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Supplemental Material

Text. Description of the Studied Large River Basins.

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Text. Description of the Studied Large River Basins.

Asia's largest river, the Yangtze River, is divided into three segments (Fig. 3). The headwater Jinsha River originates from the Qiangtang terrane with valleys deeply incising along the thrust contact between the Qiangtang and Songpan–Garze terranes where outcropping rocks are composed of Triassic turbidites and Triassic–Cenozoic granitic intrusions. Entering the Yangtze Craton, the mainstream is joined by several tributaries in the Sichuan Basin that is filled with Triassic–Cretaceous clastic rocks. After passing through the Three Gorges, the slope rapidly decreases and the middle reaches meander across the flat lowlands overlying the Jiangnan Basin that is confined by the Qinling–Dabie Orogen to the north and the Jiangnan Orogen to the south. Farther eastward, the lower Yangtze River drains into the Subei Basin and the East China Sea.

The mainstream of the Pearl River, Xi River, originates from the southeastern plateau margin and confluent with two small tributaries (Bei and Dong Rivers) at the estuary, emptying into the northern South China Sea (Fig. 3). The upper Xi River mainly drains the Upper Paleozoic carbonates and Triassic turbidites of the Nanpanjiang Basin. The lower Xi River is joined by tributaries from the Neoproterozoic Jiangnan Orogen and flows through the Paleozoic strata and Yunkai Massif. To the east, the Bei and Dong Rivers drain Jurassic–Cretaceous intrusions of the East Cathaysia.

The Red River catchment is largely confined to the narrow, southeast-striking Ailao Shan–Red River fault zone (Fig. 3). Surrounding terranes, including Yangtze Craton and West Cathaysia to the northeast and Indochina to the south, render a complex range of outcropping rocks within the catchment. In general, the upstream region is dominated by the Paleozoic and Mesozoic sedimentary and metamorphic rocks, whereas the fault zone is characterized by several elongated high-grade metamorphic massifs and calc-alkaline to alkaline magmatic rocks. The mainstream confluent with two tributaries (Lo and Da Rivers) in the Hanoi Basin before discharging into the northwesternmost South China Sea.

Figures S1–S11.

Figure S1. Detrital zircon U-Pb age distributions of modern river sediments from East Asia, shown as a three-dimensional nonmetric multidimensional scaling (MDS) plot. Large and small rivers from different areas are color-coded. The orientation of the three-dimensional MDS plot is shown by projecting an individual sample (red circle) into three perpendicular planes. *s*—number of compiled samples. See Fig. 1B for sample localities and Table S1 for sampling information and references.

Figure S2. Detrital zircon U-Pb age distributions of modern river sediments from coastal South China (Chen et al., 2019; Li et al., 2021; Xu et al., 2007, 2014a, 2014a; Zhang et al., 2017a), shown as (A) a nonmetric multidimensional scaling (MDS) plot and (B–E) plots of kernel density estimation (KDE) spectra. In the MDS plot, solid and dashed lines mark the closest and the second closest neighbors, respectively. In the KDE plot, samples are grouped as a function of relative similarities of age signatures revealed by the MDS statistics. *s*—number of compiled samples, *n*—number of concordant analyses. See Fig. 1B for sample localities and Table S1 for sampling information and references.

Figure S3. Detrital K-feldspar Pb isotopic compositions of Upper Oligocene–Neogene strata from the northern Yinggehai Basin (Wang et al., 2019a; Zhang et al., 2021). Error bars are not shown for brevity. See Fig. 1B for sample localities and Table S2 for sampling information and references.

Figure S4. Schematic drainage reconstruction for Southeast China before (A) and after (B–D) the provenance shift of the Oligocene–Lower Miocene strata of West Taiwan (modified from Deng et al., 2017; Lan et al., 2016; Wang et al., 2018). Arrows stand for potential sediment dispersal routes. The provenance shift has been related to a drainage expansion of small rivers of coastal Southeast China (e.g., Min River; Lan et al., 2016) and an addition of a paleo–Yangtze River source either by long-distance sediment transport (Deng et al., 2017) or by sediment reworking of the East China Sea Shelf Basin (Wang et al., 2018).

Figure S5. Detrital zircon U-Pb age distributions of Upper Oligocene–Neogene strata from the margin of Hainan Island (Jiang et al., 2015; Lei et al., 2019; Lyu et al., 2021; Wang et al., 2015a, 2016, 2019b; Yan et al., 2011), shown as (A) a nonmetric multidimensional scaling (MDS) plot and (B–F) plots of kernel density estimation (KDE) spectra. In the MDS plot, solid and dashed lines mark the closest and the second closest neighbors, respectively. In the KDE plot, samples are grouped as a function of relative similarities of age signatures revealed by the MDS statistics. *s*—number of compiled samples, *n*—number of concordant analyses. See Fig. 1B for sample localities and Table S1 for sampling information and references.

Figure S6. Detrital zircon U-Pb age distributions of modern sediments from small rivers surrounding the northwestern South China Sea (Cao et al., 2015; Fyhn et al., 2019; Gong et al., 2021; He et al., 2020; Jonell et al., 2017; Usuki et al., 2013; Wang et al., 2015b, 2018a; Xu et al., 2014b), shown as (A) a nonmetric multidimensional scaling (MDS) plot and (B–D) plots of kernel density estimation (KDE) spectra. See Fig. S2 for legend description.

Figure S7. Detrital zircon U-Pb age distributions of Upper Cretaceous–Lower Paleogene strata from a range of inland basins along the eastern and southeastern margin of the Tibetan Plateau, using the compiled dataset (Chen et al., 2017b; Li et al., 2018; Wang et al., 2014b; Zhao et al., 2021) of Zhao (2021). Note that their raw dataset is slightly modified by excluding replicate samples (CX-25-2 and CX-31-2; Zhao et al., 2021) and implementing the strategy of this study in data treatment. Samples are grouped into stratigraphic intervals and compared by kernel density estimation (KDE) spectra, with grey bars marking different ages or magnitudes of key peaks within individual basins. s—number of compiled samples, n—number of concordant analyses.

