

Anisotropy-revealed change in hydration along the Alaska subduction zone

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DR1. Table of Split shear wave splitting measurements

DR2. Table of Null shear wave splitting measurements

DR3. Table of station orientations.

Shear Wave Splitting

Shear wave splitting provides constraints on seismic anisotropy in the upper mantle and crust through the fast splitting direction and the delay time [Long and Silver, 2009]. When incoming shear waves interact with anisotropic earth materials, they split into two orthogonally propagating waves that travel at different velocities. The polarization of the faster traveling wave denotes the fast splitting direction and is directly related to the orientation of seismic anisotropy. The time delay between the fast and slow waves is the delay time and relates to both the strength of anisotropy and the layer thickness [Silver and Chan, 1991]. Shear wave splitting is a path integrated measurement sensitive to any anisotropy along each ray path, so splitting alone does not provide mineral source or depth information on anisotropy.

Layered Anisotropy

Multiple layers of anisotropy are likely present in subduction zones: including the sub-slab mantle, the subducting slab itself, the mantle wedge, and the upper plate. Along the AACSE, different stations are sensitive to different portions of the subduction system. OSNAP stations, for example, sample anisotropy in the subducting slab, overriding plate, and in the sub-slab mantle. OSNAP stations sample little-to-no mantle wedge material removing it as a potential influence on shear wave splitting. Constraining different layers of anisotropy with shear wave splitting requires sufficient backazimuthal coverage to model the impacts of 2 or 3 layers of anisotropy [Eakin and Long, 2013]. The high noise levels and short deployment time of the AACSE OBS instruments prevents sufficient coverage of shear wave splitting measurements to model multiple layers of anisotropy, with the majority of AACSE stations returning fewer than 3 measurements.

Previous source-side shear wave splitting measurements, which isolate anisotropy in the sub-slab mantle, show predominantly plate motion aligned splitting beneath the OSNAP stations, both in the Shumagin Gap and in the Semidi segment [Lynner and Long, 2014]. This suggests anisotropy in the sub-slab mantle does not greatly vary beneath the AACSE deployment, removing the sub-slab mantle as the source of the observed splitting change at OSNAP stations. Only the subducting slab and the upper plate, therefore, are the potential drivers of OSNAP splitting. I ascribe the change in OSNAP splitting to the subducting slab since it follows a known tectonic structural change that has previously been suggested to be well hydrated [Shillington et al., 2015]. With the existing splitting measurements, however, I cannot rule out the upper plate as a contributing, or the sole, source of the change in splitting at OSNAP stations. This will be the focus of a future study.

Other layers of anisotropy may exist along the margin. Stress-induced shape preferred anisotropy, for example, has been suggested to contribute to shear wave splitting at convergent margins with high coupling. Changes in interplate coupling between the Shumagin gap and the Semidi segment may lead to variations in stress-induced cracks and subsequent anisotropy. Stress-induced anisotropy, however, is unlikely to be the primary source of the change in anisotropy at OSNAP stations. Stress-induced anisotropy is typically restricted to the upper-crust, which produces ~0.1 to ~0.3 seconds of delay time [Crampin and Chastin, 2003]. This is far below the ~1.0s dt I observe at the OSNAP stations. Stress-induced anisotropy may play a role in the pattern of shear wave splitting along the margin but is unlikely the primary cause of the along strike change in splitting.

