

Pohl, A., Donnadieu, Y., Godderis, Y., Lanteaume, C., Hairabian, A., Frau, C., Michel, J., Laugie, M., Reijmer, J.J.G., Scotese, C.R., and Borgomano, J., 2020, Carbonate platform production during the Cretaceous: GSA Bulletin, <https://doi.org/10.1130/B35680.1>.

Data Repository

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Table S1. Compilation of neritic carbonate preservation rates.

Table S2. Table of the 89 climatic experiments conducted with the FOAM model. Each experiment was run until deep-ocean equilibrium (≥ 2000 model years). The past 50 years were used to build the monthly climatology files used for the analyses. Model output is available in numerical (netcdf) format at <https://doi.org/10.1594/PANGAEA.904255>.

S1. Geological database of Cretaceous carbonate preservation rates

The database content is shown in Figure 1 in the main text. Mean preservation rates calculated using 5-m.y. time bins are given in Fig. S1. Raw data can be found in Table S1. This new database for carbonate preservation rates was created from the literature. To calculate the rate of preservation we divided the thickness by the stratigraphic interval duration (Bosscher and Schlager, 1993). The overall stratigraphic thicknesses were used. Compaction was not corrected for because it was not sufficiently well constrained. As a result, rates were surely underestimated. Duration estimates came from the stratigraphic boundaries described in the literature (Gradstein et al., 2012).

S2. The GEOCLIM model

The GEOCLIM model is a climate-carbon cycle model that asynchronously couples a carbon cycle box model with a general circulation model. Because the ocean-atmosphere general circulation model could not be run for several million years due to excessive computation time, we first ran climatic simulations for each time slice/paleogeography with several CO₂ levels (70 - 8960 ppm) and gathered them in a catalog of simulations (one catalog per time slice/paleogeography). We then interpolated between the different members of the lookup table all along the 10-million-year-long integration of the carbon cycle model to obtain climatic variables required for the computation of continental weathering (i.e., temperature and runoff) for any CO₂ value within the covered CO₂ range.

The carbon cycle box model is the COMBINE model upgraded following Donnadieu et al. (2006). The oceans are modeled by 9 “boxes”. Four boxes represent the photic and deep ocean zones for the northern and southern high-latitude regions. Three boxes represent the low to mid-latitude open-ocean photic zone, thermocline, and deep ocean. Two boxes represent the low to mid-latitude photic zone and deep epicontinental reservoirs. A 10th box represents the atmosphere.

The climatic component of GEOCLIM is the FOAM ocean-atmosphere model version 1.5 (Jacob, 1997), a mixed-resolution general circulation model. The atmospheric component of FOAM is a parallelized version of the National Center for Atmospheric Research’s (NCAR) Community Climate Model 2 (CCM2) with radiative and hydrologic physics upgrades after CCM3 version 3.2. It uses a R15 spectral resolution ($4.5^\circ \times 7.5^\circ$) with 18 levels in the vertical dimension. The ocean model is the Ocean Model version 3 (OM3). It is a 24-level z-coordinate ocean general circulation model, run at a $1.4^\circ \times 2.8^\circ$ resolution.

Our setup of the GEOCLIM model is very similar to the one previously used by Donnadieu et al. (2006), with the following exceptions. We used the version of the FOAM model that resolves ocean dynamics, whereas Donnadieu et al. (2006) used the slab mixed-layer model. We followed the recent study by Suchéras-Marx et al. (2019) by considering that calcifying plankton was already well-established in open-ocean settings during the Early Cretaceous (Ridgwell, 2005; Ridgwell and Zeebe, 2005). Lastly, we did not rely on Walker et al. (2002) to provide an estimate of the area of shallow-marine carbonates (see Equation 1 in main text), rather, we assumed that the area of the carbonate platforms varied proportionally with the area of the shallow-marine environments available to carbonate deposition during the Cretaceous (Pohl et al., 2019) and derived A_{platform} from the paleogeographical maps.