Figure S8. Kernel density estimation (KDE) spectra of detrital zircon U-Pb ages from Paleogene strata of (A) the Jianchuan Basin (Clift et al., 2020; Feng et al., 2021; He et al., 2021; Wissink et al., 2016; Yan et al., 2012; Zheng et al., 2020) and (B) the Simao Basin (Chen et al., 2017b; Wissink et al., 2016). Samples from known formations are grouped. In the Jianchuan Basin, note the spectral difference between the Baoxiangsi–Jinsichang formations and the Shuanghe–Jianchuan formations. Samples Shi-12-02, Shi-12-01, Lij-12-01 have no age control and are interpreted to be equivalents of the Shuanghe–Jianchuan formations due to the presence of Paleogene zircons therein. In the Simao Basin, note that the age KDE spectrum of the Mengla Formation does show a multimodal pattern, contrasting with the original illustration of Chen et al. (2017b) that generally displays a unimodal pattern of age probability density spectrum peaking at ca. 220 Ma. Sample Jing-13-01 is tentatively assigned to the Mengla Formation because of their highly comparable age spectra.

Figure S9. Comparisons of detrital zircon age signatures among the modern sediments of the upper Yangtze River and Red River and the Paleogene formations of areas stretching from the Jianchuan Basin in the northwest to the Na Duong–Cao Bang Basins in the southeast, shown as (A) a nonmetric multidimensional scaling (MDS) plot and (B) plots of kernel density estimation (KDE) spectra. See Fig. S5 for legend description and Fig. 17 for the difference between the Na Duong–Cao Bang Basins (Clift et al., 2020) and the other datasets shown here.

Figure S10. Comparisons of detrital zircon age signatures of Upper Oligocene–Neogene strata among the Red River Basin (Clift et al., 2020; Hoang et al., 2009; Wissink et al., 2016), the Yinggehai Basin, and the Qiongdongnan Basin, shown as (A) a nonmetric multidimensional scaling (MDS) plot and (B–D) plots of kernel density estimation (KDE) spectra. See Fig. S5 for legend description and Fig. 14 for the dataset of the Yinggehai and Qiongdongnan Basins.

Figure S11. Cenozoic variation of sedimentation rate in (A) the East China Sea Shelf Basin (Clift, 2006), (B) the Pearl River Mouth Basin (Clift, 2006), and (C) the Yinggehai Basin (Hoang et al., 2010). Eo.–Eocene; Olig.–Oligocene; Mio.–Miocene; Plio.–Pliocene; L.–Lower; M.–Middle; U.–Upper.

Tables S1–S4.

Table S1. Compilation of detrital zircon U-Pb geochronology data from Cenozoic strata of key inland basins of East Asia and offshore basins of the China Seas as well as modern sediments in different rivers and offshore marine. Samples are listed in an order based on geographic localities in individual basins and rivers. For Cenozoic strata of the East and South China Seas, data are derived from the Xihu Sag of the East China Sea Shelf Basin (Wang et al., 2018c; Yang et al., 2006; Zhang et al., 2018), the Western Foothills and Hsuehshan Range of West Taiwan (Chen et al., 2017a; Lan et al., 2016; Yan et al., 2018), the northern sags of the Pearl River Mouth Basin (Cao et al., 2018; Fu et al., 2021; He et al., 2020; Shao et al., 2016, 2018; Wang et al., 2018b, 2017; Zhao, 2015), the western and central Qiongdongnan Basin (Lei et al., 2019; Lyu et al., 2021; Shao et al., 2016, 2018; Yan et al., 2011), as well as the Yinggehai Basin (Cao et al., 2015; Jiang et al., 2015; Wang et al., 2014a, 2015a, 2016, 2018a, 2019b, 2020b; Xie et al., 2016; Yan et al., 2011). For Cenozoic strata of inland basins, data are derived from the Jiangnan Basin (Sun et al., 2018; Wang et al., 2014c; Yang et al., 2019), Nanjing area of the Subei Basin (Zheng et al., 2013), the Jianchuan Basin (Clift et al., 2020; Feng et al., 2021; He et al., 2021; Wissink et al., 2016; Yan et al., 2012; Zheng et al., 2020), the Simao Basin (Chen et al., 2017b; Wissink et al., 2016), the Red River Basin (Clift et al., 2020; Hoang et al., 2009; Wissink et al., 2016), and the Na Duong–Cao Bang Basins (Clift et al., 2020). For river sediments, data are derived from the Yangtze River (He et al., 2013, 2014; Hoang et al., 2009; Huang et al., 2020; Kong et al., 2009, 2011; Wang et al., 2020a; Yang et al., 2012), Southeast South China (Chen et al., 2019; Li et al., 2021; Xu et al., 2007, 2014a, 2014a; Zhang et al., 2017a), West Taiwan (Chen et al., 2019; Deng et al., 2017), the Pearl River (He et al., 2020; Liu et al., 2017; Zhao et al., 2015), Southwest South China (He et al., 2020), Hainan Island (Cao et al., 2015; Gong et al., 2021; Wang et al., 2015b; Xu et al., 2014b), the Red River (Clift et al., 2006a; Fyhn et al., 2019; Hoang et al., 2009; Wang et al., 2018a; Wissink et al., 2016), and Central Vietnam (Fyhn et al., 2019; Jonell et al., 2017; Usuki et al., 2013; Wang et al., 2018a). For marine sediments, data are derived from the East China Sea (Huang et al., 2020) and the northern South China Sea (Huang et al., 2020; Xu et al., 2016; Zhong et al., 2017). No extra common Pb correction is applied for the raw data. $^{206}\text{Pb}/^{238}\text{U}$ age is adopted for grains younger than 1000 Ma with discordance test calculated by $[1 - (^{206}\text{Pb}/^{238}\text{U} \text{ age} / ^{207}\text{Pb}/^{206}\text{Pb} \text{ age})] * 100$, and $^{207}\text{Pb}/^{206}\text{Pb}$ age for those older than 1000 Ma with discordance test calculated by $[1 - (^{206}\text{Pb}/^{238}\text{U} \text{ age} / ^{207}\text{Pb}/^{235}\text{U} \text{ age})] * 100$.

