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GSA Data Repository Item 2019203

## Data Repository

### Tables

DR1. New (U-Th)/He and fission track thermochronology data.

DR2. Coded thermal modeling results and timing of onset of cooling ( $t_{cool}$ ). This table shows all samples with data from > 1 thermochronometer. Modeling was attempted for all samples meeting this criteria. The coded results from the models in Figure DR2 are shown in the section of the table, 'Thermal Model Data', and include the range of acceptable modeled inflection points as constrained by the thermochronology data MAX  $t_{cool}$ , MEAN  $t_{cool}$ , and MIN  $t_{cool}$  derived from Figure DR2. If samples record a change in Neogene cooling rate *after* the change documented by the acceleration in cooling ( $t_{cool}$ ), we denote this with YES/NO/MAYBE (Y/N/M) in the column 'Neogene change in cooling rate'. The columns after this designate whether this change was an INCREASE/DECREASE (I/D) in column 'Neogene Cooling Increase or Decrease' and show the timing of this change in the columns 'Max onset change in cooling' and 'Min onset change in cooling'.

### Figure

DR1. A)  $t_{cool}$  vs. trench distance; B)  $t_{cool}$  vs. longitude; C) sample elevation vs.  $t_{cool}$

DR2. HeFTy inverse modeling results and example of coded outputs from HeFTy thermal models.

### Extended Methods

Three new samples were collected and analyzed using three low-temperature thermochronology dating techniques, (U-Th-Sm)/He thermochronology on apatite and zircon (AHe, ZHe) and apatite fission track thermochronology (AFT). We specifically targeted collecting of new samples from the hinterland and thrust belt domains as these regions were underrepresented south of 52°S. Samples (1-2 kg) were collected at outcrops exposed near sea level in Seno Skyring accessed by zodiac watercraft. Samples FPT17-68 and FPT17-59 were collected for ZHe, AFT, and AHe from the Jurassic Tobífera Formation in the hinterland domain. One granitic conglomerate clast, FPT17-45 was collected in the thrust belt domain from the Upper Cretaceous Escarpada Formation for AFT and AHe thermochronology.

We integrated our new data with all existing thermochronology samples from the ZHe, AFT, and AHe systems published in the literature between 47° and 54° S (Figure 1; Thomson et al., 2001; 2010; Haschke et al., 2006; Fosdick et al., 2013; Guillaume et al., 2013; Georgieva et al., 2016; Christeleit et al., 2017). Zircon fission track (ZFT) data were omitted from this compendium as published ZFT dates have Late Cretaceous - Paleocene dates reflecting and earlier cooling history than the target of this study. All samples (55) with data from >2 thermochronometric systems (ZHe, AFT, and AHe) were chosen for thermal modeling (Table DR2).

### *(U-Th-Sm)/He thermochronology*

Apatite and zircon grains were picked for (U-Th-Sm)/He thermochronology at the Basin Analysis and Helium Thermochronology Lab at the University of Connecticut. Whole euhedral grains were selected for clarity and size using a Leica M165 binocular microscope. Single grain aliquots were secured in Nb tubes. For each sample, four grains were analyzed for ZHe thermochronology and three grains were analyzed for AHe thermochronology at the University of Colorado Thermochronology Research and Instrumentation Lab (TRaIL). Nb-packed aliquots were loaded in an ASI Alphachron™ to extract and measure He gas. A 25W diode laser heated the aliquot to ca. 800-110°C for 5 to 10 minutes to extract 4He gas. A spike of 13 ncc  $^3\text{He}$  was used to spike extracted  $^4\text{He}$  gas and analyzed on a Balzers PrismaPlus QME 220 quadrupole mass spectrometer. Apatite grains and Nb tubes are dissolved in  $\text{HNO}_3$  spiked with a  $^{235}\text{U}$ ,  $^{230}\text{Th}$ ,  $^{145}\text{Nd}$  tracer. Zircon grains and niobium tubes are dissolved in a series of HF solutions - the first spiked with a  $^{235}\text{U}$ ,  $^{230}\text{Th}$ ,  $^{145}\text{Nd}$  tracer - and a final  $\text{HNO}_3$ :HF mixture. Both apatite and zircon samples solutions are analyzed for U, Th and Sm on a Thermo Element 2 magnetic sector ICP-MS. Results are reported in Table DR1.