References

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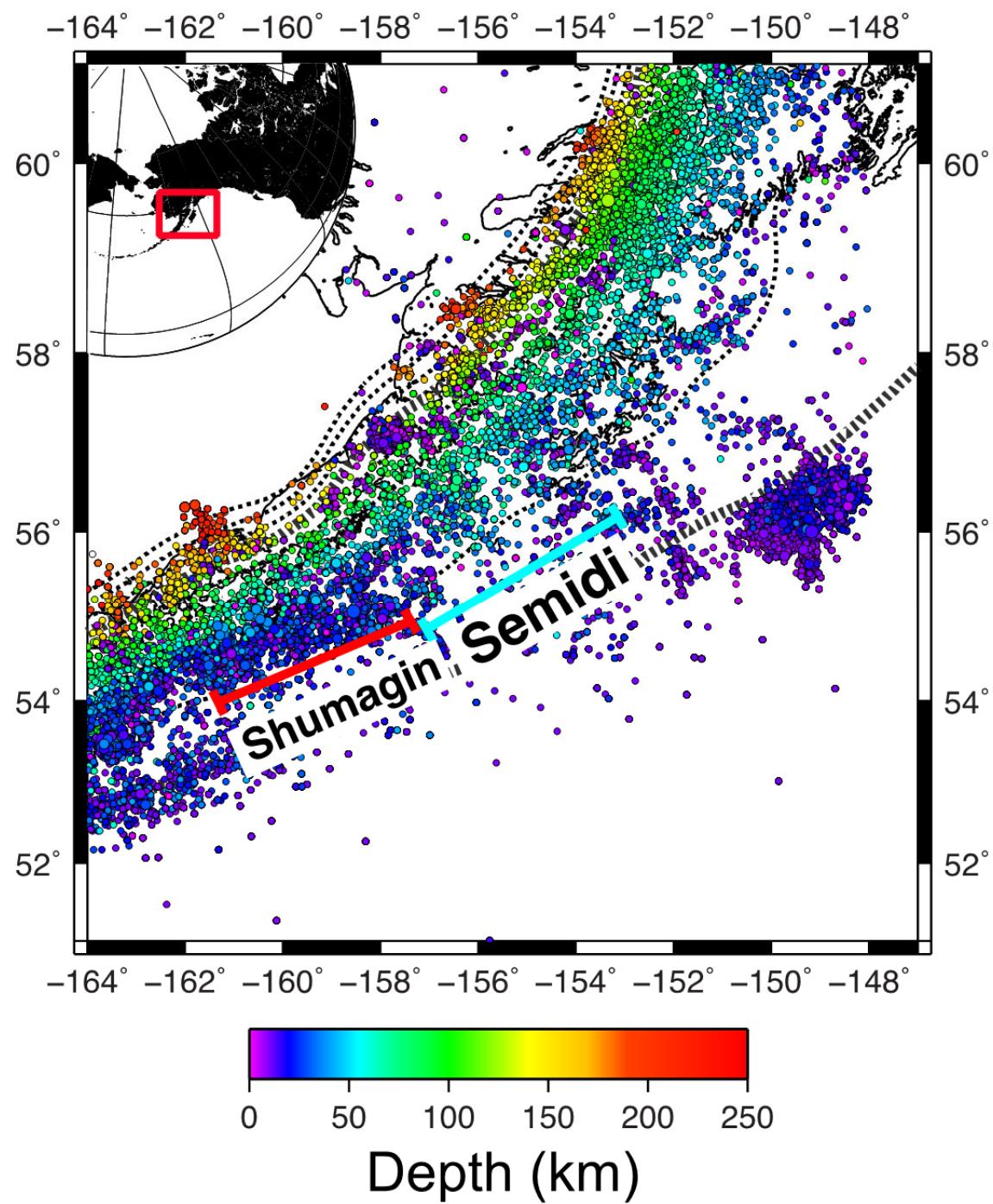


Figure S1. Seismicity along the Alaska-Aleutians subduction zone in the region of study from the National Earthquake Information Center (NEIC).

