 0.28325$ and $^{176}\text{Lu}/^{177}\text{Hf} = 0.0384$ (Griffin et al., 2004) were used to define the
76 depleted mantle line. Single-stage Hf model ages (T_{DM}) were calculated with
77 $^{176}\text{Lu}/^{177}\text{Hf} = 0.0384$ (Griffin et al., 2004) relative to the depleted mantle that is
78 assumed to have linear isotopic growth ranging between $^{176}\text{Hf}/^{177}\text{Hf} = 0.279718$ at
79 4.55 Ga and 0.283250 at present, whereas two-stage Hf model ages (T_{DM}^{C}) were
80 calculated assuming a mean value of $^{176}\text{Lu}/^{177}\text{Hf} = 0.015$ for the average continental
81 crust (Griffin et al., 2002).

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Figure S1. Cathodoluminescence images of representative zircons for U-Pb dating and Lu-Hf isotope analyses for the Mangling intrusive complex.

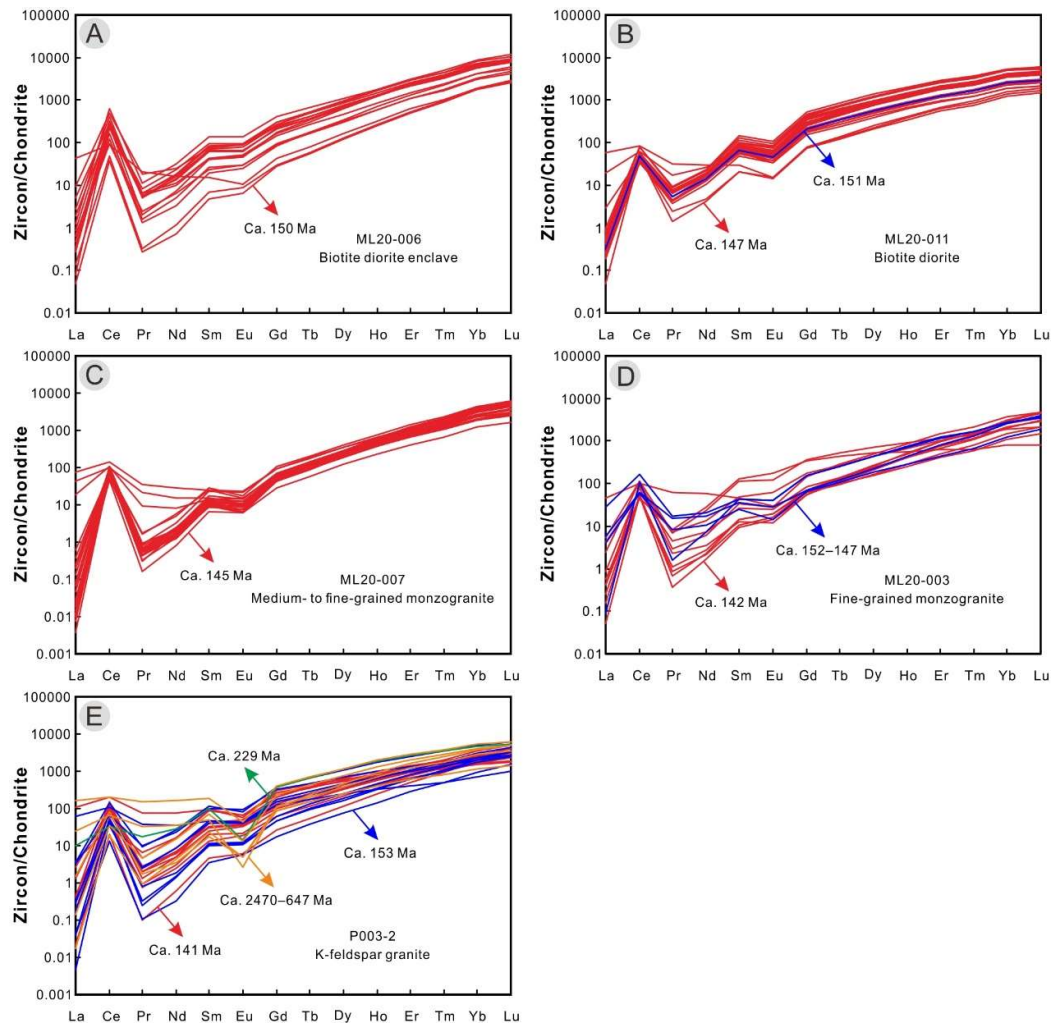


Figure S2. Chondrite-normalized rare earth element patterns of analyzed zircons for the Mangling intrusive complex. Normalizing values are from [Sun and McDonough \(1989\)](#).

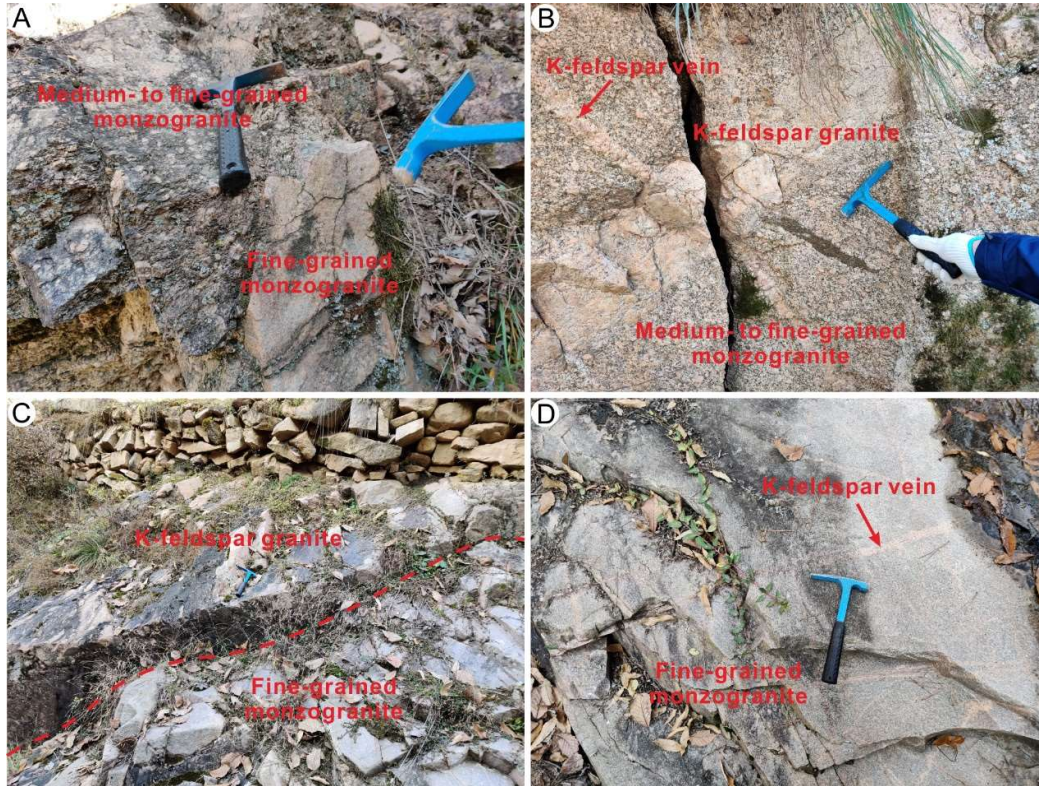


Figure S3. Field photographs showing relationships of the granitic rocks of the Mangling intrusive complex. (A) Medium- to fine-grained monzogranite was intruded by the fine-grained monzogranite. (B) Medium- to fine-grained monzogranite was crosscut by the K-feldspar granite and a K-feldspar vein. (C) K-feldspar granite intrudes into the fine-grained monzogranite. (D) Fine-grained monzogranite was crosscut by K-feldspar veins.