### *Apatite Fission Track Thermochronology*

Apatite fission track samples were prepared following the methods of Donelick et al. (2005) by mounting apatite grains in epoxy, polishing the mounts, and etching samples with 5.5M nitric acid for 20 seconds at 21°C. A mica sheet was affixed to the epoxy mount and irradiated at Oregon State University. Upon return from the reactor, the mica sheets were etched in 49% HF for ~ 15 minutes at 23°C. Apatite fission track thermochronology was conducted at the University of Arizona Fission Track Laboratory using the external detector method using 1600x magnification. Both track density ratios and average track etch pit diameter ( $D_{\text{par}}$ ) was recorded for between 24 and 30 grains per sample. Results are reported in Table DR1.

### *Inverse Thermal Modeling*

Results from ZHe, AFT, and AHe from new and existing data sets were modeled using HeFTy v.1.9.3 (Ketcham, 2005). For ZHe samples (using all grains available from our data and published data), we entered single grain data individually applying the helium diffusion model from Guenthner et al., (2013) and the ejection correction for alpha calculation. For a given sample, all available single-grain AHe dates were modeled together using the radiation damage accumulation and annealing model from Flowers et al., (2013) and the ejection correction for alpha calculation. Raw counting ratios were entered for AFT samples along with  $D_{\text{par}}$  measurements, when available. We used the apatite fission track annealing model from Ketcham et al., (2007).

Inverse model constraints for the 100 to 0 Ma time period were left purposely broad to explore all cooling histories, with constraints only at model initiation (100 and 300 °C between 80 and 50 Ma) and modern surface temperatures of  $10 \pm 5$  °C. These constraints were designed to be consistent with four existing thermal models generated by Fosdick et al. (2013) using the inverse modeling tool in HeFTy from samples in the hinterland and internal thrust belt domains between 51° and 51.5° S. Although the model parameters described above explore only monotonic t-T paths, preliminary use of alternative model parameters that allow reheating failed to identify evidence of non-monotonic t-T paths. Each model was required to explore possible t-T paths until 100 paths with a goodness of fit  $>0.5$  were identified.

Models were initially run with analytical error from ZHe, AFT, and AHe analyses; however, in 22/54 cases, models with these parameters failed to yield any t-T paths with a goodness of fit > 0.5 after exploring >100,000 t-T paths. In this case, it is reasonable to assume that there is additional error not due to analytical uncertainty, but from variations in annealing kinetics not captured by the annealing models (Flowers et al., 2009; Guenthner et al., 2013). For these samples, we increased individual grain error to 20% for the ZHe and/or AHe systems, altering no other input parameters, and reran the model. Twelve of the 22 samples with 20% ZHe and/or AHe grain error in HeFTy identified 100 t-T paths with a goodness of fit > 0.5 were recorded, and we designated these results as lower confidence (Table DR2). For ten samples, HeFTy's exploration of >100,000 possible time temperature paths failed to identify any good fit paths. We attribute this to a non-geologic discrepancy between thermochronometric data, (e.g. helium implantation, mineral inclusions, U-Th rich grain boundary phases; Peyton et al., 2012; Murray et al., 2014) and these samples were excluded from further analysis.