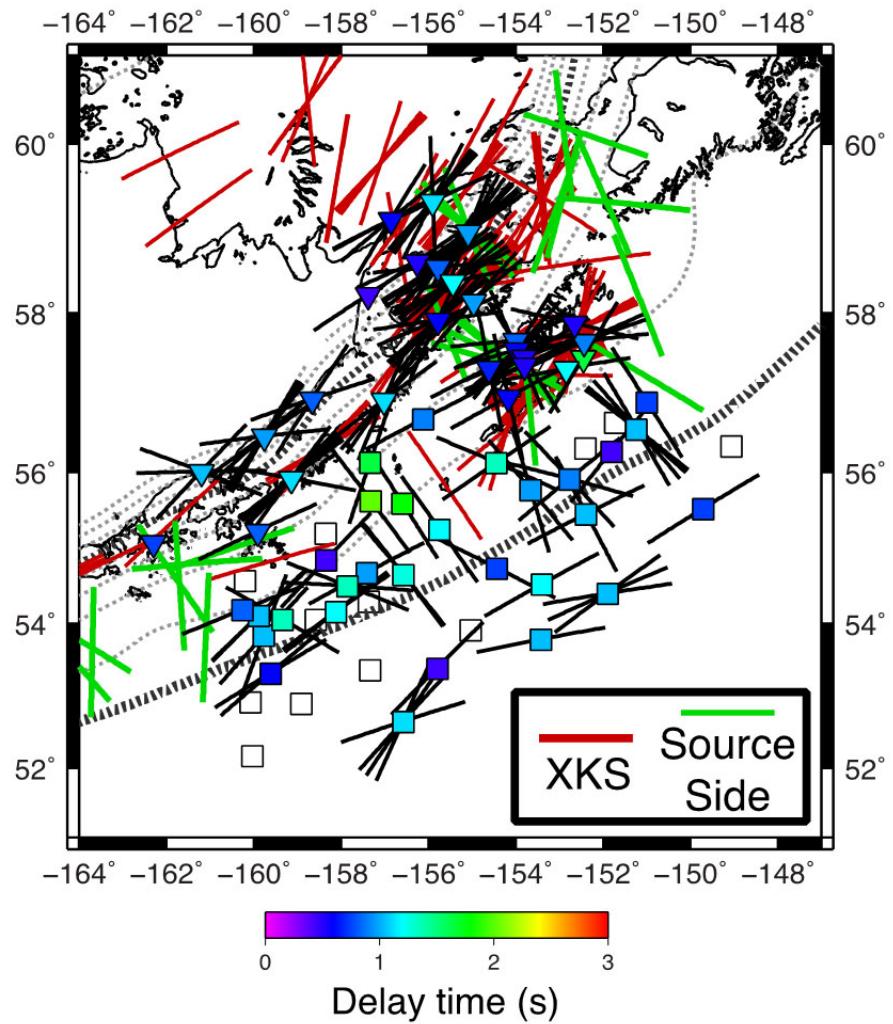


Figure S2. Individual shear wave splitting measurements at AACSE stations along with previous shear wave splitting measurements in the region [Hanna and Long, 2012; Lynner and Long, 2014; Venereau et al., 2019; McPherson, 2020]. Fast splitting directions for AACSE stations are shown by the orientation of the black bars. AACSE station averaged delay times are shown by the station color. Previous XKS splitting measurements are shown as red lines. Source-side measurements, which sample only the sub-slab mantle near the slab, are plotted in yellow.