Table S2. Compilation of detrital K-feldspar Pb isotope data of modern sediments from the upper Yangtze River (Clift et al., 2008; Zhang et al., 2014, 2017b, 2020) and Red River (Bodet and Schärer, 2001; Clift et al., 2008), and Cenozoic strata from the Hanoi Basin (Clift et al., 2008) and northern Yinggehai Basin (Wang et al., 2019a; Zhang et al., 2021).

Table S3. Compilation of whole-Nd isotope data of Cenozoic strata from West Taiwan (Lan et al., 2014), the southern Pearl River Mouth Basin (Shao et al., 2015; Yan et al., 2018), and the adjacent continental-ocean transition zone (Li et al., 2003), the joint area of the Yinggehai–Qiongdongnan Basins (Yan et al., 2007), as well as the Hanoi Basin (Clift et al., 2006b).

Table S4. A summary of references for thermochronometric and paleoaltimetric results obtained from the eastern and southeastern margin of the Tibetan Plateau.

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Figure S1

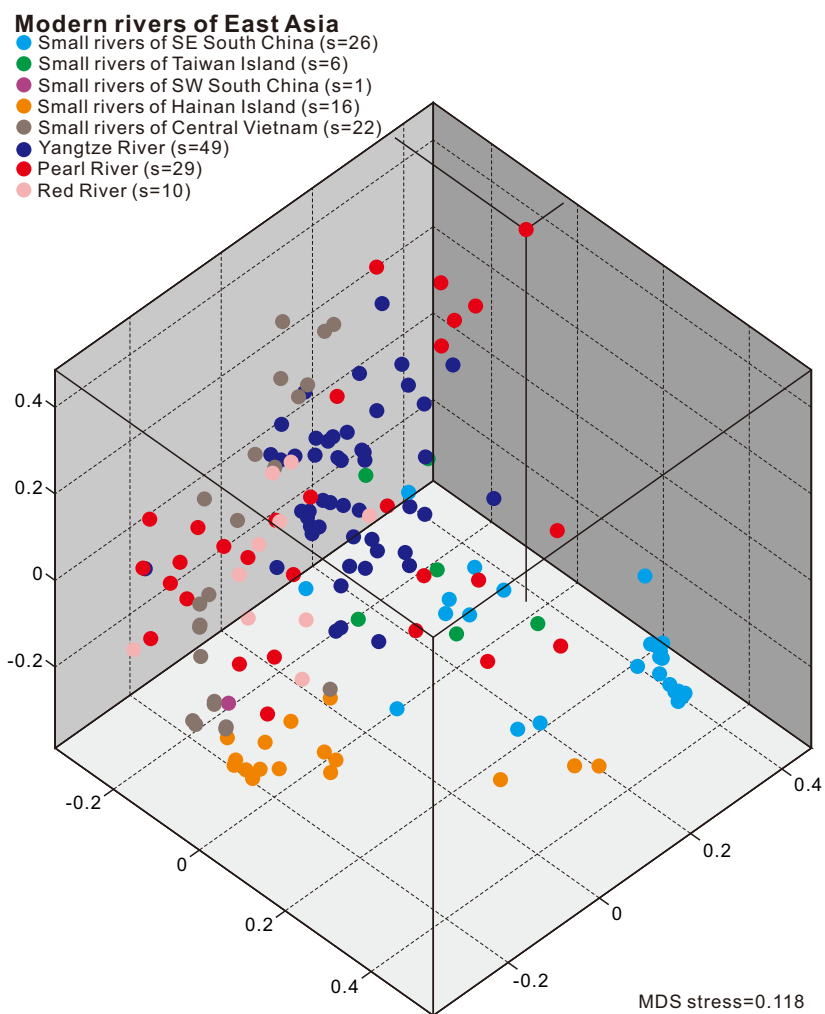