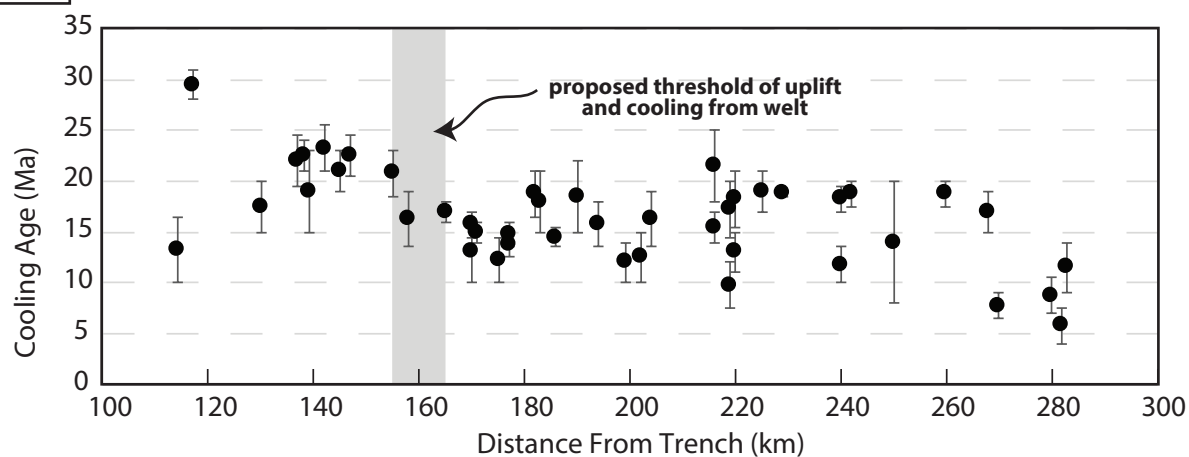
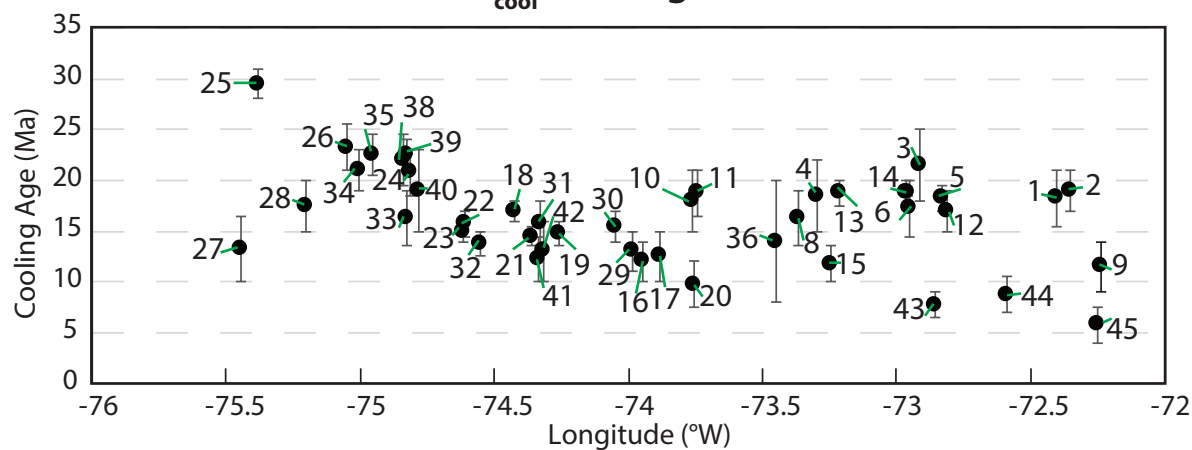
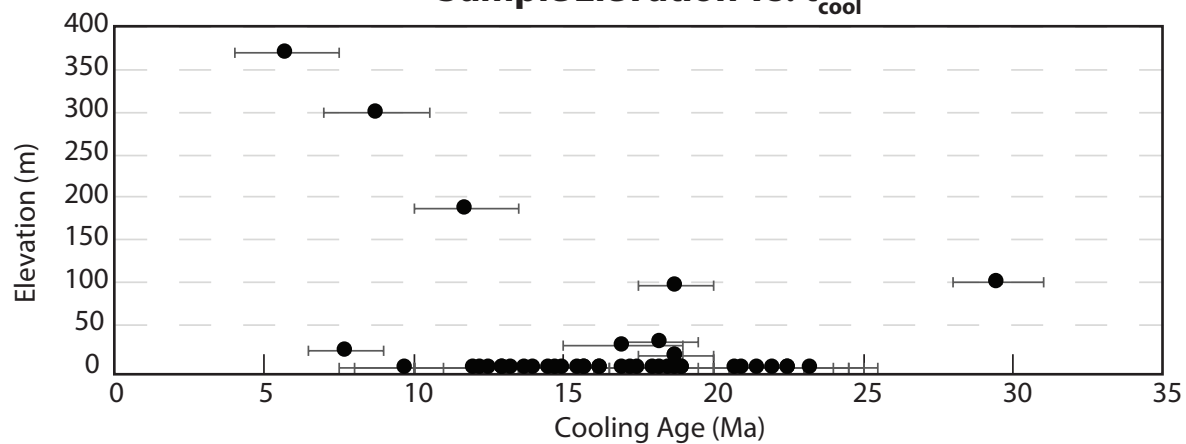
We coded the temporal range of the onset of Neogene cooling by identifying two points (Table DR2; Figure DR2): 1) the time and temperature of the inflection point where the earliest possible cooling was required by the suite of 100 good fit paths ( $t_{cool}$ ); and 2) the latest possible time where cooling was required by the model for the same temperature identified in point one. The mean of these two constraints is plotted in Figure 2B.