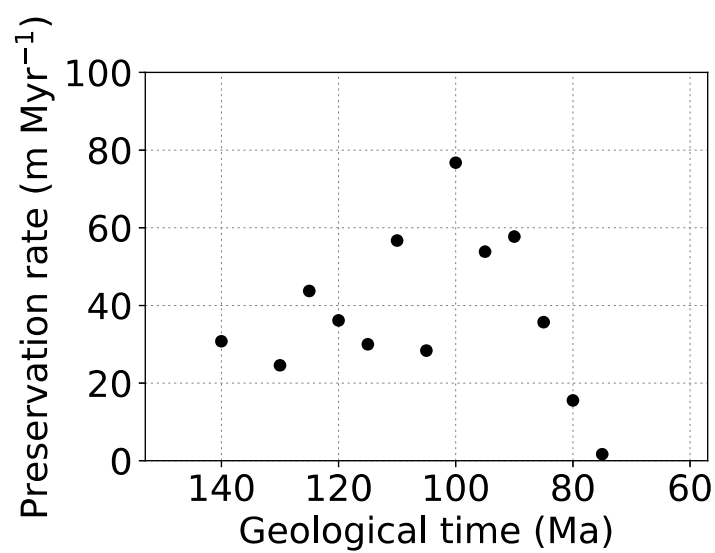


Figure S1. Mean preservation rates calculated using time bins of 5 m.y.

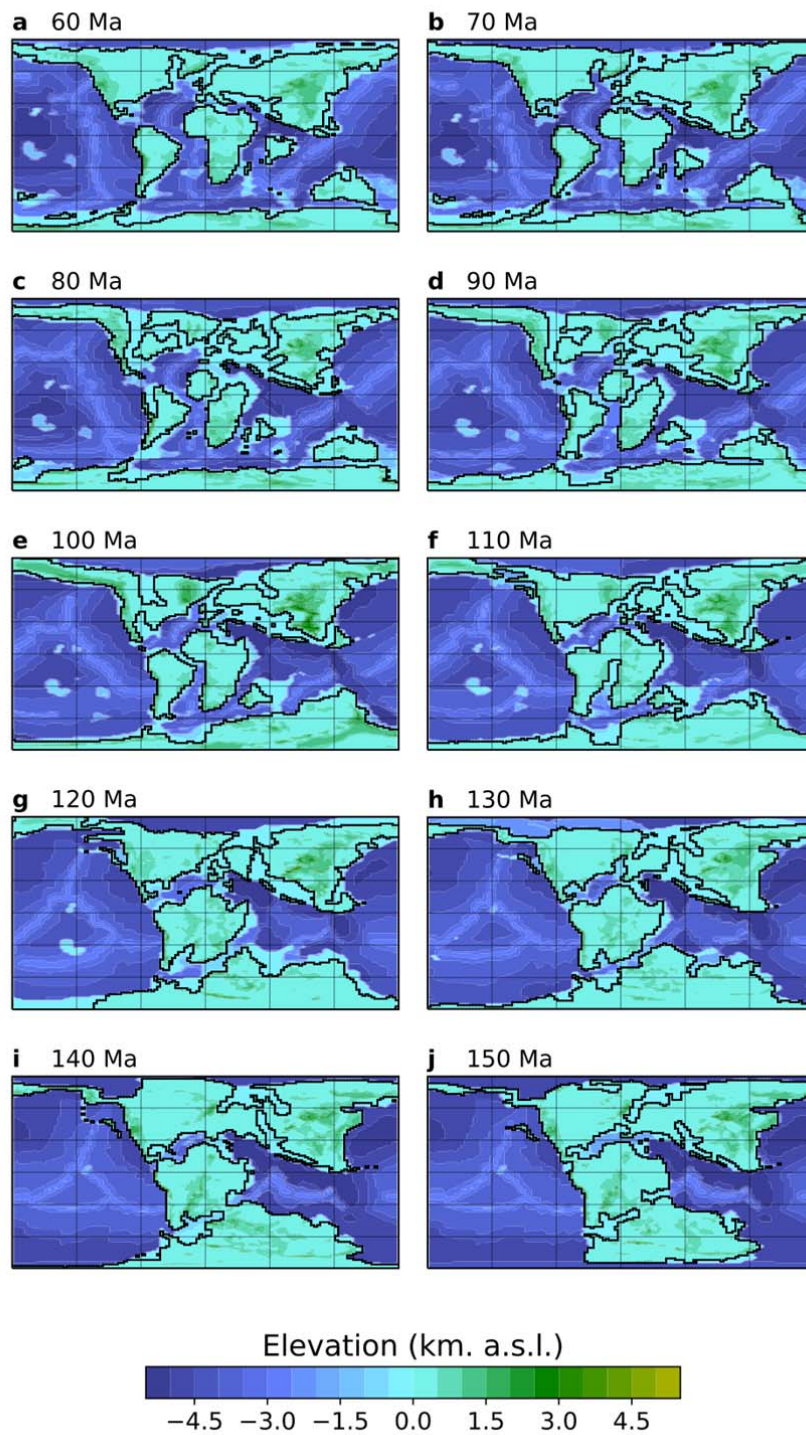
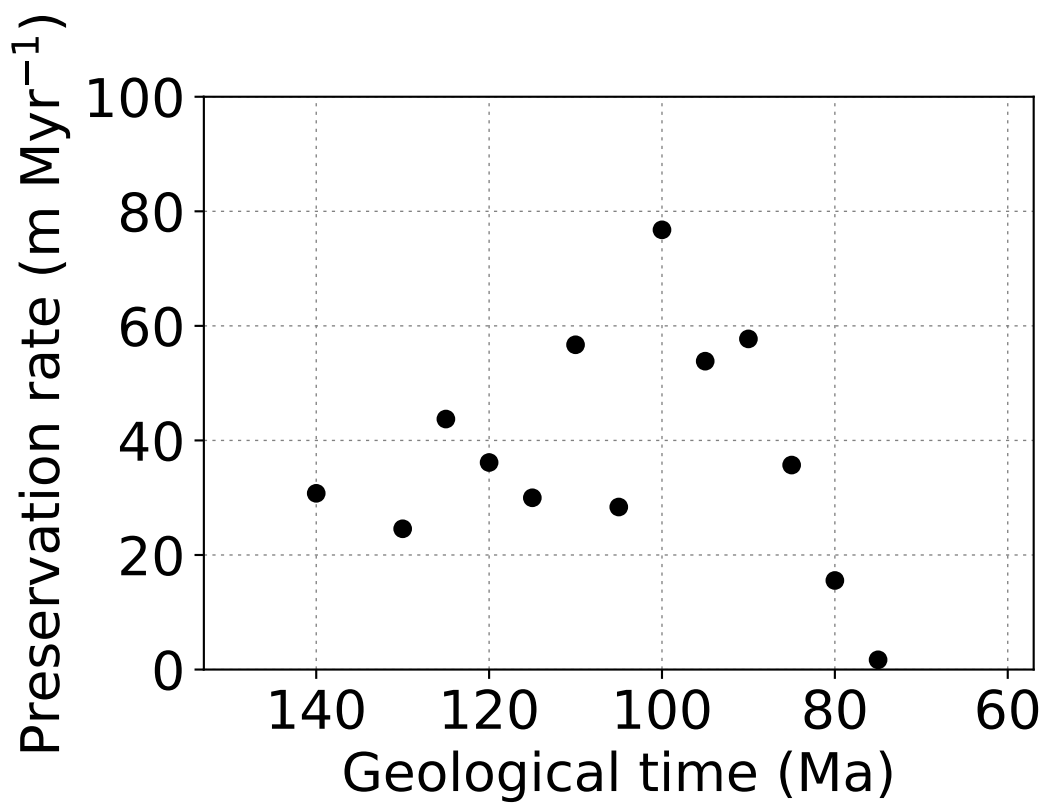