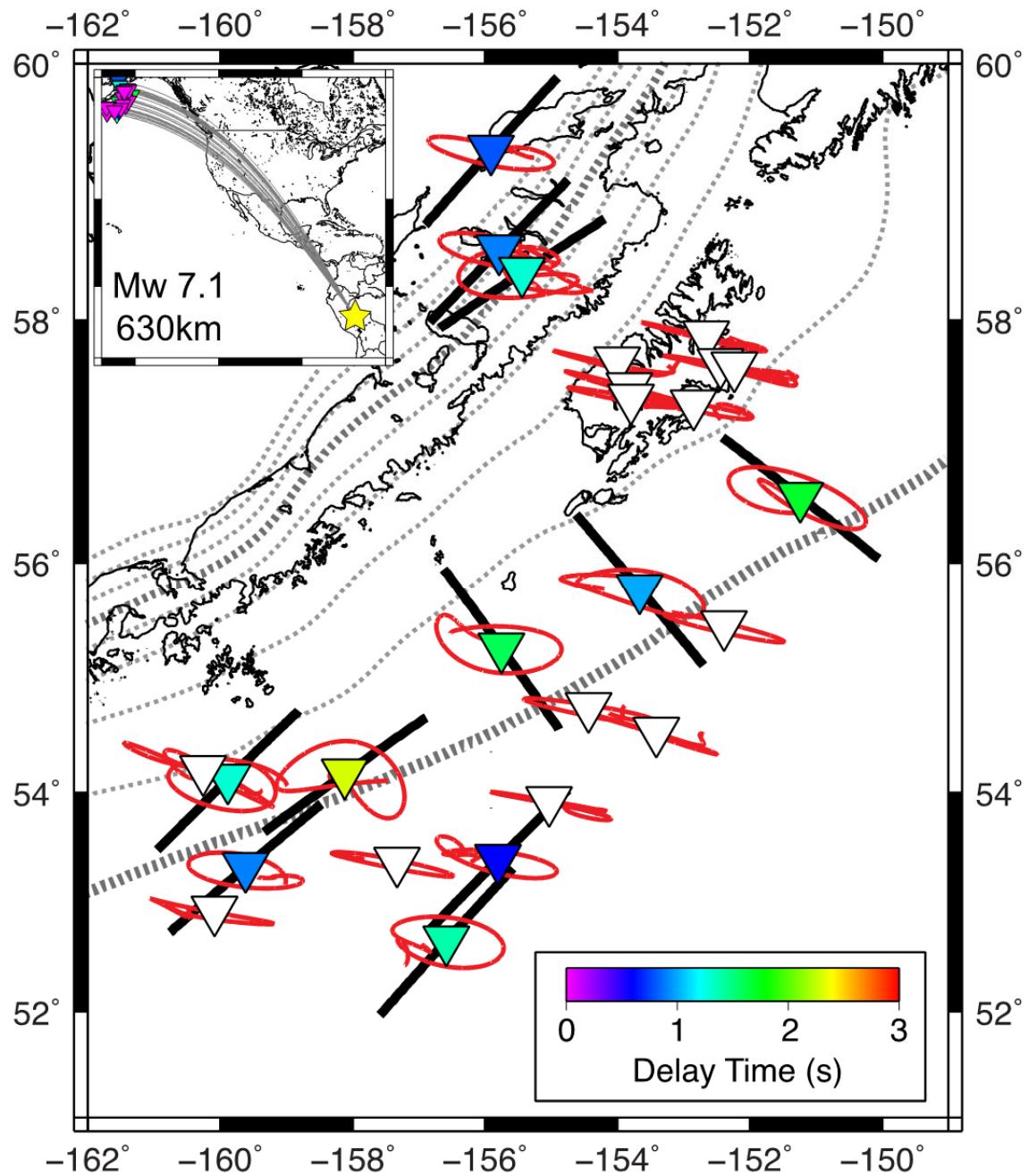


Figure S3. Shear wave splitting at AACSE stations recorded from a South American event on August 24th, 2018. Orientations of the black bars denote the measured fast splitting direction, and the symbol color shows the measured delay time. White stations represent a null measurement. Particle motions of each measurement are shown. Split measurements are characterized by elliptical particle motions, while null measurements have linear particle motions.

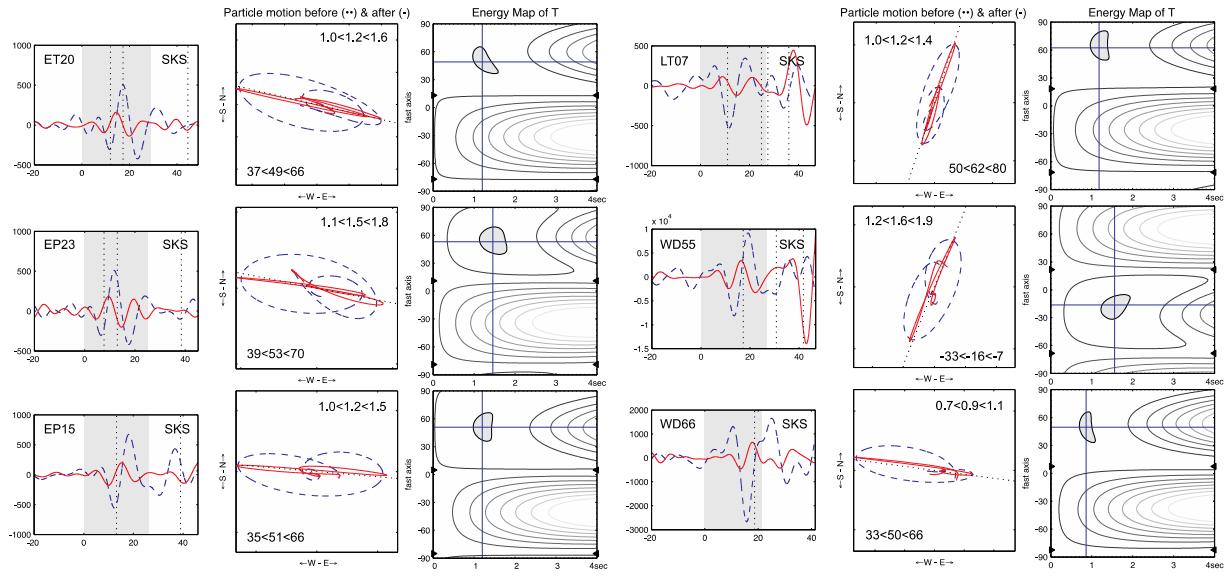


Figure S4. Examples of individual splitting measurements at 6 AACSE stations. (left) Plots of radial (blue) and transverse (red) components for each measurement. The measurement window is highlighted for each. (middle) Initial (blue) and corrected (red) particle motions for the splitting measurements. The resultant fast direction and delay time with associated error bounds are shown in the top and bottom, respectively. (right) Error space plots of the splitting measurements with the 2σ error spaces highlighted. Well resolved split measurements are apparent for several fast directions suggesting the observed differences in splitting directions across AACSE are robust.