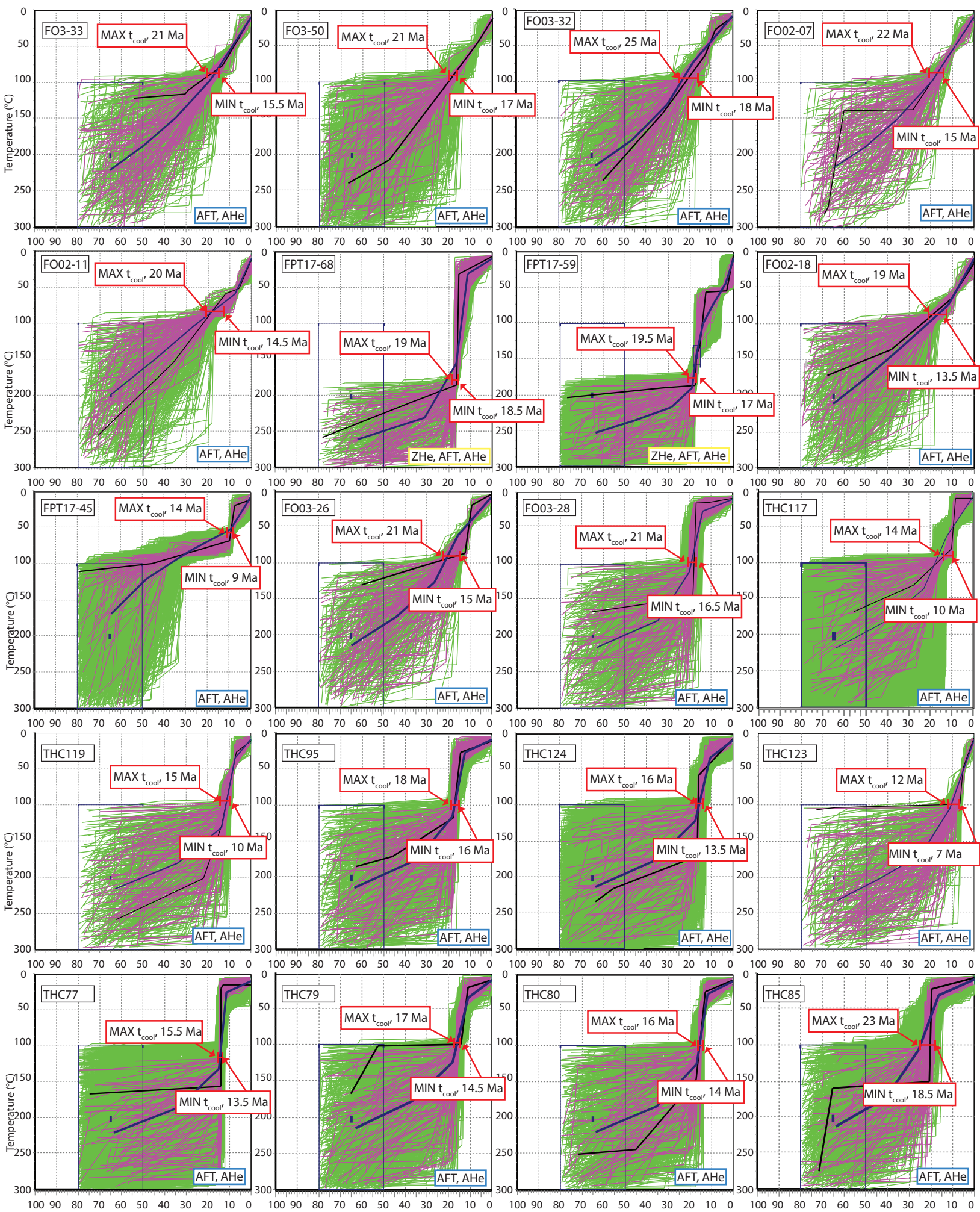
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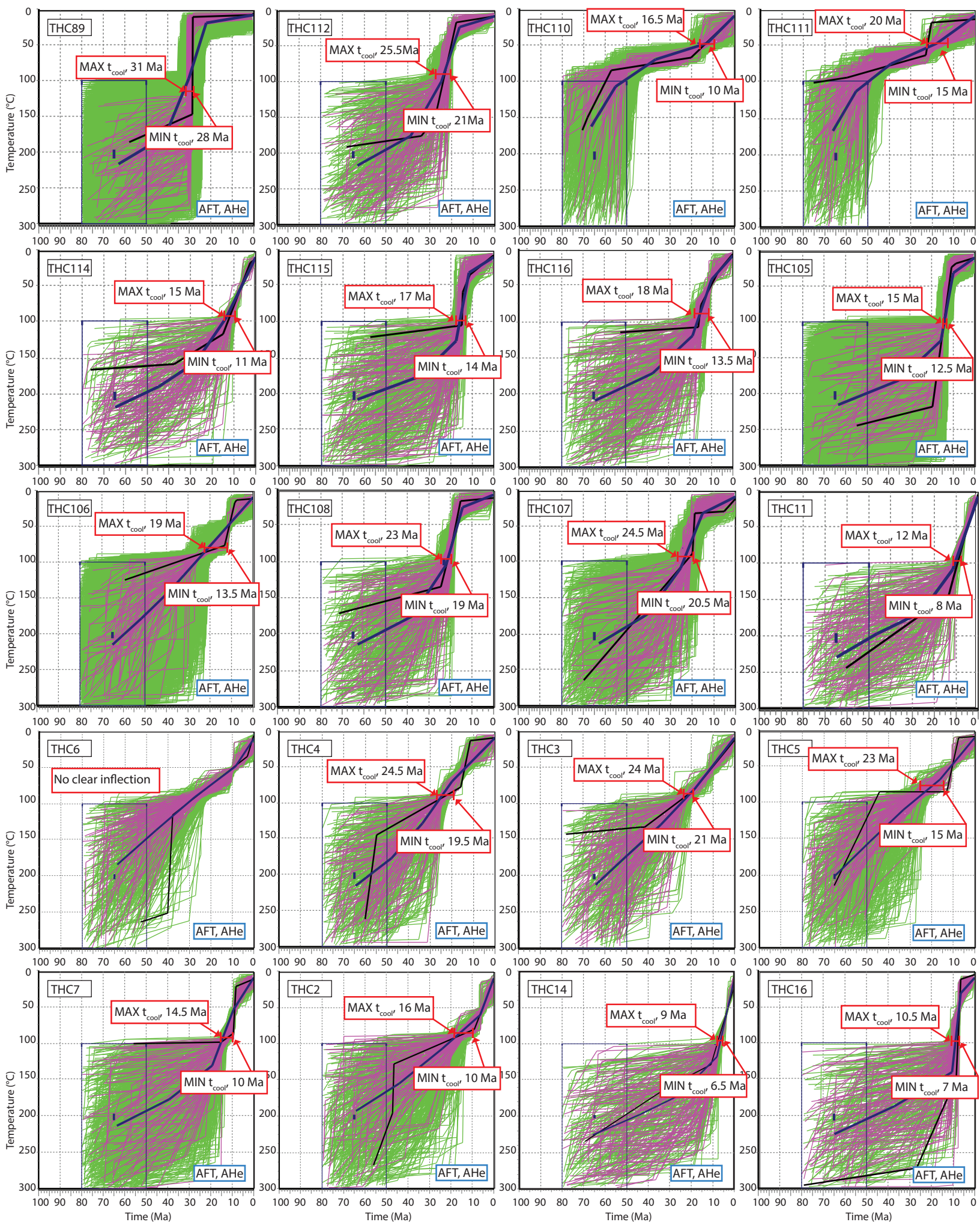
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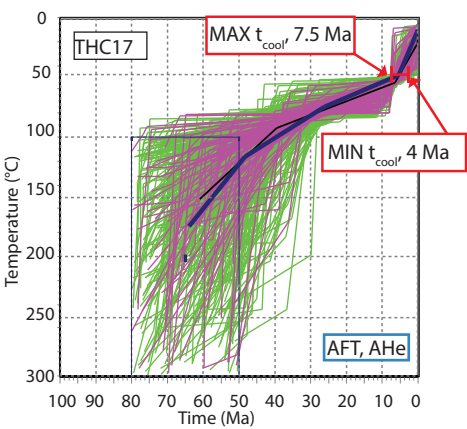