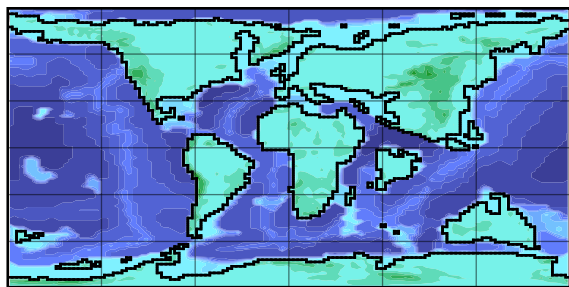
Figure S2. Continental reconstructions used in the GEOCLIM model (Scotese, 2016; Scotese and Wright, 2018).

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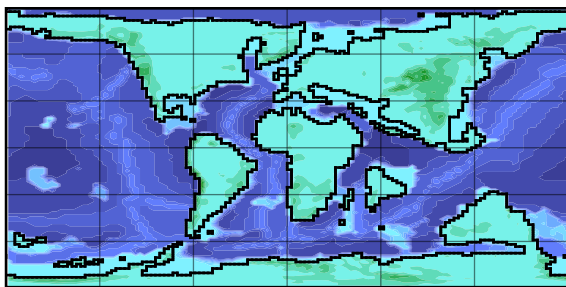
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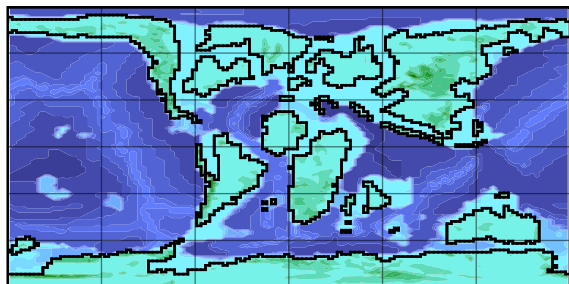
a 60 Ma