Figure S2

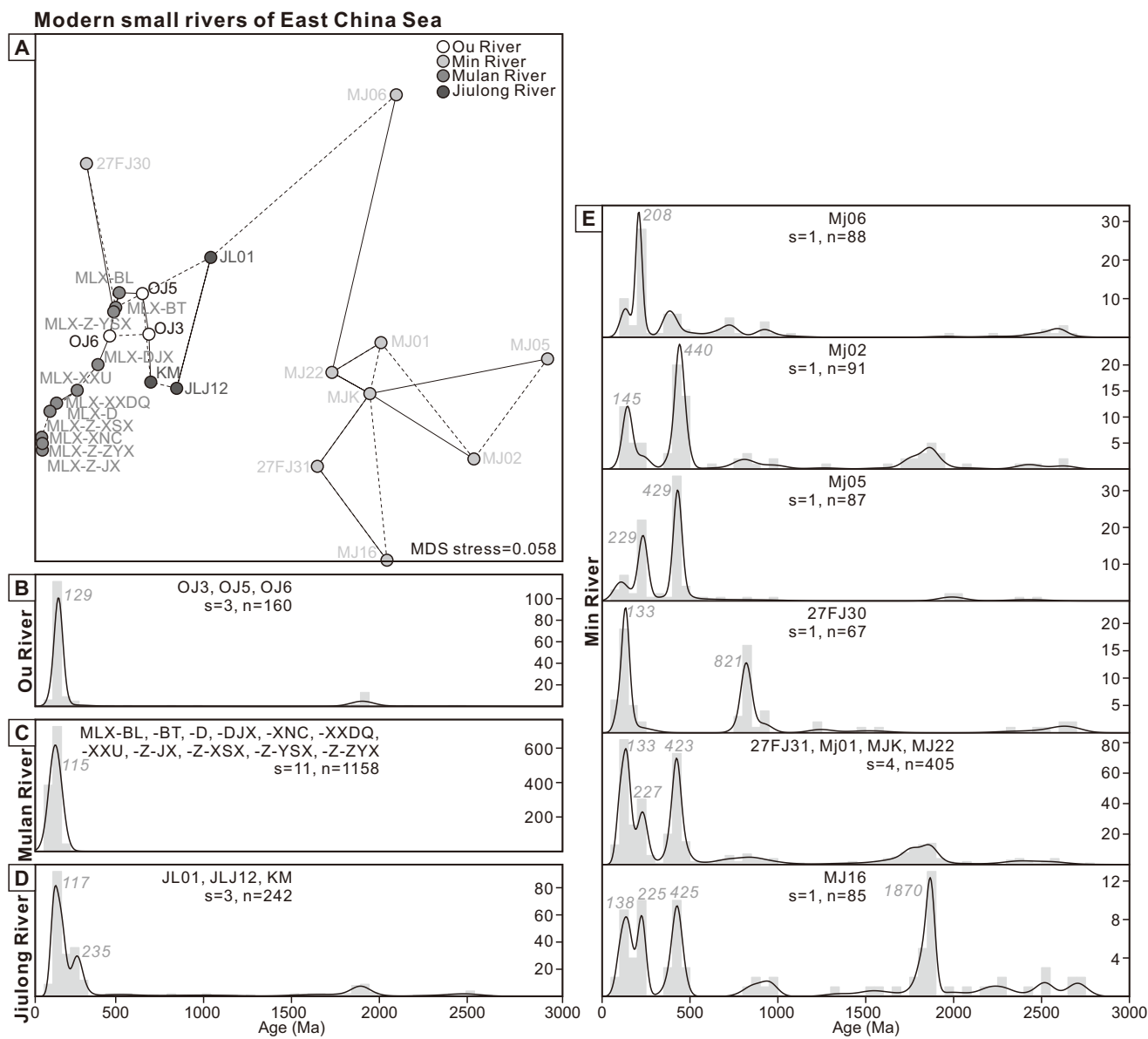


Figure S3

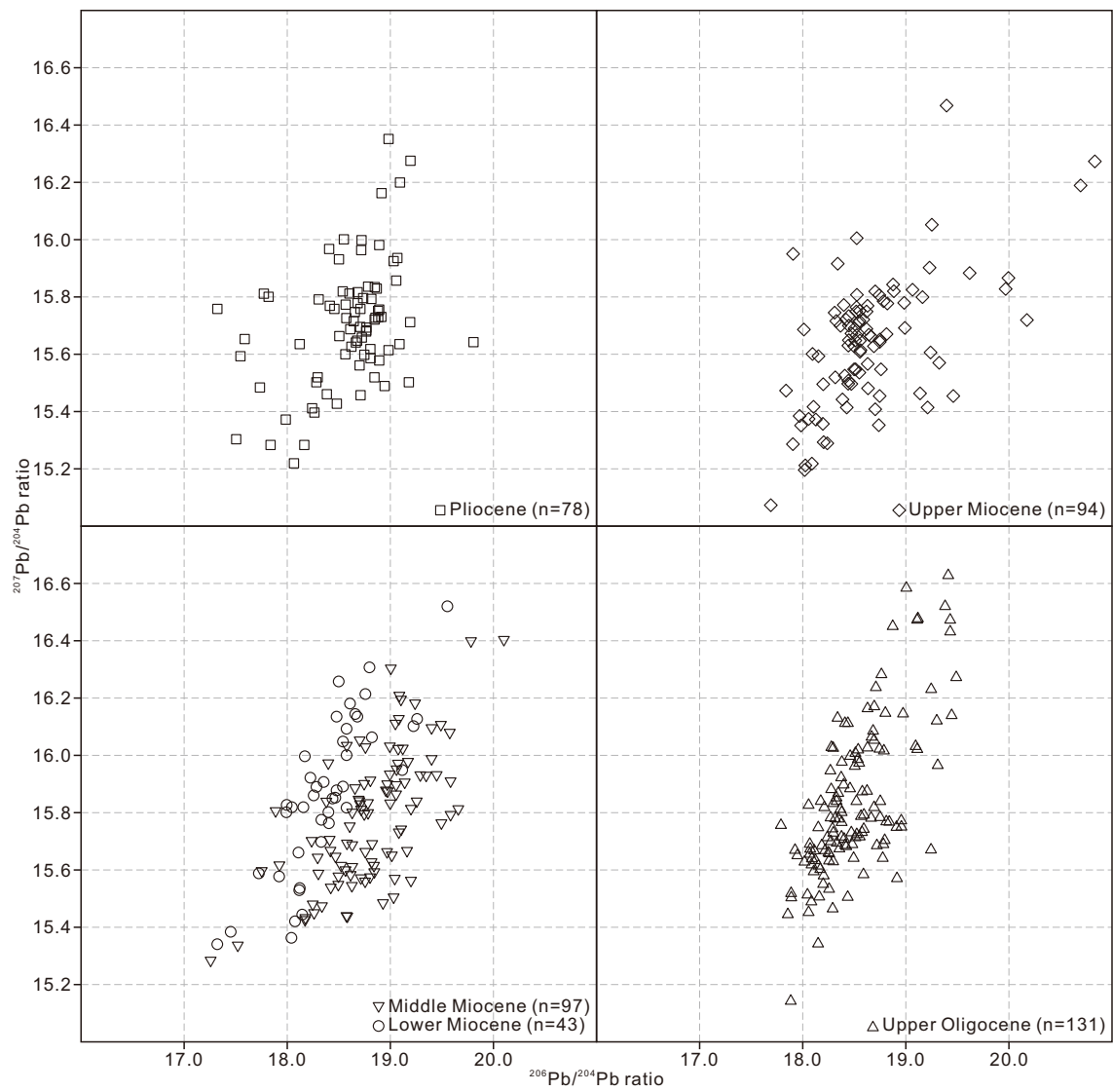


Figure S4

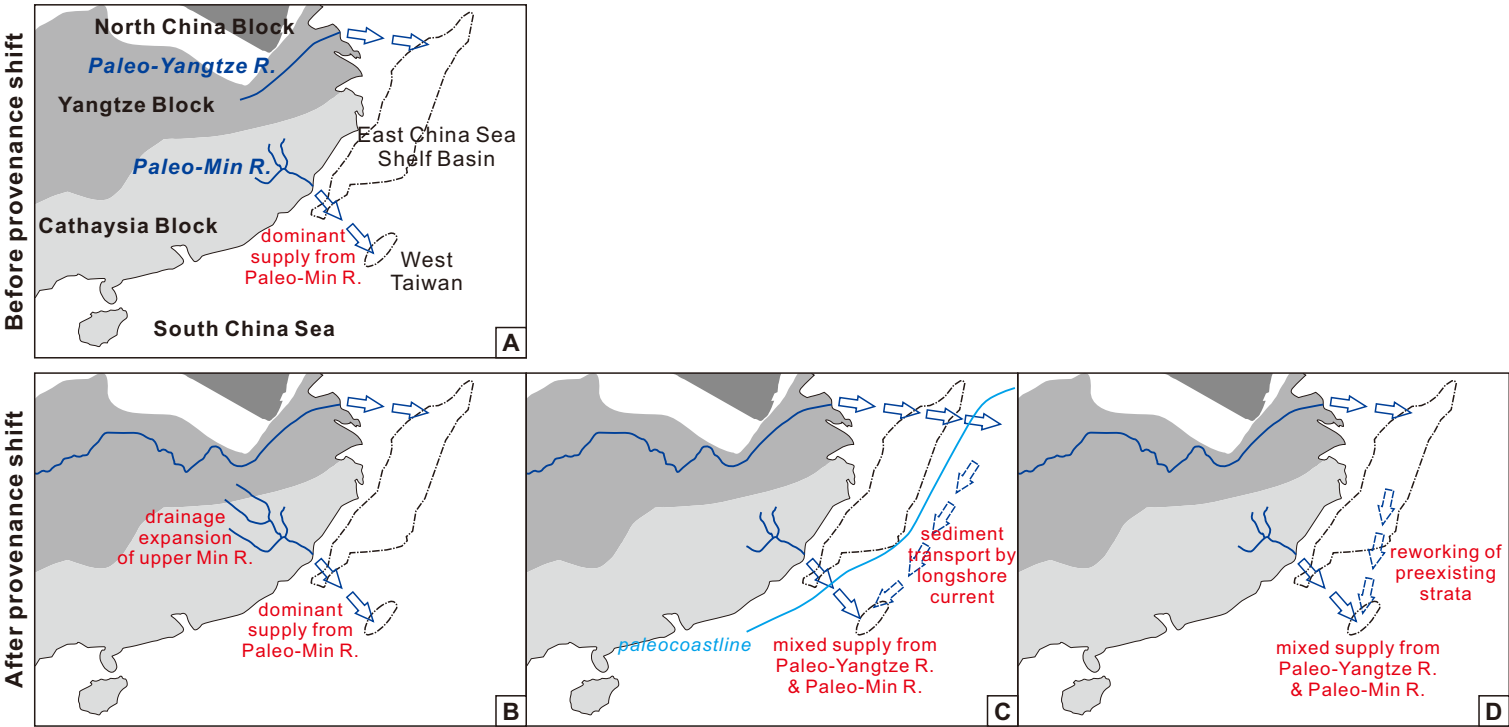


Figure S5

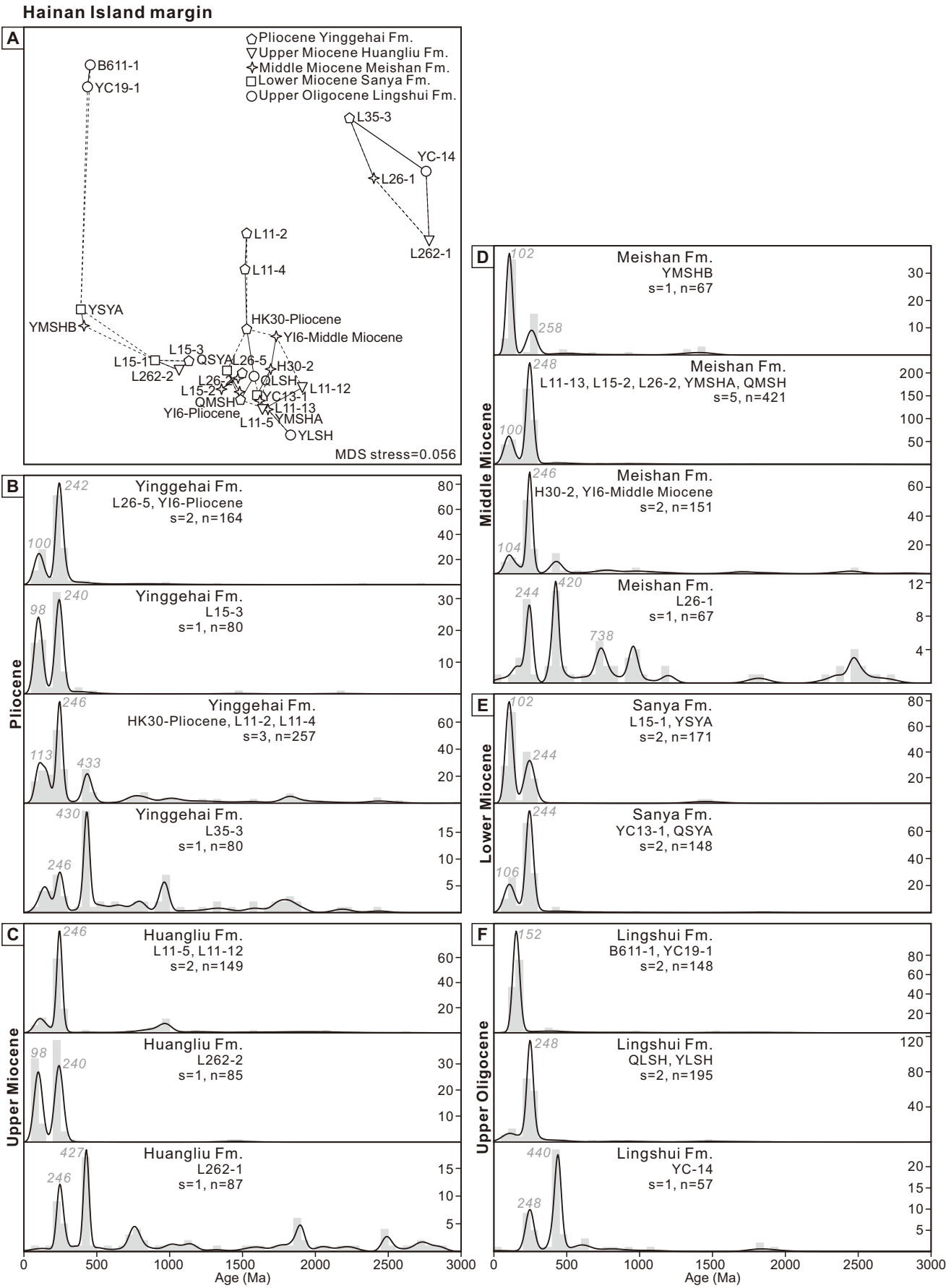


Figure S6

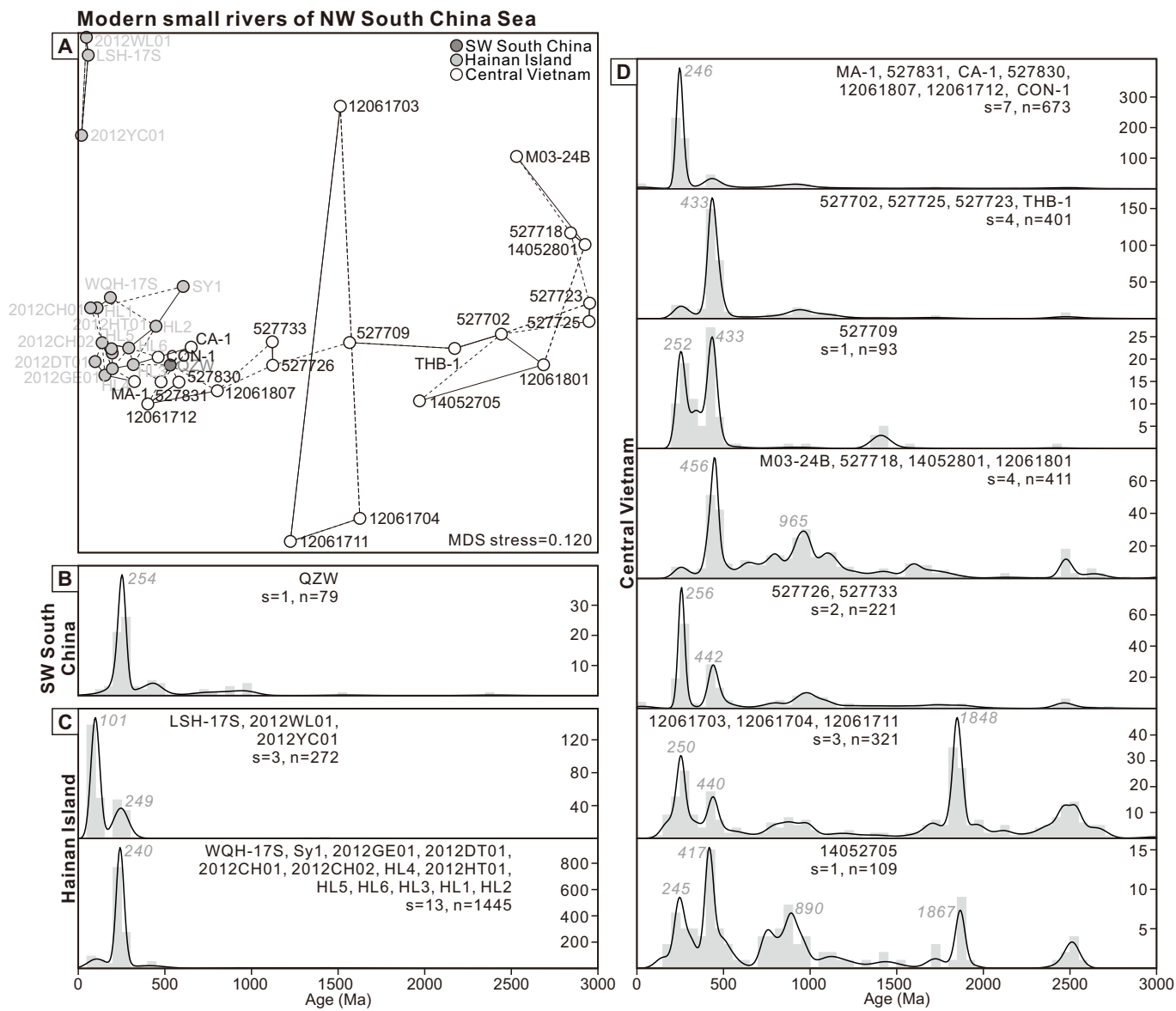