**A** **$t_{cool}$  vs. Trench Distance****B** **$t_{cool}$  vs. Longitude****C****Sample Elevation vs.  $t_{cool}$** 



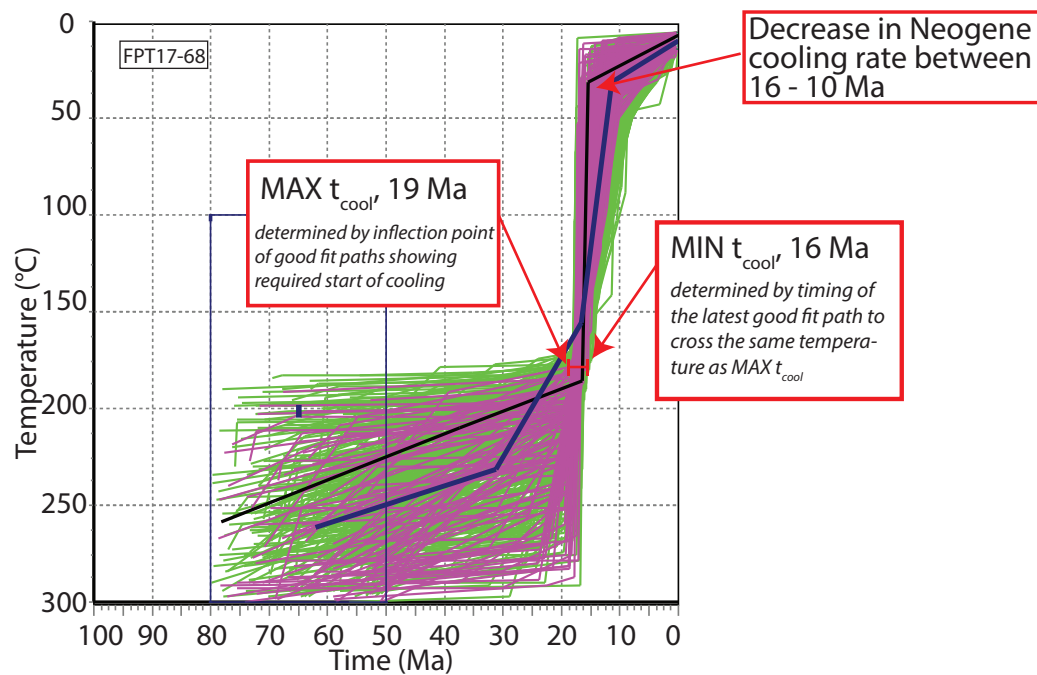






## DR 2. HeFTy Inverse Modeling Results

Example of data coding from HeFTy thermal model output



- Best Fit t-T path
- Good Fit t-T path
- Acceptable Fit t-T path
- Mean t-T path
- Model Constraints