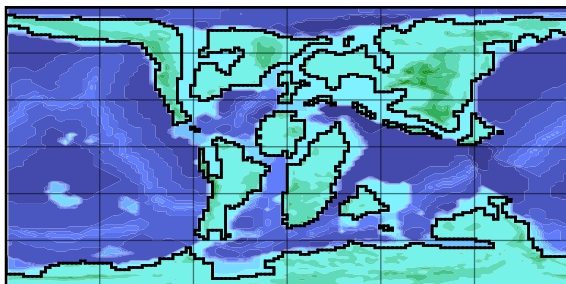
b 70 Ma



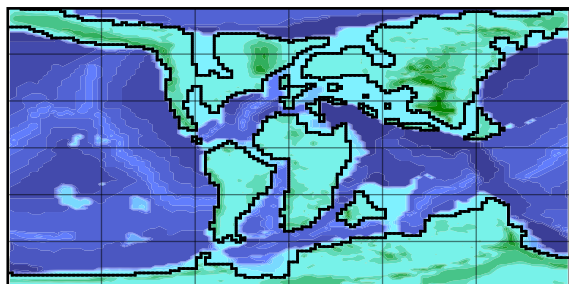
c 80 Ma



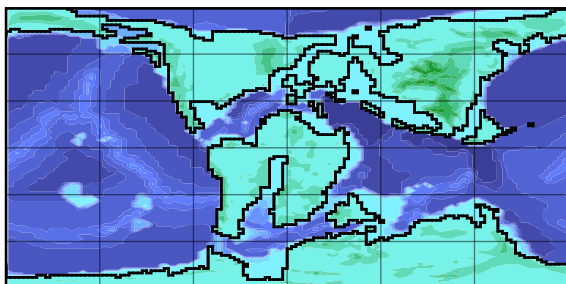
d 90 Ma



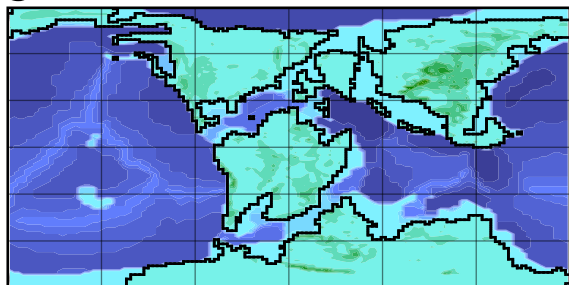
e 100 Ma



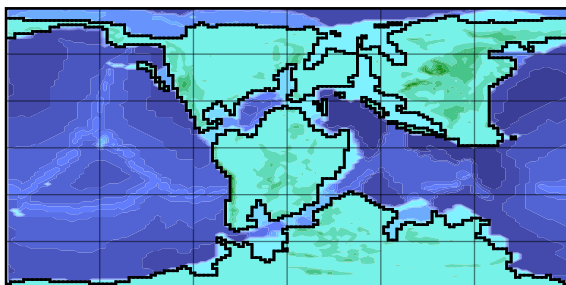
f 110 Ma



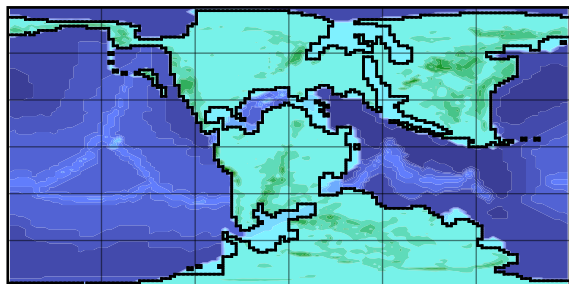
g 120 Ma



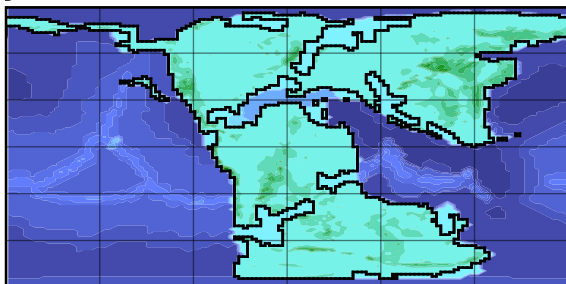
h 130 Ma



i 140 Ma



j 150 Ma



Elevation (km. a.s.l.)