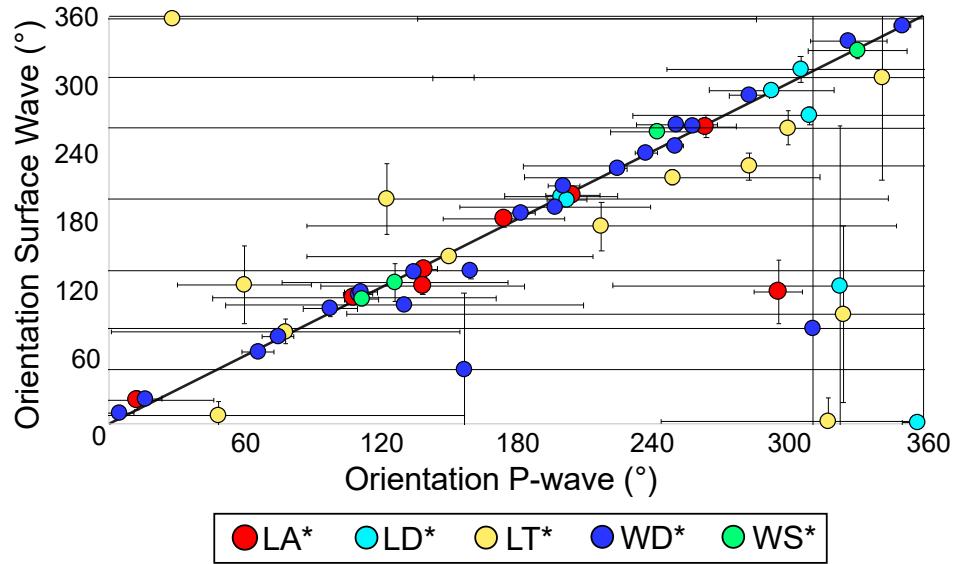


Figure S5. Calculated orientations for each of the AACSE OBS stations using Rayleigh wave [Braunmiller et al., 2020] and P-wave [Doran and Laske, 2017] polarizations with associated errors for the different instrument types. While the calculated orientations typically agree, the surface waves provided better resolved orientations.

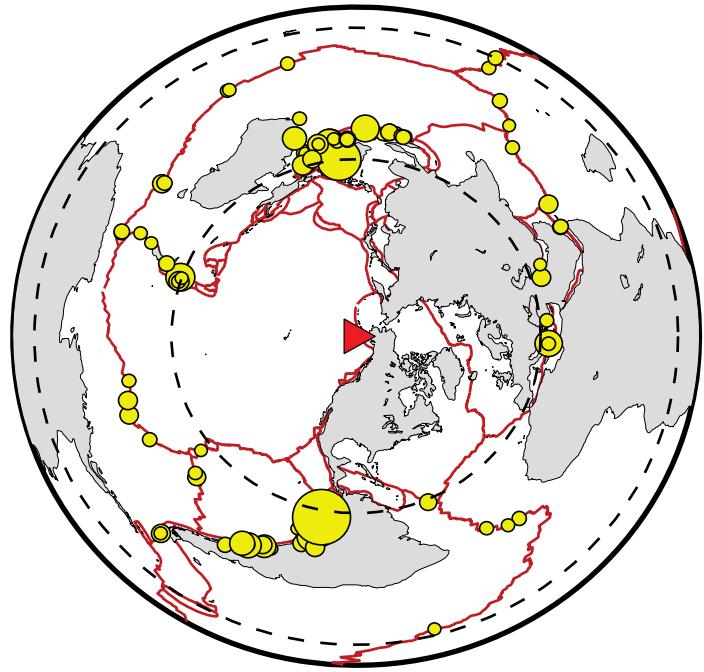


Figure S6. Map of the earthquakes used in this study.

WD70	56.541073	-151.234407	-8.8	114.5	80	268.7	99.7	SKS	-65.7	-41.3	-23.3	1.1	1.7	2.4
WD70	56.541073	-151.234407	0.4	97.7	19	288	101.1	SKS	-49.6	-30	-7.1	0.7	1.1	1.5
LA39	56.8827	-151.0011	-14.7	-70.2	267	105.6	97.4	SKS	-61.7	-32.4	-3	0.6	1	1.7
WD46	55.521393	-149.703732	-0.2	119.8	20	270.2	90.5	SKS	43.5	58.2	71.8	0.8	1	1.4