Figure S7

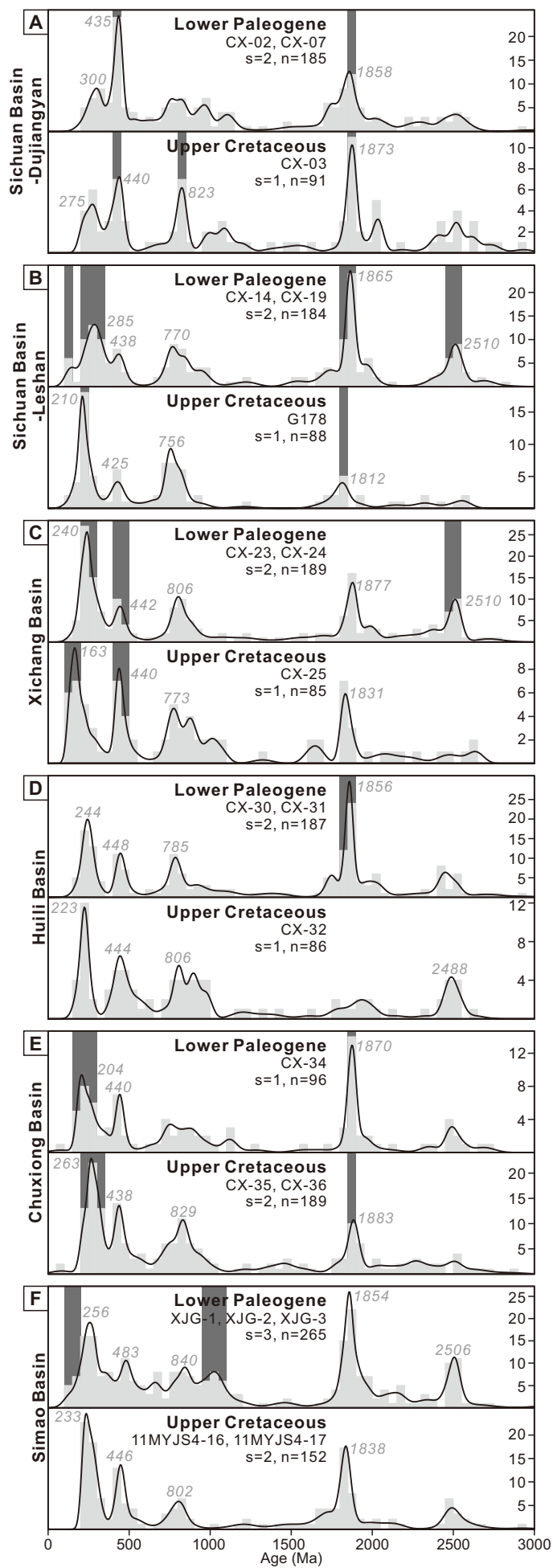


Figure S8

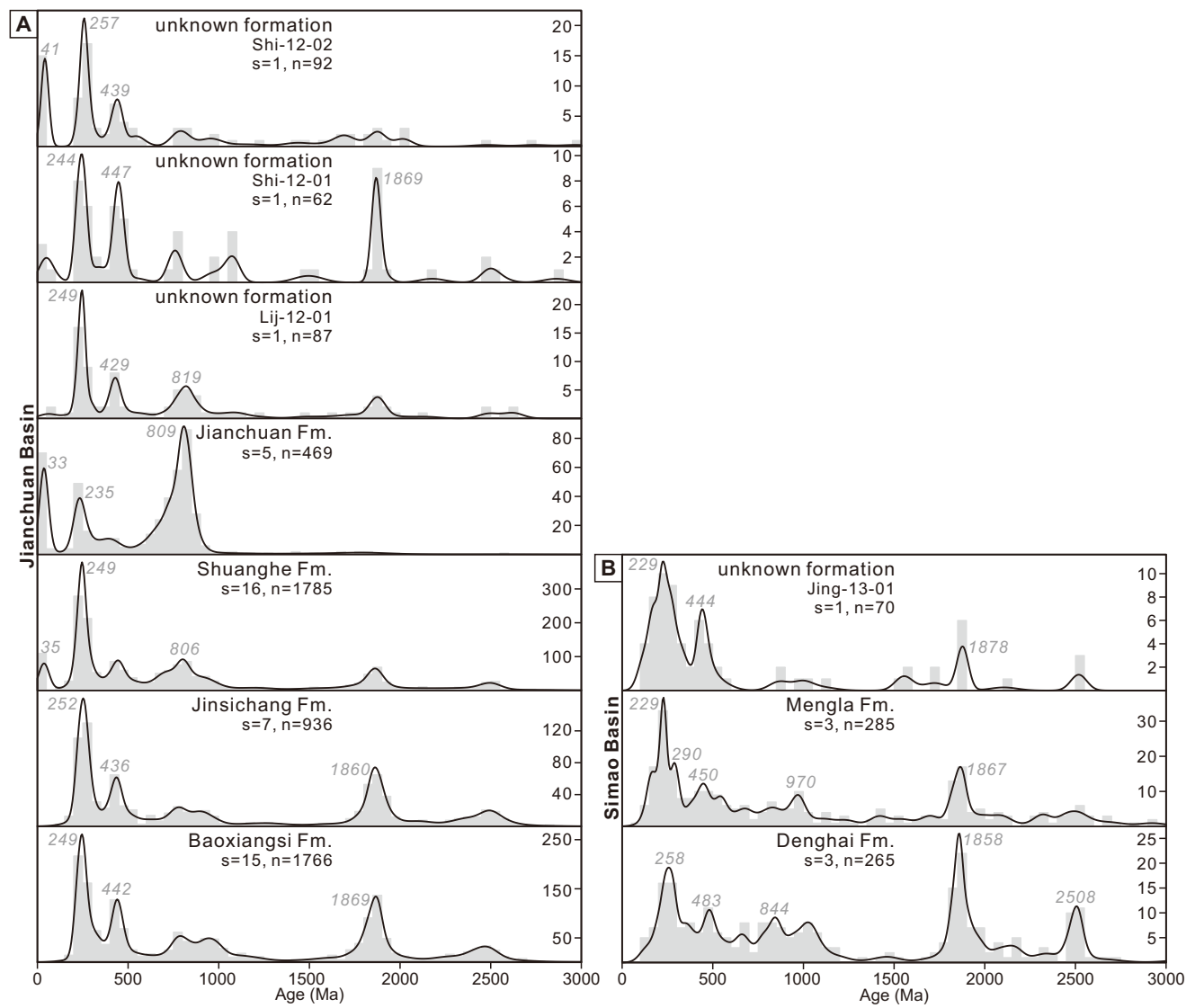


Figure S9

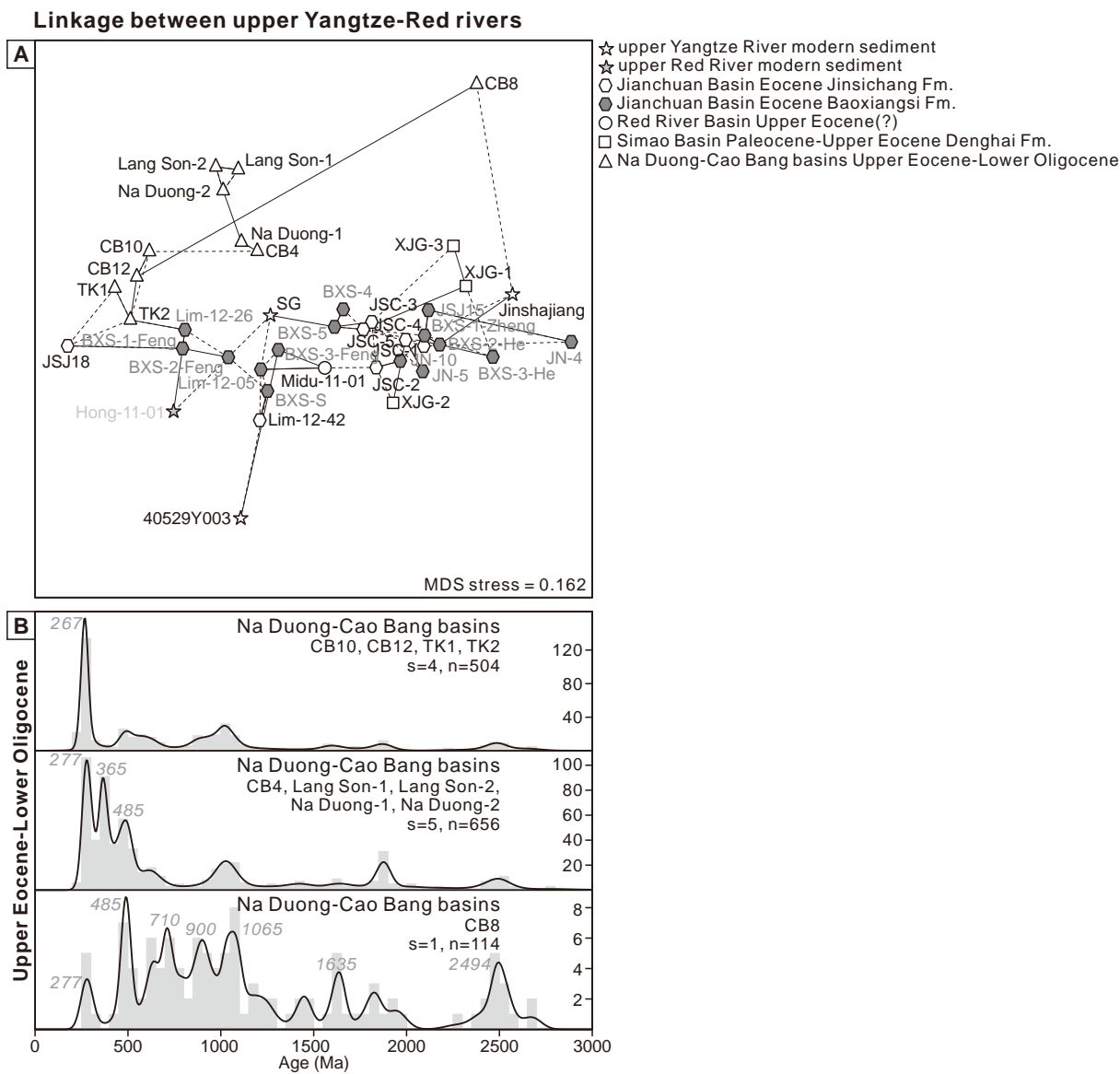


Figure S10

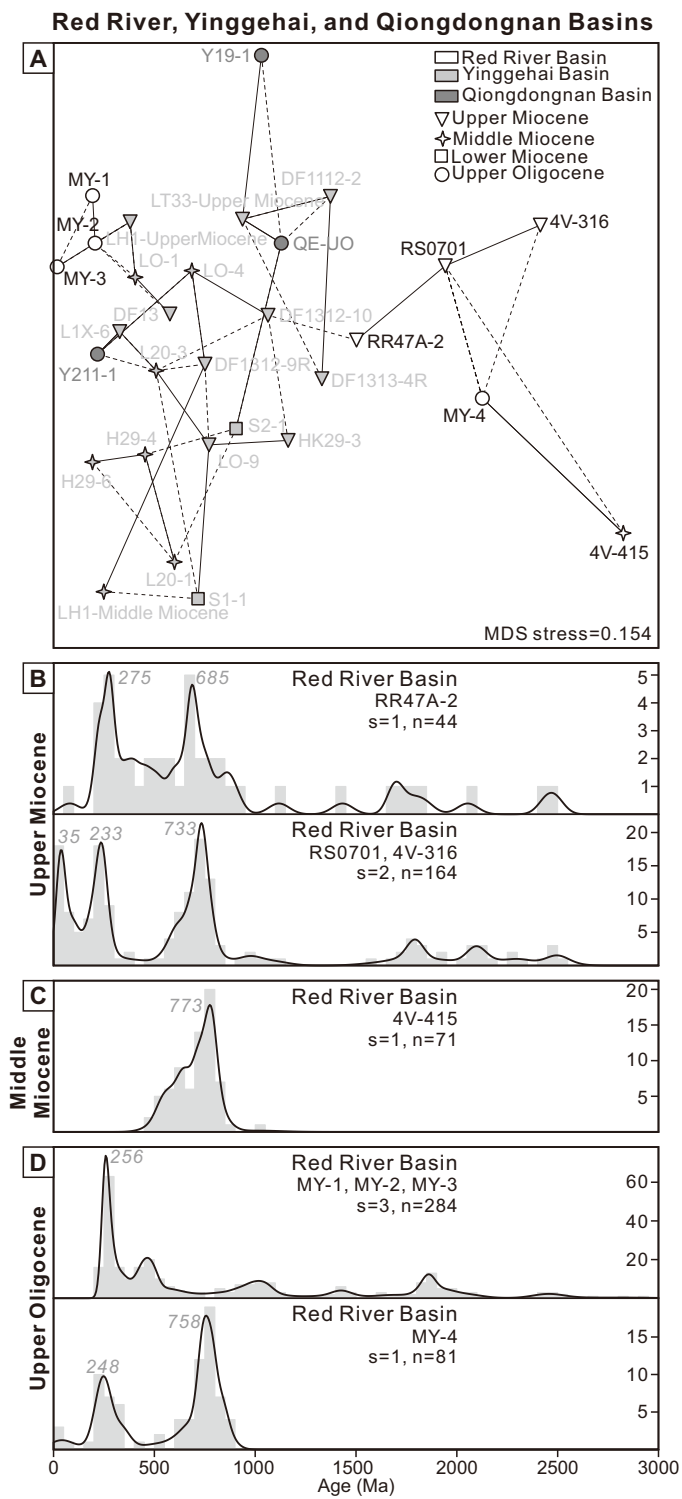


Figure S11