| Experiment name | Globally-averaged, mean annual surface air temperature (°C) |
|-----------------|---|
| | |
| 60Ma_140ppm | 2.042 |
| 60Ma_210ppm | 4.612 |
| 60Ma_280ppm | 15.43 |
| 60Ma_420ppm | 17.18 |
| 60Ma_560ppm | 18.36 |
| 60Ma_840ppm | 20.27 |
| 60Ma_1120ppm | 21.47 |
| 60Ma_2240ppm | 24.61 |
| 60Ma_4480ppm | 28.6 |
| | |
| 70Ma_280ppm | 5.737 |
| 70Ma_420ppm | 6.936 |
| 70Ma_560ppm | 16.63 |
| 70Ma_840ppm | 18.98 |
| 70Ma_1120ppm | 20.38 |
| 70Ma_2240ppm | 23.83 |
| 70Ma_4480ppm | 27.5 |
| | |
| 80Ma_140ppm | 7.369 |
| 80Ma_210ppm | 9.103 |
| 80Ma_280ppm | 15.41 |
| 80Ma_420ppm | 16.7 |
| 80Ma_560ppm | 18.43 |
| 80Ma_840ppm | 21.65 |
| 80Ma_1120ppm | 22.99 |
| 80Ma_2240ppm | 26.08 |
| 80Ma_4480ppm | 29.3 |
| | |
| 90Ma_70ppm | 1.581 |
| 90Ma_140ppm | 14.94 |
| 90Ma_210ppm | 16.24 |
| 90Ma_280ppm | 17.26 |
| 90Ma_420ppm | 18.6 |
| 90Ma_560ppm | 19.56 |
| 90Ma_840ppm | 20.93 |
| 90Ma_1120ppm | 22.36 |
| 90Ma_2240ppm | 25.84 |
| 90Ma_4480ppm | 29.52 |
| | |
| 100Ma_140ppm | 0.8807 |
| 100Ma_210ppm | 16.01 |
| 100Ma_280ppm | 16.86 |
| 100Ma_420ppm | 18.51 |
| 100Ma_560ppm | 19.41 |
| 100Ma_840ppm | 21.11 |
| 100Ma_1120ppm | 22.23 |
| 100Ma_2240ppm | 25.78 |
| 100Ma_4480ppm | 29.42 |
| | |

| Experiment name | Globally-averaged, mean annual surface air temperature (°C) |
|-----------------|---|
| | |
| 110Ma_140ppm | 3.921 |
| 110Ma_210ppm | 6.52 |
| 110Ma_280ppm | 14.1 |
| 110Ma_420ppm | 16.33 |
| 110Ma_560ppm | 17.77 |
| 110Ma_840ppm | 19.4 |
| 110Ma_1120ppm | 20.77 |
| 110Ma_2240ppm | 24.1 |
| 110Ma_4480ppm | 27.92 |
| | |
| 120Ma_140ppm | 2.022 |
| 120Ma_210ppm | 6.712 |
| 120Ma_280ppm | 14.26 |
| 120Ma_420ppm | 16.12 |
| 120Ma_560ppm | 17.33 |
| 120Ma_840ppm | 19.09 |
| 120Ma_1120ppm | 20.51 |
| 120Ma_2240ppm | 24.18 |
| 120Ma_4480ppm | 28.1 |
| | |
| 130Ma_560ppm | 8.58 |
| 130Ma_650ppm | 14.66 |
| 130Ma_740ppm | 15.81 |
| 130Ma_840ppm | 16.75 |
| 130Ma_1120ppm | 17.78 |
| 130Ma_2240ppm | 21.84 |
| 130Ma_4480ppm | 26.38 |
| 130Ma_8960ppm | 30.81 |
| | |
| 140Ma_560ppm | 6.4 |
| 140Ma_840ppm | 9.089 |
| 140Ma_930ppm | 9.787 |
| 140Ma_1020ppm | 15.98 |
| 140Ma_1120ppm | 16.6 |
| 140Ma_2240ppm | 20.38 |
| 140Ma_4480ppm | 24.5 |
| 140Ma_8960ppm | 29.45 |
| | |
| 150Ma_560ppm | 6.842 |
| 150Ma_840ppm | 8.407 |
| 150Ma_1120ppm | 11.23 |
| 150Ma_1210ppm | 15.06 |
| 150Ma_1305ppm | 16.79 |
| 150Ma_1400ppm | 17.32 |
| 150Ma_1680ppm | 18.39 |
| 150Ma_1960ppm | 19.75 |
| 150Ma_2240ppm | 20.37 |
| 150Ma_4480ppm | 24.35 |
| 150Ma_8960ppm | 28.87 |
| | |