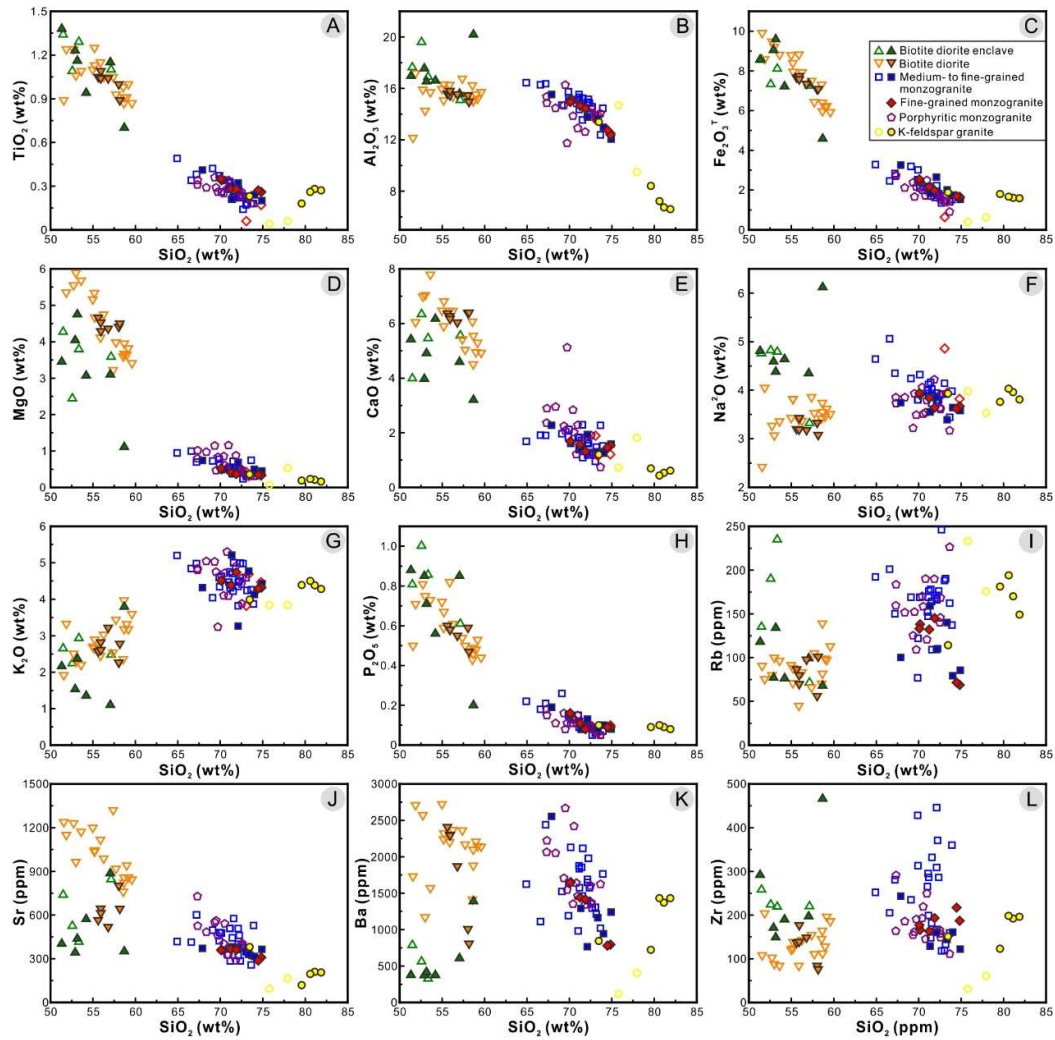


Figure S4. Harker diagrams of selected major and trace elements against SiO_2 for the Mangling intrusive complex. Published data of Mangling dioritic and granitic rocks are same as Fig. 5.

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